

Deep-Water Volcaniclastic Fans: What Can We Learn from the Past?*

Andrea Di Capua¹, Gianluca Gropelli³, and Federica Barilaro²

Search and Discovery Article #51487 (2018)**

Posted June 18, 2018

*Adapted from oral presentation given at AAPG Asia Pacific Region, Geosciences Technology Workshop, Influence of Volcanism and Associated Magmatic Processes on Petroleum Systems, Oamaru, New Zealand, March 14-16, 2017

**Datapages © 2018 Serial rights given by author. For all other rights contact author directly.

¹Italian National Research Council – Institute for the Dynamics of Environmental Processes (CNR –IDPA), Italy; University of Insubria – Como, Italy (andrea.dicapua@unimib.it)

²ETHZ – Geological Institute, Switzerland; University of Insubria – Como, Italy

³Italian National Research Council – Institute for the Dynamics of Environmental Processes (CNR –IDPA), Italy

Abstract

Volcanism is a surface process widely diffuse in space and time on the Earth, whose episodically, short-lived activation deeply impacts both proximal and distal environments. Despite of the increasing knowledge on its capability to produce, transport, and deposit detritus, volcanism is often not considered as an Earth surface process that might have shaped the sedimentary basins in the past, overall when volcanic edifices in source-to-sink systems are not preserved.

We present a resume of the state of art of the investigations that we are carrying out on the foreland and foredeep basins developed during the Oligocene times at the boundary between the European and the Adriatic plates. In these basins, wide turbidite fans have been fed by huge volumes of primary and secondary dispersal discharged by volcanic centers located from the SE France to the embryo of the Alpine Belt. Such fans, now structured in the Helvetic Nappes of the Western Alps (Taveyenne Sandstones) and the Sub-Liguri Nappes of the Northern Apennines (Val d'Aveto Formation), preserve peculiar features that have been investigated in terms of geometries, sedimentary structures, petrography, mineralogy, and chemistry, through the combination of different techniques (field clast counts, image analyses, sandstone petrography, XRD, SEM-EDS). The result is a multidisciplinary sketch of the capability of volcanism to strongly influence the sedimentation of distal (tens to hundreds of kilometers), deep-water basins independently from the tectono-climatic boundary conditions that either favor or not the production, transport, and deposition of detritus. As a consequence, this sketch highlights the role of volcanism in shaping sedimentary basins, as well as its importance as a “complete sedimentary factory”, able to instantaneously deliver large amounts of sediments far away from the main source. This study also underlines the importance of the combination of the sedimentological and the volcanological approach to the identification of primary and secondary volcanic deposits, as their different physical, chemical, and mineralogical properties control primary and diagenetic features that largely influence fabric type textures and consequent rock petrophysical properties. All these findings must be considered the key break-through for a comprehensive evaluation of volcaniclastic reservoir properties.

Introduction

The study of the influence of volcanism on the deep-water sedimentary systems is one of the most intellectually stimulating frontiers in sedimentology and petroleum geology. The main challenge on the study of volcanoclastic sequences concerns the identification of the influence that volcanic events had on the source-to-sink. During a volcanic event, in fact, huge volumes of volcanic particles are produced and instantaneously transferred across the environment for tens to thousands of kilometers, being either directly discharged into basins or temporarily dispersed in transfer zones (Smith, 1991; Manville et al., 2009). The capability of volcanoes to induce the accumulation of thick sequences despite the tectono-climatic conditions in the surrounding environments is regularly either underestimated or not considered in sedimentary models (e.g., Catuneanu et al., 2011), and a systematic characterization of volcanic/volcanoclastic deposits is still missing. This is particularly evident in approaching ancient volcanoclastic sequences of turbidite fans that lack a clear stratigraphic correlation with volcanic source areas (Di Capua and Groppelli, 2014).

In petroleum geoscience, volcanoclastic turbidite fans have been recognized as potential hydrocarbons reservoirs, but the presence of high portions of mechanically and mineralogically unstable components makes their primary and secondary features difficult to understand for reservoir quality predictions (Seeman and Scherer, 1984). Thus, an accurate definition of the volcano-sedimentary features in deep-water environments could represent a break-through for the prediction and evaluation of reservoir quality in volcanoclastics.

Through the development of a multidisciplinary standardized workflow, based on the recognition of the main sedimentological, petrographic and petrophysical features, integrated with a combination of high-resolution techniques approach (SEM-EDS and XRD), this study carries out an innovative on volcanoclastic deposits in deep-sea water systems, and their correlation with their generating events and post-depositional history. The Northern Apennines foredeep basin (Northern Apennines - Italy) and the Northern Alpine foreland basin (Western Alps - France and Switzerland) have been the object of this study. For the first time, a classification of the deep-water deposits has been proposed and three main categories have been recognized. These include: 1) non-, 2) syn-, and 3) post-volcanic deposits. In non-volcanic deposits, volcanic detritus may be present as particles eroded from ancient volcanic sequences, so volcanism had no influence on sedimentation. In syn-volcanic deposits, volcanic detritus is directly provided during or soon after a volcanic event. In post-volcanic deposits, volcanic detritus is supplied by the resedimentation of dispersed, unconsolidated volcanic detritus, so the progressive lithification of onland volcanoclastic deposits tends to proportionally decrease the amounts of volcanic particles in the deposits. For their peculiar features, the presented work mainly focuses on a detailed description of syn-volcanic deposits.

Syn-Volcanic Deposits as Potential Reservoir Rocks

Syn-volcanic deposits are primary and secondary dispersals settled during or soon after a main eruption event. Particles generally move and settle under extreme physical conditions (e.g., hot temperatures, high pressures). This implies strong lateral and vertical variations of the nature of their deposits. In addition, their interaction with the surrounding environmental conditions strongly affects their physical properties. A further distinction between syn-eruptive and post-eruptive deposits can be made on the base of the influence of volcanism on sedimentation.

Syn-Eruptive, Syn-Volcanic Deposits

Syn-eruptive deposits are all the deposits directly settled during eruptive episodes by pyroclastic density currents (PDCs). They are of great importance for petroleum systems, as they are featured by great thicknesses, strong lateral variations, and primary and secondary porosity controlled by the generating volcanic process, the conditions of the depositional environment (e.g., presence of water, presence of paleosurfaces), and their post-depositional history. PDCs, in fact, are hot flows that mix with cold water entering a subaqueous basin. The physical behavior of such flows determines their modification and eventually disaggregation during their underwater motion before settling.

If PDCs are disaggregated, the resulted deposits consist of a thick, well sorted basal deposit, enriched in angular volcanic particles, and a thinner mud cap of a mixed volcanic/non-volcanic composition. This body anatomy is common both in modern and ancient settings (e.g., Taveyanne Sandstones – Di Capua and Gropelli, 2016a). Even though such types of deposits are alike to the sandy debris flow deposits of Shanmugam (2002), petrographic analyses reveal that glass can develop in the intergranular spaces, decreasing the primary porosity and permeability. Secondary porosity and permeability is the result of glass devitrification and dissolution.

If PDCs are not disaggregated, the resulted deposits show primary volcanic textures (e.g., porphyric), hydrothermal modifications, with the growth of secondary mineralogical phases (e.g., pyrite), and secondary fractures generated by the fluidization of underlying sediments. Such fractures are the most prominent porosity features (Di Capua et al., 2016). In addition, glassy-welded eutaxitic texture might also develop in-between close accidental clasts, as the result of increasing internal pressures and shear along their boundaries (Di Capua and Gropelli, 2016b).

Post-Eruptive, Syn-Volcanic Deposits

Post-eruptive, syn-volcanic deposits derive from the rapid resedimentation of unwelded volcanic deposits soon after volcanic events. They are settled by flash flood events, such as lahars, whose motion is controlled by the kinetic interaction of particles. At the microscale, they are characterized by volcanic particles that plastically embed accidental detritus to form a pseudomatrix (sensum Dickinson, 1970). Thus, porosity and pore interconnection are generally low, but might increase thanks to the devitrification of glass particles and dissolution of minerals.

Conclusion

The presented work and classification show the importance of a multidisciplinary and multiscale approach in the identification of the temporal relationship between volcanism and sedimentation, considered the main factor controlling the physical, minero-chemical, diagenetic, and petrophysical properties of volcanoclastic sequences. This innovative approach provides a comprehensive evaluation of the reservoir properties of volcanoclastic sequences.

Selected References

Allen, P.A., 2008, From Landscapes into Geological History: *Nature*, v. 451, p. 274-276.

- Catuneanu, O., 2002, Sequence Stratigraphy of Clastic Systems: Concepts, Merits, and Pitfalls: *Journal of African Earth Sciences*, v. 35, p. 1-43.
- Catuneanu, O., W.E. Galloway, C.G.S.C. Kendall, A.D. Miall, H.W. Posamentier, A. Strasser, and M.E. Tucker, 2011, Sequence Stratigraphy: Methodology and Nomenclature: *Newsletters on Stratigraphy*, v. 44/3, p. 173-245.
- Critelli, S., and R.V. Ingersoll, 1995, Interpretation of Neovolcanic Versus Palaeovolcanic Sand Grains: An Example from Miocene Deep-Marine Sandstone of the Topanga Group (Southern California): *Sedimentology*, v. 42, p. 783-804.
- Di Capua, A., and G. Groppelli, 2014, What Does “Volcanoclastic” Mean in a Distal Sedimentary Succession?, *in* R. Rocha, J. Pais, J.C. Kullberg, and S. Finney (eds.), *At the Cutting Edge of Stratigraphy: STRATI 2013*, Springer International Publishing Switzerland.
- Di Capua, A., and G. Groppelli, 2016a, Application of Actualistic Models to Unravel Primary Volcanic Control on Sedimentation (Taveyanne Sandstones, Oligocene Northalpine Foreland Basin): *Sedimentary Geology*, v. 336, p. 147-160.
- Di Capua, A., and G. Groppelli, 2016b, Emplacement of Pyroclastic Density Currents (PDCs) in a Deep-Sea Water Environment: The Val d’Aveto Formation Case (Northern Apennines, Italy): *Journal of Volcanology and Geothermal Research*, v. 328, p. 1-8.
- Di Capua, A., G. Vezzoli, and G. Groppelli, 2016, Climatic, Tectonic and Volcanic Controls of the Sediment Supply to an Oligocene Foredeep Basin: The Val d’Aveto Formation (Northern Italian Apennines): *Sedimentary Geology*, v. 332, p. 68-84.
- Dickinson, W.R., 1970, Interpreting Detrital Modes of Graywacke and Arkose: *Journal of Sedimentary Petrology*, v. 40, p. 695-707.
- Douillet, G.A., D.A. Pacheco, U. Kueppers, J. Letort, È. Tsang-Hin-Sun, J. Bustillos, M. Hall, P. Ramòn, and D.B. Dingwell, 2013, Dune Bedforms Produced by Dilute Pyroclastic Density Currents from the August 2006 Eruption of Tungurahua Volcano, Ecuador: *Bulletin of Volcanology*, v. 75, p. 1-20.
- Garzanti, E., and M.G. Malusà, 2008, The Oligocene Alps: Domal Unroofing and Drainage Development During Early Orogenic Growth: *Earth and Planetary Science Letters*, v. 268, p. 487-500.
- Graham, U.M., and H. Ohmoto, 1994, Experimental Study of Formation of Mechanisms of Hydrothermal Pyrite: *Geochimica et Cosmochimica Acta*, v. 58/10, p. 2187-2202.
- Manville, V., K. Nemeth, and K. Kano, 2009, Source to Sink: A Review of Three Decades of Progress in the Understanding of Volcaniclastic Processes, Deposits, and Hazards: *Sedimentary Geology*, v. 220, p. 136-161.

- Manville, V., E.H. Newton, and J.D.L. White, 2005, Fluvial Response to Volcanism: Resedimentation of the 1800a Taupo Ignimbrite Eruption in the Rangitaiki River Catchment, North Island, New Zealand: *Geomorphology*, v. 65, p. 49-70.
- Ogniben, L., M. Parotto, and A. Pratlun, 1975, Structural Model of Italy: Maps and Explanatory Notes: Consiglio Nazionale delle Ricerche 1, 502 p.
- Ruffini, R., M.A. Cosca, A. d'Atri, J.C. Hunziker, and R. Polino, 1995, The Volcanic Supply of the Taveyanne Turbidites (Savoie, France): A Riddle for the Tertiary Alpine Volcanism, *in* R. Polino and R. Sacchi (eds.), *Rapporti Alpi-Appennino e Guide alle Escursioni*: Accademia Nazionale delle Scienze detta dei XL, p. 359-376.
- Seemann, U., and M. Scherer, 1984, Volcaniclastics as Potential Hydrocarbon Reservoirs: *Clay Minerals*, v. 9, p. 457-470.
- Shanmugam, G., 2002, Ten Turbidite Myths: *Earth-Science Review*, v. 58, p. 311-341.
- Smith, G.A., 1991, Facies Sequences and Geometries in Continental Volcaniclastic Sediments, *in* R.V. Fisher and G.A. Smith (eds.), *Sedimentation in Volcanic Settings*: SEPM Special Publication 5, p. 109-121.
- Trofimovs, J., R.S.J. Sparks, and P.J. Talling, 2008, Anatomy of a Submarine Pyroclastic Flow and Associated Turbidity Current: July 2003 Dome Collapse, Soufrière Hills, Volcano, Montserrat, West Indies: *Sedimentology*, v. 55, p. 617-634.
- von Aulock, F.W., F.B. Wadsworth, Y. Lavallé, and J. Vasseur, 2014, Sintering of Glass in Hydrous Atmospheres and Its Implication for Welding of Volcanic Deposits: AGU Fall Meeting 2014.

Andrea Di Capua^{1,2},
Gianluca Groppelli¹,
Federica Barilaro^{2,3}
(andrea.dicapua@idpa.cnr.it)

1. Italian National Research Council – Institute for the Dynamics of Environmental Processes (CNR – IDPA)
2. University of Insubria – Como (Italy)
3. ETHZ – Geological Institute

Deep-water volcanoclastic fans: what can we learn from the past?



Volcaniclastics as potential hydrocarbon reservoirs

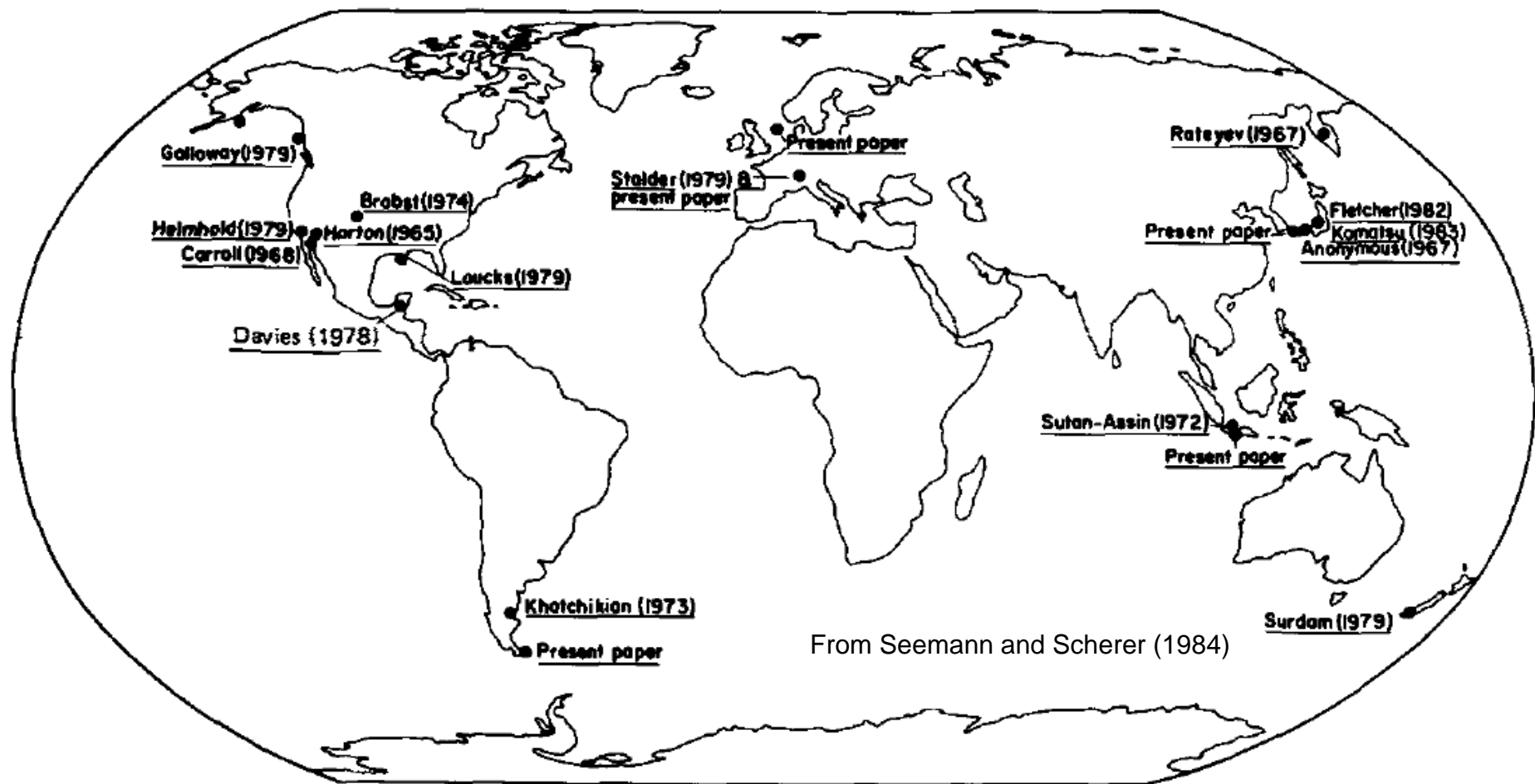
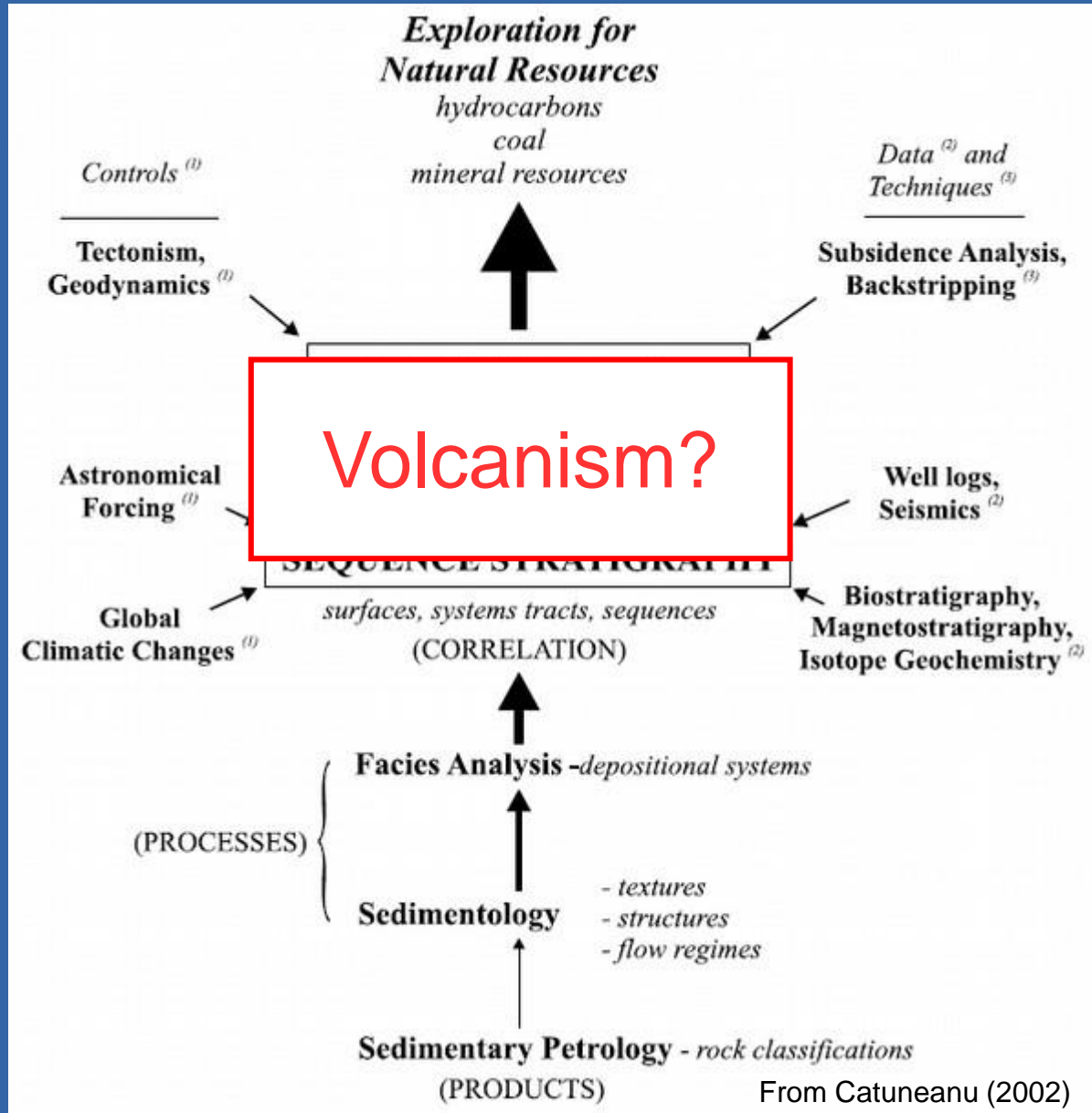
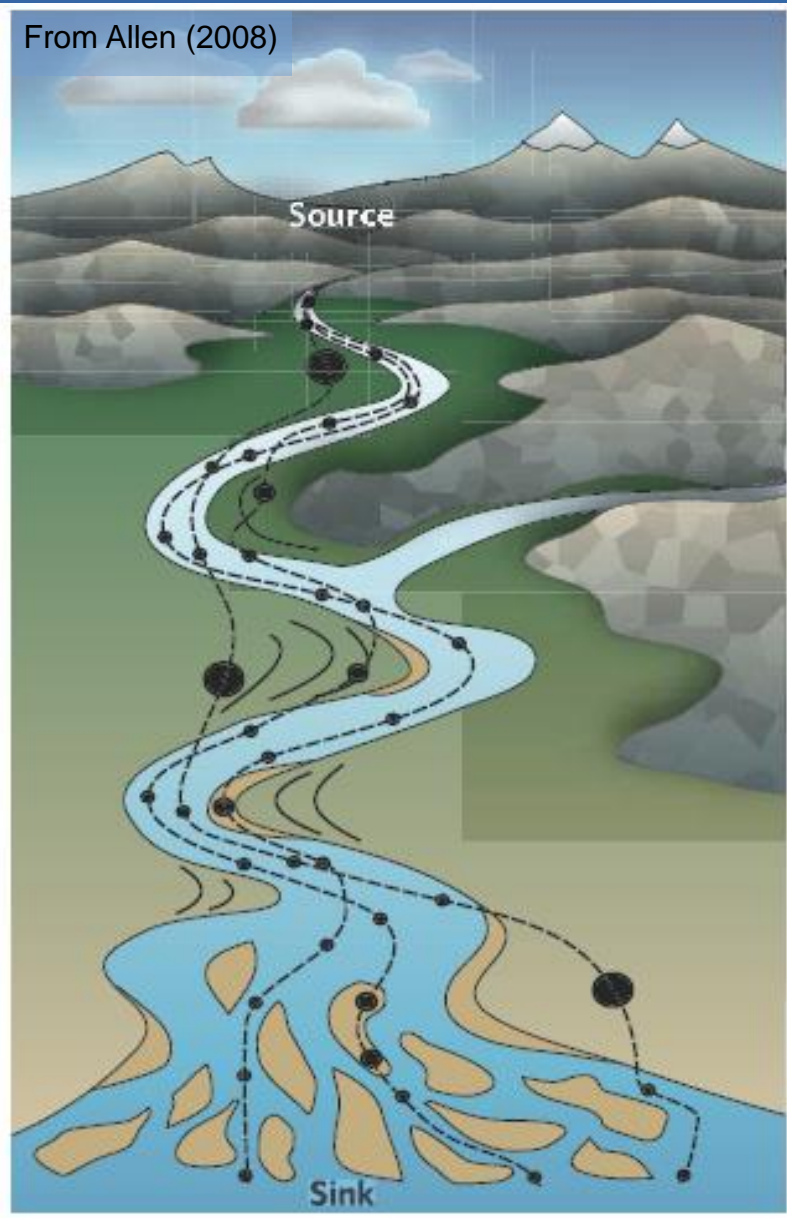


FIG. 1. Areas in which volcanics and/or volcaniclastics have been described in connection with hydrocarbon exploration or production. For the sake of clarity, only the first-named author of each reference is indicated.

Clastic Sedimentation

From Allen (2008)



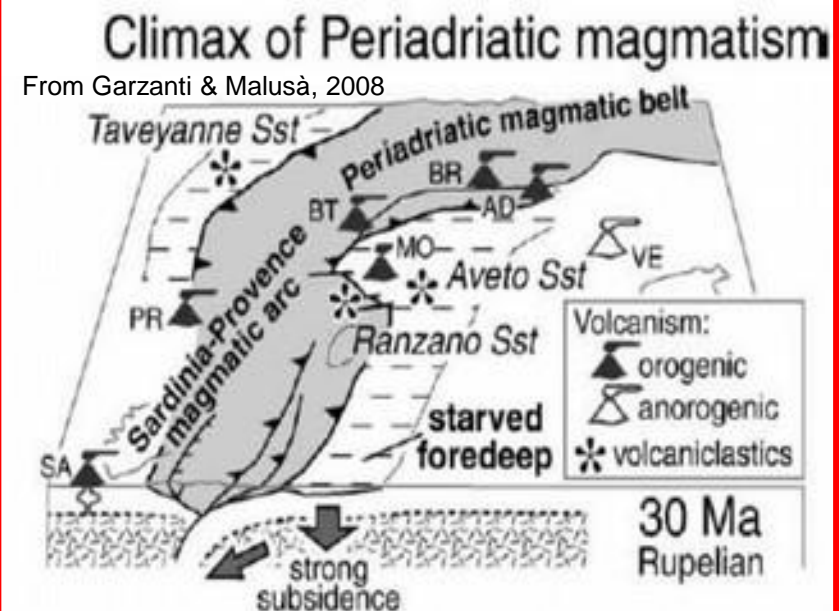
Geological setting: Oligocene volcanoclastic deposits around the Alps

From Ogniben et al., 1975

Taveyanne Sandstones

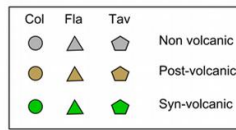
Milan

Val d'Aveto Formation

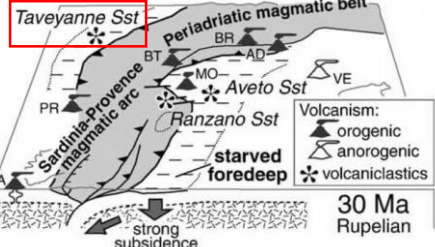


Taveyanne Sandstones (foreland basin)

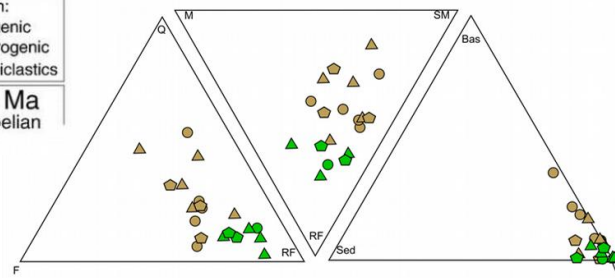
From Di Capua & Groppelli (2016a)



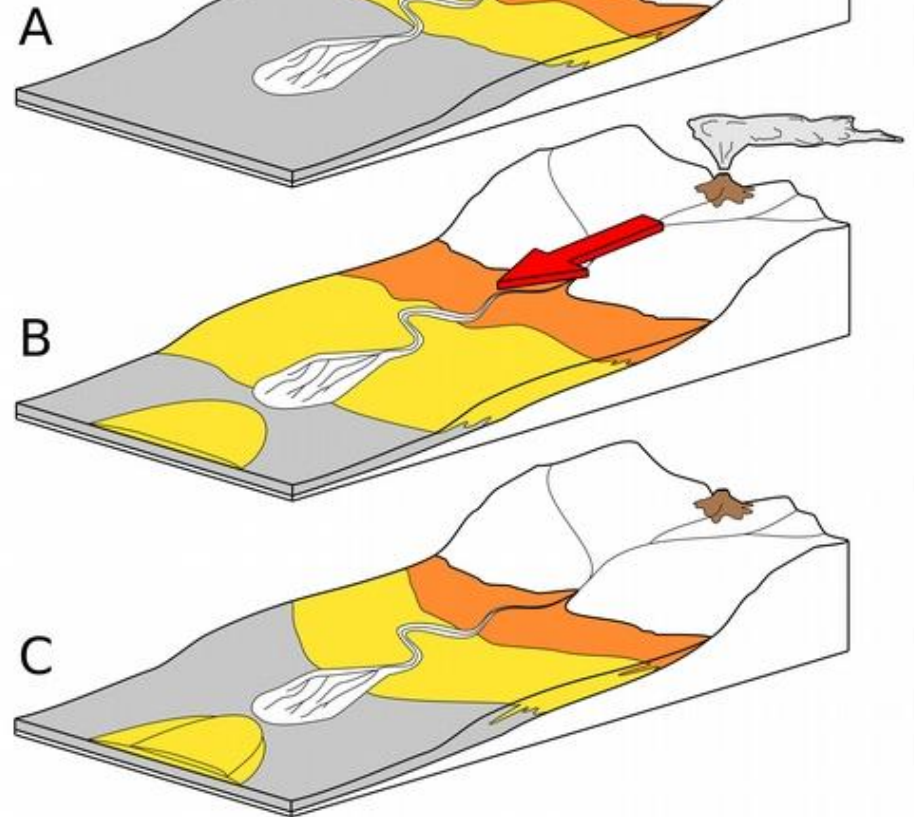
Climax of Periadriatic magmatism



From Garzanti and Malusà (2008)



"Actualistic" sedimentary model

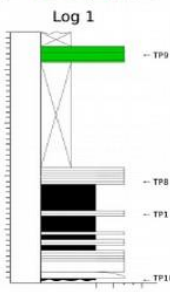
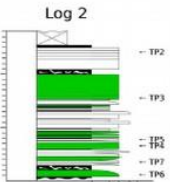


From Di Capua & Groppelli (2016a)

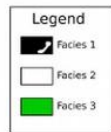
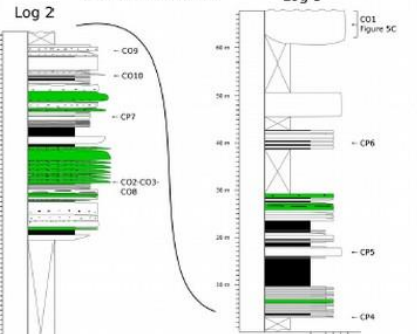
From Di Capua & Groppelli (2016a)



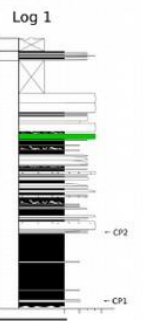
B) Taveyanne



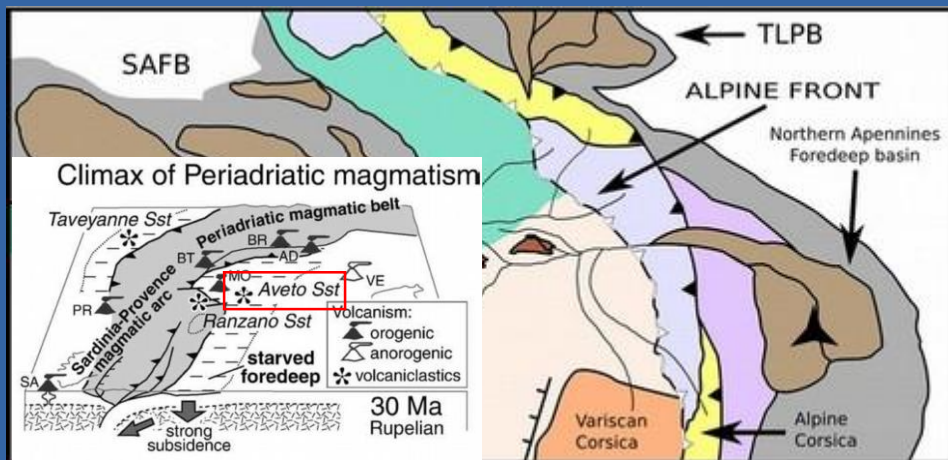
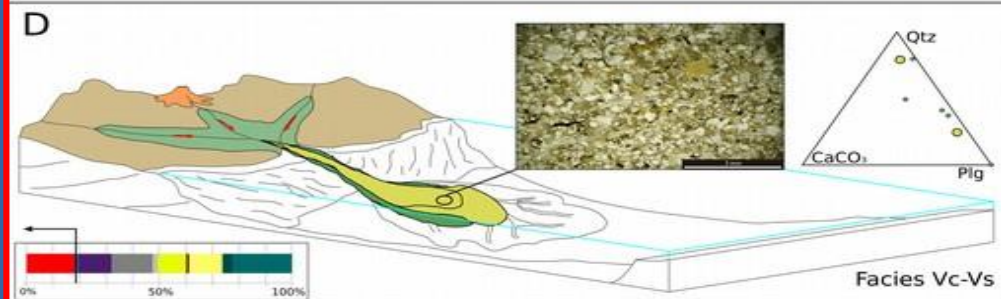
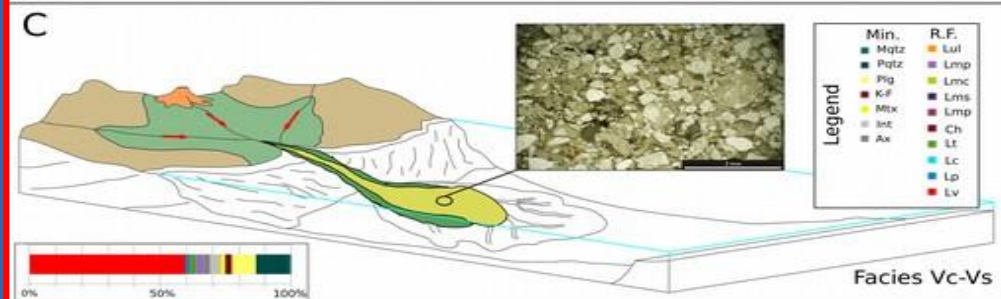
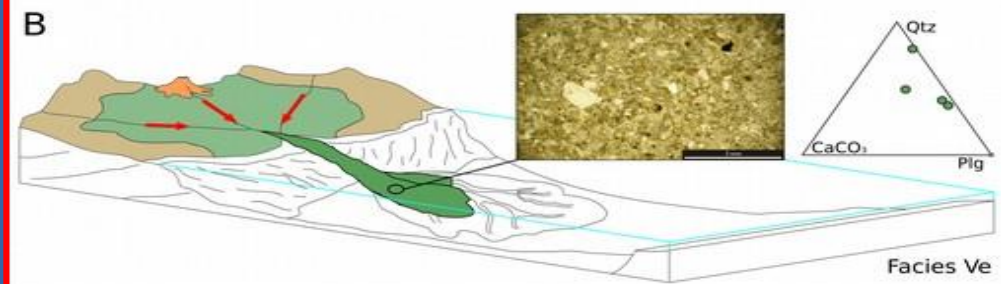
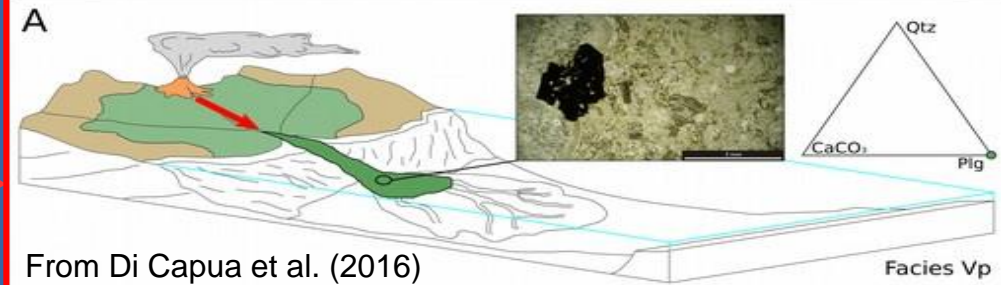
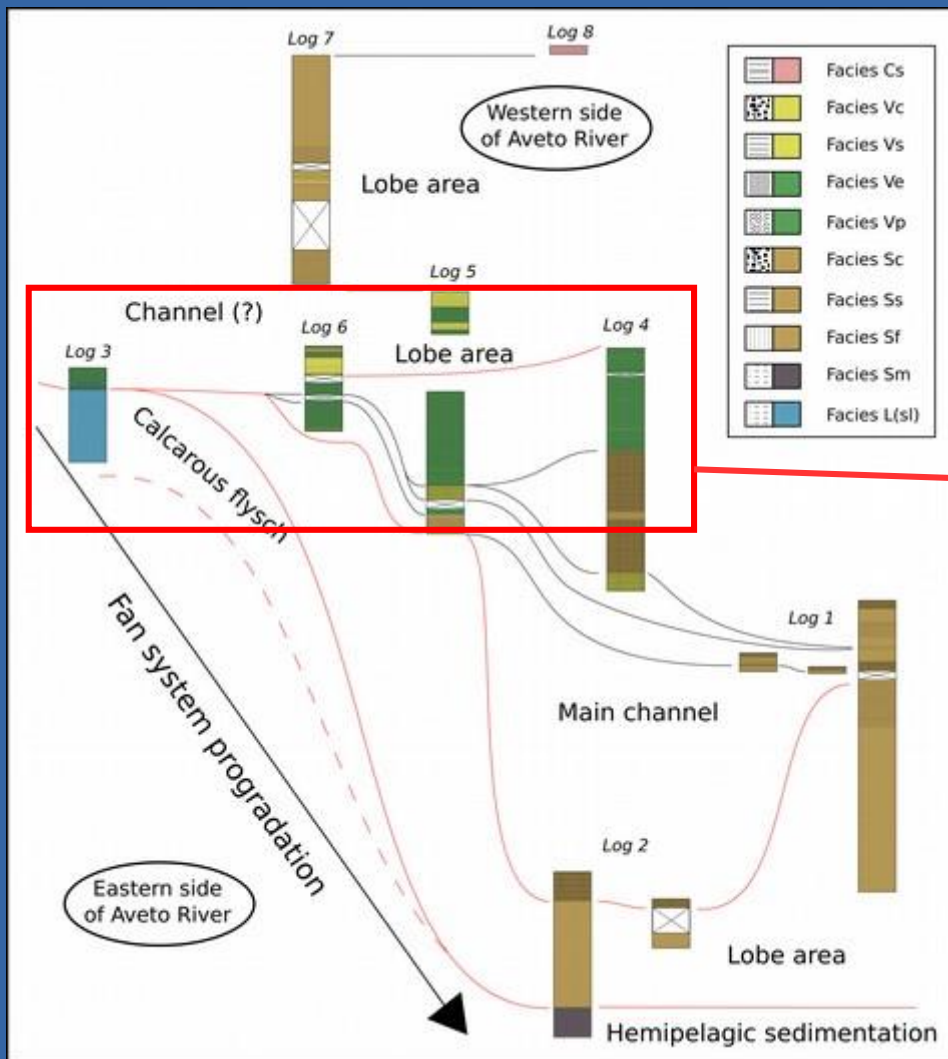
Col de l'Oulette



From Di Capua & Groppelli (2016a)



Val d'Aveto Formation (foredeep basin)



Approach and methodology

Multidisciplinary standardized workflow based on:

.Intensive fieldwork activities (geological mapping, structural analyses, log measurements and correlations, facies analyses)

.Laboratory activities (petrography, sandstone point-counts, X-ray powder diffraction analyses, SEM-EDS, image analyses, effective porosity analysis, paleomagnetism)

Classification of clastic deposits

CATEGORIES	DEFINITION	MAIN FEATURES	EFFECTIVE POROSITY
Non-volcanic deposits	Volcanic detritus may be present as particles eroded from ancient volcanic sequences, so volcanism had no influence on sedimentation	Lack of pseudomatrix; volcanoclastic detritus (if present) = palaeovolcanic fragments (sensu Critelli & Ingersoll, 1995)	From 1,1 to 8,8% Function of hydraulic sorting
Syn-volcanic deposits <ul style="list-style-type: none"> • Syn-eruptive • Post-eruptive 	Volcanic detritus is directly provided during or soon after a volcanic event	Porphyritic to eutaxitic textures; abundance of pseudomatrix; volcanoclastic detritus = neovolcanic fragments (sensu Critelli & Ingersoll, 1995)	From 1,7 to 11,6 % Function of volcanic event
Post-volcanic deposits	Volcanic detritus is supplied by the resedimentation of dispersed, unconsolidated volcanic detritus, so the progressive lithification of onland volcanoclastic deposits tends to proportionally decrease the amounts of volcanic particles in the deposits	Lack of pseudomatrix; volcanoclastic detritus = neovolcanic + palaeovolcanic fragments (sensu Critelli & Ingersoll, 1995)	From 1,8 to 11,4% Function of hydraulic sorting

Based on the temporal relationship between volcanism and sedimentation

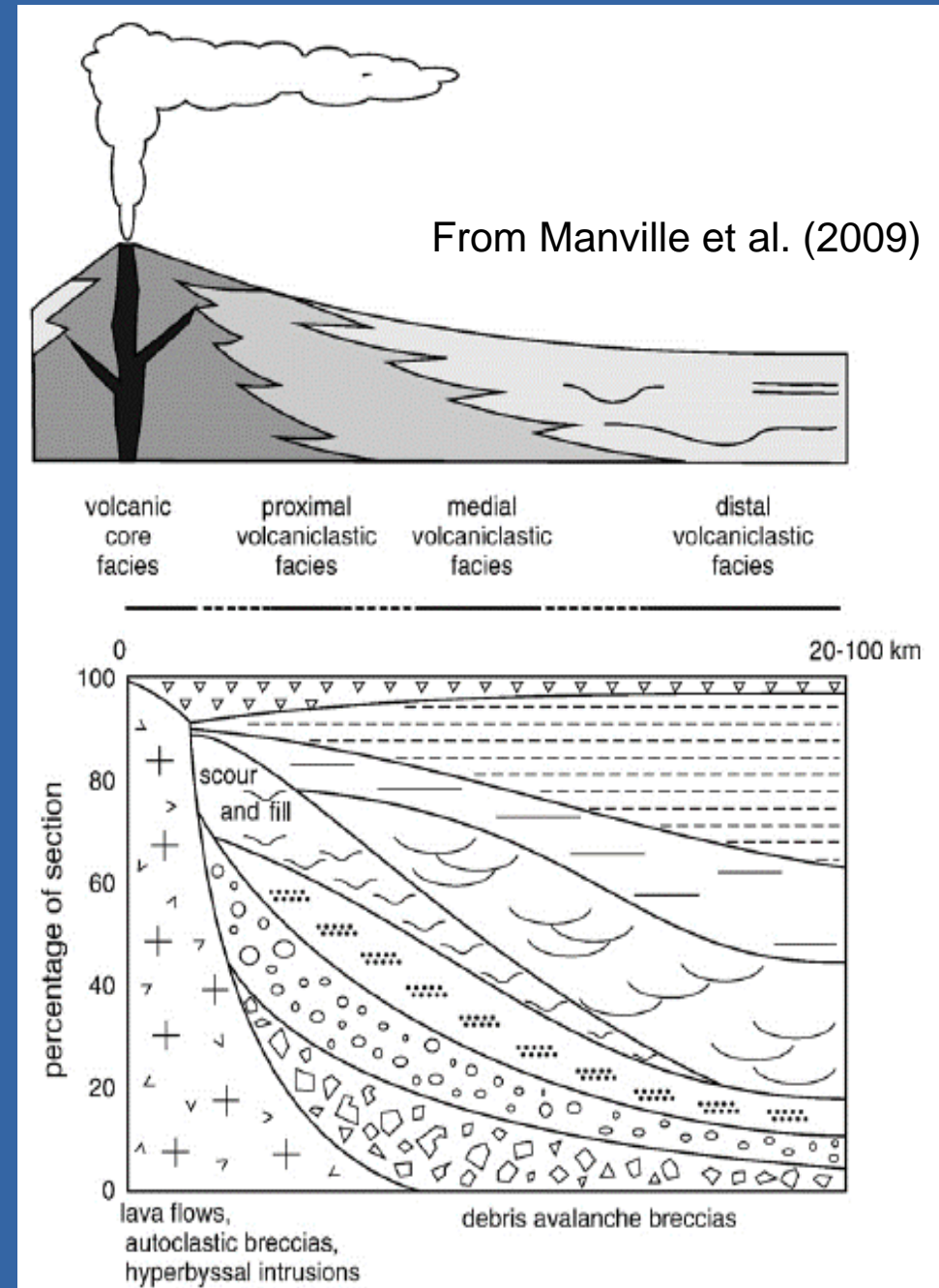
Syn-volcanic deposits

Features varying on the base of: 1) the type of explosive volcanic event and 2) the environmental conditions

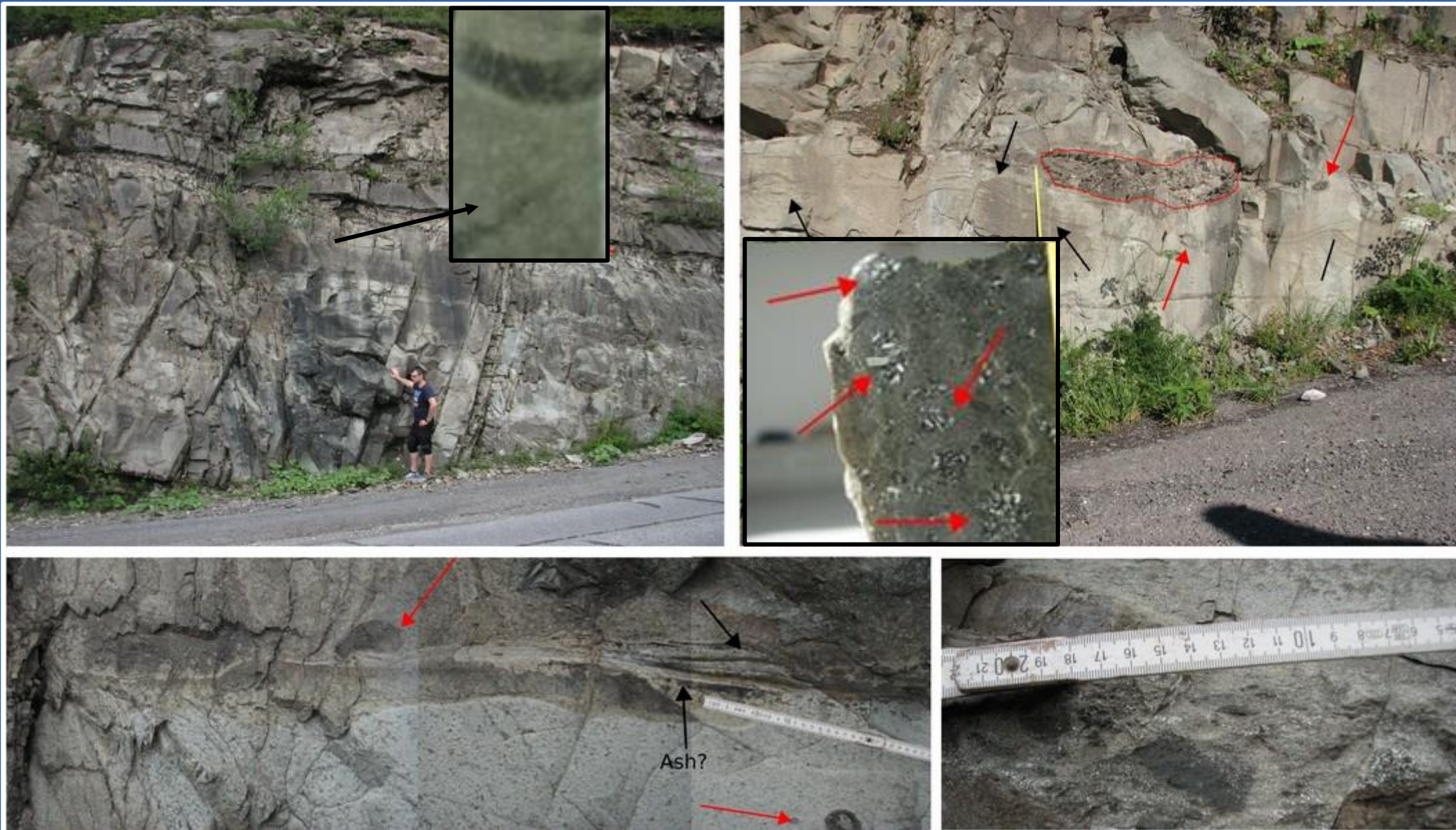
Grain-size distribution = function of the physical behavior of the eruptive event

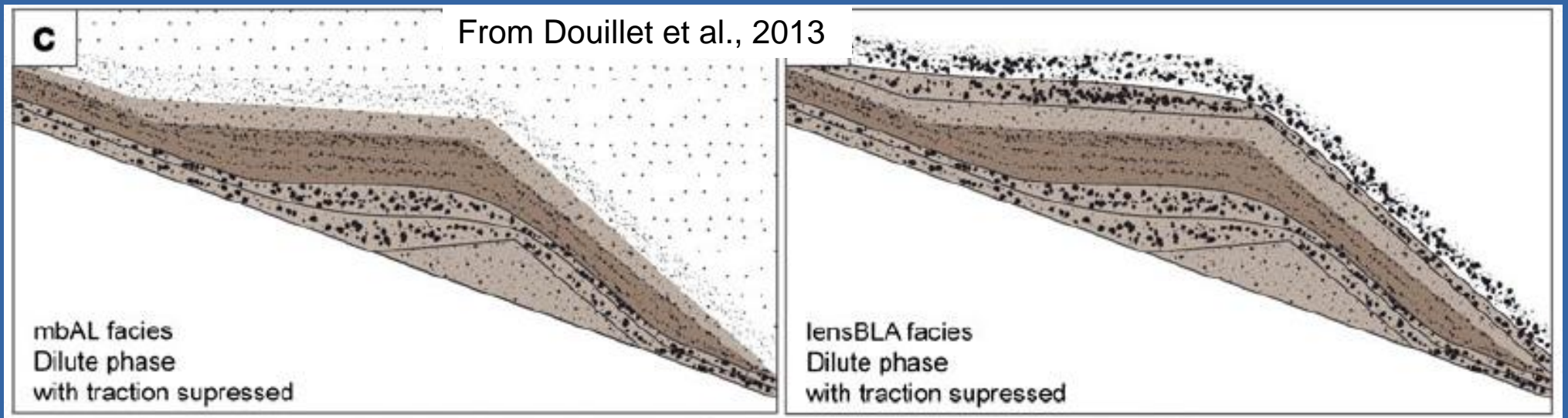
Detritus produced and instantaneously transported for tens to hundreds of kilometers

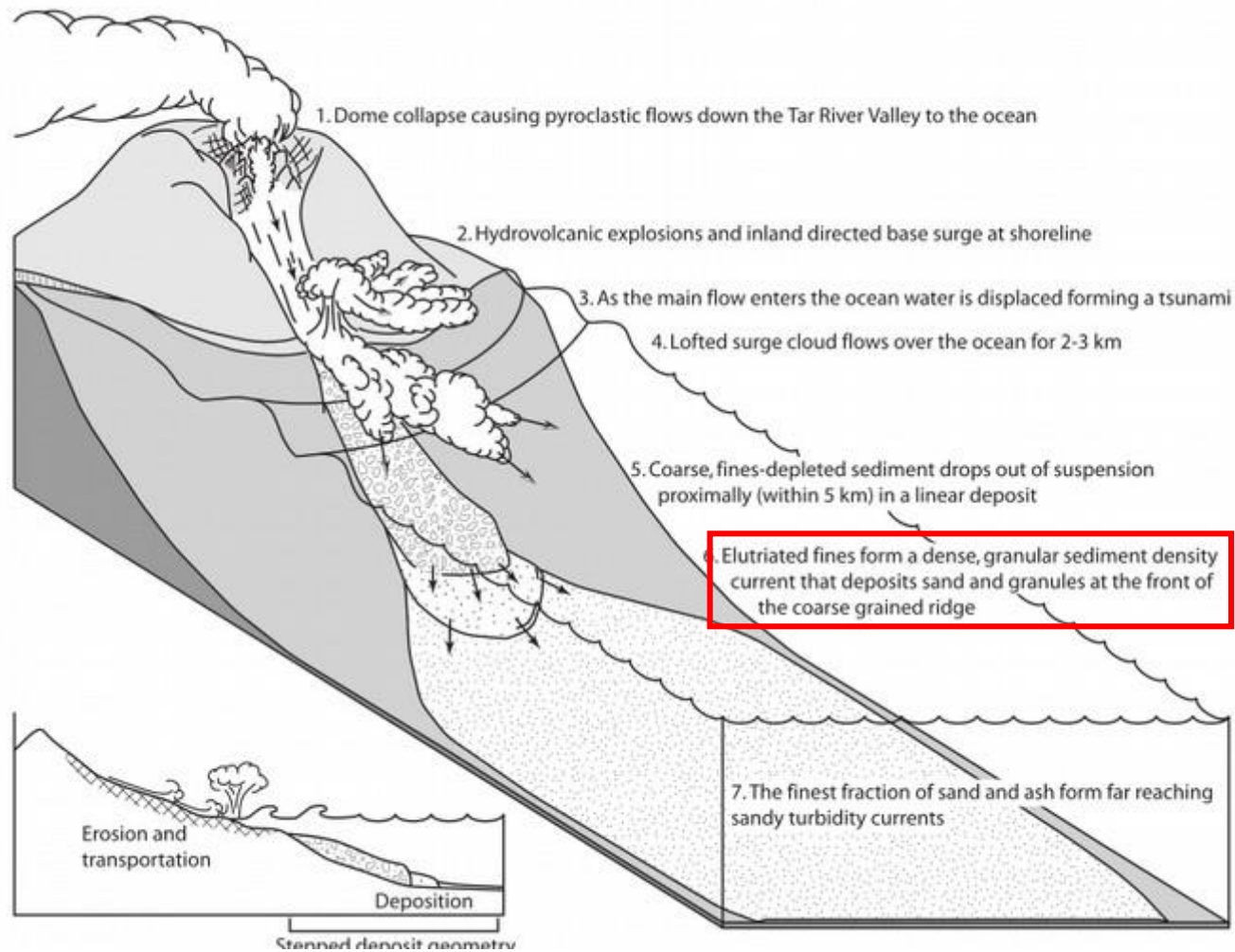
Syn-Volcanic Syn-eruptive or Post-eruptive deposits



Syn-volcanic syn-eruptive deposits: disaggregated PDCs



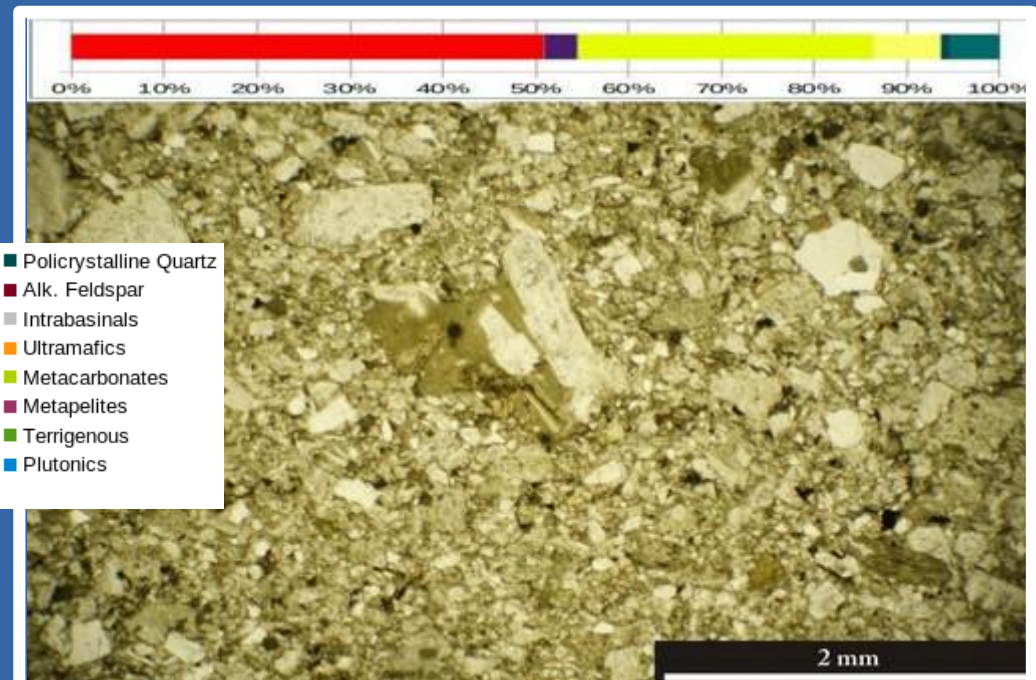
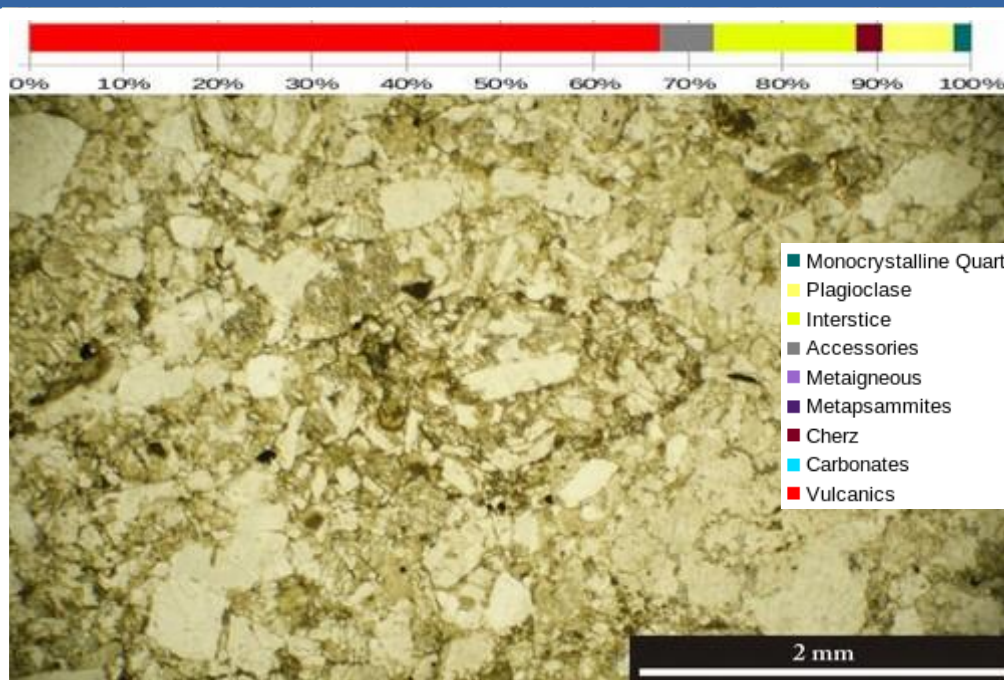
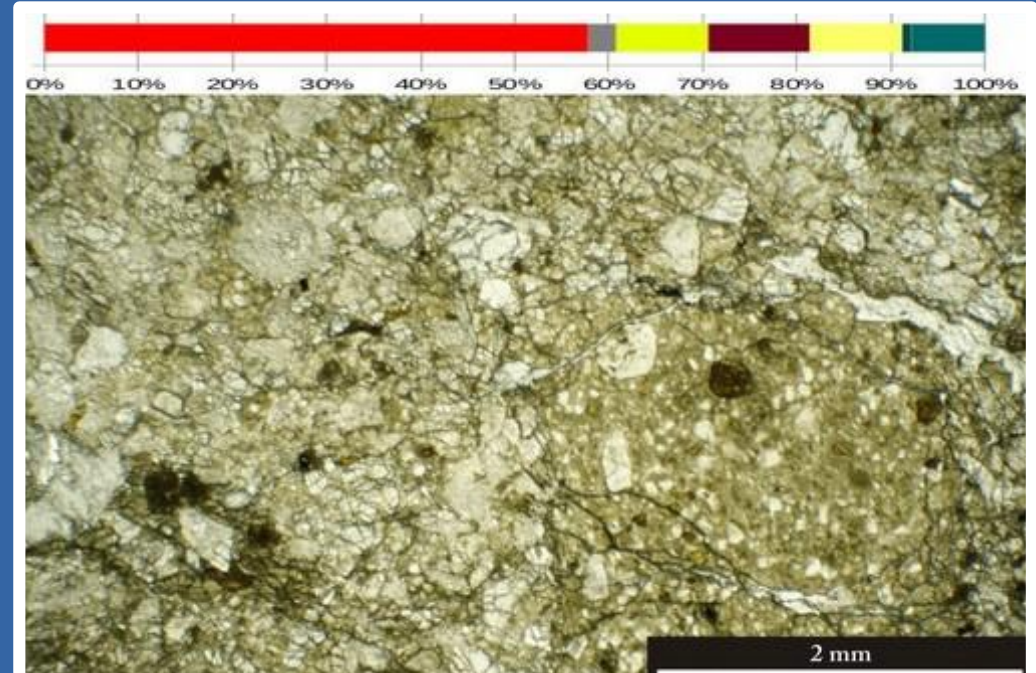
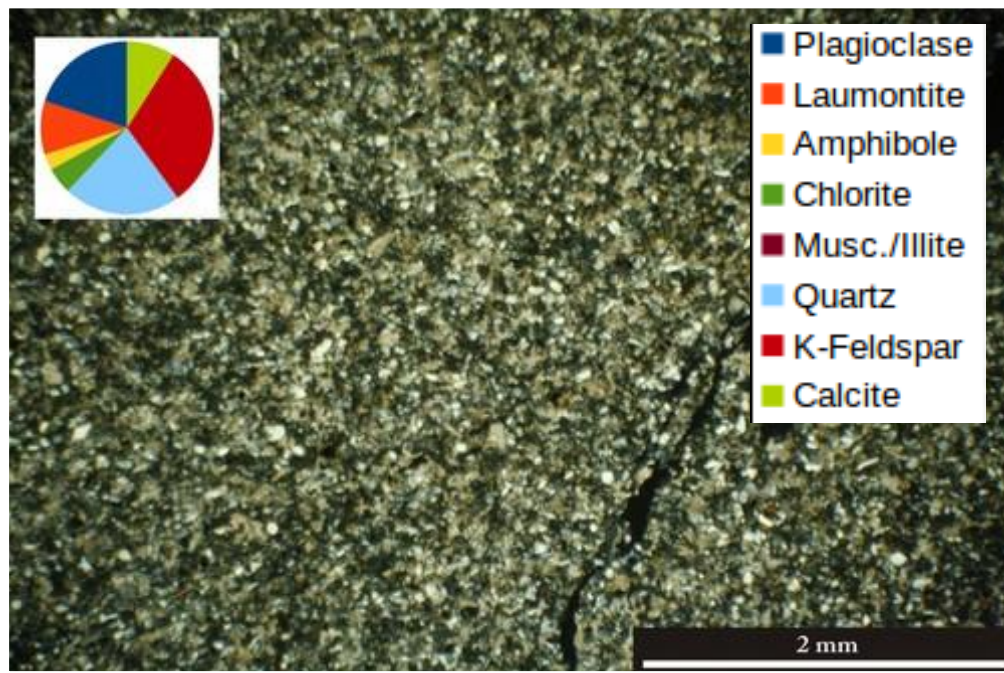


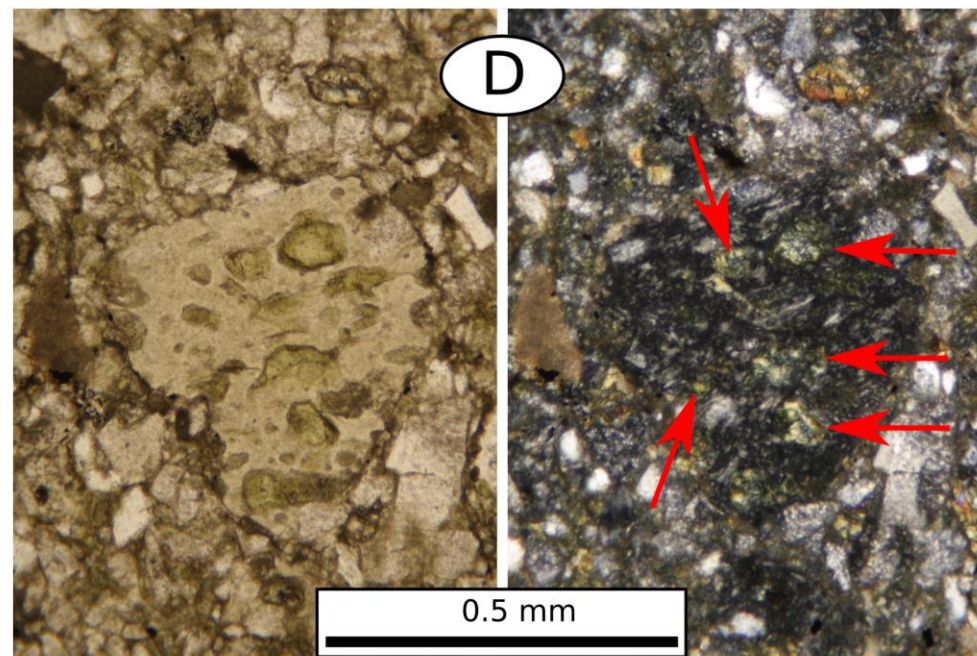
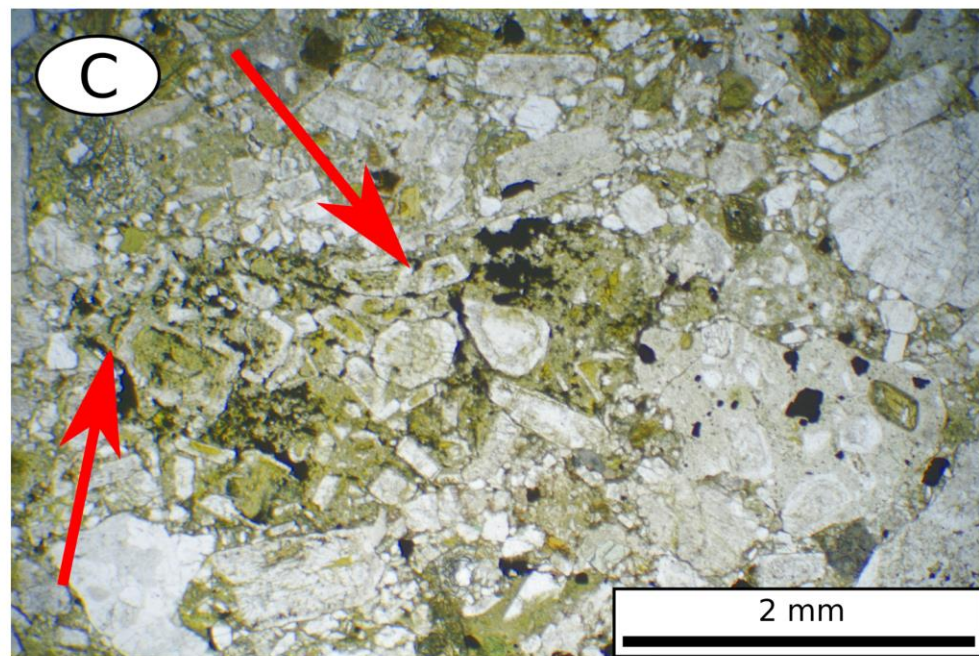
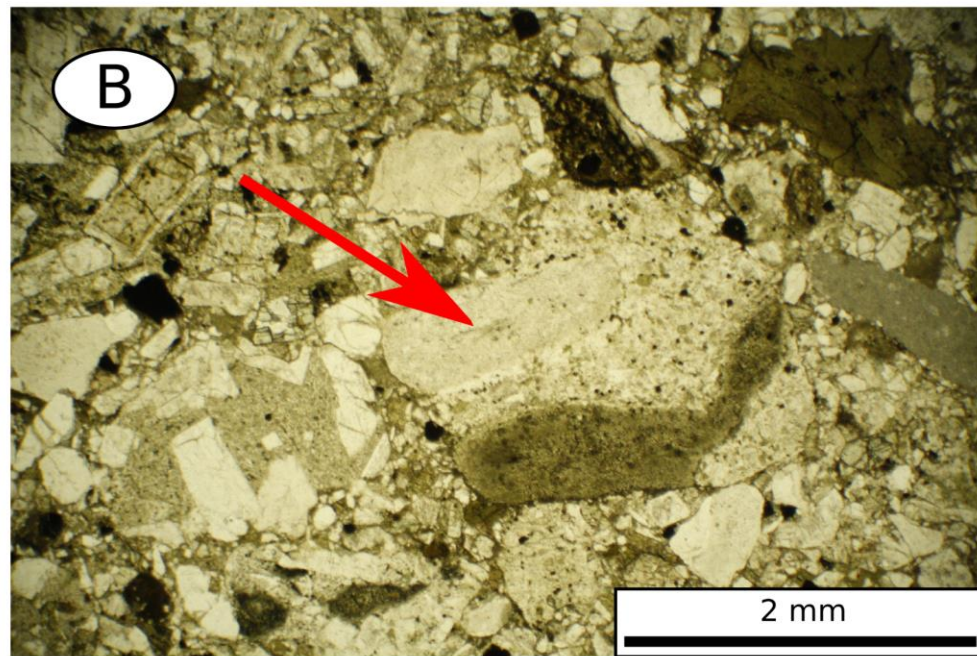
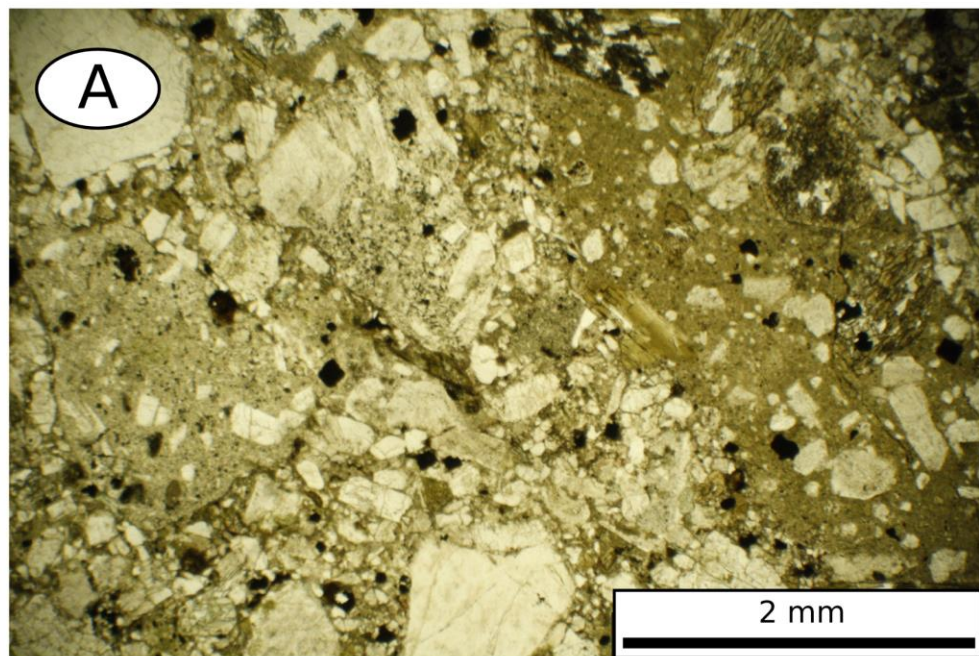


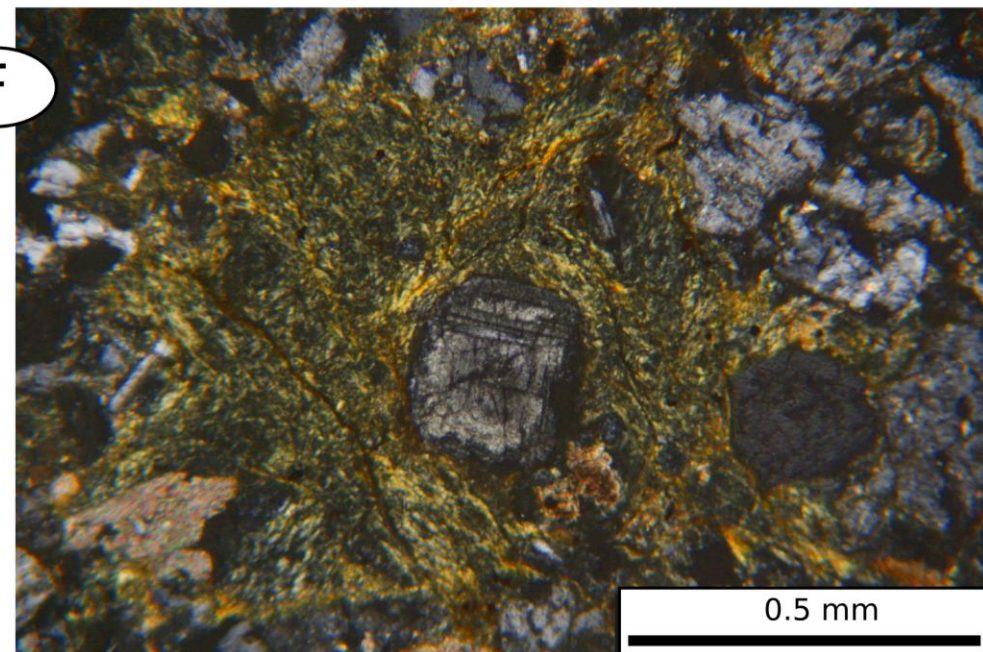
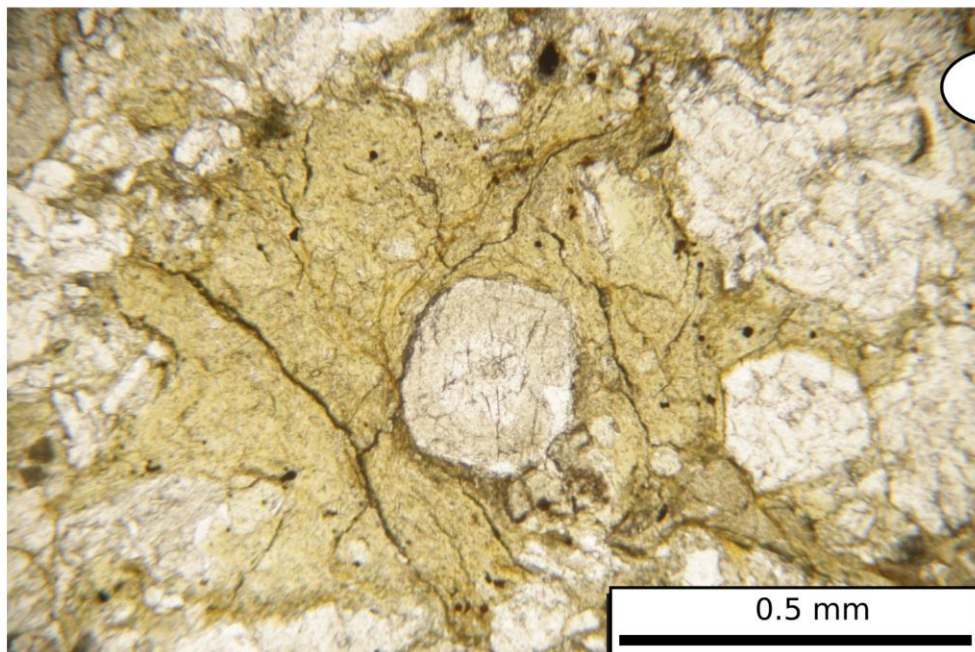
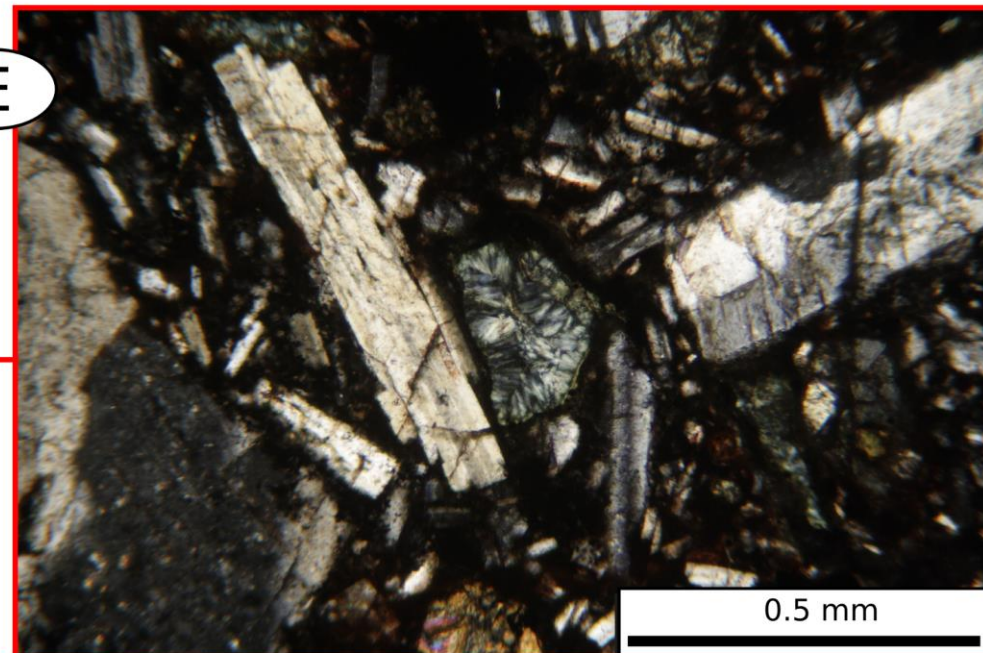
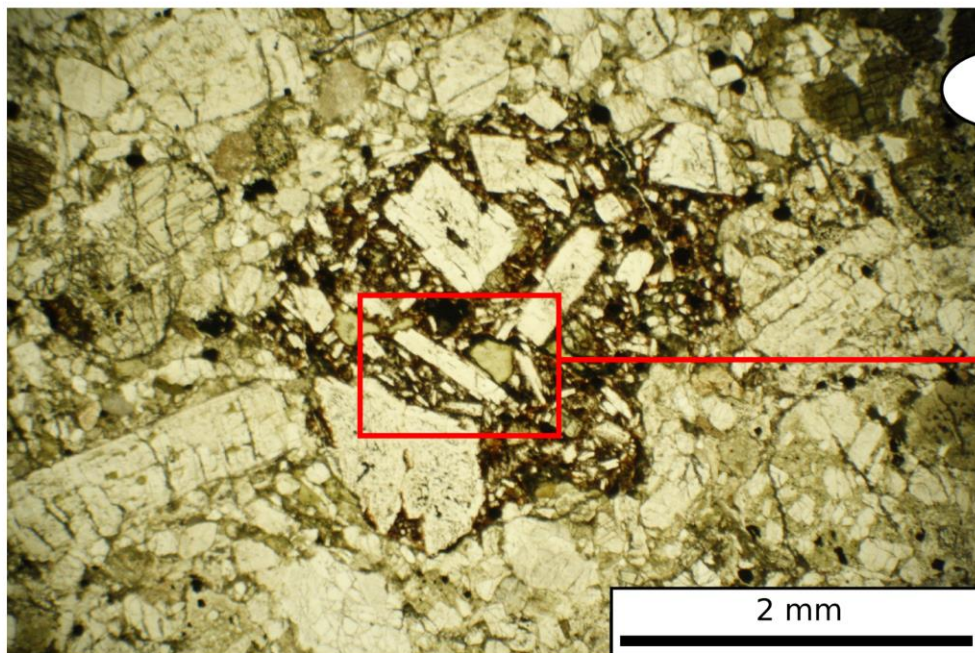
From Trofimovs et al., 2008

mbAL facies
Dilute phase
with traction suppressed

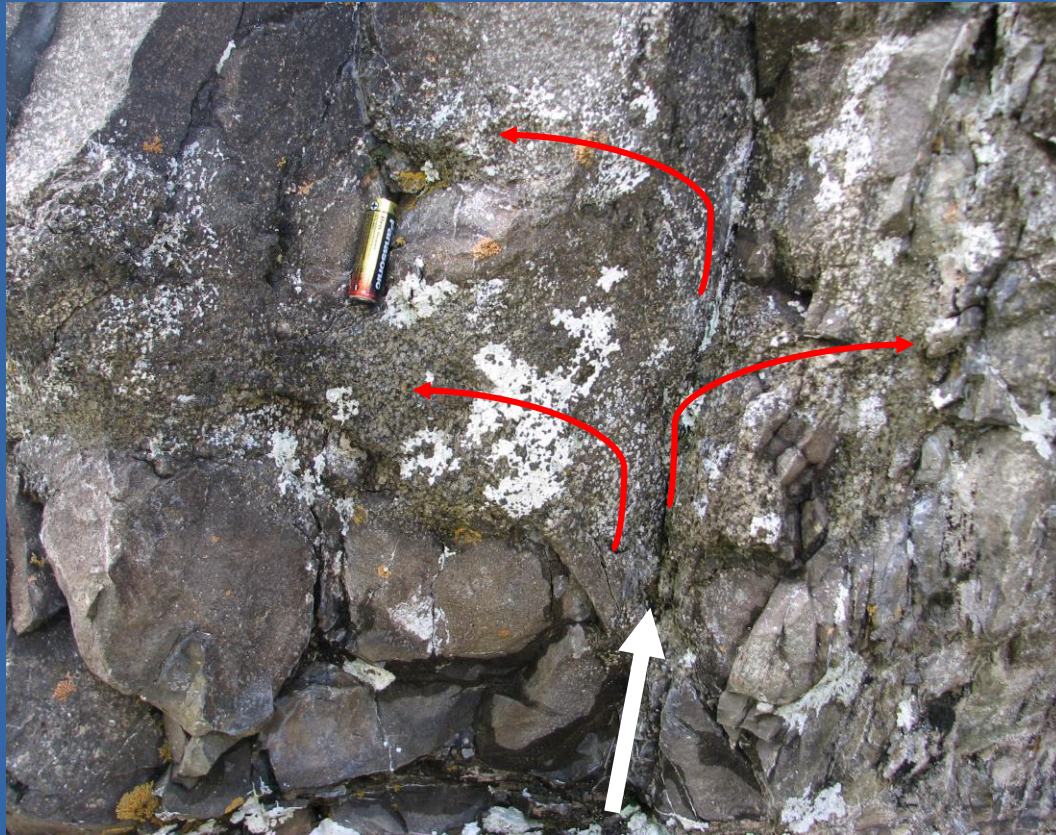
lensBLA facies
Dilute phase
with traction suppressed







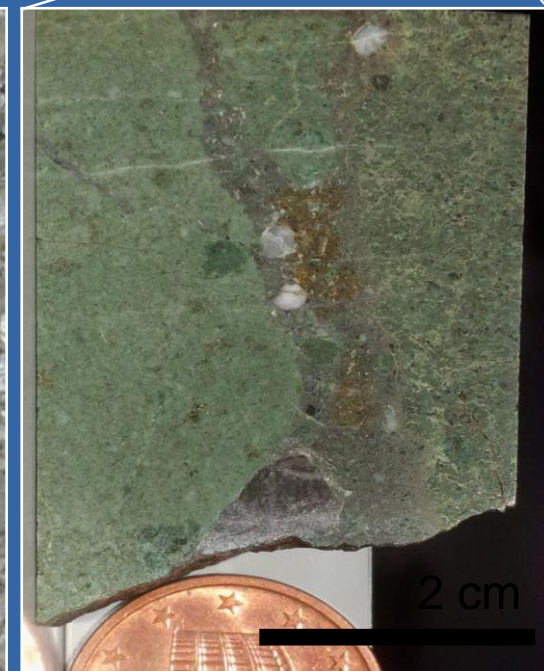
Laumontite in Taveyanne Sandstone



Laumontite cement creating «bee-structures»
(Ruffini et al., 1995), substituting glass and
plagioclase.

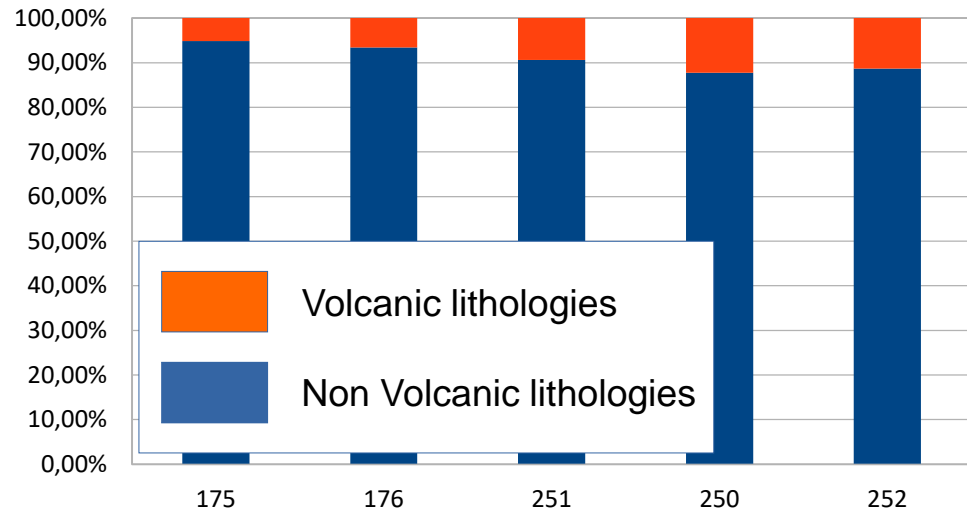
Preservation of primary porosity (e.g., Zhu et al.,
2014)

Syn-volcanic syn-eruptive deposits: preserved PDCs

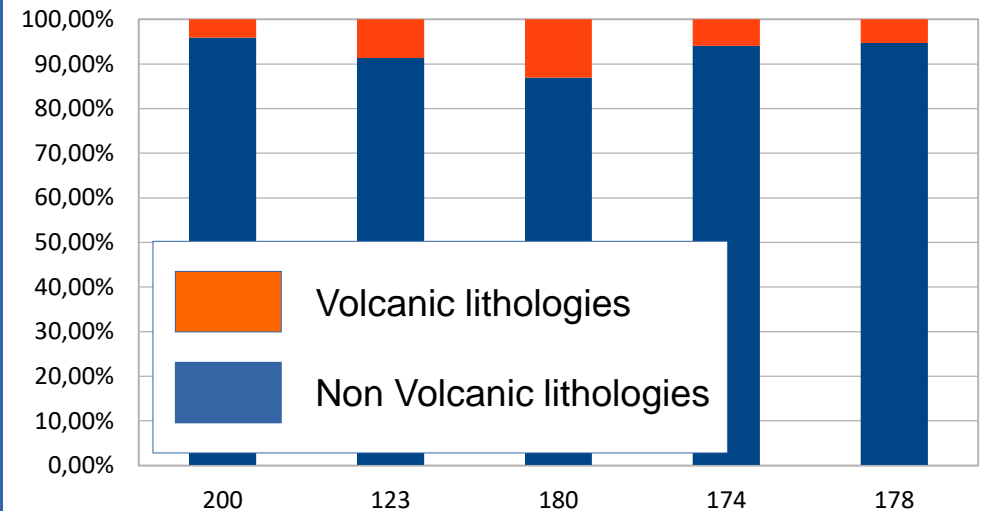


Gravel-size detritus

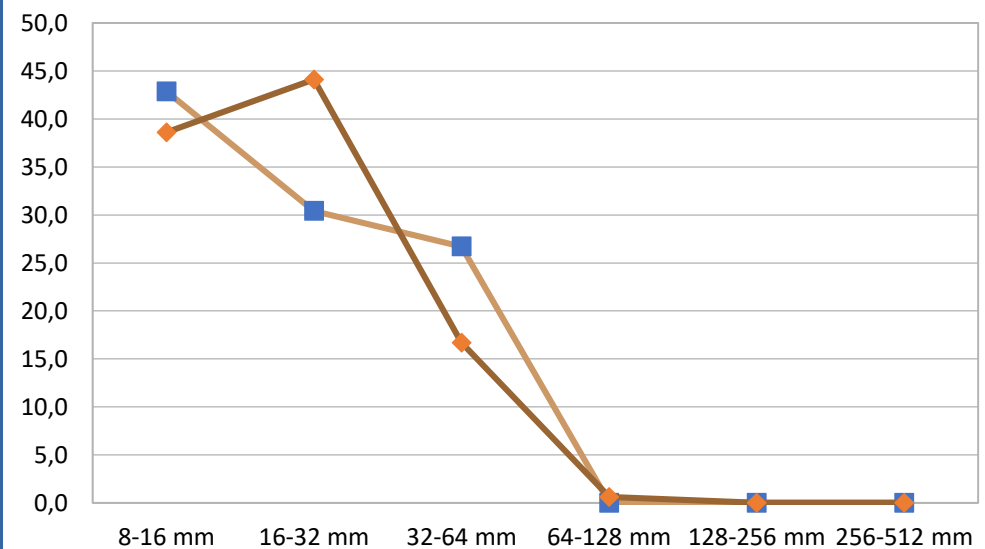
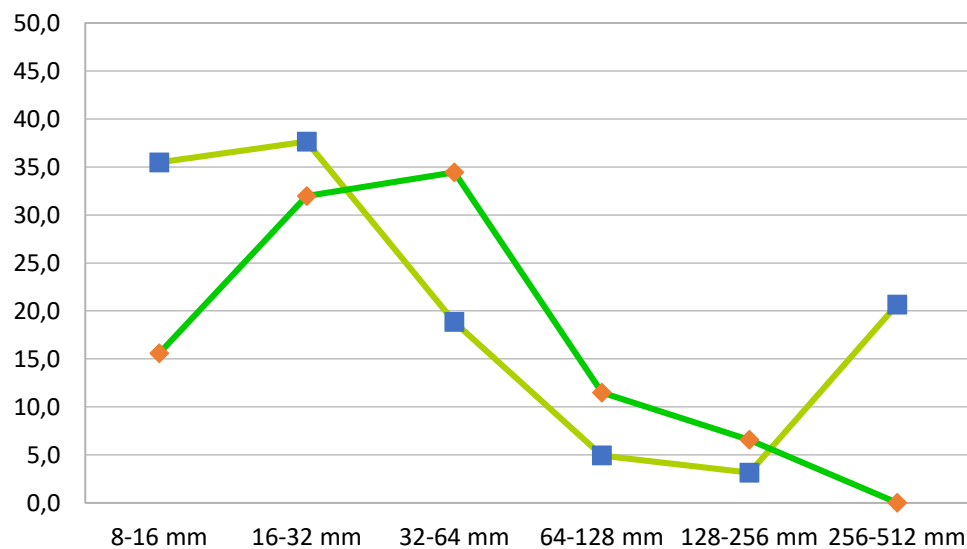
Syn-volcanic deposits



Non-volcanic deposits

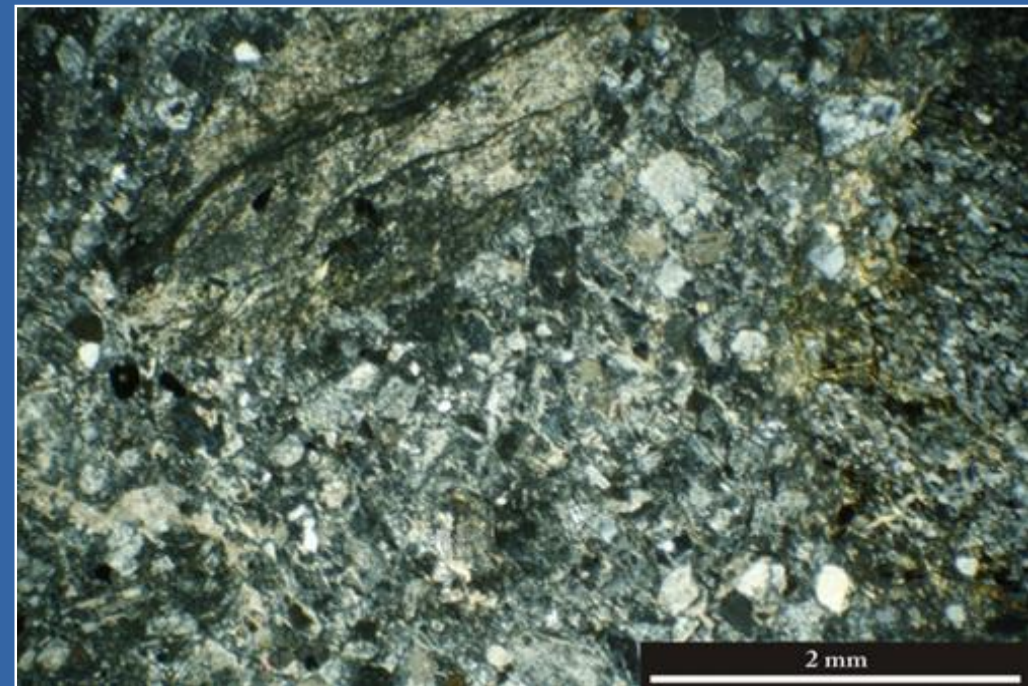
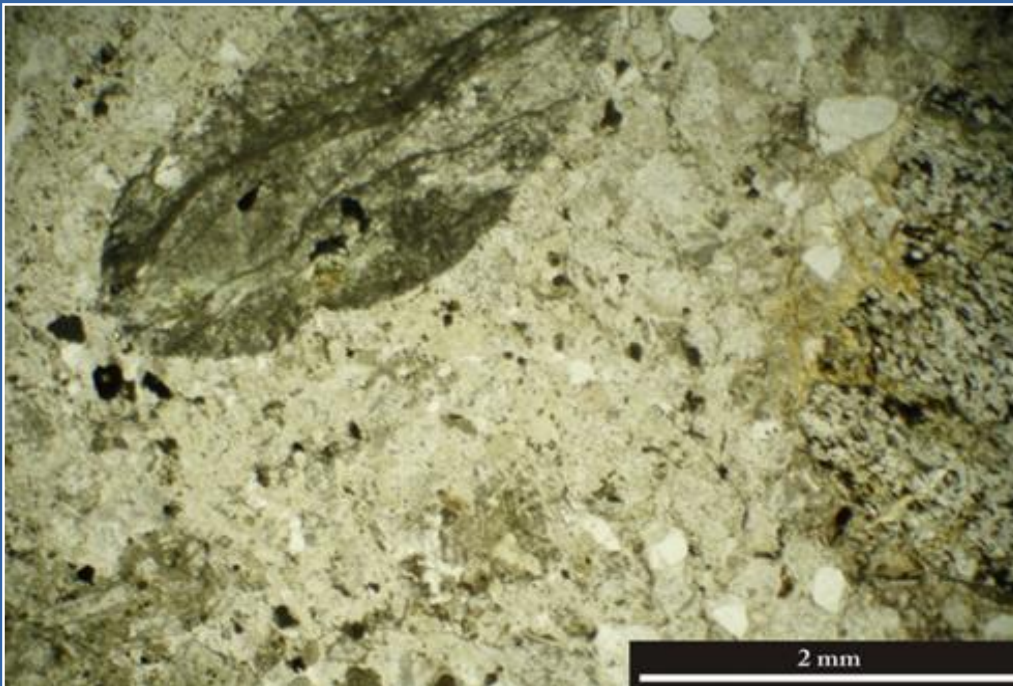
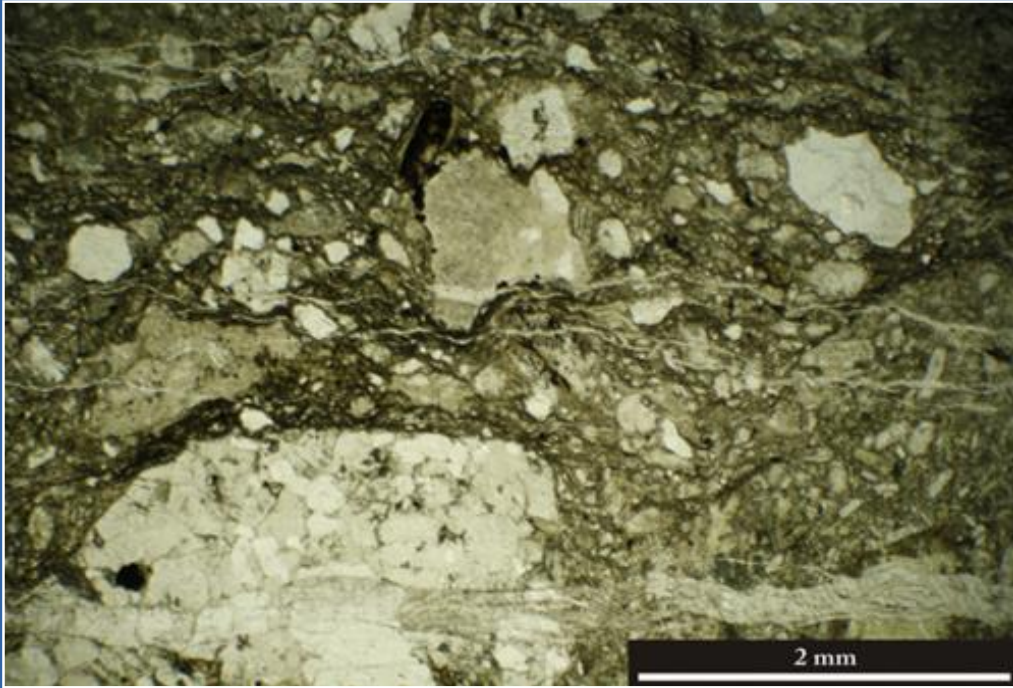


Lithological association of the gravel-size detritus, resulting from clast counts in the field

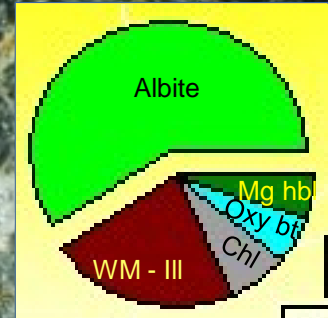
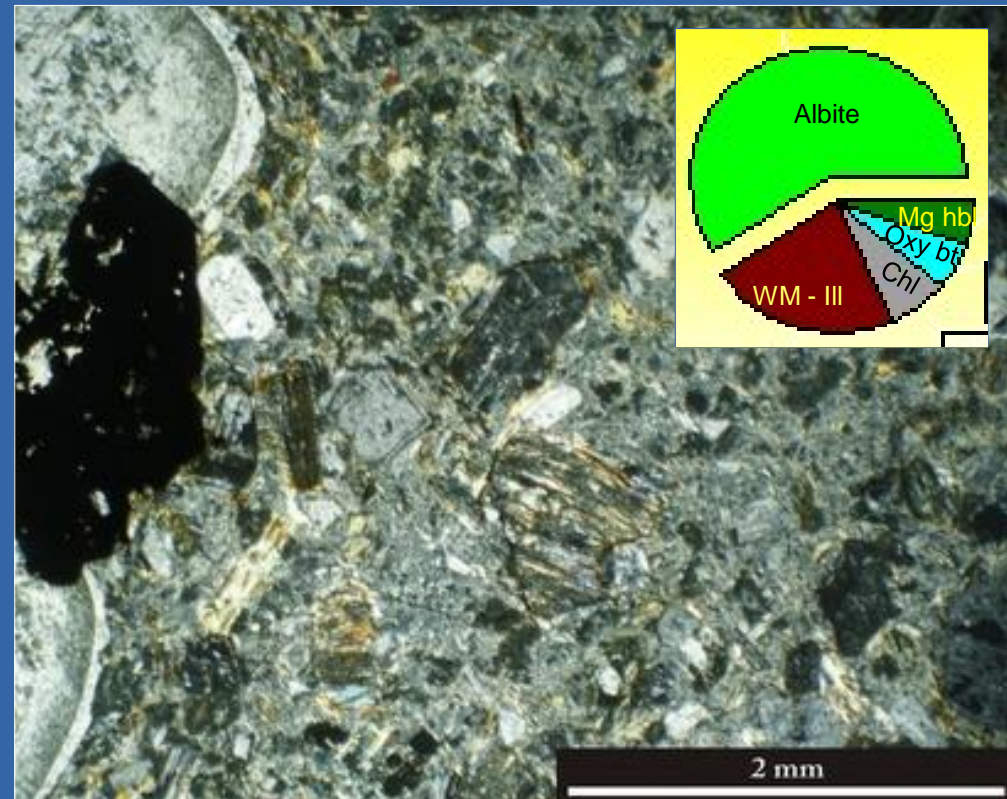
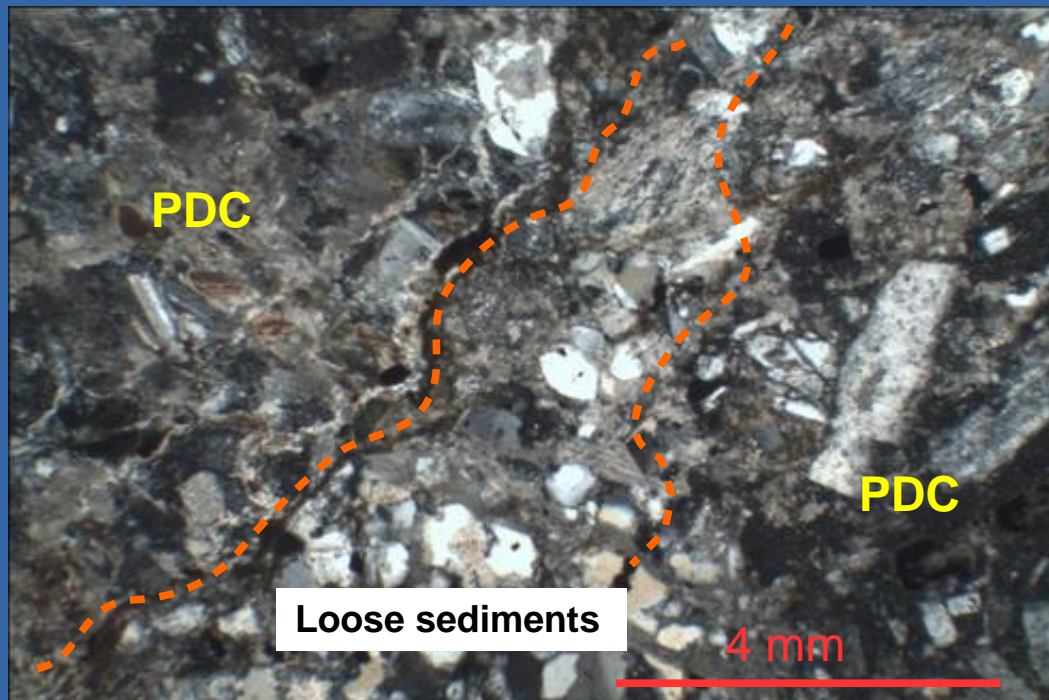
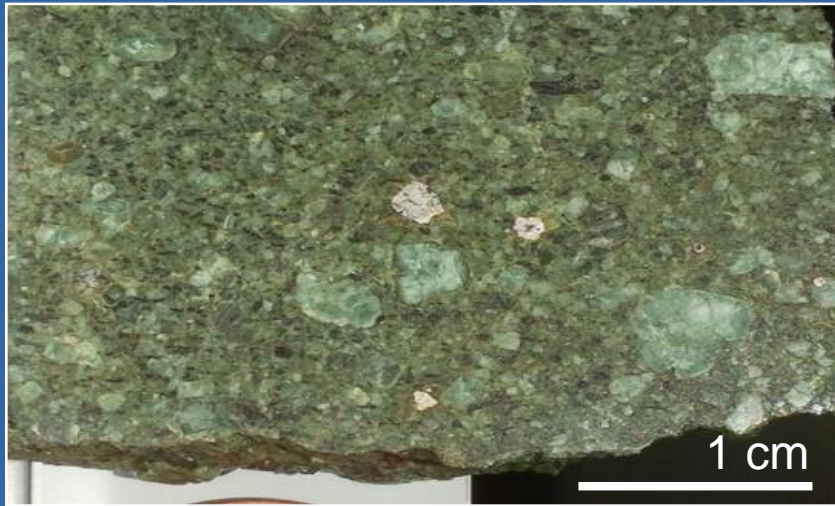


Grain size distribution of the gravel-size detritus resulting from image analyses

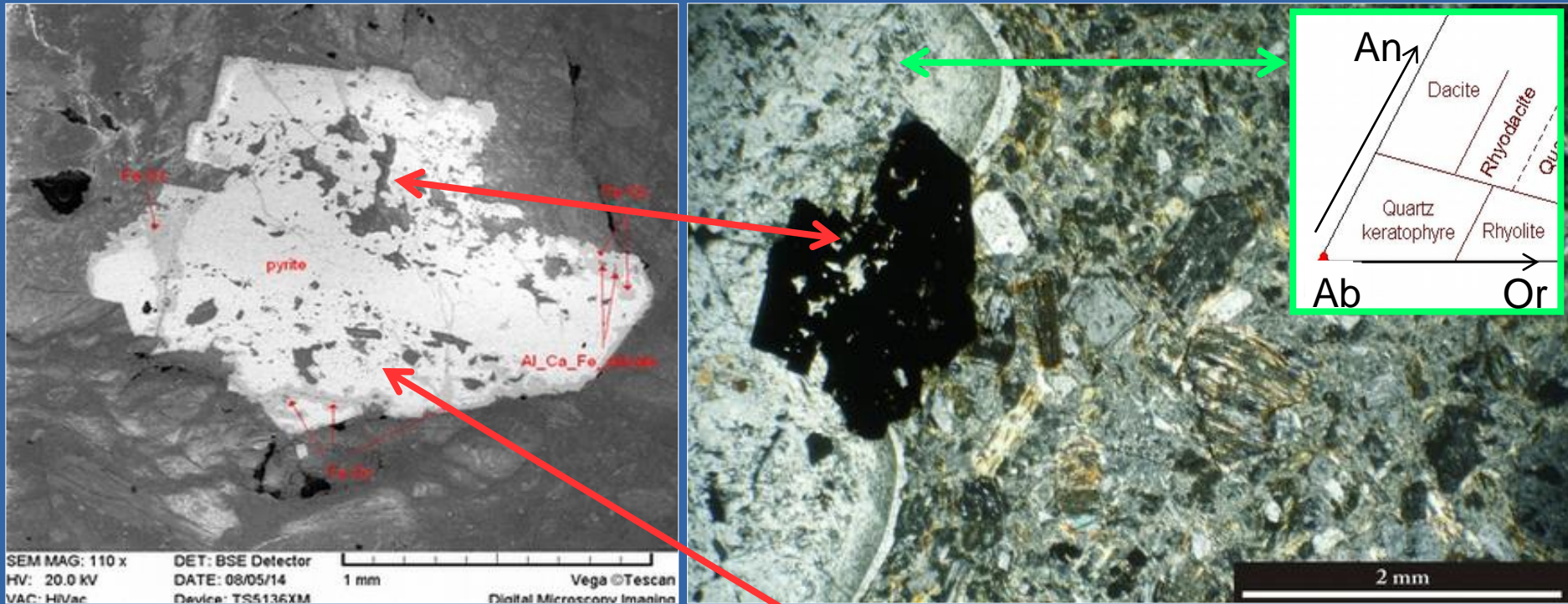
Syn-volcanic syn-eruptive deposits: preserved PDCs



Syn-volcanic syn-eruptive deposits: preserved PDCs



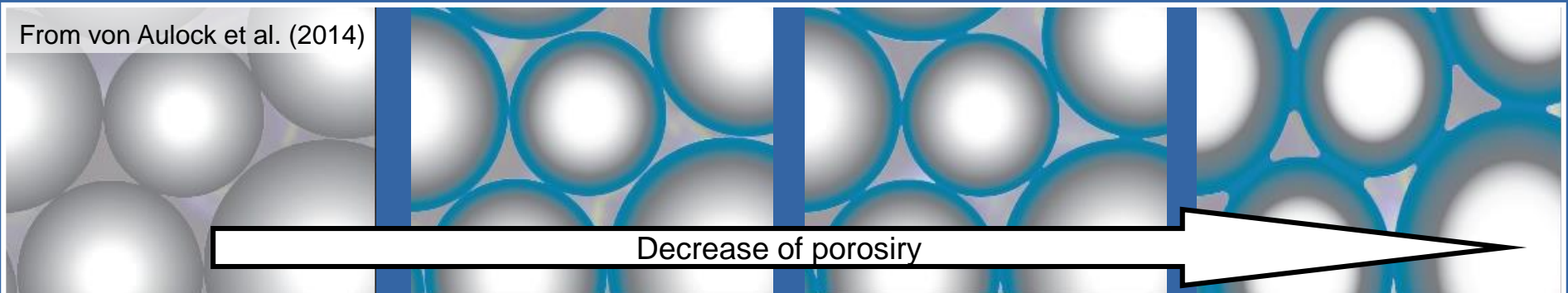
Syn-volcanic syn-eruptive deposits: preserved PDC



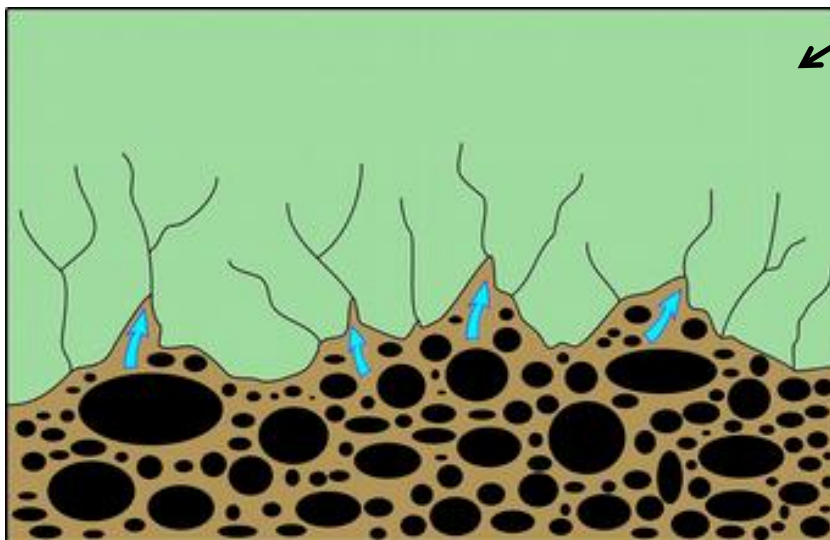
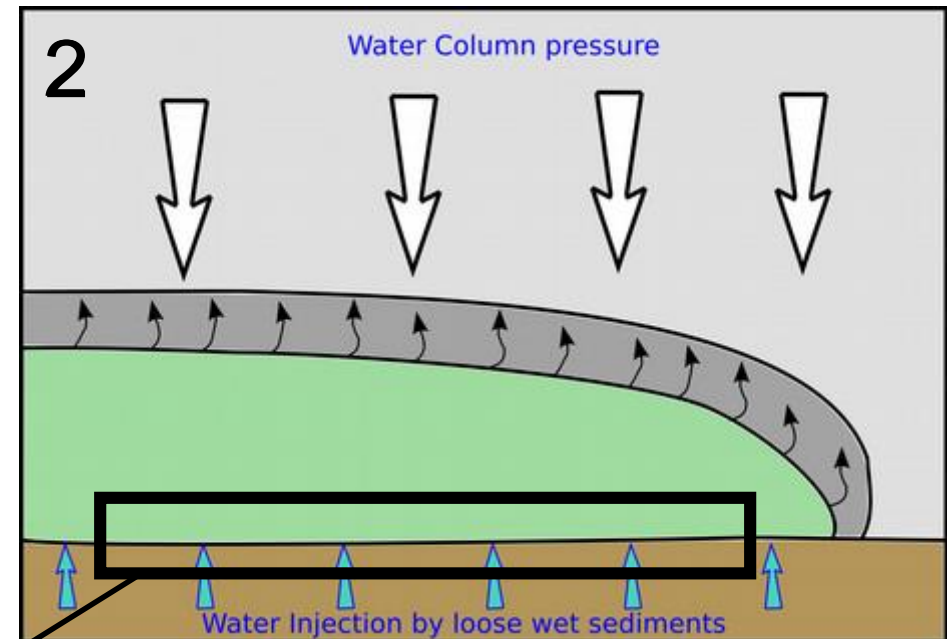
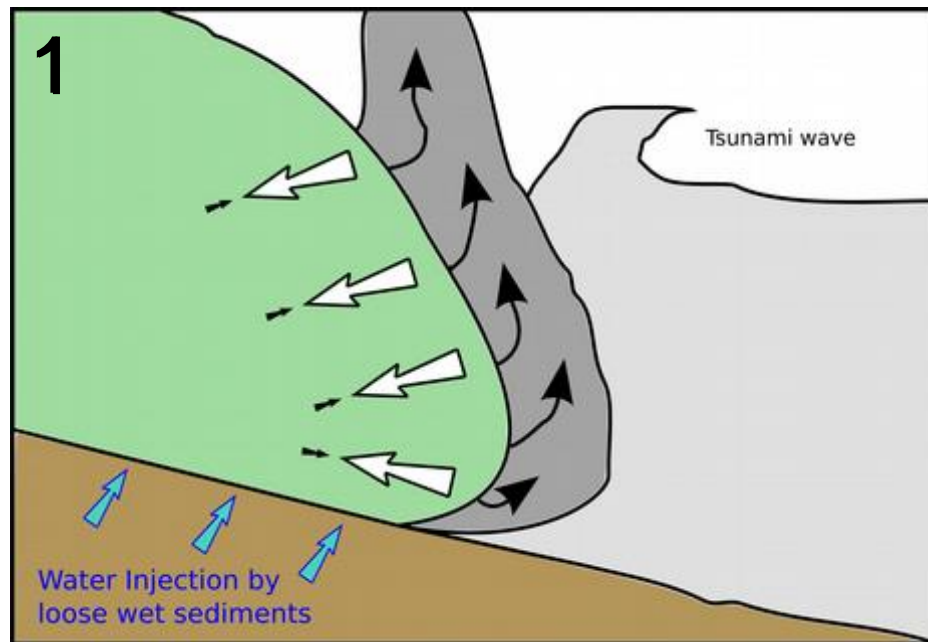
Type	Surface *1	Coexist. Phases	Grain Size	Habit	Faces *2	Habit Type *3	Abundance	Temperature	Experiments *4	Remarks
Pyrite C1	S°	S°	<1μ - <50μ	cubo - octa	100/111	1, 2, 19	very minor	250 - 350	II, IV	growth on top of sulfur droplet surface
Pyrite C2	S°	S°	<1μ - <50μ	pyritohedral	210/100/111	4	> Py C1	300 - 350	II	growth between two sulfur droplets
Pyrite C3	S°, Po	S°, Po, Py A	1μ - 10μ	pyritohedral	210/100/111	4, 7	Py C1 < Py C3 < Py C2	300 - 350	II	growth between Py A and sulfur droplets
Pyrite D	---	hematite	< 1μ - 0.3 mm	pyritohedral	210/100/111	4, 7, 6	> Py B1	250 - 350	II, IV	formation of clusters

From Graham and Ohmoto, 1994

From von Aulock et al. (2014)



Syn-volcanic syn-eruptive deposits: preserved PDC



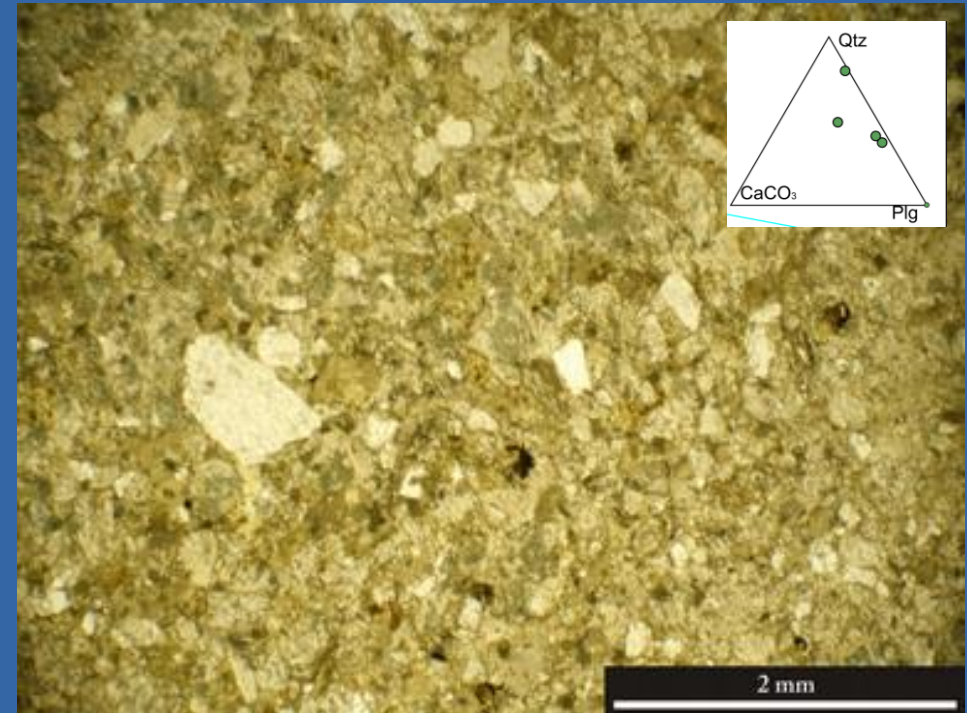
2,1%

Decrease of effective porosity

4,4%

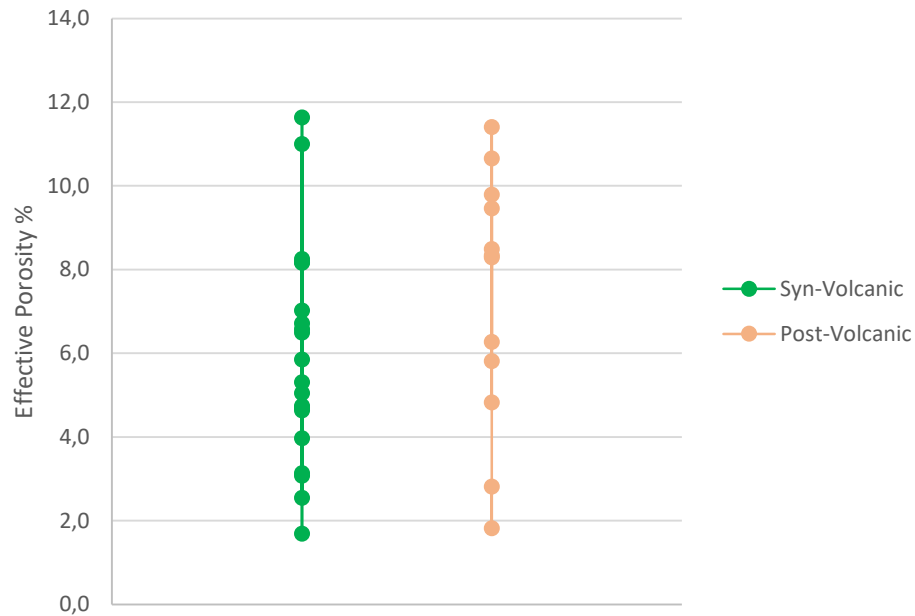
From Di Capua & Groppelli (2016b)

Syn-volcanic post-eruptive deposits

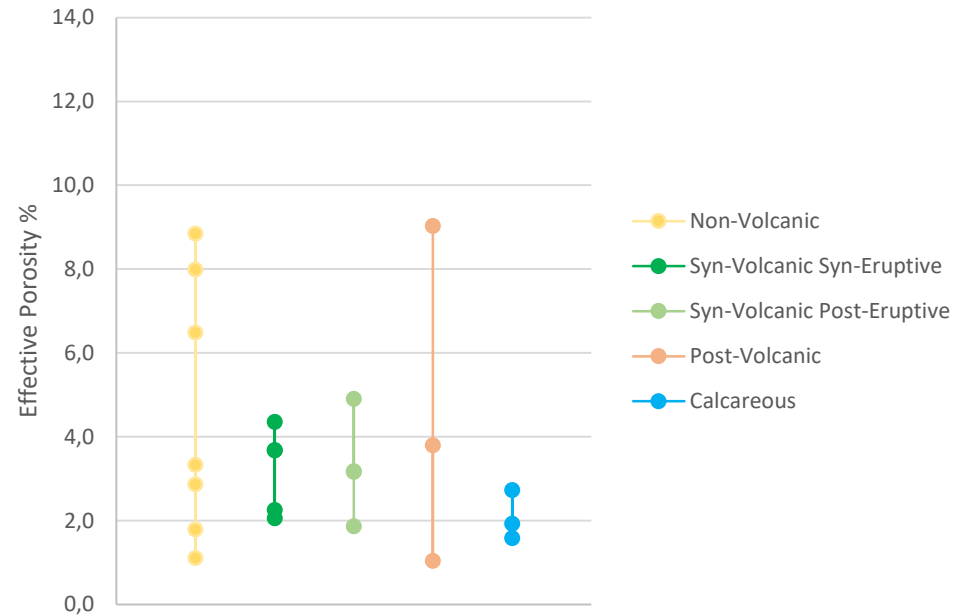


Effective porosity

Taveyanne Sandstone

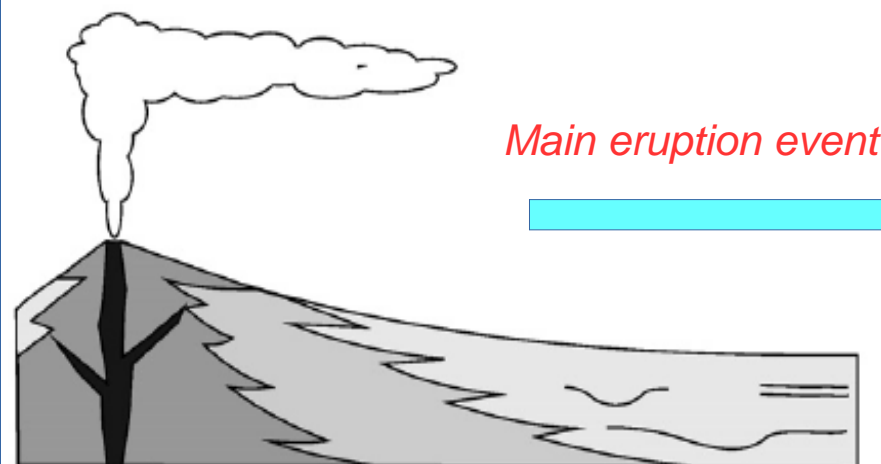


Val d'Aveto Formation



Function of the behavior of eruptive event

Post-depositional history (diagenesis)



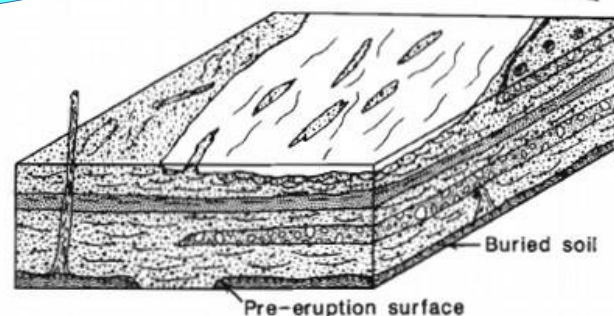
Main eruption event

volcanic core facies proximal volcaniclastic facies medial volcaniclastic facies distal volcaniclastic facies

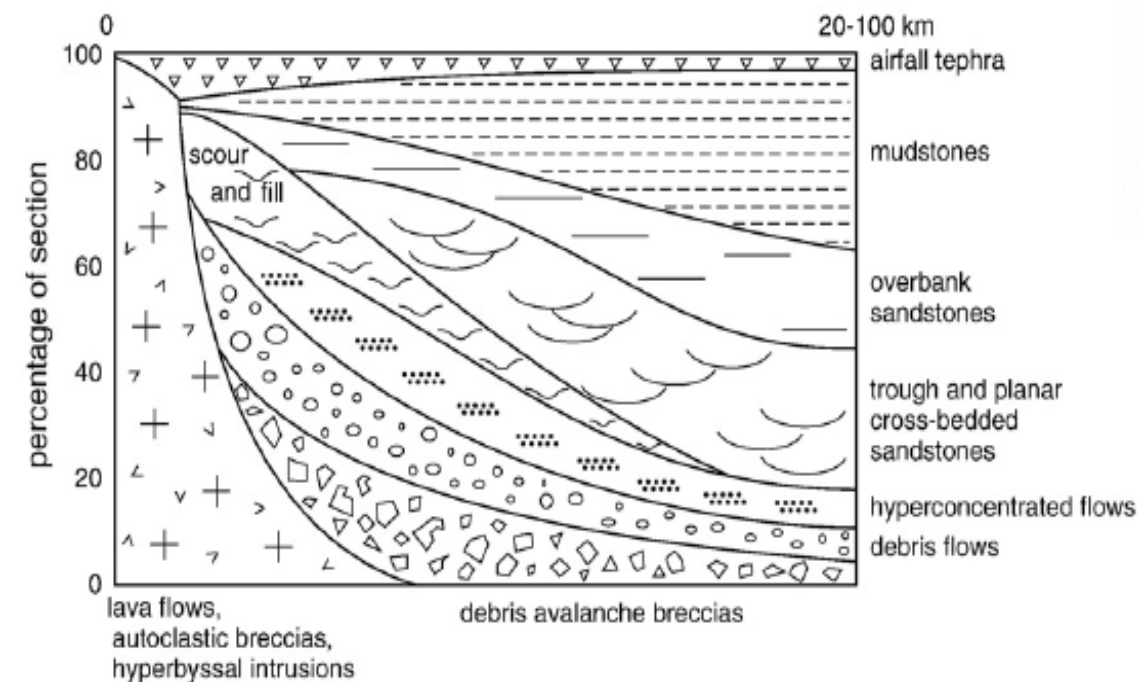
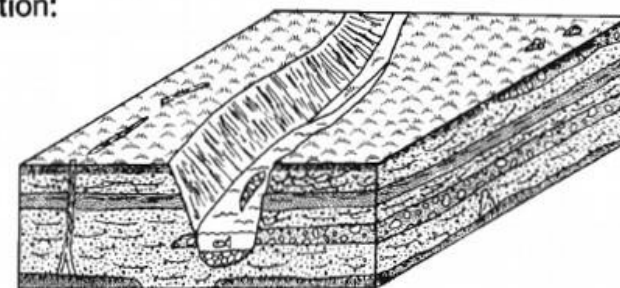
LANDSCAPE AND SEDIMENTATION:

From Smith (1991)

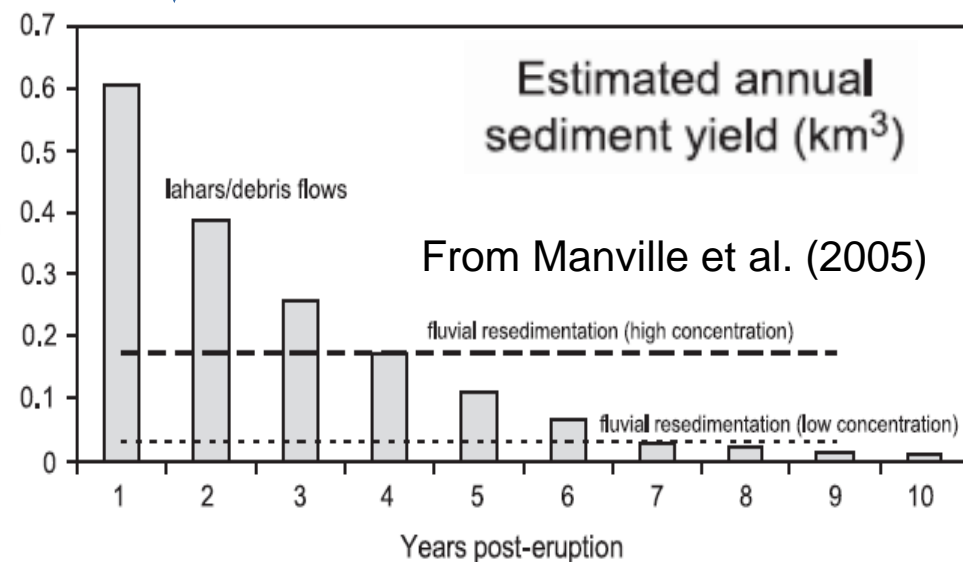
Syneruption:



Inter-eruption:

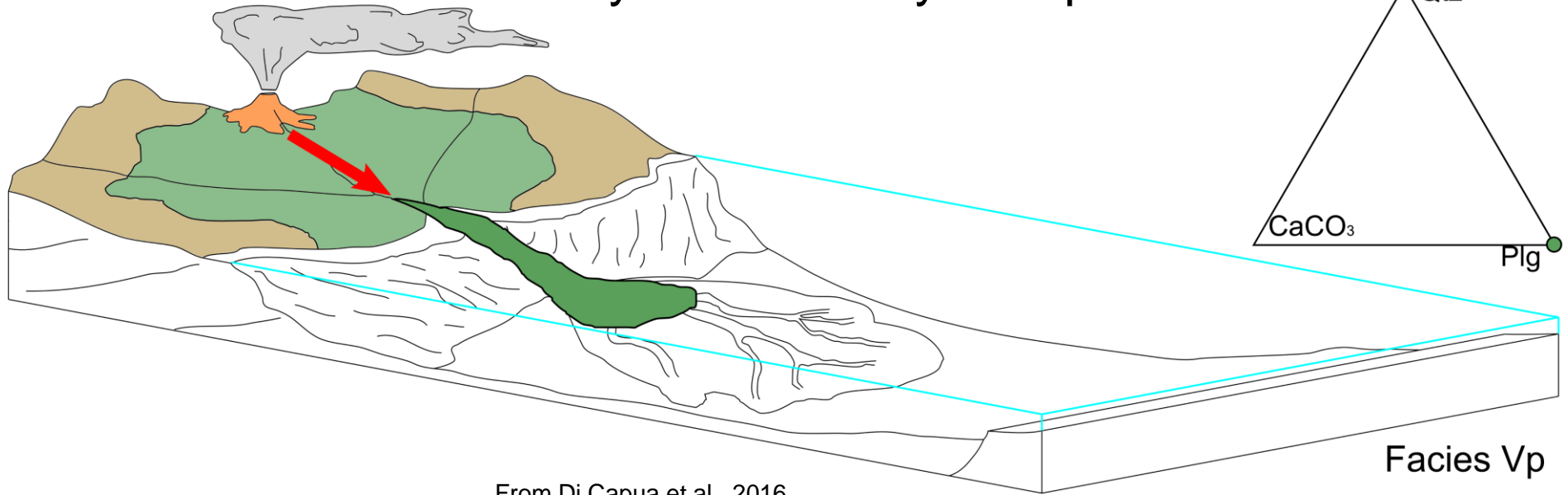


From Manville et al. (2009)

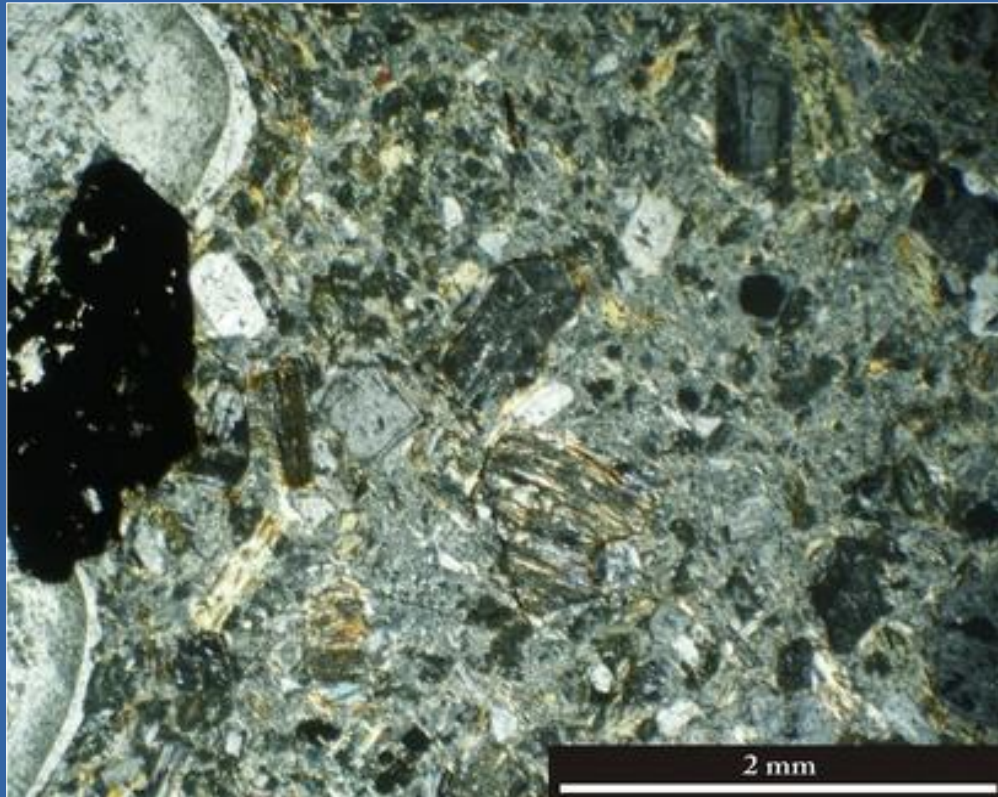


From Manville et al. (2005)

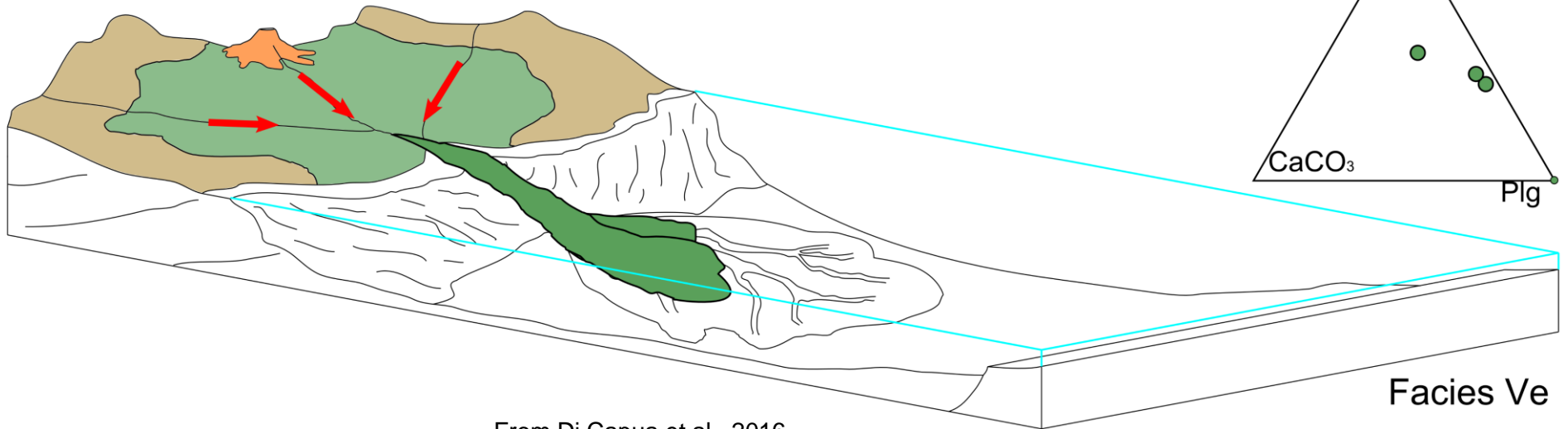
Syn-volcanic syn-eruptive facies



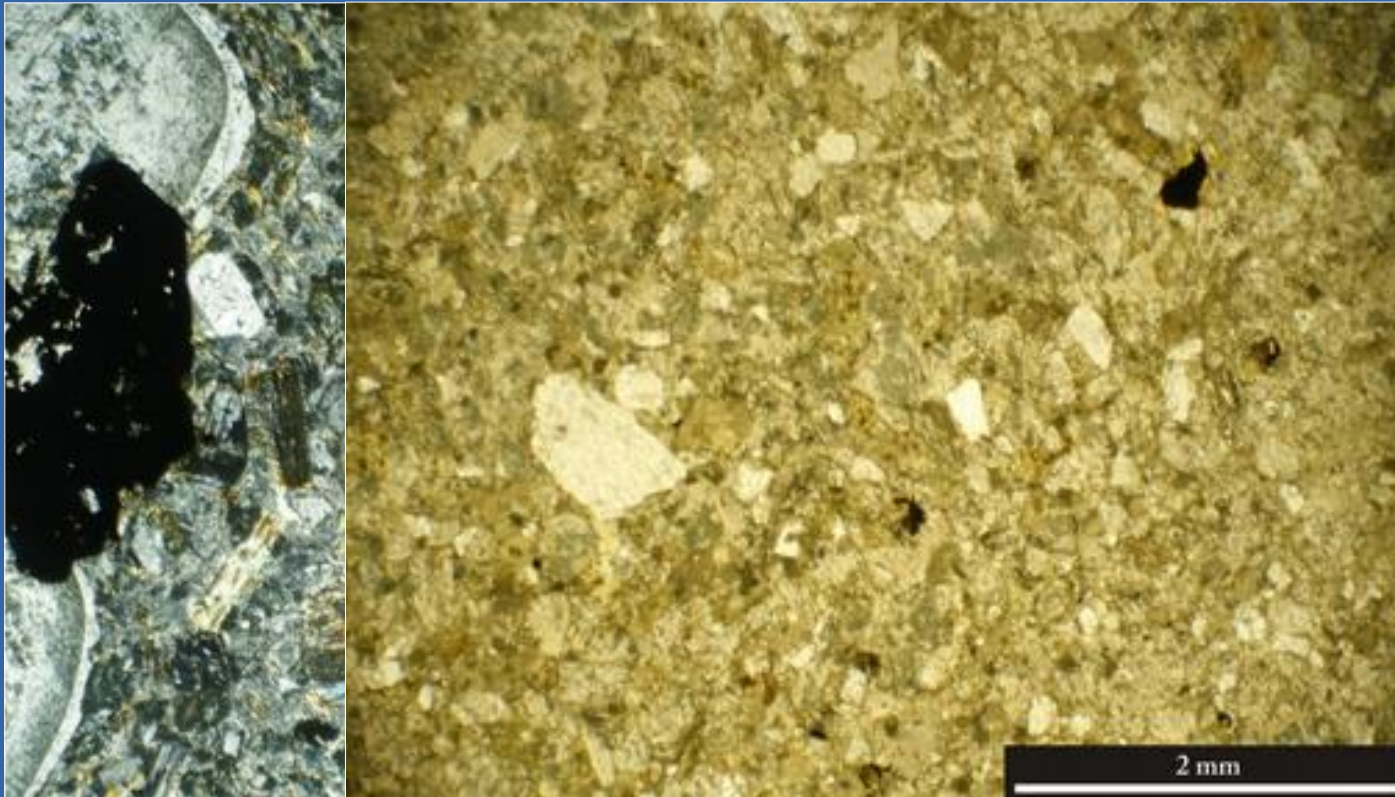
From Di Capua et al., 2016



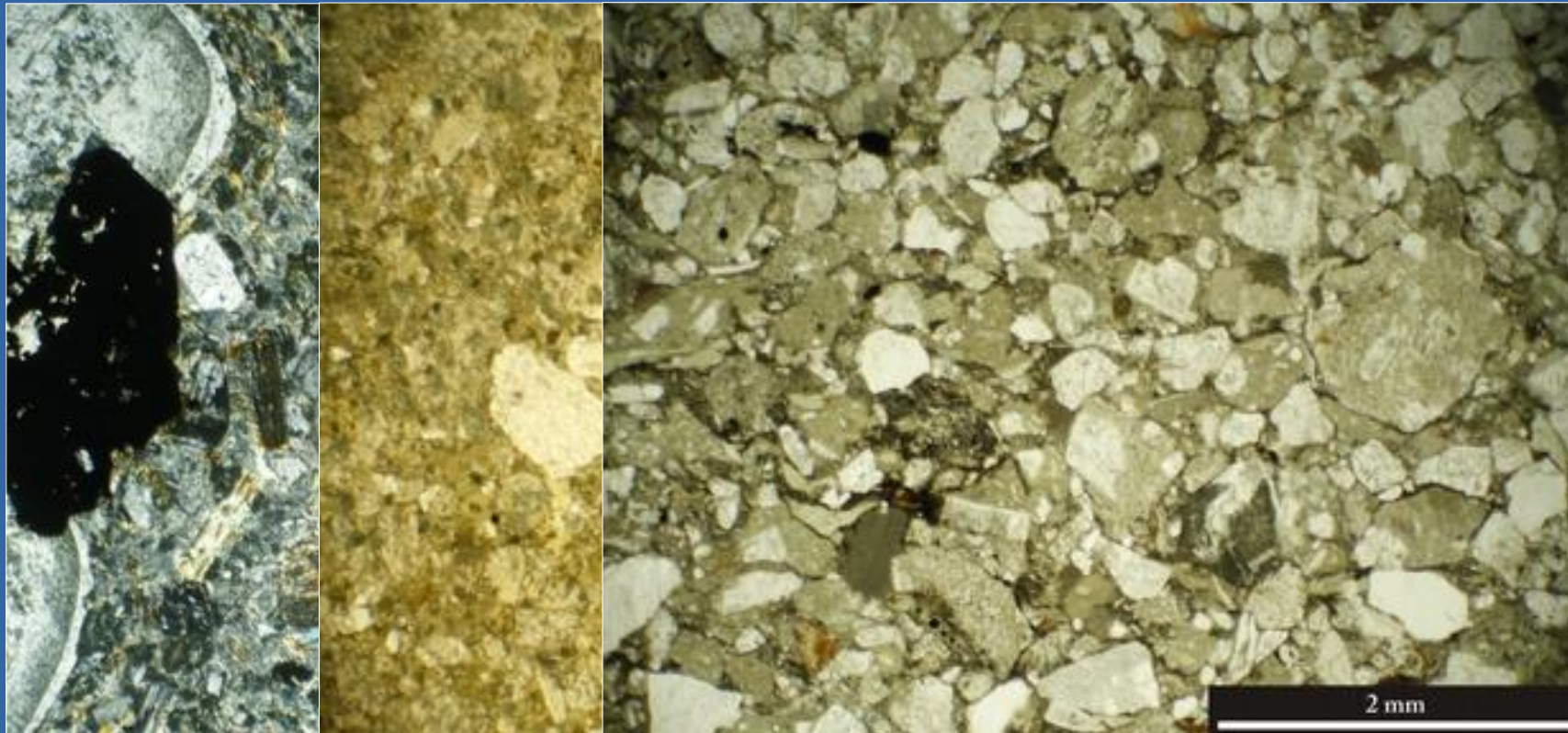
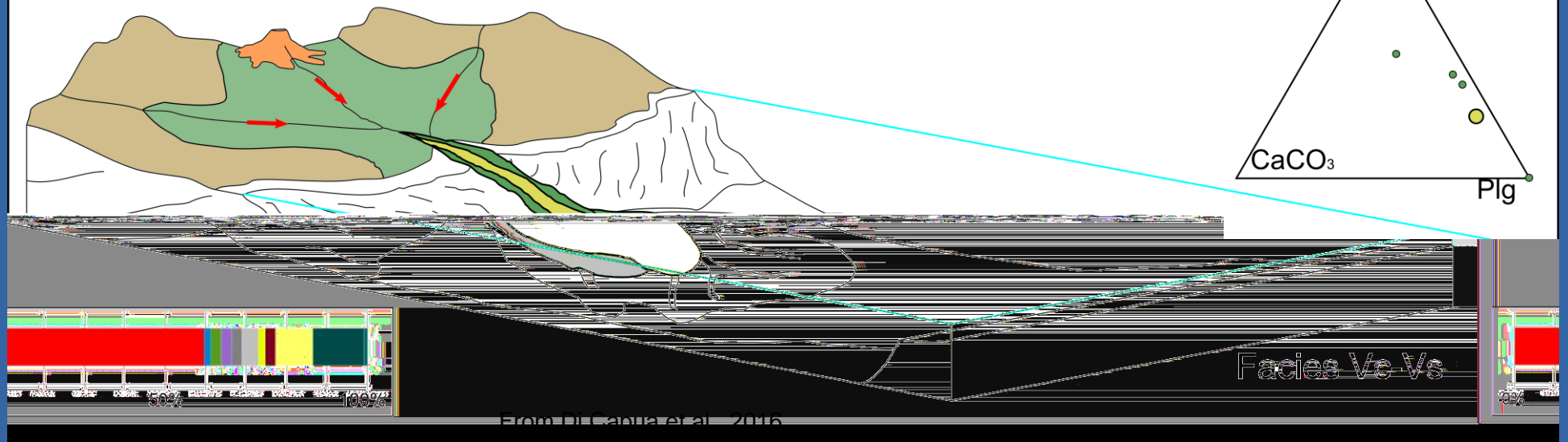
Syn-volcanic post-eruptive facies



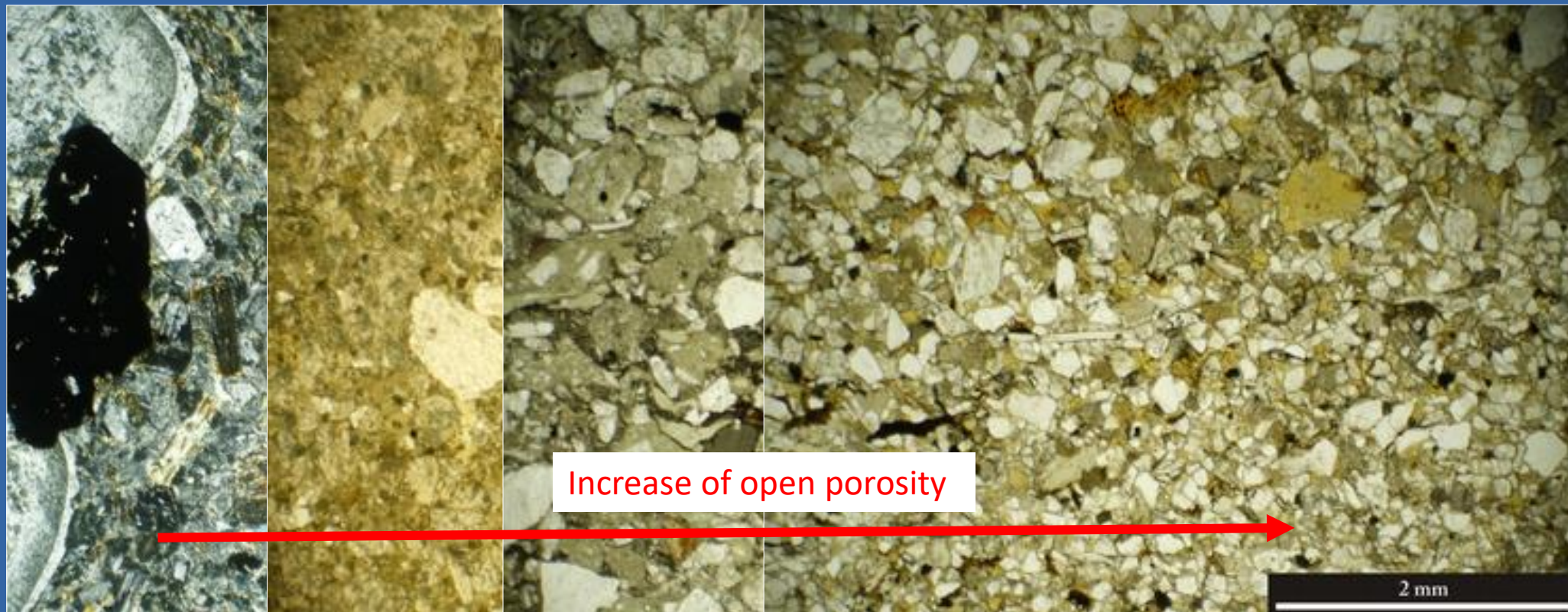
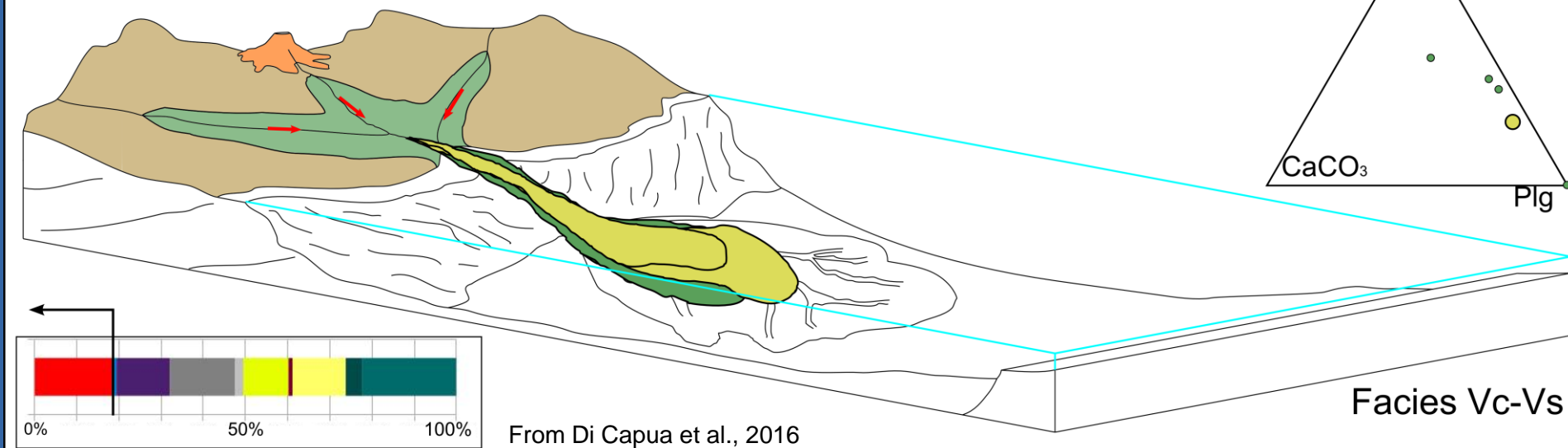
From Di Capua et al., 2016



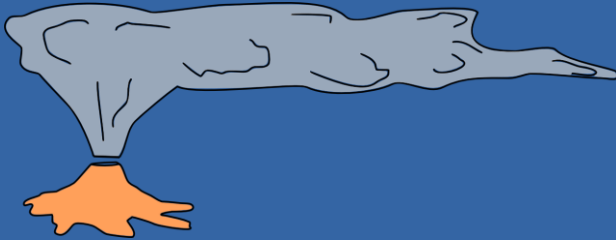
Post-volcanic facies



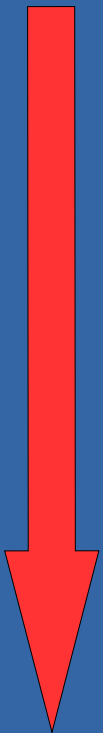
Post-volcanic facies



Take-away points



The presented work and classification show the importance of a multidisciplinary and multiscale approach in the identification of the temporal relationship between volcanism and sedimentation, considered the main factor controlling the physical, minero-chemical, diagenetic and petrophysical properties of volcanoclastic sequences.



Thank you for your attention



Ongoing projects:

- .Characterization of the microstructures of volcaniclastic deposits (petrography and SEM)
- .Injection of water in PDCs: numerical modeling
- .Volcaniclastic sedimentation in extensional basins: facies and geometries