#### Origin of Petroliferous Dolomitic Beds in the Uteland Butte Member, Lower Green River Formation, Uinta Basin, Utah\*

Federico Rueda<sup>1</sup>, Hans G. Machel<sup>1</sup>, and Michael D. Vanden Berg<sup>2</sup>

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#### **Abstract**

The lacustrine Green River Formation is an important oil-producing formation in the Uinta Basin, Utah. Of particular interest for horizontal drilling are three beds with up to 100% dolomite in the Uteland Butte member (UBM) of the lower Green River Formation. These beds are 1.5 to 8 feet in thickness and interbedded with organic-rich limestones and shales. They have up to 30% porosity but only max 0.1 mD permeability, thus forming an unconventional reservoir. However, not all dolomitized layers, of which there are several more in the UBM, are of reservoir quality. This study attempts to determine the role of dolomitization and other diagenetic processes in reservoir development, i.e., when and how were the abnormally high porosities paired with low permeabilities created. Furthermore, delineating the regional geometry of the dolomite layers is critical for the understanding the petroleum production potential. Methods of investigation include outcrop and core petrography, thin section microscopy, SEM, XRD, CL, and isotopic analyses. Deposition of the UBM took place during three transgressiveregressive cycles that were driven by climate variations. The lake level was high during cooler and wetter periods that alternated with warmer and drier periods, which led to lower lake levels from reduced fluvial input and/or evaporation. The three reservoir layers were deposited during the first of these cycles as lime muds in lacustrine littoral to sublittoral environments as intraclastic, peloidal grainstones, and silty peloidal packstones, and in shallow littoral environments as peloidal, bioturbated mudstones and wackestones. The dolomite-bearing layers are greenish to beige in hand specimen and outcrop. Crystal sizes are <15 µm and porosity is mainly intercrystal. There is no discernible relationship between dolomitization and depositional environments. Permeability is low due to irregular and commonly disconnected pore throats. In addition, post-dolomitization silicification commonly formed nodules and layers of length-slow chalcedony and equigranular quartz, as well as blocky ferroan calcite and equant to blocky calcite cement, which reduced secondary porosity and permeability. Dolomitization took place very early, i.e., almost syndepositionally, from lake water that was moderately evaporated and probably enriched in Mg during drier periods. Increased fresh water input during more humid climate periods stopped dolomitization and facilitated further deposition of lime mud layers that are now interbedded with the dolomitized beds. The origin of the microporosity - or lack thereof - is as yet undetermined.

<sup>\*\*</sup>Datapages © 2017 Serial rights given by author. For all other rights contact author directly.

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#### **References Cited**

Logan, S.K., J.F. Sarg, and M.D. Vanden Berg, 2016, Lithofacies, deposition early diagenesis, and porosity of the Uteland Butte member, Green River Formation, Eastern Uinta Basin, Utah and Colorado: Open File report 652, Utah Geological Survey.

Sibley, D.F., and J.M. Gregg, 1987, Classification of Dolomite Rock Textures: Journal of Sedimentary Petrology, v. 57/6, p. 967-975.

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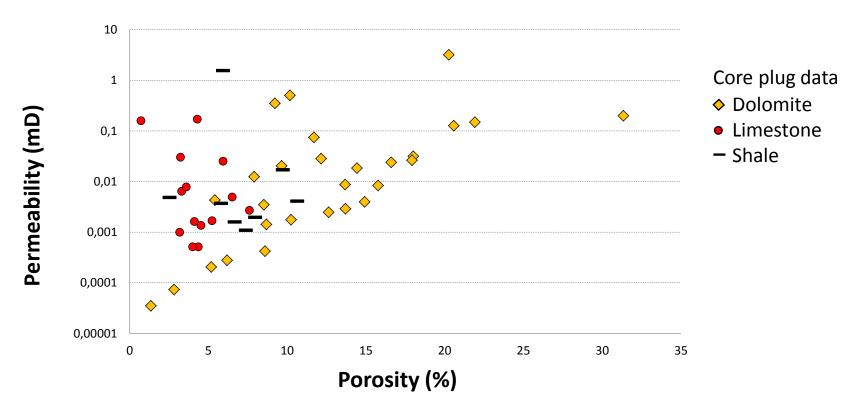




#### 1.Dolomite Reservoirs Characteristics

- 2. Objectives
- 3. Methodology
- 4. Study Area
- 5. Stratigraphy
- 6. Facies Depositional Environments
- 7. Cyclicity
- 8. Dolomite Textures
- 9. Geochemistry
- 10.Interpretations
- 11.Conclusions

#### **DOLOMITE RESERVOIR CHARACTERISTICS**



- ✓ Unconventional reservoirs
- ✓ Estimated resources: 177 MMOB / 218 BCFG
- ✓ Depth: North: ~8200 ft (~2700 m) / South: ~5500 ft (~ 1700 m)
- ✓ Dolomite bed thickness:  $^{\sim}$ **1.2 7.5 ft** ( $^{\sim}$ 0.5 m 2.5 m)
- ✓ BHT: 200°-140°F (90 60 °C)

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#### **OBJECTIVES**

1.Origin of the Uteland Butte member dolomites?

2.Implications for reservoir geology?

- 1. Dolomite Reservoirs Characteristics
- 2. Objectives

### 3. Methodology

- 4. Study Area
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#### **METHODOLOGY**

- ✓ Core + well logs + Outcrops
- ✓ Petrography, CL, SEM
- ✓ XRD, EMP

**Trace Elements** 

✓ Conventional stable isotopes

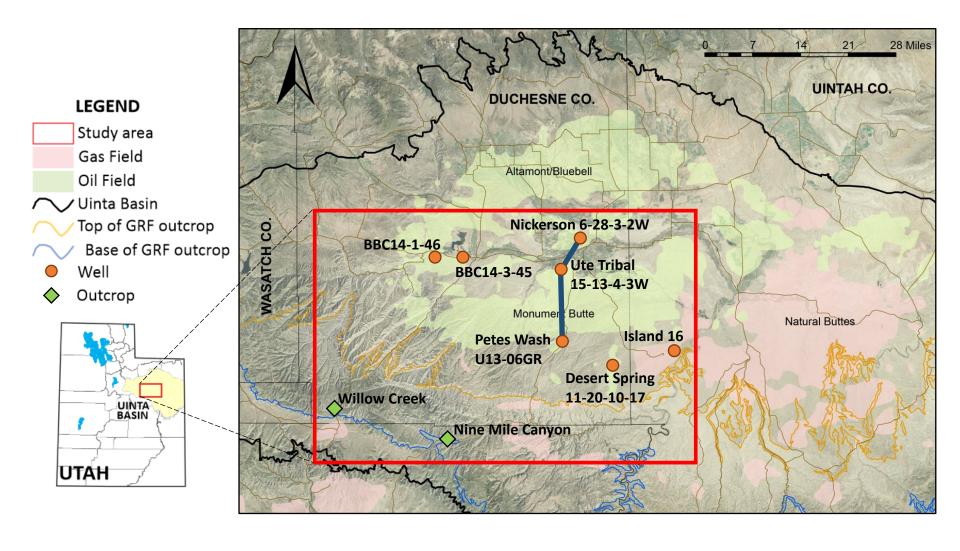
Clumped isotopes

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#### **STUDY AREA**

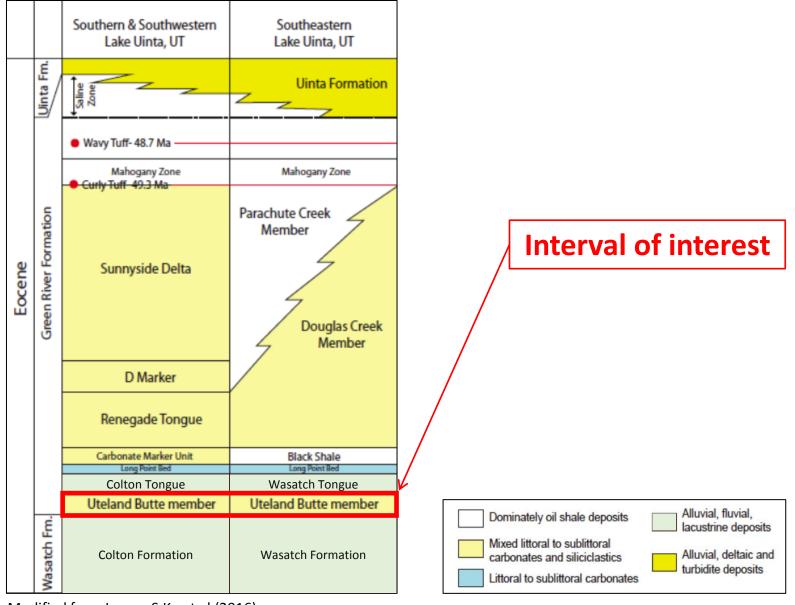


- 1. Dolomite Reservoirs Characteristics
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- 4. Study Area

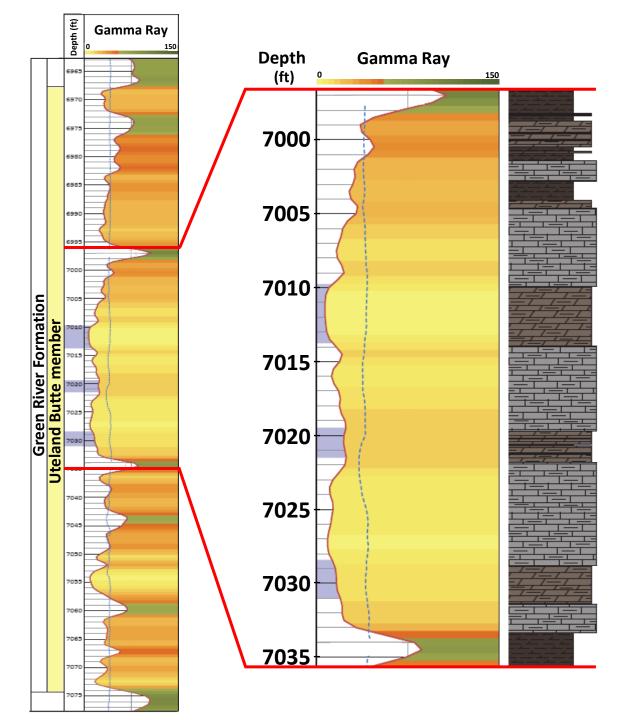
#### 5.Stratigraphy

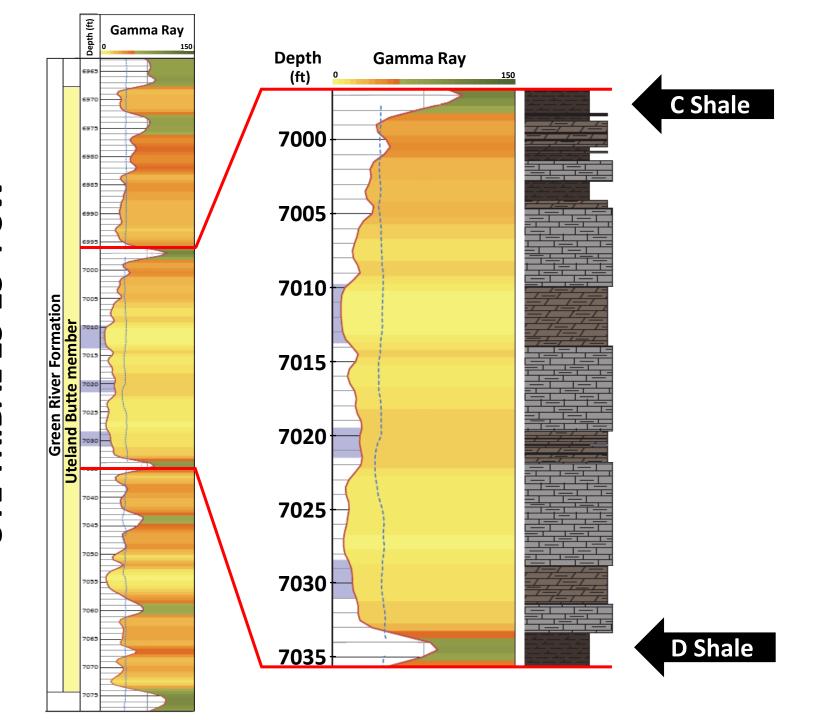
- 6. Facies Depositional Environments
- 7. Cyclicity
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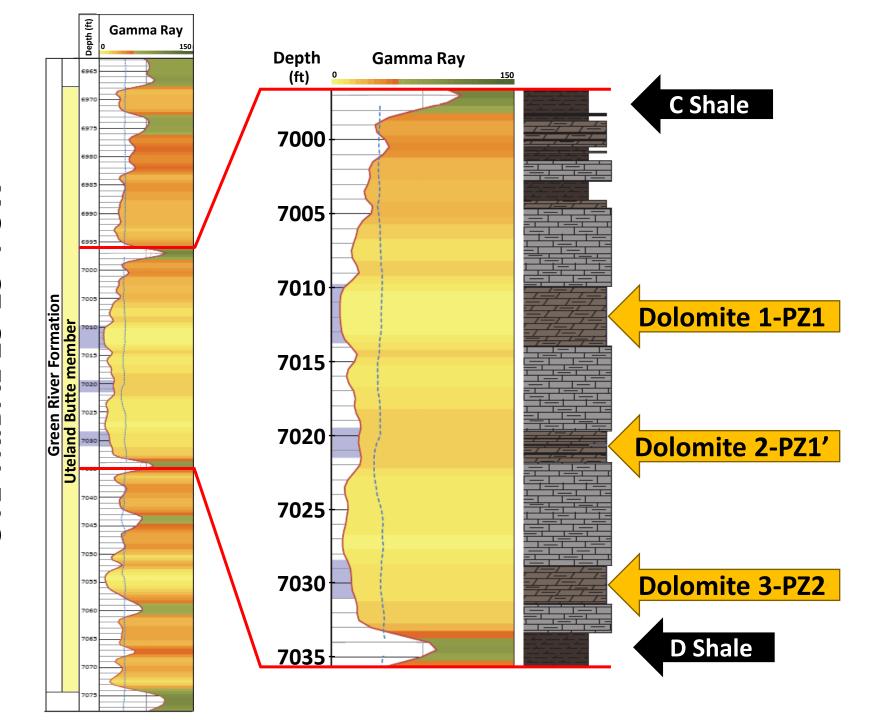
#### Generalized stratigraphic column for Uinta Basin



Modified from Logan, S.K., et al (2016)







- 1. Dolomite Reservoirs Characteristics
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#### **6.Facies - Depositional Environments**

- 7. Cyclicity
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# 7024.5 7027 7029.5 ft 1 <u>ft</u> 2 <u>ft</u>

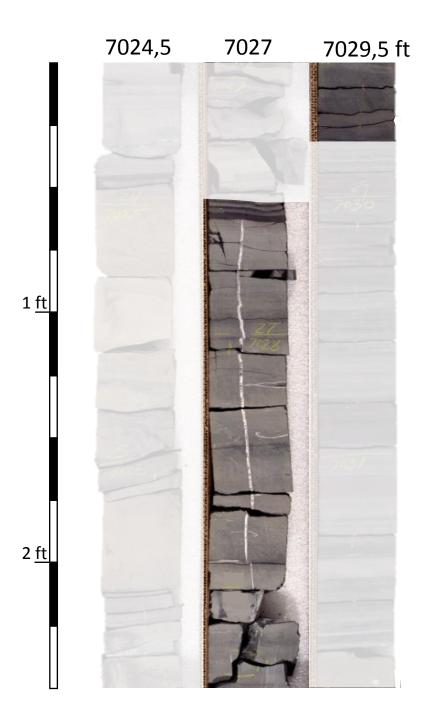
#### **FACIES**

Three basic rock types: Shale – Limestone – Dolostone



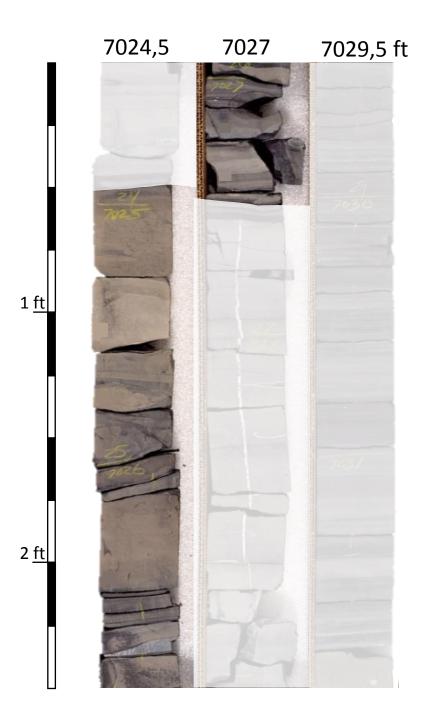
#### **FACIES**

At the base: Shale



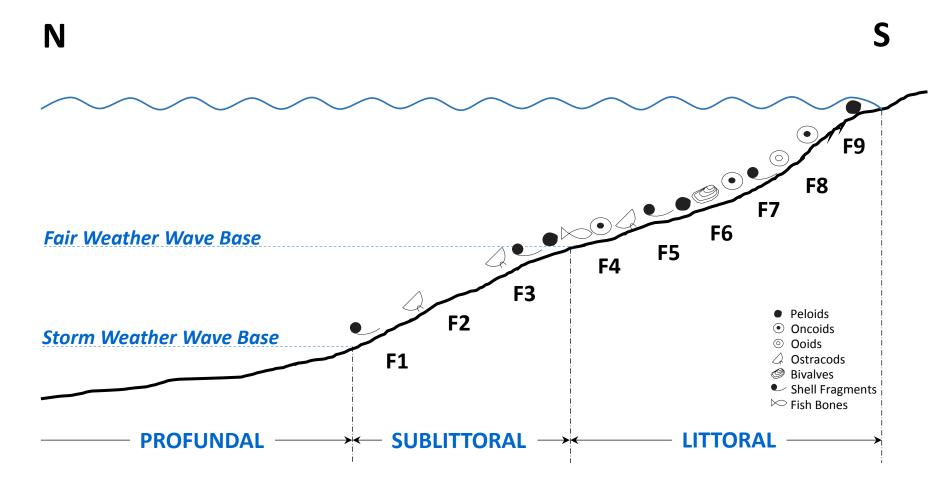
#### **FACIES**

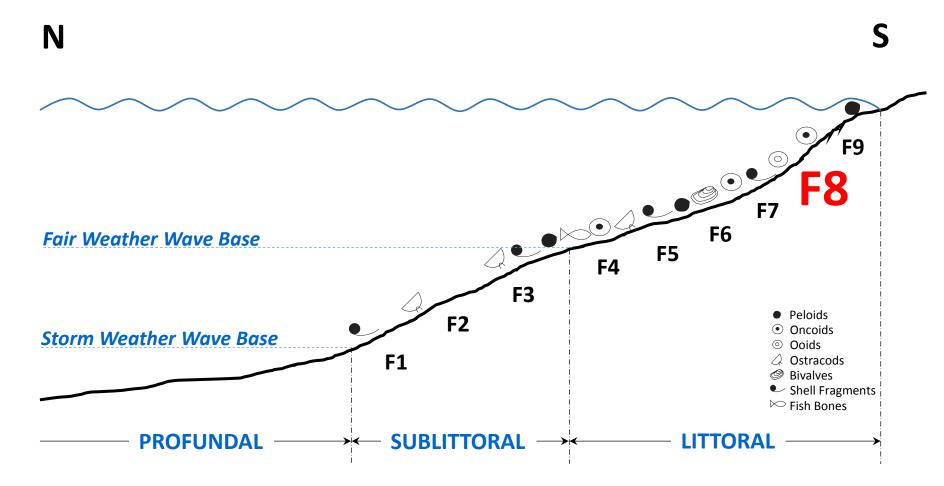
In the middle: Molluscan limestone

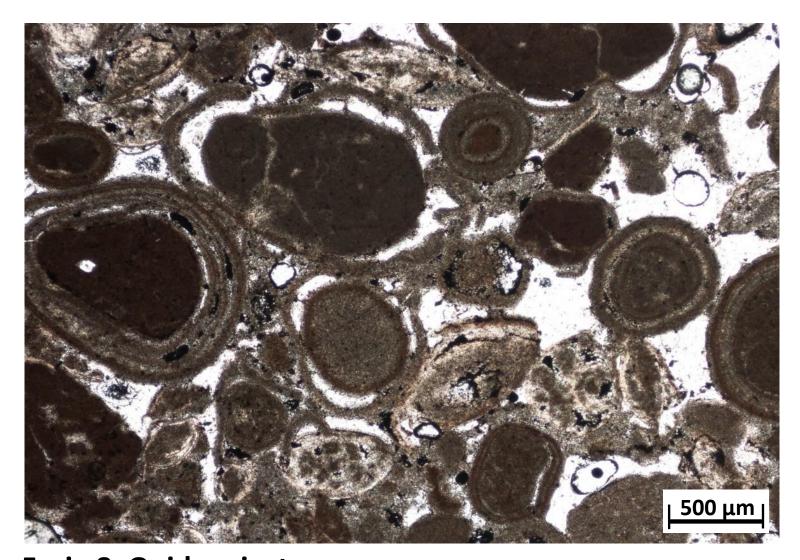


#### **FACIES**

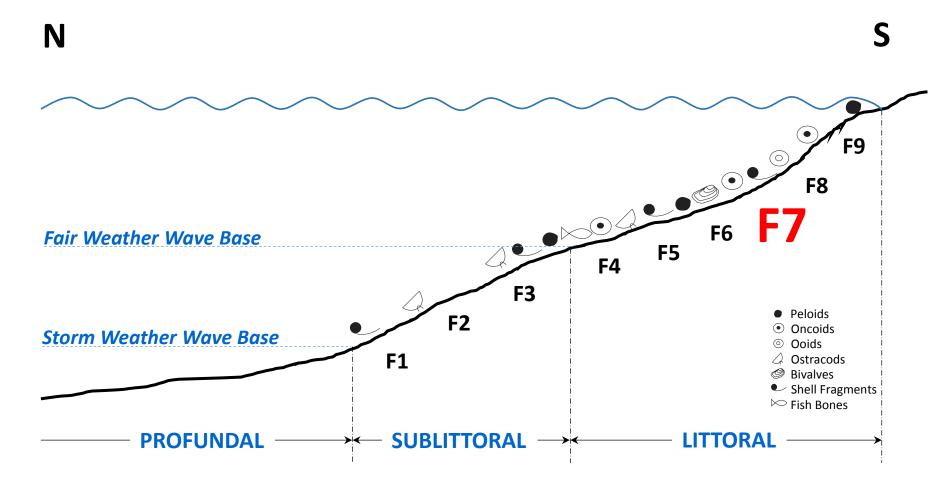
At the top: Wackestone coarsening upward to packstone grainstone

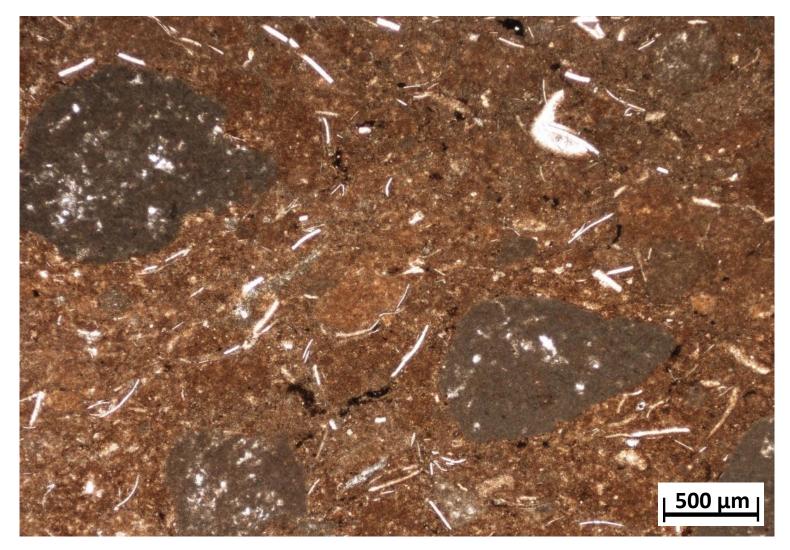






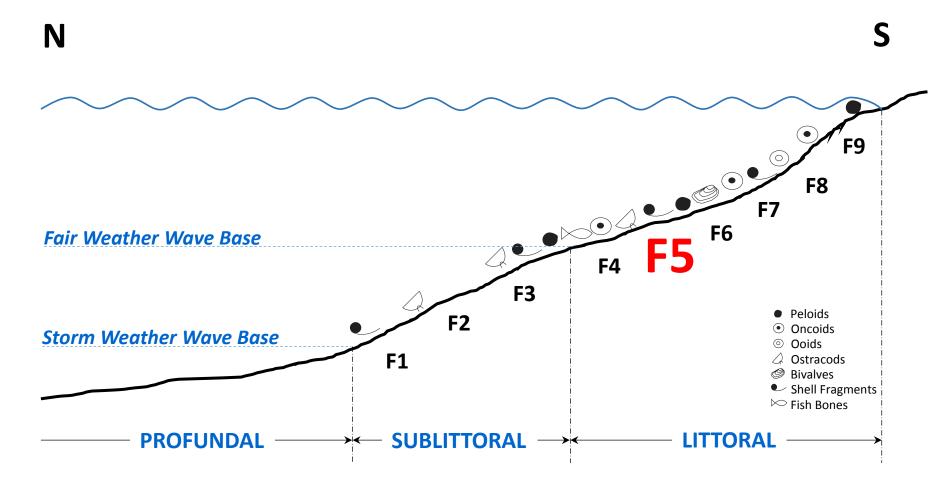
**Facie 8. Ooid grainstone**Subspherical, dolomitized, and partially dissolved ooids

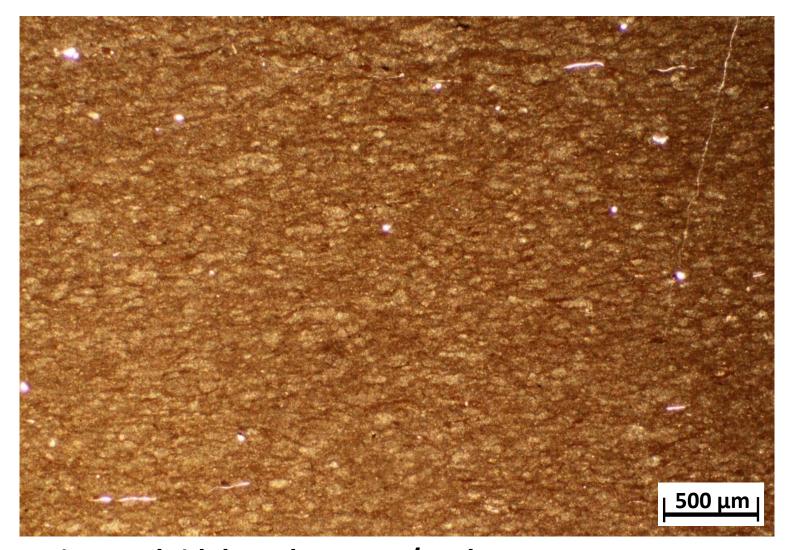




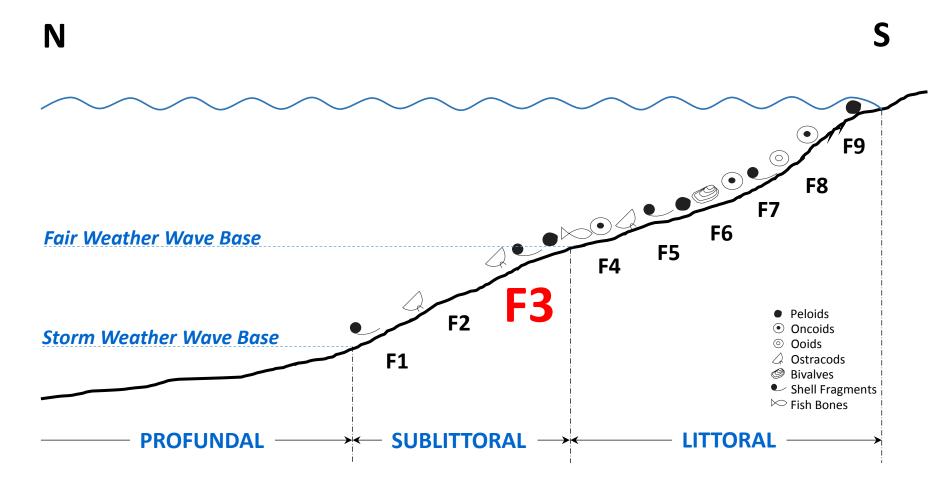
Facie 7. Intraclastic packstone

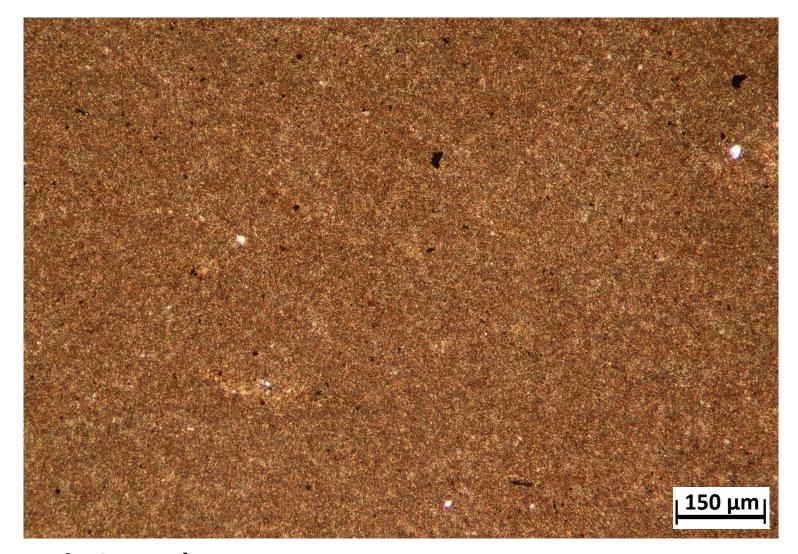
Poorly sorted, subangular, sand to gravel sized dolomitized intraclast, and molluscan shell fragments





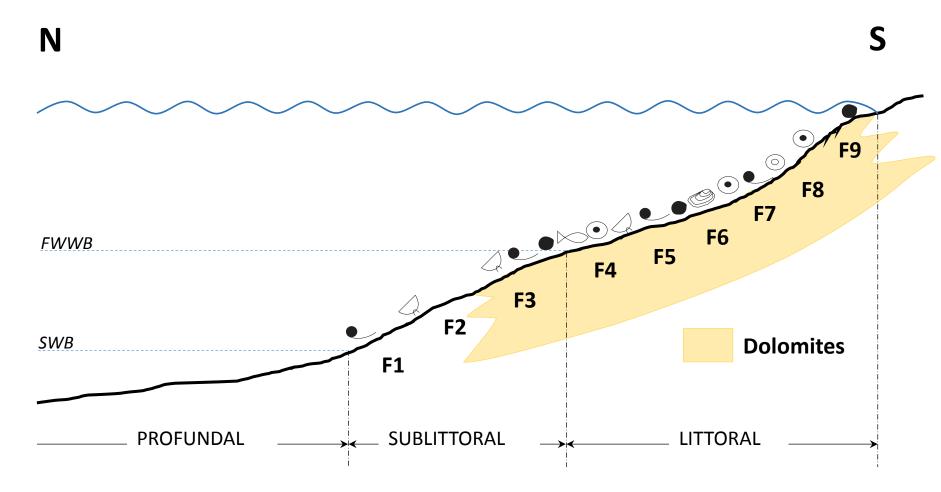
Facie 5. Peloidal wackestones/packstone
Microcrystalline dolomite matrix with abundant peloids and ostracods shell fragments





Facie 3. Mudstone

Microcrystalline dolomite matrix with scarce ostracods shell fragments

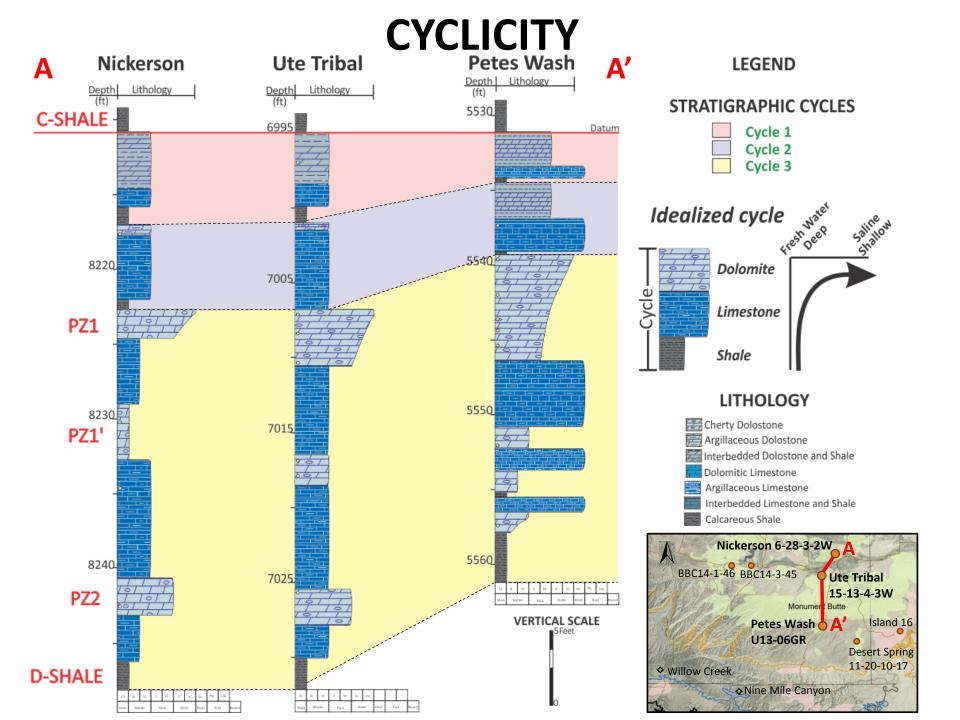


- 1. Dolomitization (yellow) cross-cut facies boundaries
- 2. Replaced micritic matrix from littoral to sublittoral environments
- 3. Excellent fabric preservation

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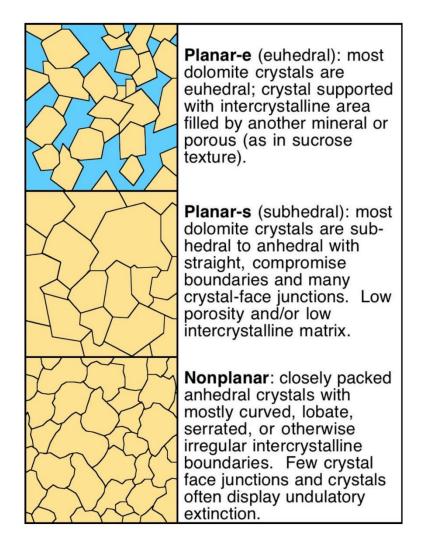


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#### 8. Dolomite Textures

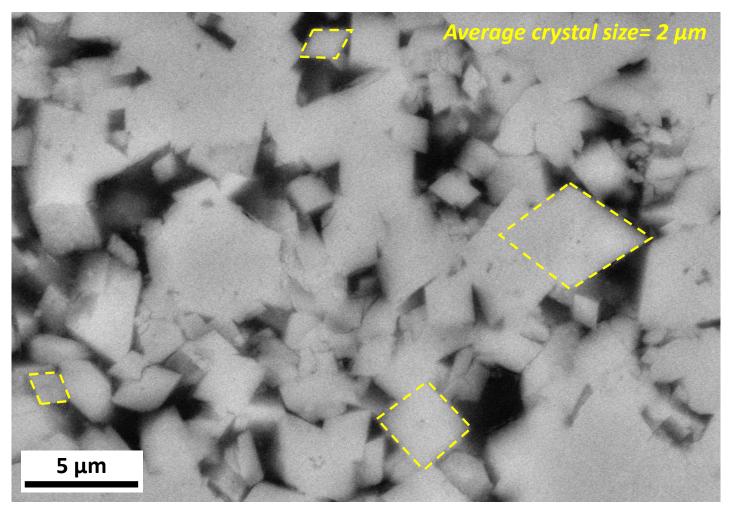
- 9. Geochemistry
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#### **TEXTURES**



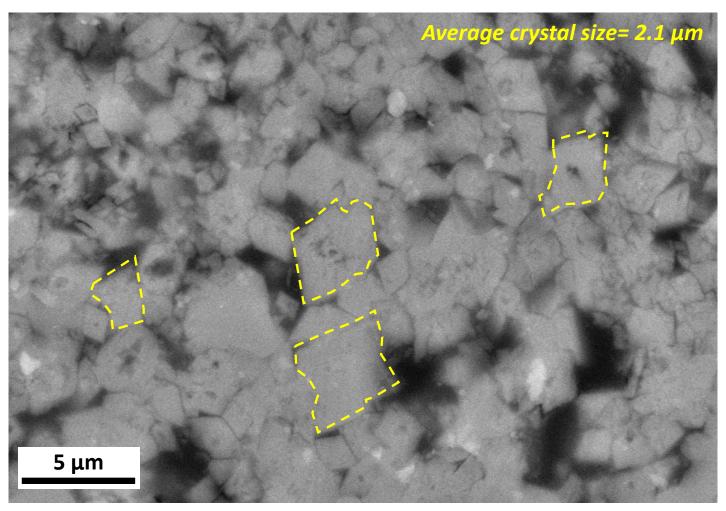
Dolomite textural classification (Sibley & Gregg, 1987)

#### **TEXTURES**



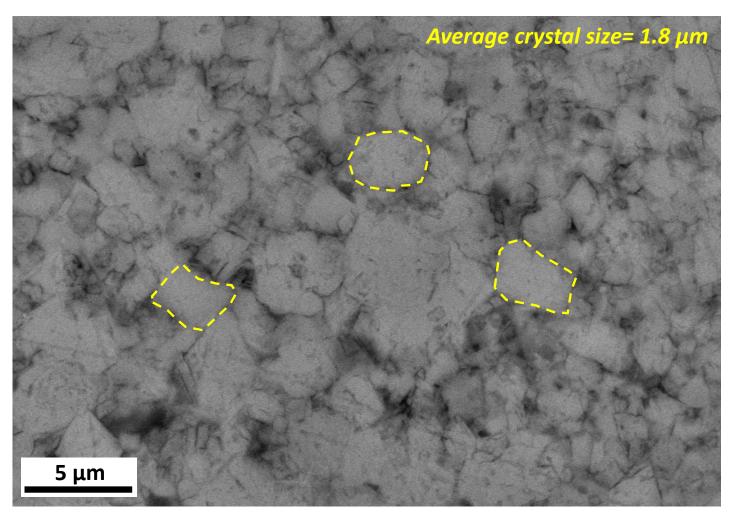
SEM image of **planar-e** dolomite texture of PZ1', well Nickerson. Porosity: **21**% (data from image analysis software)

## **TEXTURES**



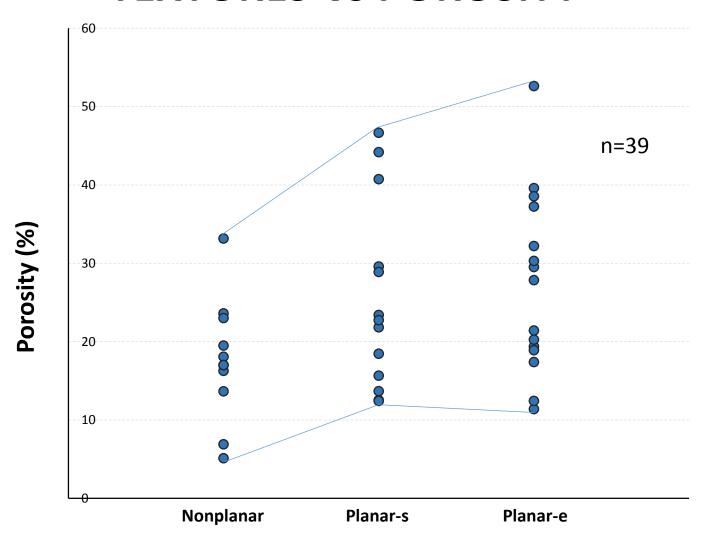
SEM image of **planar-s** dolomite texture of PZ2, well Nickerson. Porosity: **18**% (data from image analysis software)

## **TEXTURES**



SEM image of **nonplanar** dolomite texture at the base of PZ1, well Ute Tribal. Porosity: **14%** (data from image analysis software)

## **TEXTURES Vs POROSITY**



High porosity values (data from image analysis software) are associated with planar-e to planner-s textures

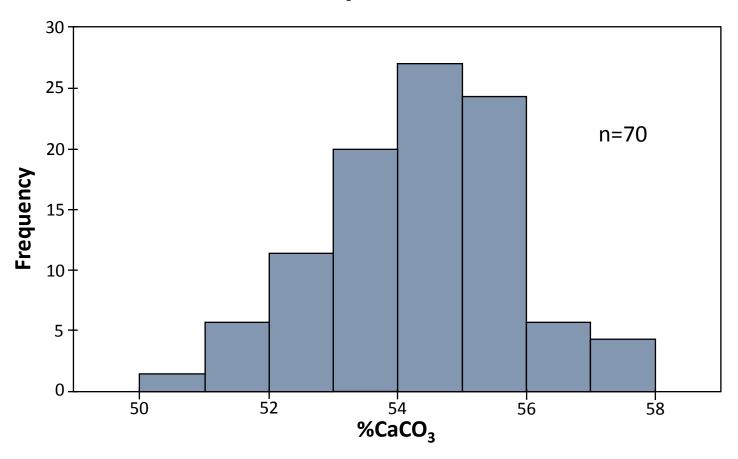
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## **STOICHIOMETRY**

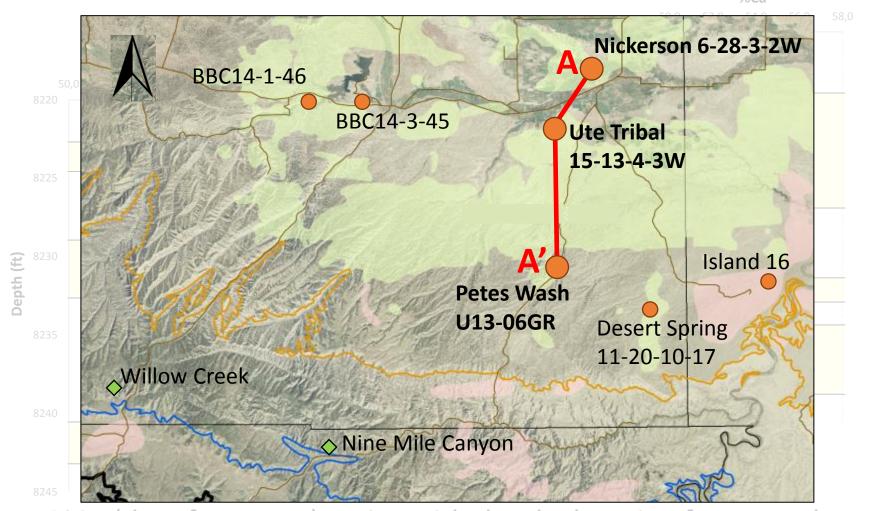
XRD / All wells



**%CaCO<sub>3</sub>** dolomite distribution suggests changes in lake water chemistry and/or variable degrees of recrystallization

## **%CaCO<sub>3</sub> Vs DEPTH** XRD data, A-A'

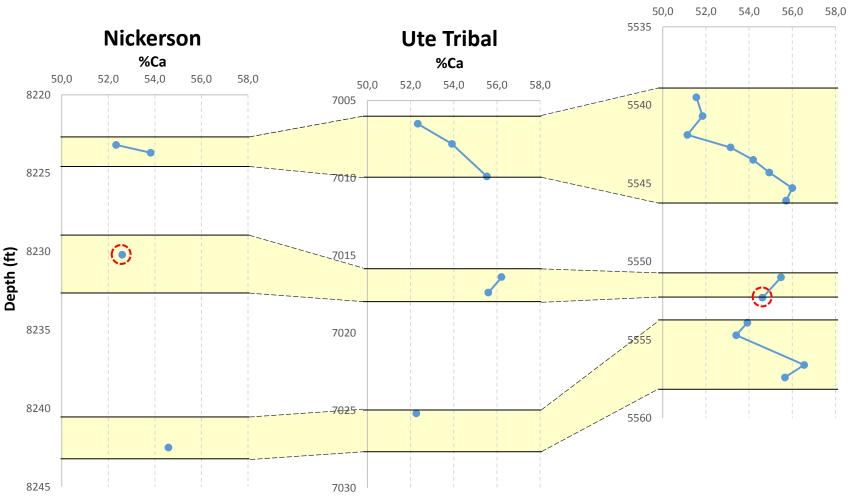
Petes Wash %Ca



**%Ca** (data from XRD) varies with depth changing from nearly stoichiometric dolomite (top) to calcium rich dolomite (bottom)

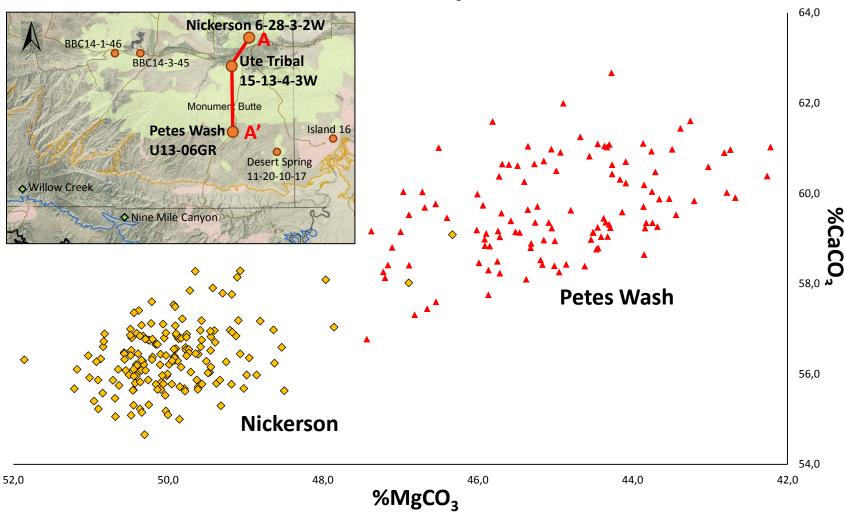
%CaCO<sub>3</sub> Vs DEPTH XRD data, A-A'





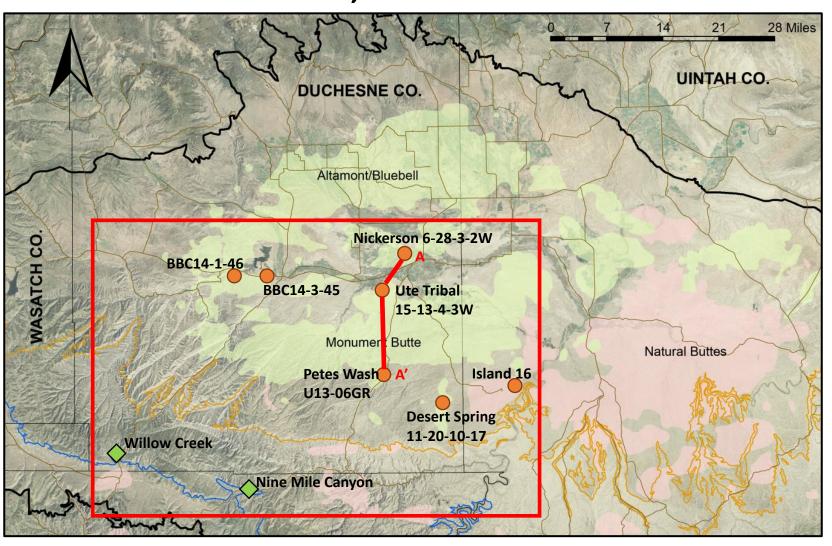
**%Ca** (data from XRD) varies with depth changing from nearly stoichiometric dolomite (top) to calcium rich dolomite (bottom)

## %CaCO<sub>3</sub> REGIONAL TREND EMP data / PZ1'

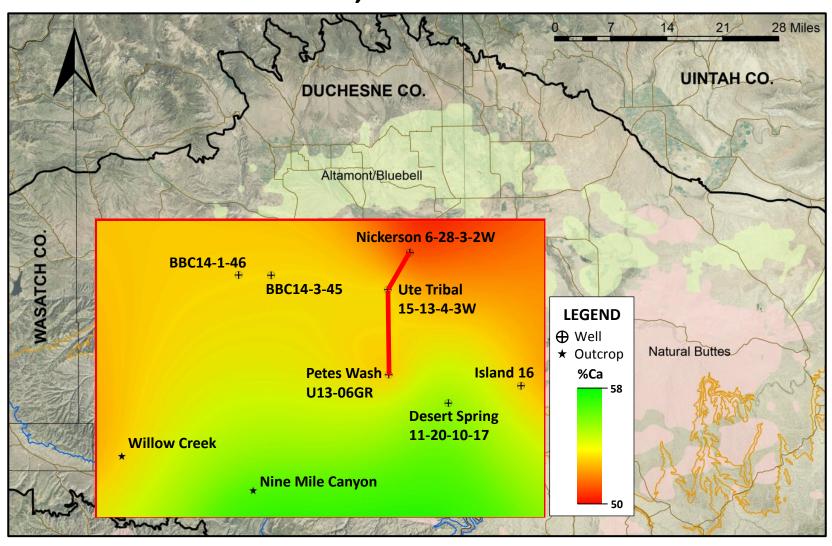


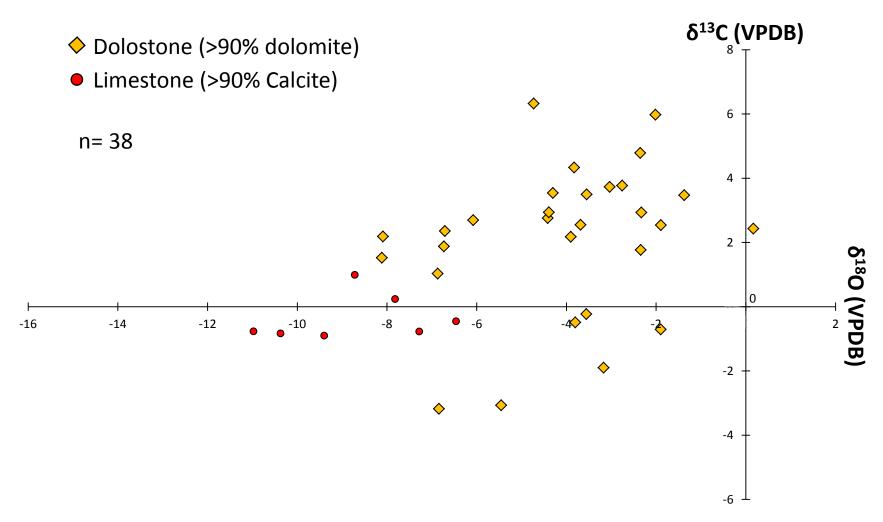
**Ca:Mg ratio** (data from EMP) exhibits a regional trend N-S regardless of depth (PZ1')

## **%CaCO<sub>3</sub> REGIONAL TREND**XRD Data, all wells for PZ1'

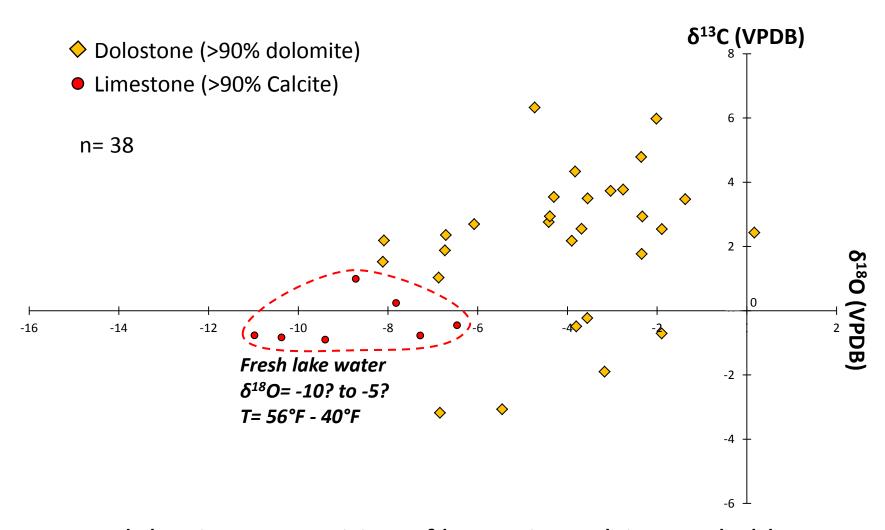


## **%CaCO<sub>3</sub> REGIONAL TREND**XRD Data, all wells for PZ1'

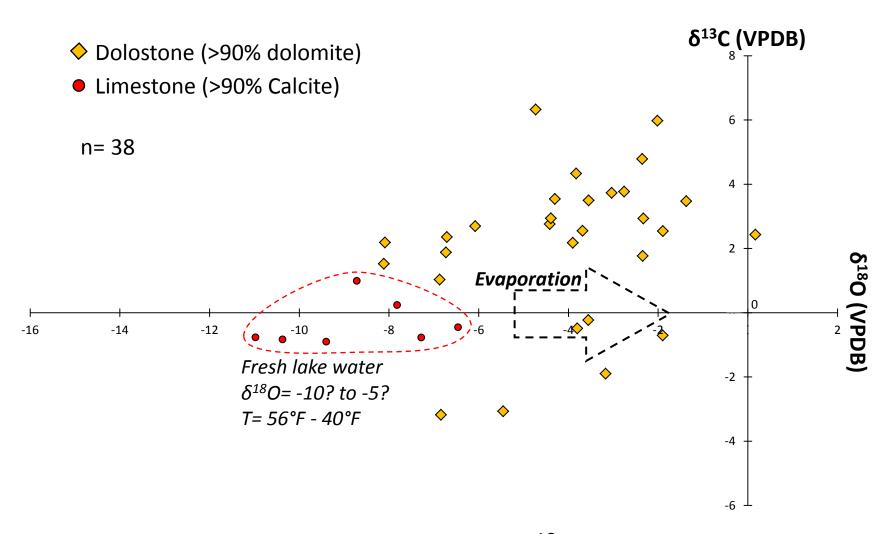




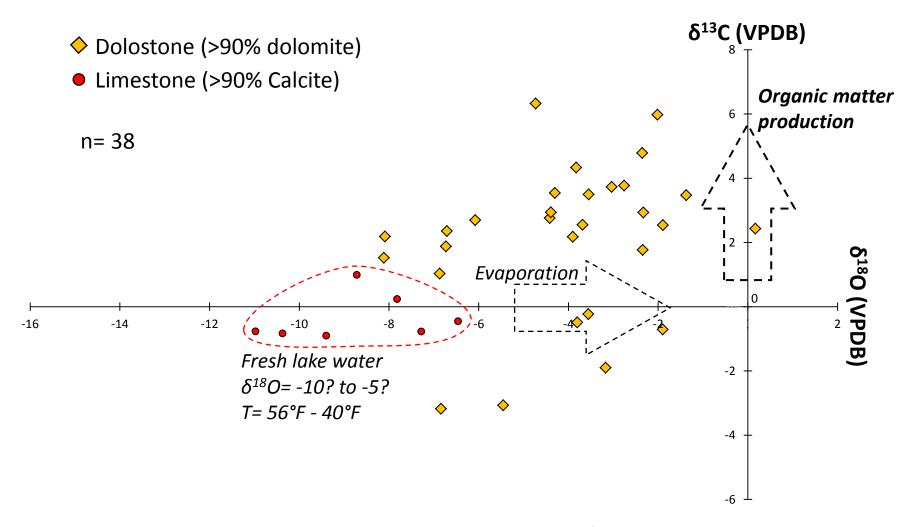
<sup>13</sup>C and <sup>18</sup>O of dolomite for PZ1, PZ1', and PZ2 and interbedded molluscan limestones



Pre-dolomite composition of lacustrine calcite, probably representing lake water chemistry



Evaporation of lake water increased <sup>18</sup>O dolomite values



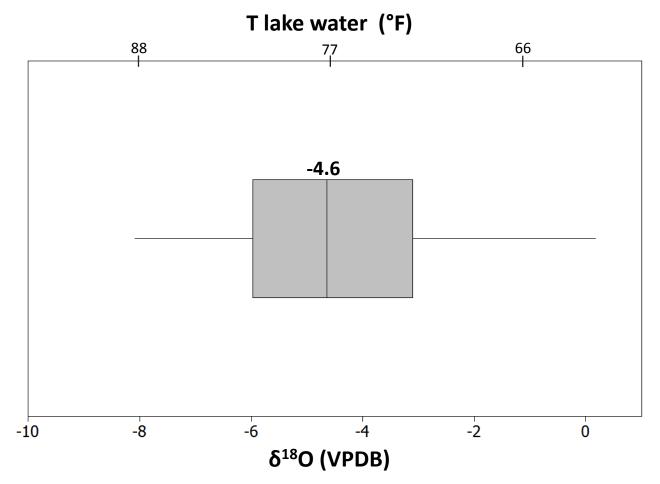
High organic matter production increased <sup>13</sup>C dolomite values

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## TEMPERATURE OF DOLOMITIZATION

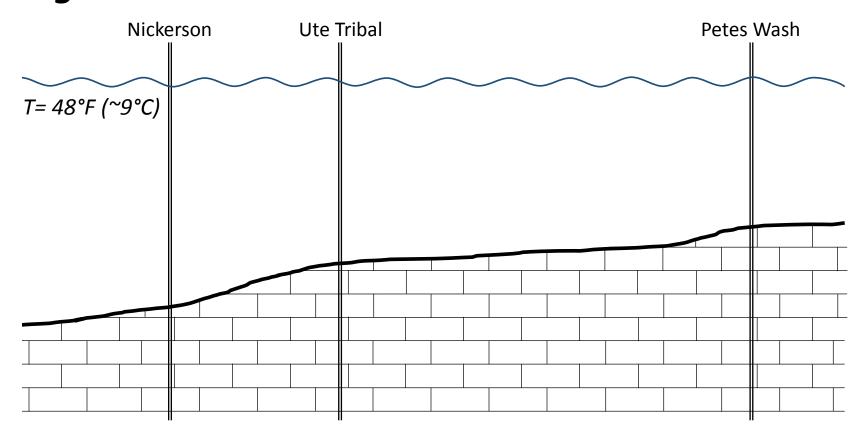


 $\delta^{18}$ O dolomite: -3.1 to -5.9

 $\delta^{18}$ O lake water: -10? to -5?

Calculated lake water temperature: 66 °to 88°F

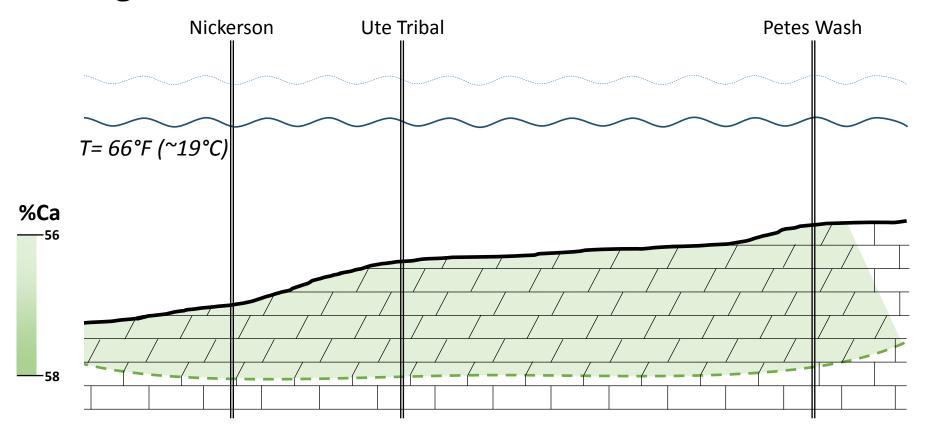
# DOLOMITIZATION DRIVEN BY CLIMATE CHANGE Stage 0:



Fresh water input - humid climate - favored deposition of wackestones/floatstone with bivalves, gastropods, and ostracods

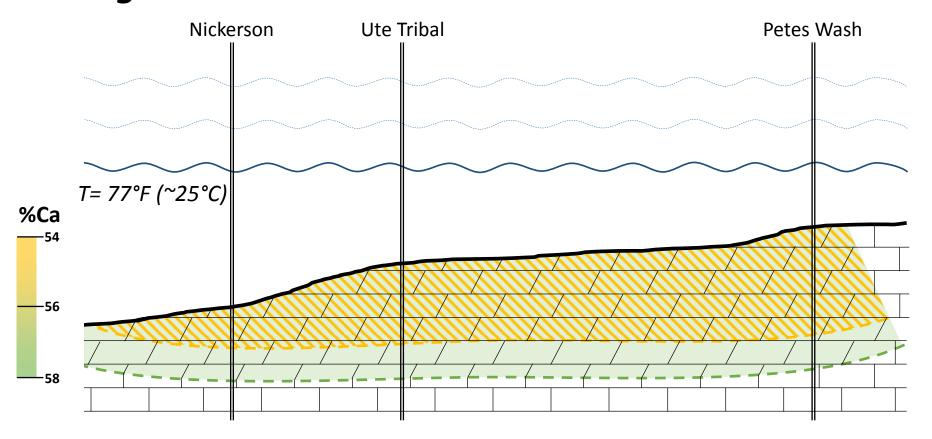
#### DOLOMITIZATION DRIVEN BY CLIMATE CHANGE

## Stage 1:



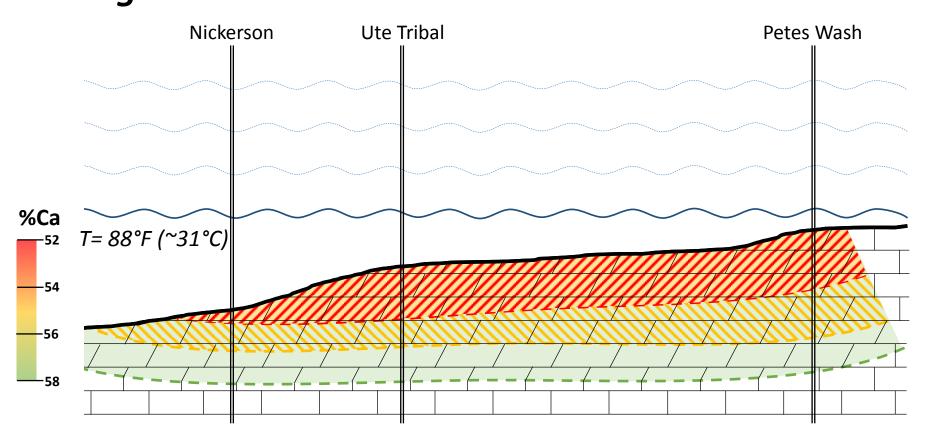
Climate changes (high evaporation + reduced fresh water input) changed lake water chemistry that promoted replacement of micrite (matrix) by high calcium dolomite

# DOLOMITIZATION DRIVEN BY CLIMATE CHANGE Stage 2:



Higher evaporation favoured dolomite recrystallization (*stage 1*) and/or micrite replacement by calcium dolomite

# DOLOMITIZATION DRIVEN BY CLIMATE CHANGE Stage 3:



At highest temperature, dolomite recrystallization (stages 1-2) persisted developing nearly stoichiometric dolomite (top). This cycle was repeated several times

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## **Origin of the Uteland Butte member dolomites**

#### **Dolomitization was driven by climate changes:**

- √ Took place in cyclically warmer climates
- ✓ Enhanced evaporation
- ✓ Reduced fresh water input
- ✓ Shallower lake water levels

## Implications for reservoir geology

- ✓ Dolomite planar-e and planar-s textures are slightly correlated with higher porosity values
- ✓ Spatial distribution of dolomite textures?
- ✓ Spatial distribution of %Ca vs Porosity?



#### REFERENCES

Logan, S. K.; Sarg, J. F.; & Vanden Berg, M. D. (2016) Lithofacies, deposition early diagenesis, and porosity of the Utelan Butte member, Green River Formation, Eastern Uinta Basin, Utah and Colorado. Open-file report 652. Utah Geological Survey

Sibley, D. F.; & Gregg, J. M. (1987). Classification of Dolomite Rock Textures. *Journal of Sedimentary Petrology, 57* (6), 967-975