

Impact of Hydrothermal Fluid Flow on Conventional and Unconventional Carbonate-Rich Reservoirs, Midcontinent USA*

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Abstract

Hydrothermal fluid flow in Ordovician, Mississippian, and Pennsylvanian reservoir rocks of the Midcontinent is partially responsible for generating the porosity in those reservoirs (Ramaker et al. 2014), as well as hydrocarbon migration and local thermal maturation. The hydrothermal fluid flow occurred in three late stages (King, 2013). Fluid flow was controlled by stratigraphic discontinuities, fault and fracture systems, and temperature-controlled density differences. These controls are critical for localization of some of the reservoirs in the Midcontinent.

Ordovician, Mississippian, and Pennsylvanian strata in Kansas all show fracturing, megaquartz, silica dissolution, carbonate dissolution, baroque dolomite, MVT minerals, and calcite after stylolitization. Cathodoluminescence petrography, fluid inclusions, $^{87}\text{Sr}/^{86}\text{Sr}$, and $\delta^{18}\text{O}$ indicate hydrothermal fluid flow affected Ordovician-through-Pennsylvanian stratigraphic units. The history is simplified into three stages of hydrothermal fluid flow (86-144°C). All show evidence of thermal pulses, suggesting tectonic valving.

Fluid inclusion data indicate *Stage 1* was from brines near seawater salinity, interpreted as connate fluids migrating out of the Anadarko and/or Arkoma basins, likely during the Pennsylvanian or Early Permian. Fluids were associated with a separate gas phase, and they precipitated megaquartz.

Stage 2 led to precipitation of baroque dolomite. Fluid inclusion data indicate high salinities (20 wt. %) and $^{87}\text{Sr}/^{86}\text{Sr}$ indicate advective fluid flow across long distances. $\delta^{18}\text{O}$ data indicate the Ordovician-Mississippian section acted as an aquifer in vertical communication, leading to warmer fluids and preferred fluid flow toward the top of the Mississippian. The shale-rich Pennsylvanian section acted as a leaky confining unit. This stage of fluid flow was associated with oil migration and likely occurred late in the Permian or after.

Stage 3 of hydrothermal fluid flow was complex and is recorded by calcite cements. Spatial variation of $\delta^{18}\text{O}$ and $^{87}\text{Sr}/^{86}\text{Sr}$ indicate cessation of advective fluid flow and initiation of localized vertical fluid flow, possibly directly out of the basement. Comparison of fluid inclusion temperature and salinity data to modern reservoir conditions indicates that this stage clearly predates the current fluid flow and thermal regime, but played a part in evolution of the reservoir system.

The first two stages of hydrothermal fluid flow are associated with fracturing, silica dissolution, and carbonate dissolution. Much of the porosity, typically assumed to originate from subaerial weathering, may have been generated by these late hydrothermal fluids. The fluids followed fracture systems and were concentrated along the tops of hydrothermal aquifers by stratigraphic discontinuities and temperature-controlled density differences. This model for hydrothermal porosity formation helps to explain the spatial variation in reservoir quality in the Mississippian and leads to an enhanced model for locating the best producers. The Mississippian to Cambrian-Ordovician section acted as a regional aquifer and Pennsylvanian acted as a leaky confining unit. In the regional aquifer cross-formational connections allowed lower density, warmer fluids to concentrate at the top of the aquifer. Reservoir porosity is partially controlled by hydrothermal fluid migration, enhancing the porosity in areas where fractures and faults led to preferred hydrothermal fluid flow, especially close to the top of the regional aquifer. Better porosity is related to late structure and stratigraphic control on fluid flow.

For the third stage of hydrothermal fluid flow, further study is necessary. A driver could be localized faulting and fracturing associated with Laramide or other deformation. Fracturing and vertical fluid flow clearly could have had an impact on late hydrocarbon migration. As these systems are highly localized, their identification may be key in predicting location of some of the best producers.

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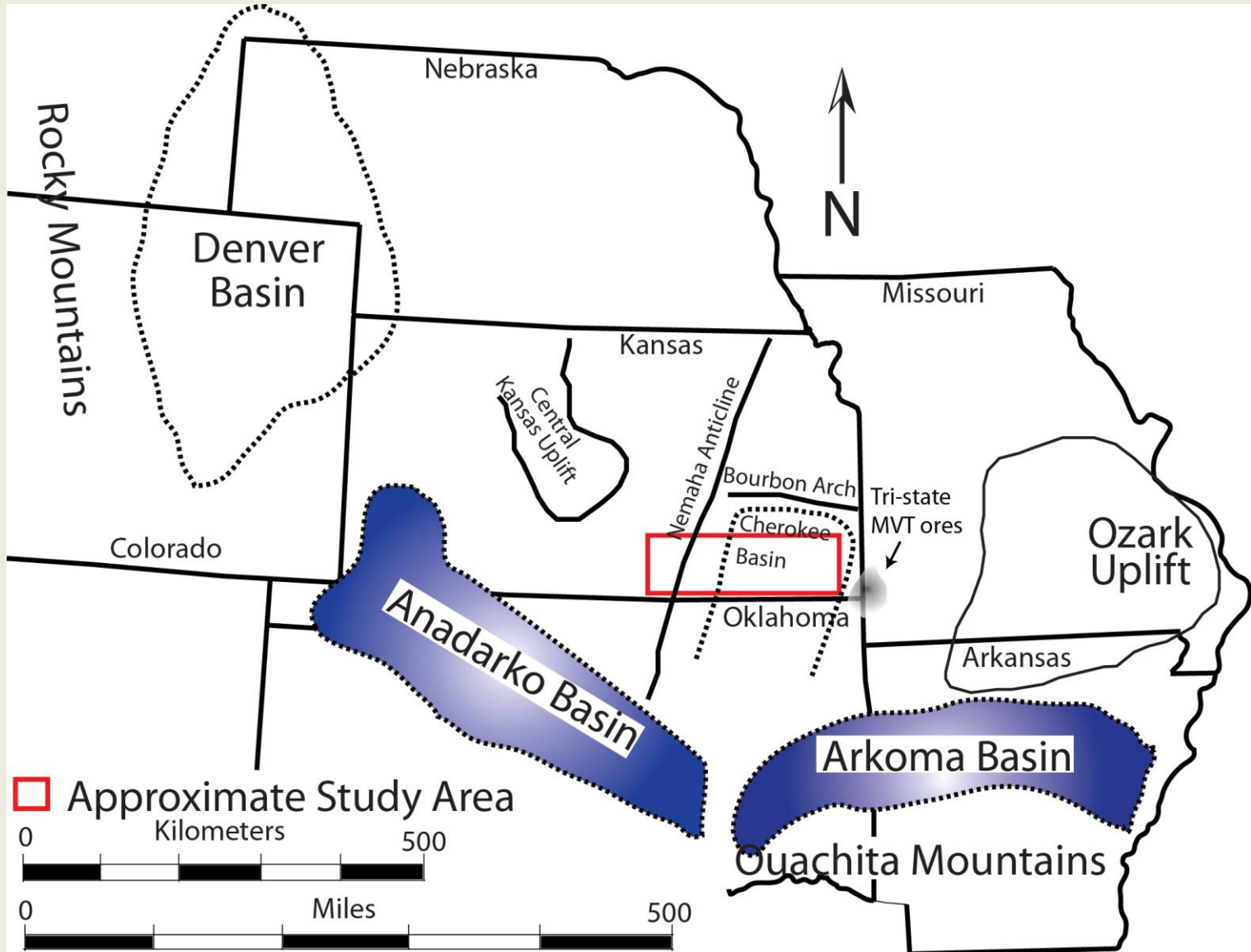
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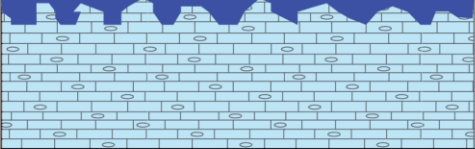
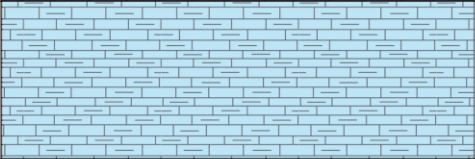

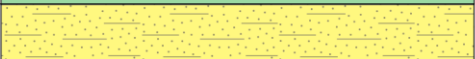
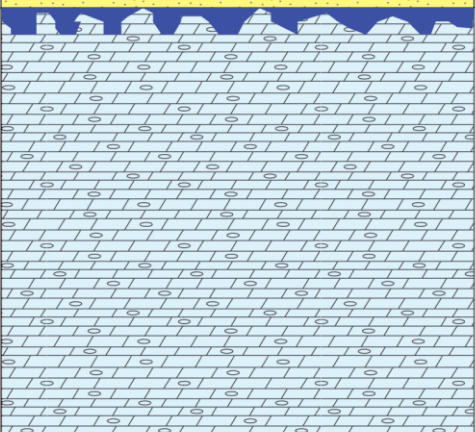


Purpose

- Study of basement through Pennsylvanian in Midcontinent, USA, shows three-stage evolution of hydrothermal system
- Concentration of hydrothermal alteration immediately below unconformities caused by density-controlled advective fluid flow
- Study demonstrates that porosity associated with unconformities is much later than development of meteoric karst and caused by a predictable fluid distribution in hydrothermal aquifers
















Study Area



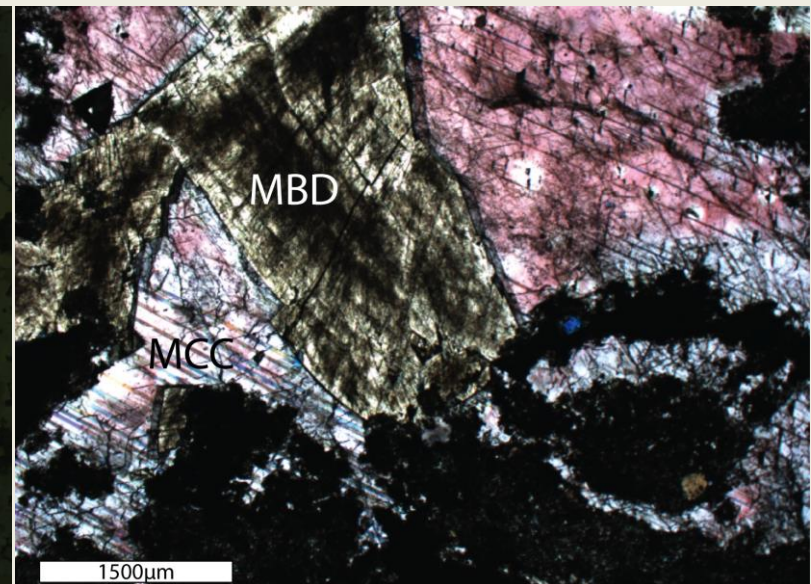
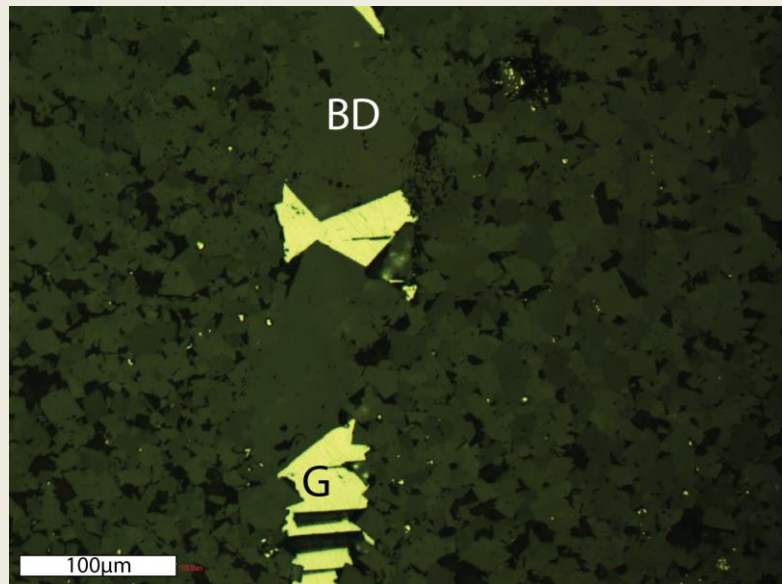
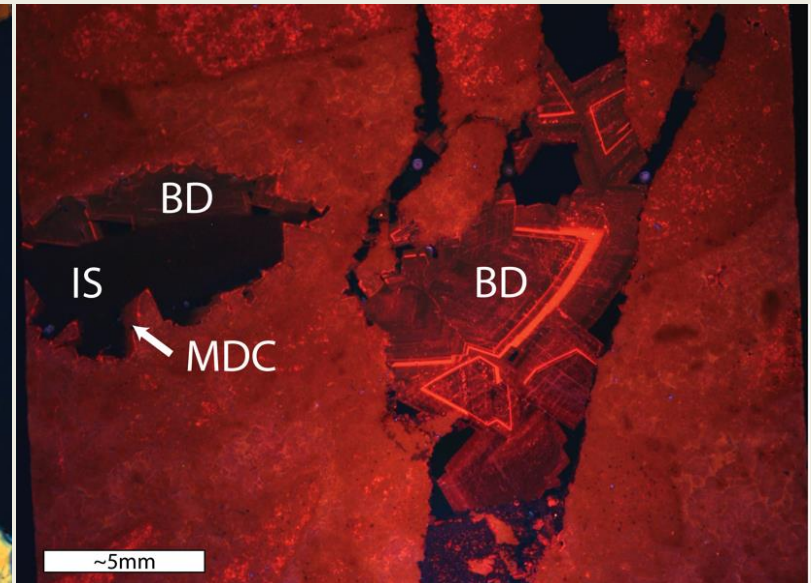
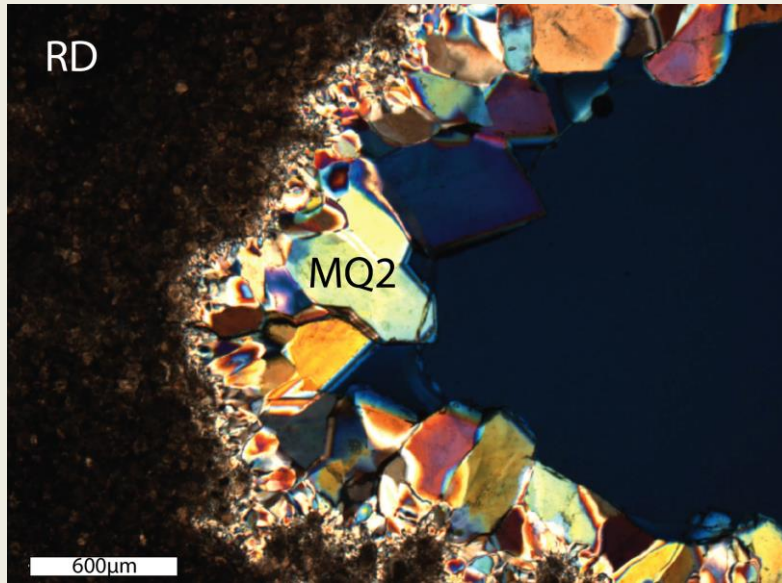
Stratigraphy

Sys.	Relevant Unit	Lithology	Depths
Perm.	Sumner Group		590-1197 ft
Penn.	Various Ls & Sh		
Mississippian	Upper Mississippian Series		3658 ft
	Lower Mississippian Series		3891 ft
D-M	Chattanooga Shale		4063 ft
Cambrian-Ordovician	Simpson Group		
	Arbuckle Group		4165 ft
			
pЄ	Granite Basement		5164 ft

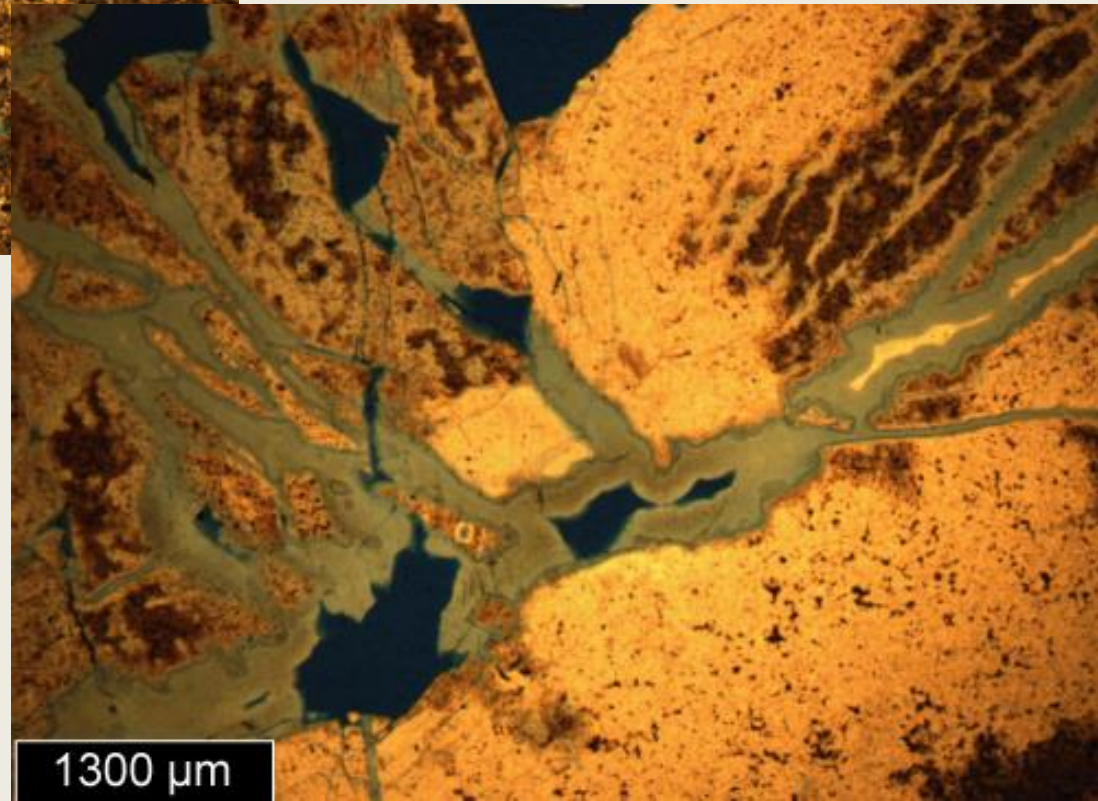
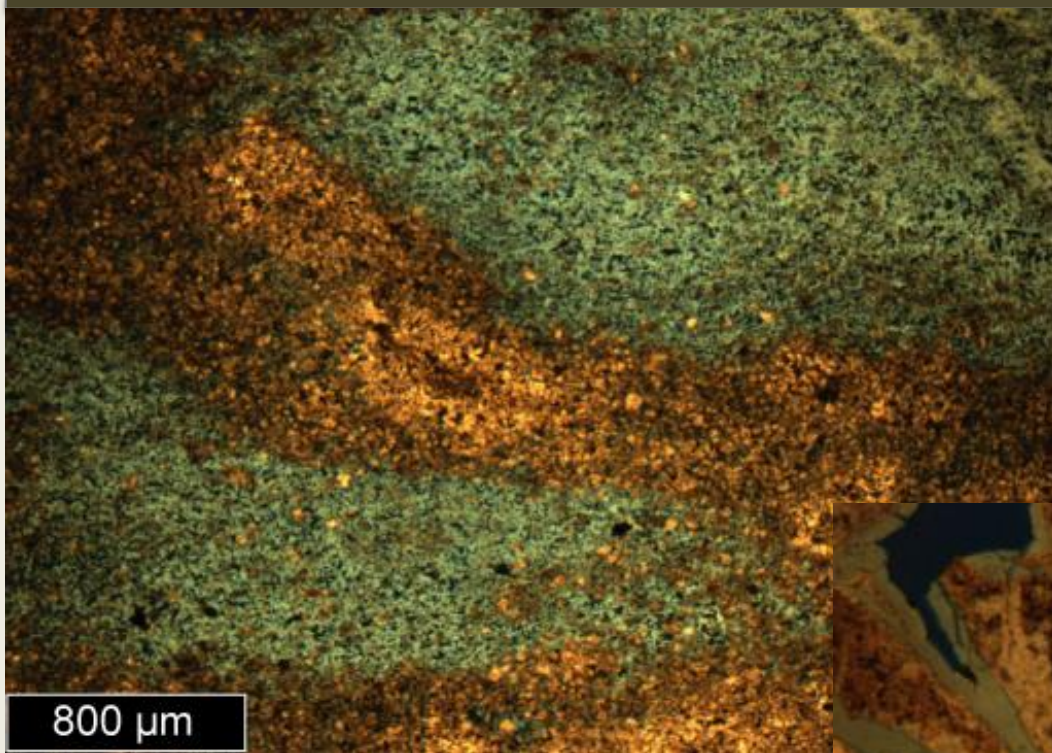
Arbuckle Group Late Paragenesis

Diagenetic Events	Early Stage	Late Stage
1. Original Deposition 2-3-4. Early Dissolution		
2-3-4. Replacement Dolomite (RD) 2-3-4. Anhydrite (A)		
5. Early Dolomite Cements (EDC) 6. Silicification (RC)		
7. Chalcedony (Ch) 8. Karsting (Carbonate Dissolution)		
9. Brecciation and collapse features 10-11. Middle Dolomite Cements (MDC)		
10-11. Pyrite (P) 12. Megaquartz 1 (MQ1)		
13. Internal Sediment (IS) 14-15. Stylolitization & emanating fractures		
14-15. Fracturing (F) 16-17. Silica Dissolution		
16-17. Carbonate Dissolution 18. Megaquartz Cement 2 (MQ2)		
19. Baroque Dolomite (BD) 20. Petroleum Migration		
21-22-23. Galena (G) 21-22-23. Sphalerite (S)		
21-22-23. Calcite Cement (CC)		

Late-Stage Paragenesis associated with fractures and porosity enhancement

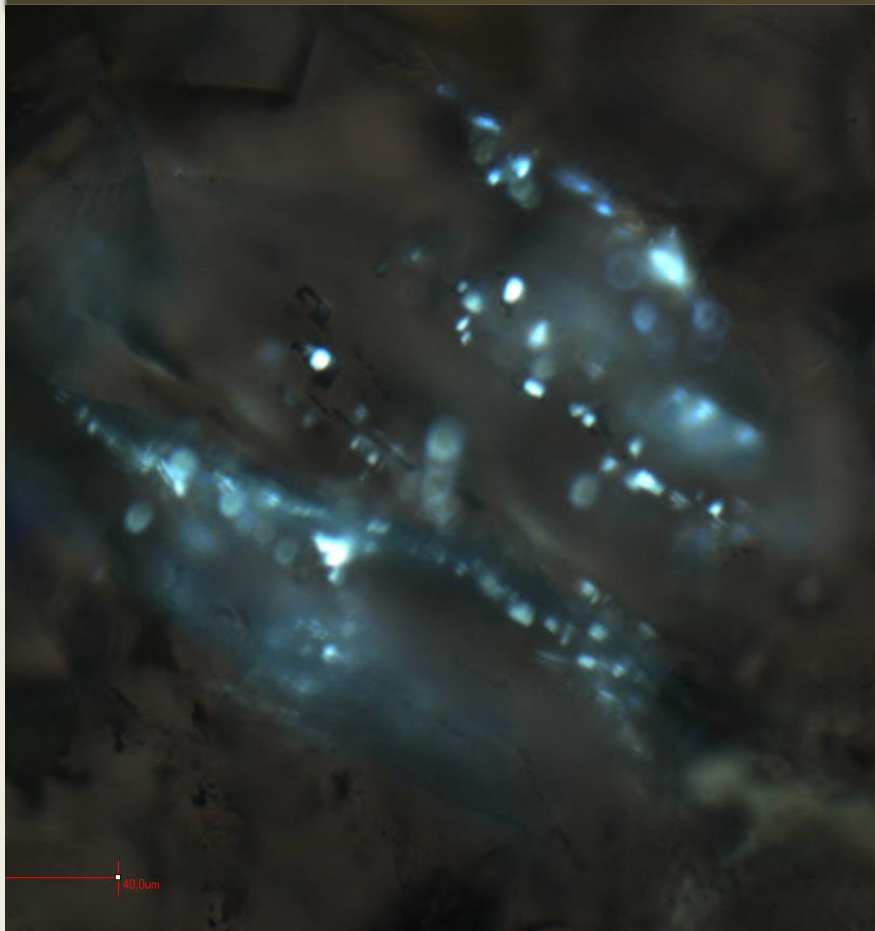


Porosity Enhancement in Silica Phases

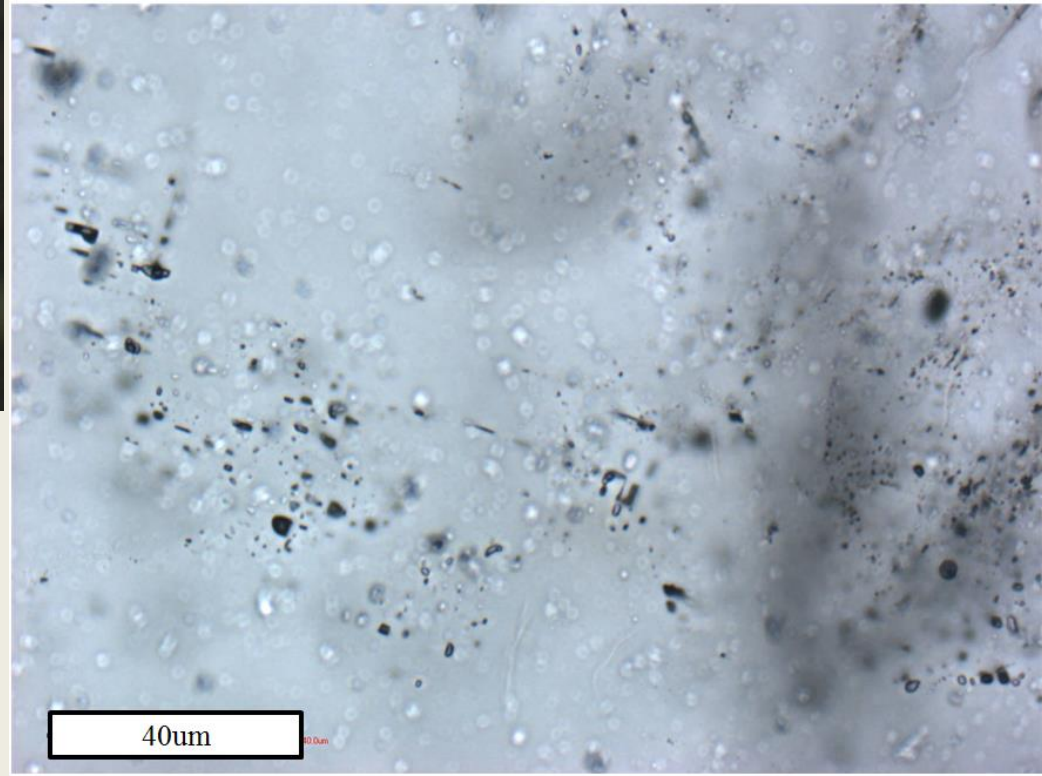


Late Mineral Phases - Hydrocarbon Migration

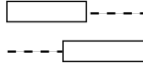

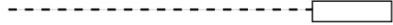
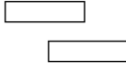


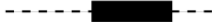






← Oil during baroque dolomite



Gas during megaquartz →



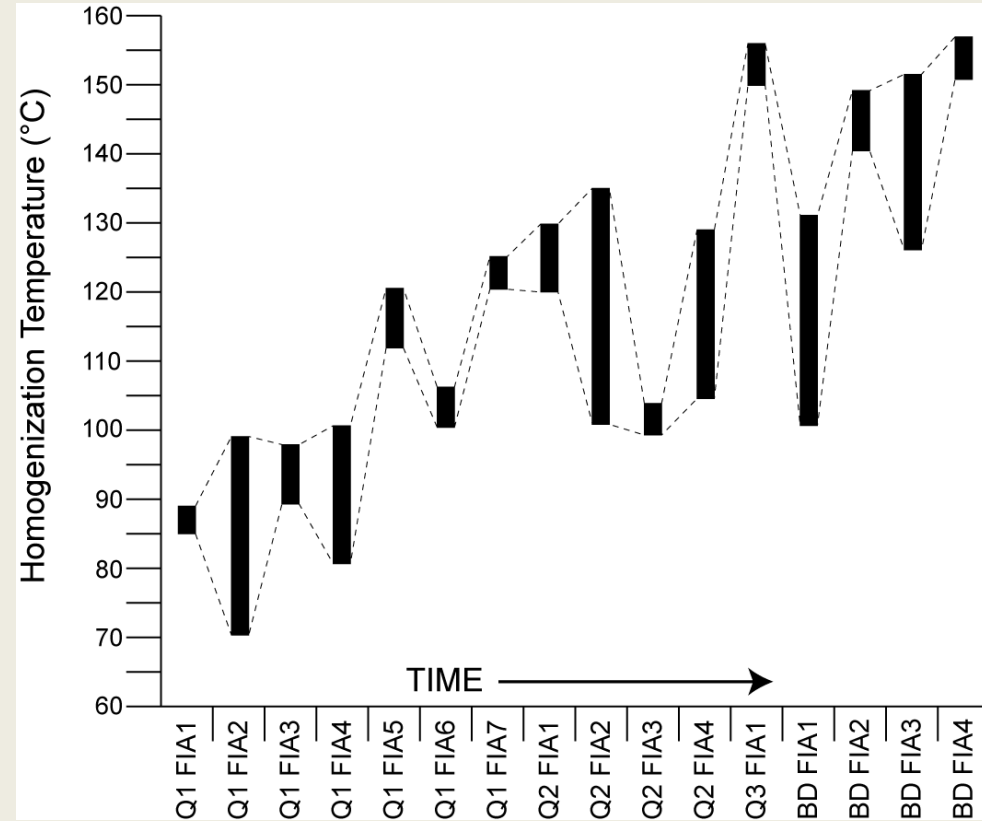
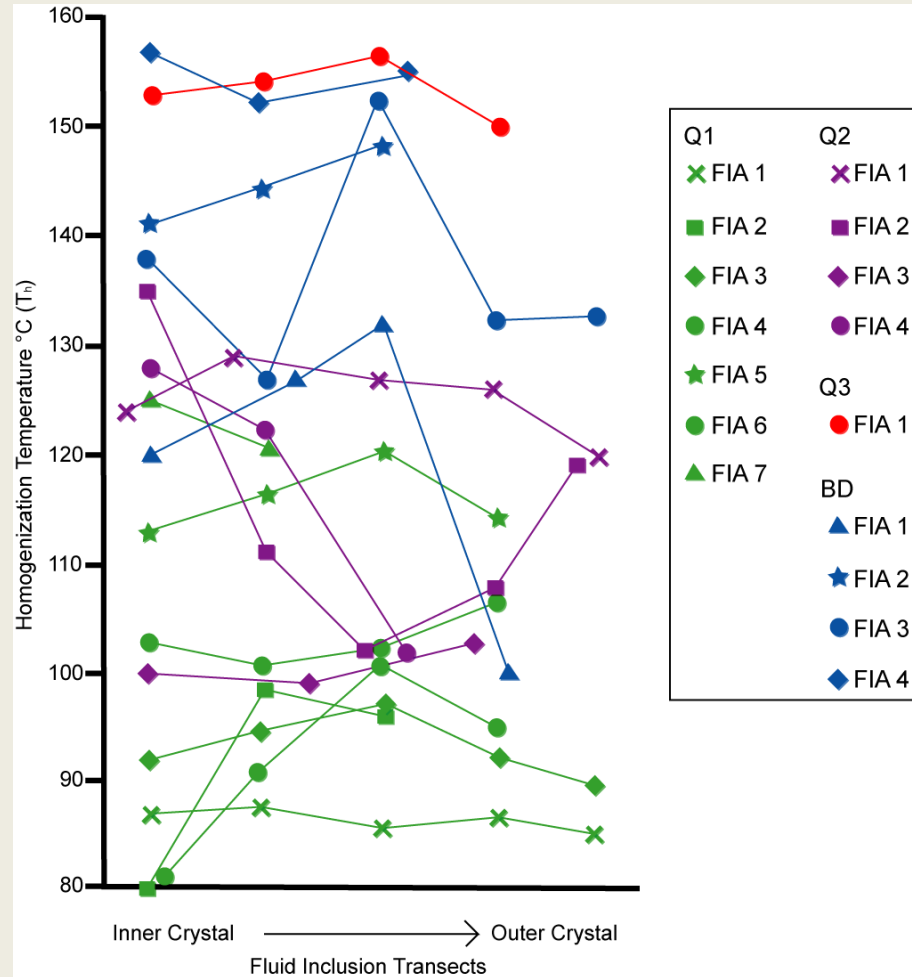
Similar Late-Stage Paragenesis in All Units

A Stratigraphic Unit	Diagenetic Events	Late Stage	
Middle Ordovician Simpson Group	1-2. Fracturing (SF1) 1-2. Dissolution		
	3. Megaquartz cement (SMQ) 4. Baroque dolomite (SBD)		
	5. Fracturing (SF2)		
	B Stratigraphic Unit	Diagenetic Events	Late Stage
	Mississippian (Upper and Lower Series)	1. Dissolution 2. Brecciation	
3. Megaquartz cement (MMQ) 4-5-6-7. Chalcedony (MCh)			
4-5-6-7. Baroque dolomite (MBD)			
4-5-6-7. Petroleum migration			
4-5-6-7. Fracturing (MF) 8-9. Calcite cement (MCC)			
8-9. Anhydrite (MA)			
C Stratigraphic Unit		Diagenetic Events	Late Stage
Middle Pennsylvanian Cherokee Group	1. Dissolution 2-3. Baroque dolomite (PBD)		
	2-3. Petroleum Migration		
	4. Fracturing (PF)		
	5. Calcite cement (PCC)		

Decreasing Age of Stratigraphic Unit

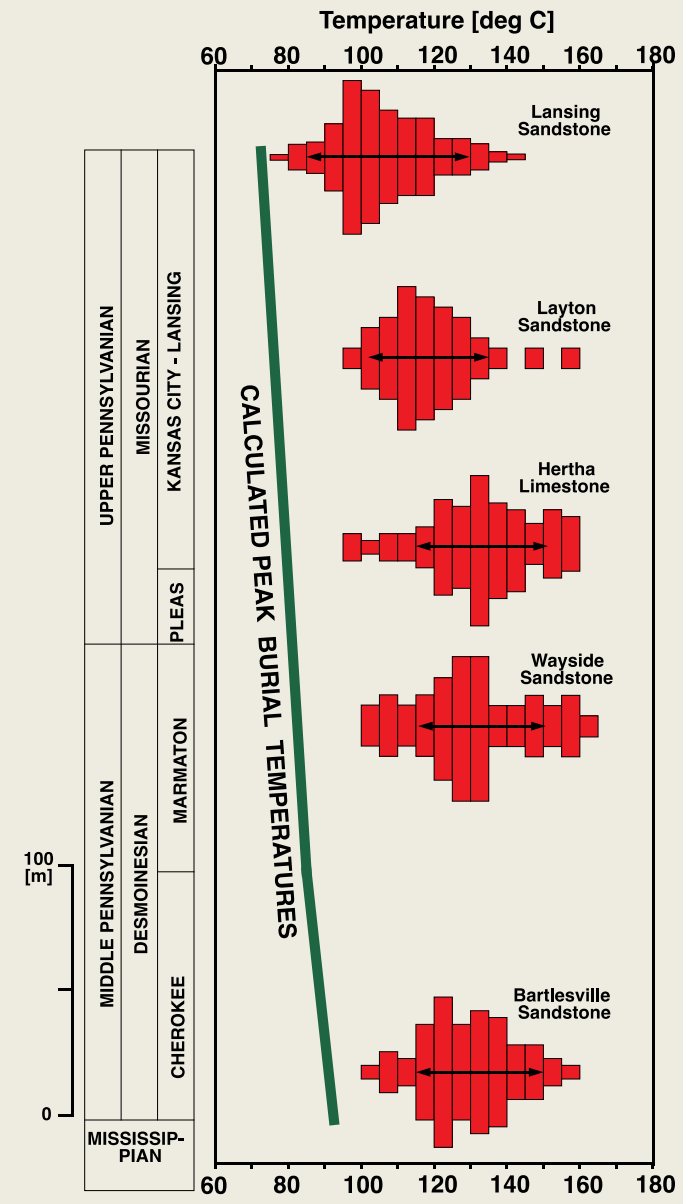
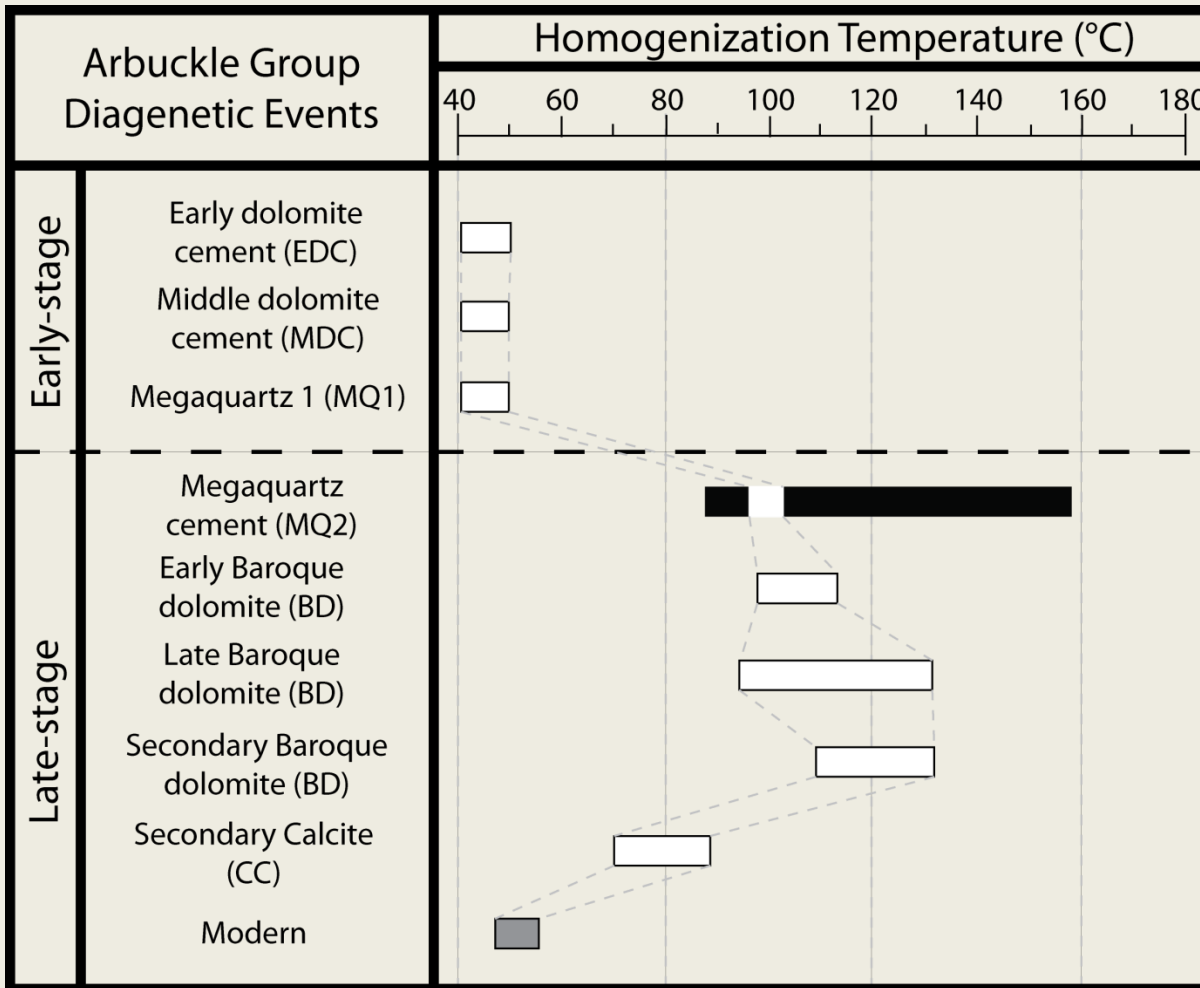


Fluid Inclusion Data – Hydrothermal



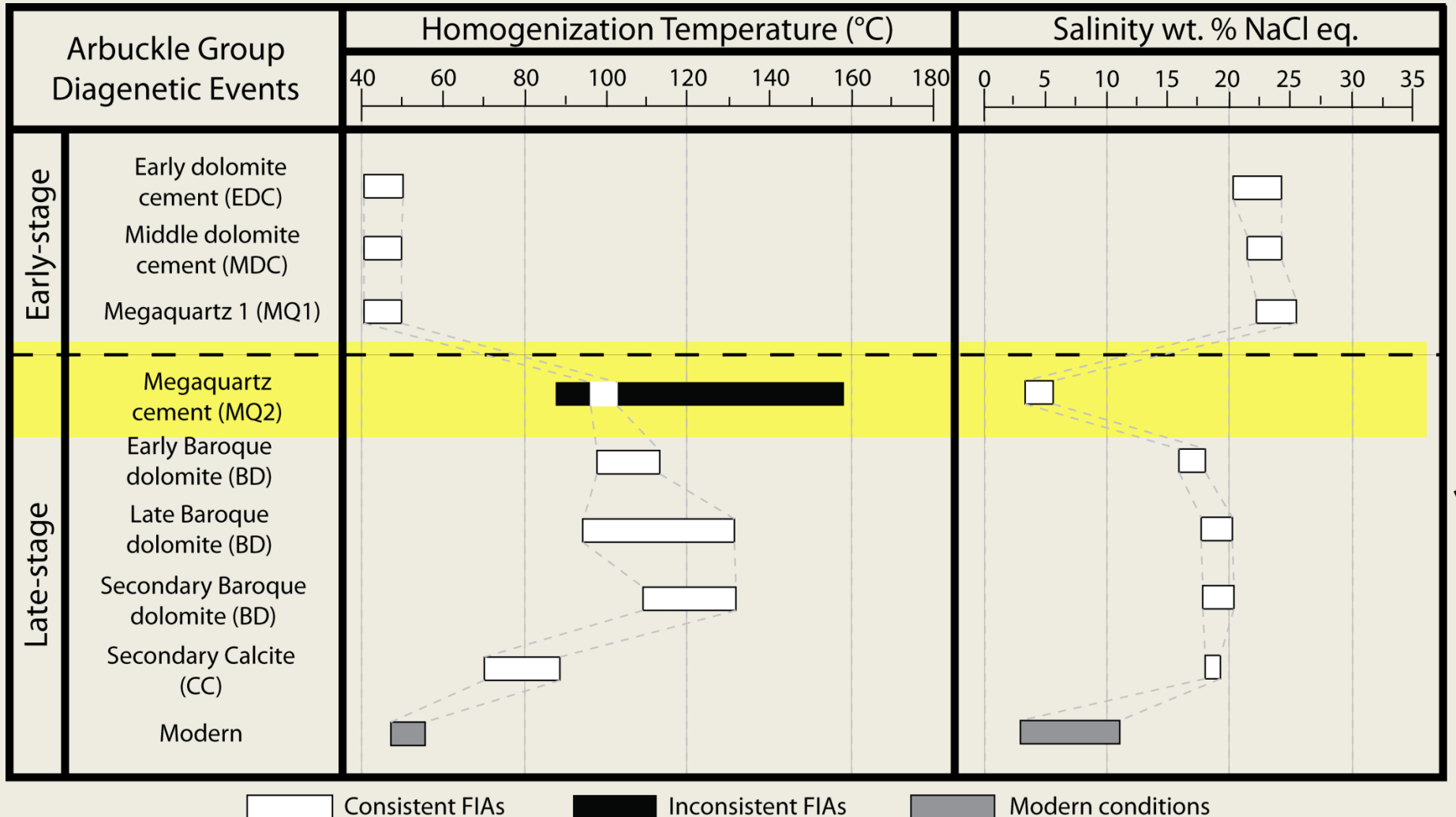
- Homogenization temperatures rise and fall through time - pulsed fluid flow

Fluid Inclusion Data – Hydrothermal

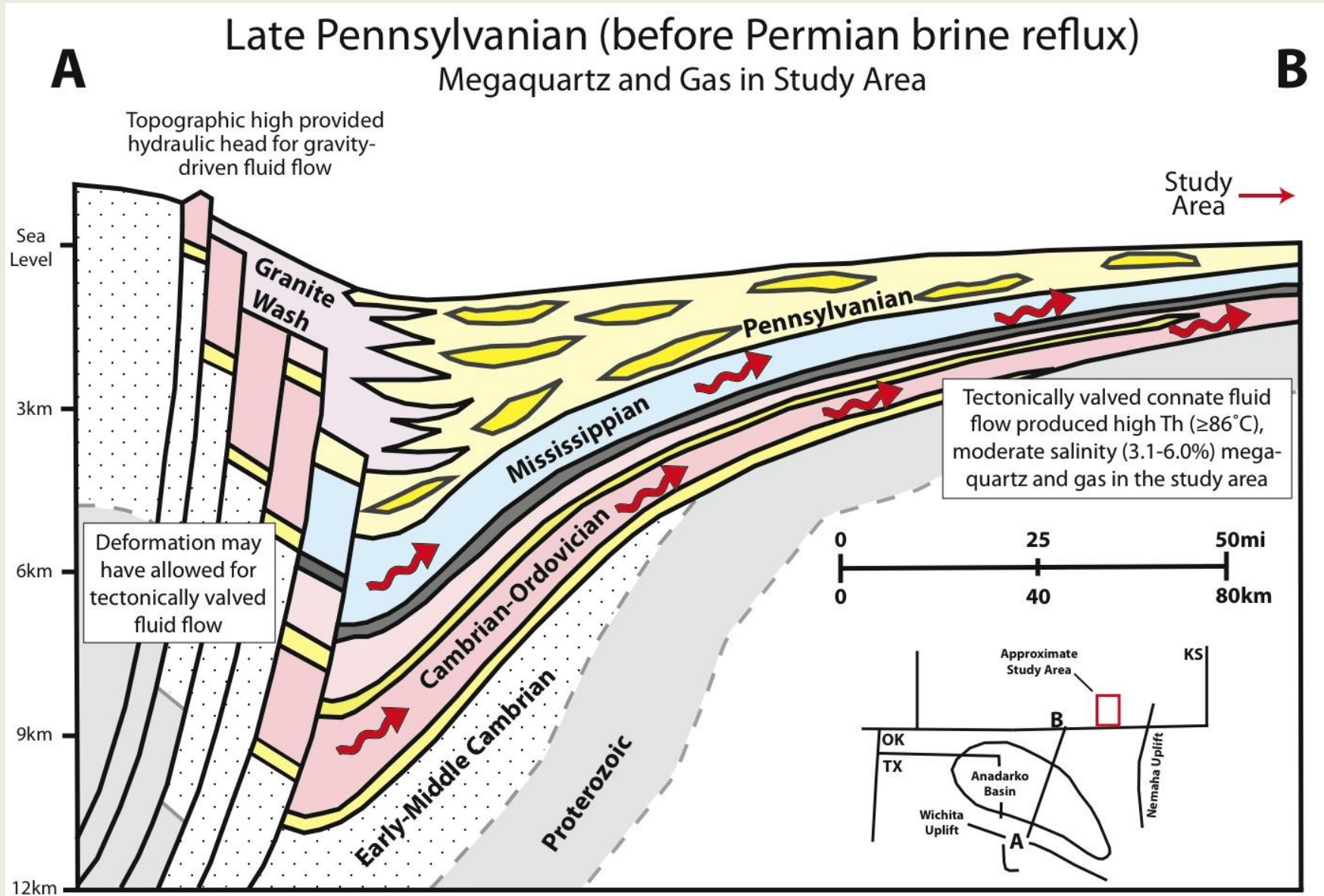


- Homogenization temperatures much higher than burial history allows
- Paleogeothermal gradients inconsistent with normal burial conductive heating

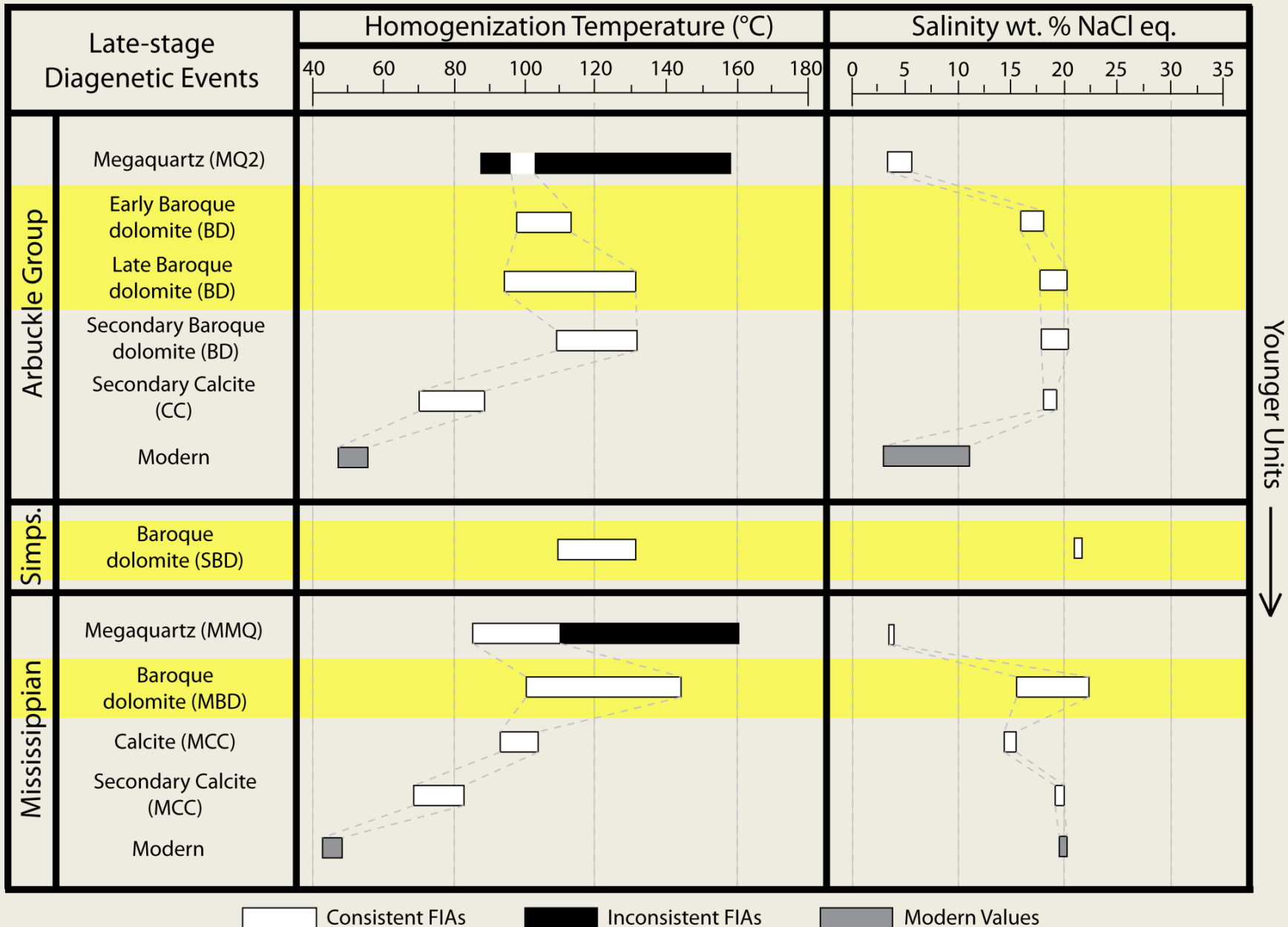
Stage 1 - Hydrothermal Flow



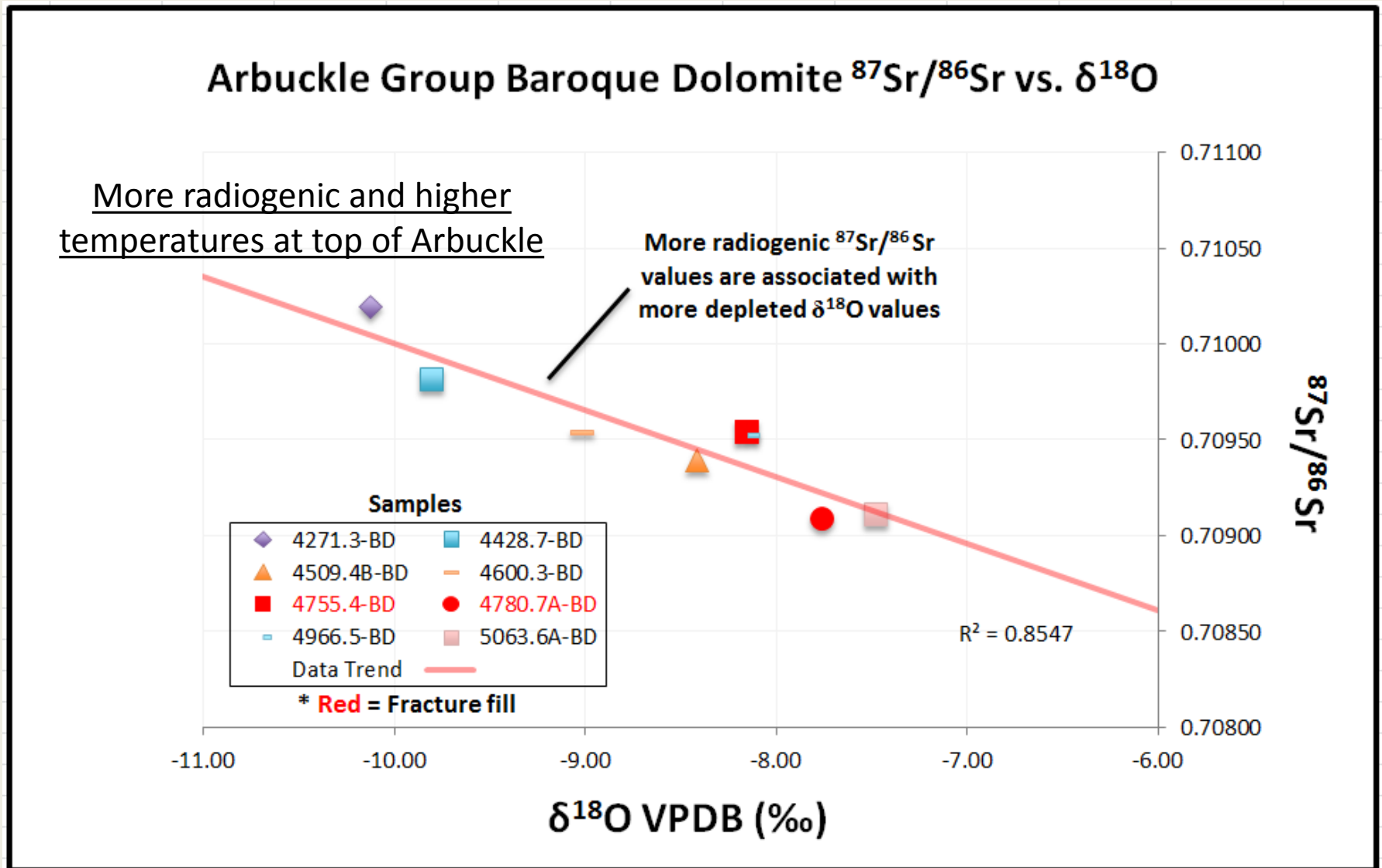
Stage 1 - Advective Fluid Flow - Megaquartz



Stage 2 - Hydrothermal Flow



Arbuckle Dolomite – $^{87}\text{Sr}/^{86}\text{Sr}$ vs. Temp.



Radiogenic values, higher temp = less rock-water interaction with carbonates

Thermal Structure - Stage 2 Regional Fluid Flow

High Temp.

Low Temp.

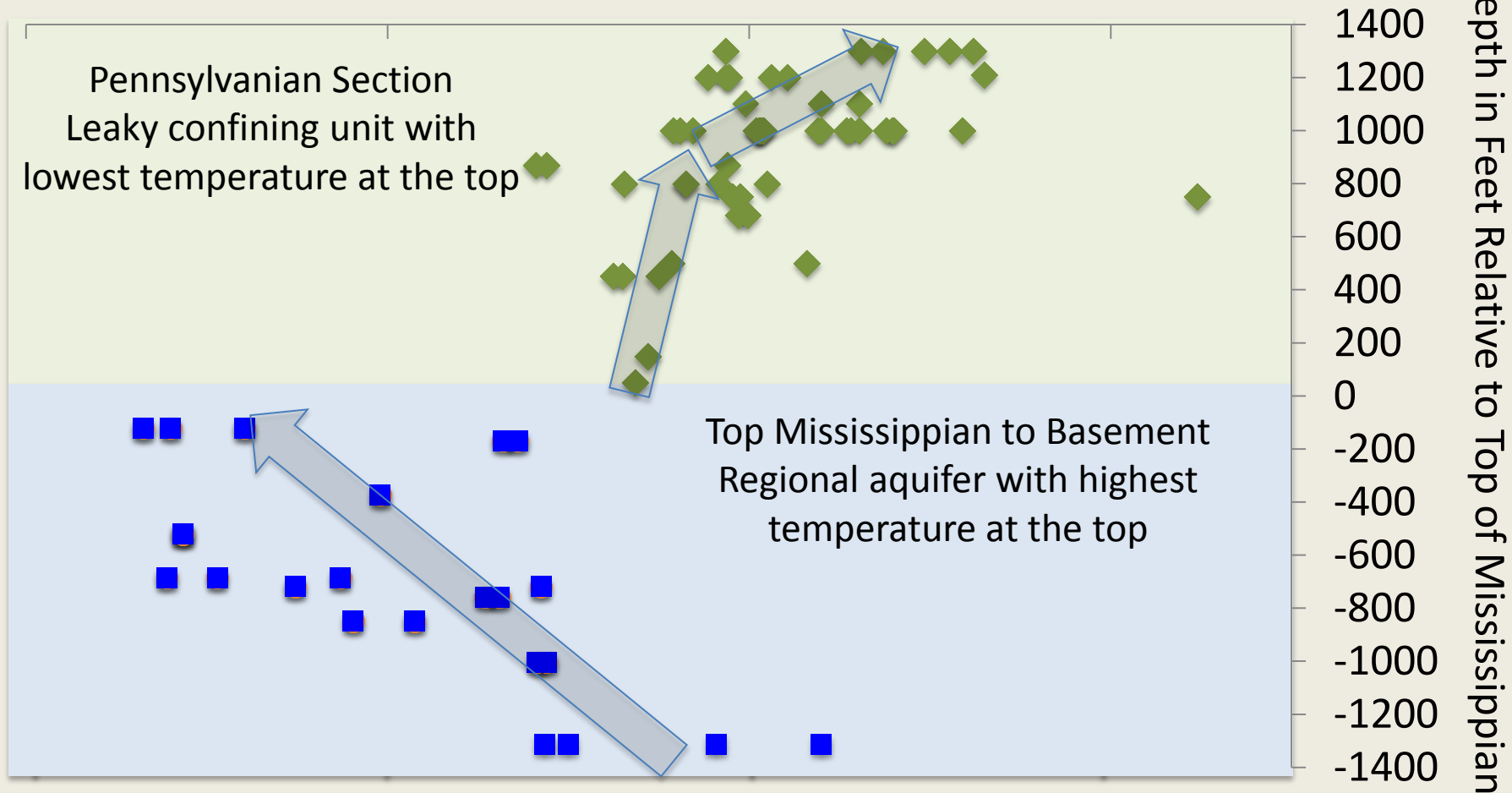
Dolomite $\delta^{18}\text{O}$

-11.00

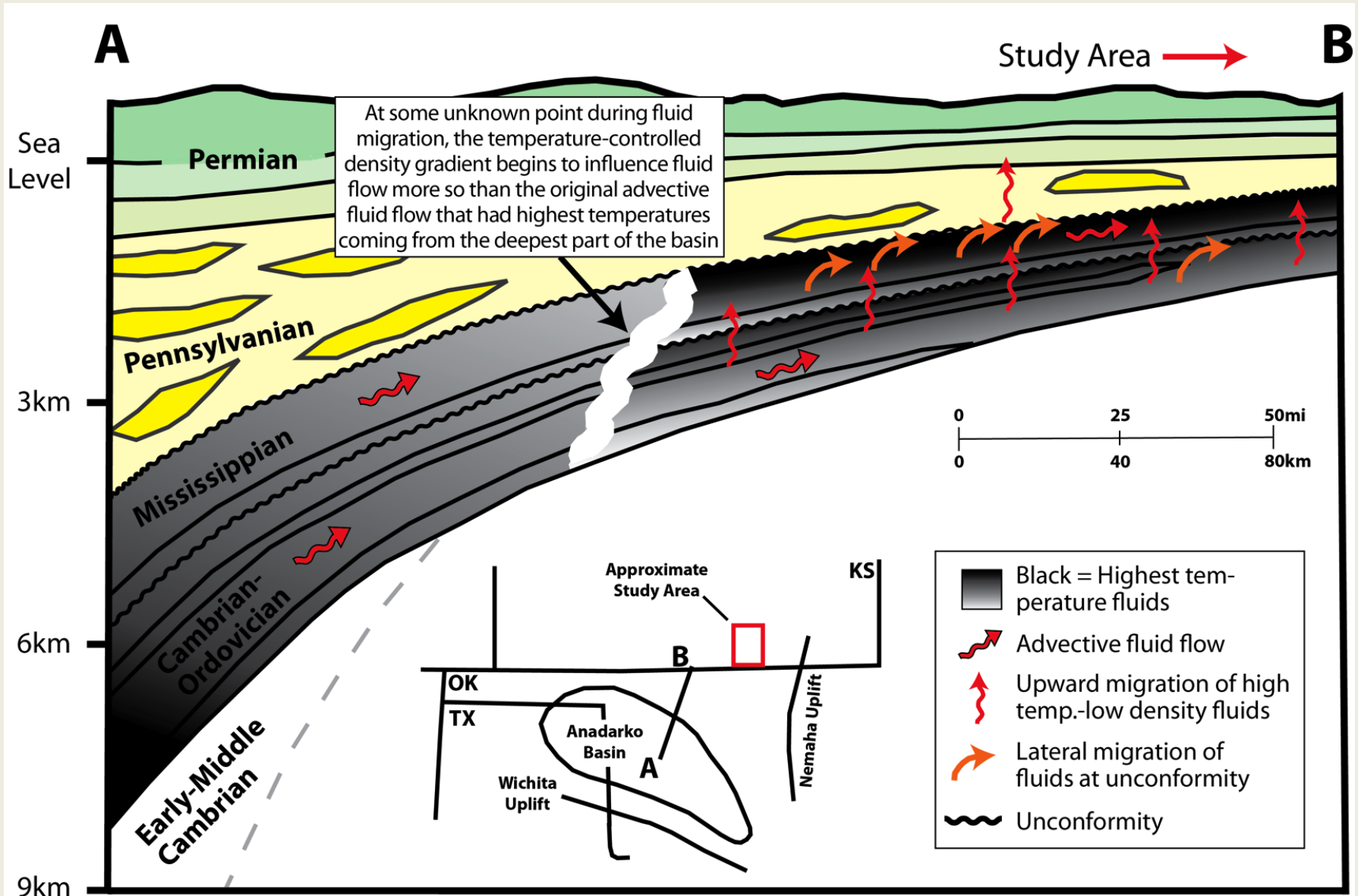
-9.00

-7.00

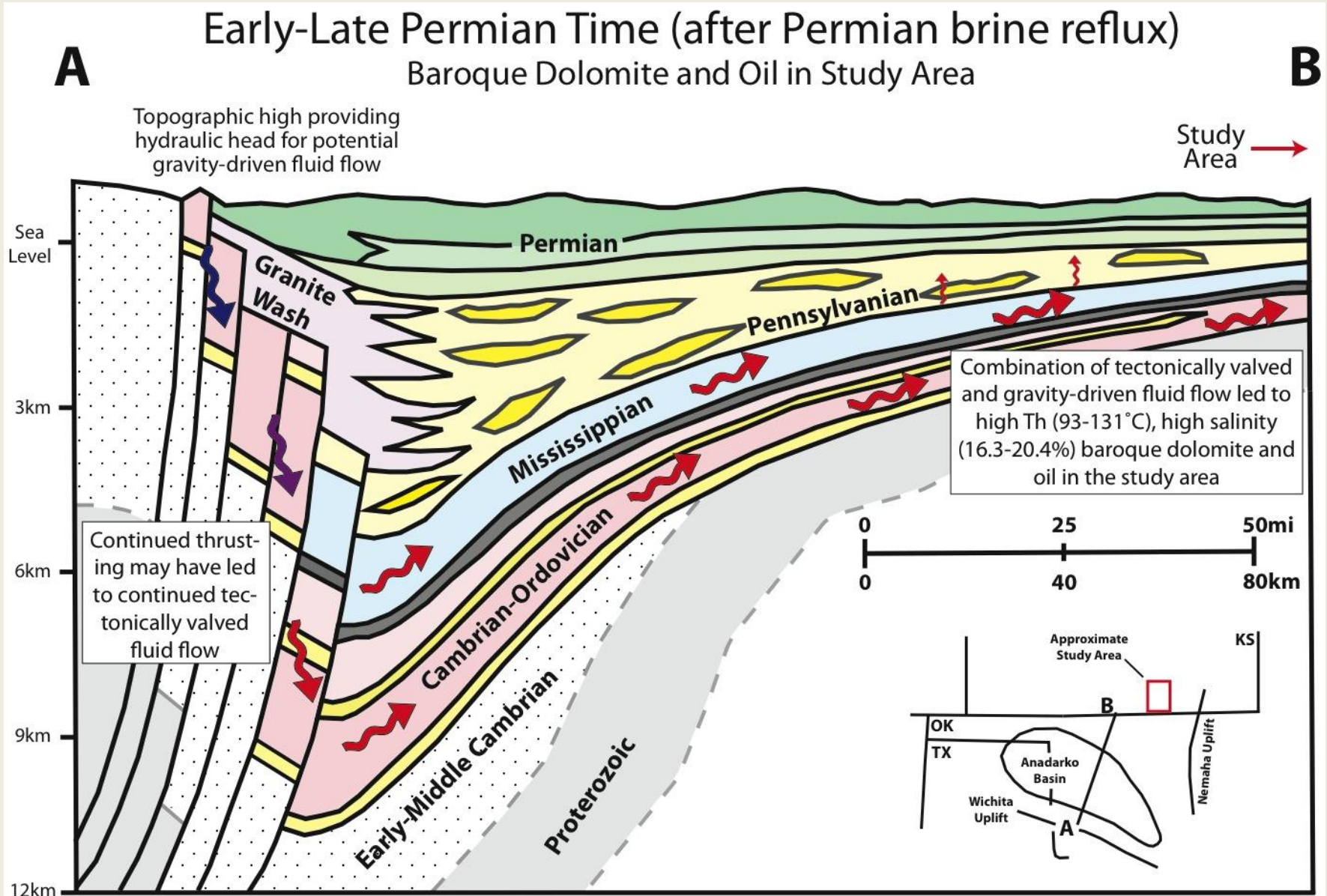
-5.00



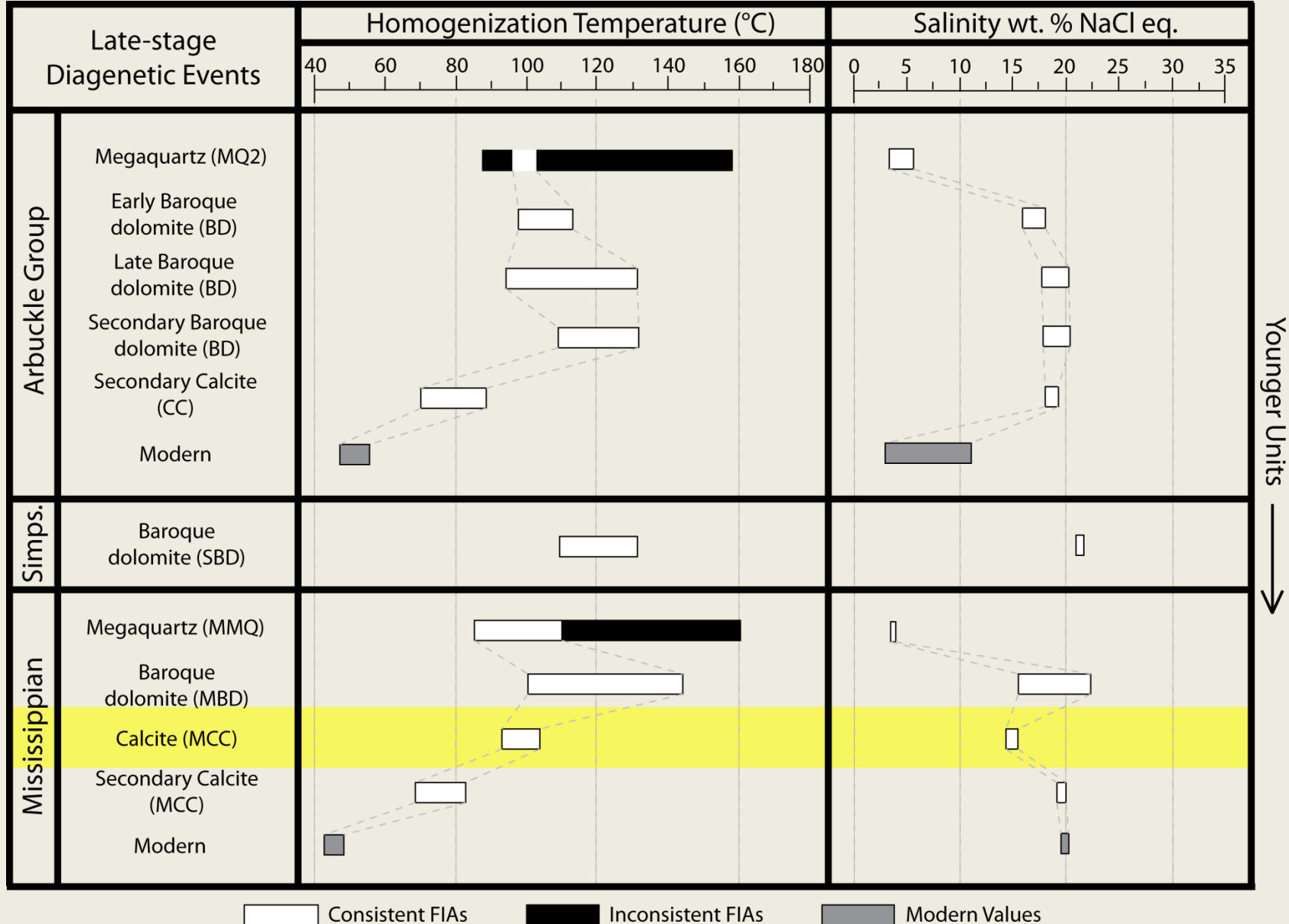
Stage 2 – Warmer fluids at top of Aquifer



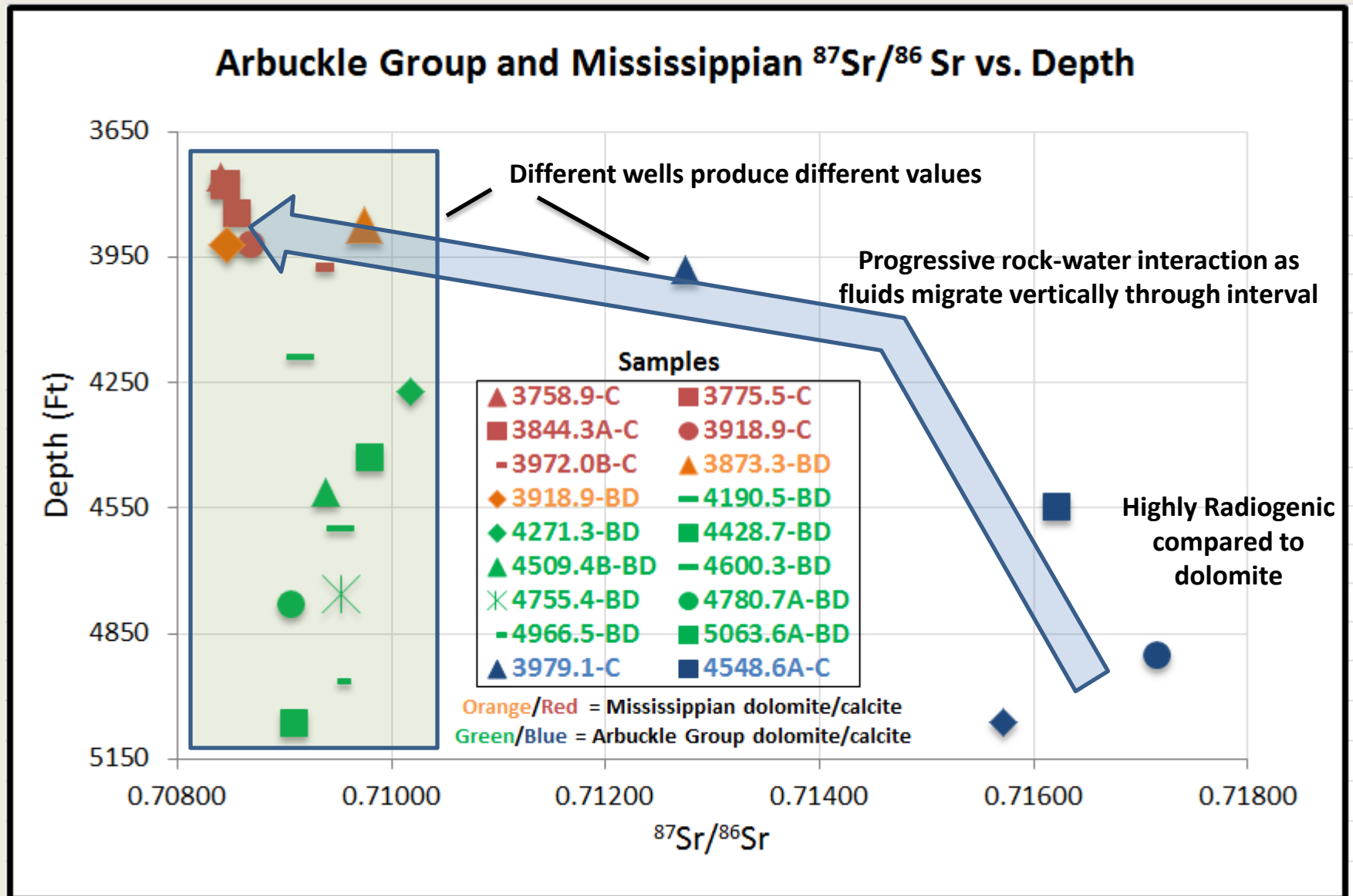
Advective Fluid Flow – Stage 2



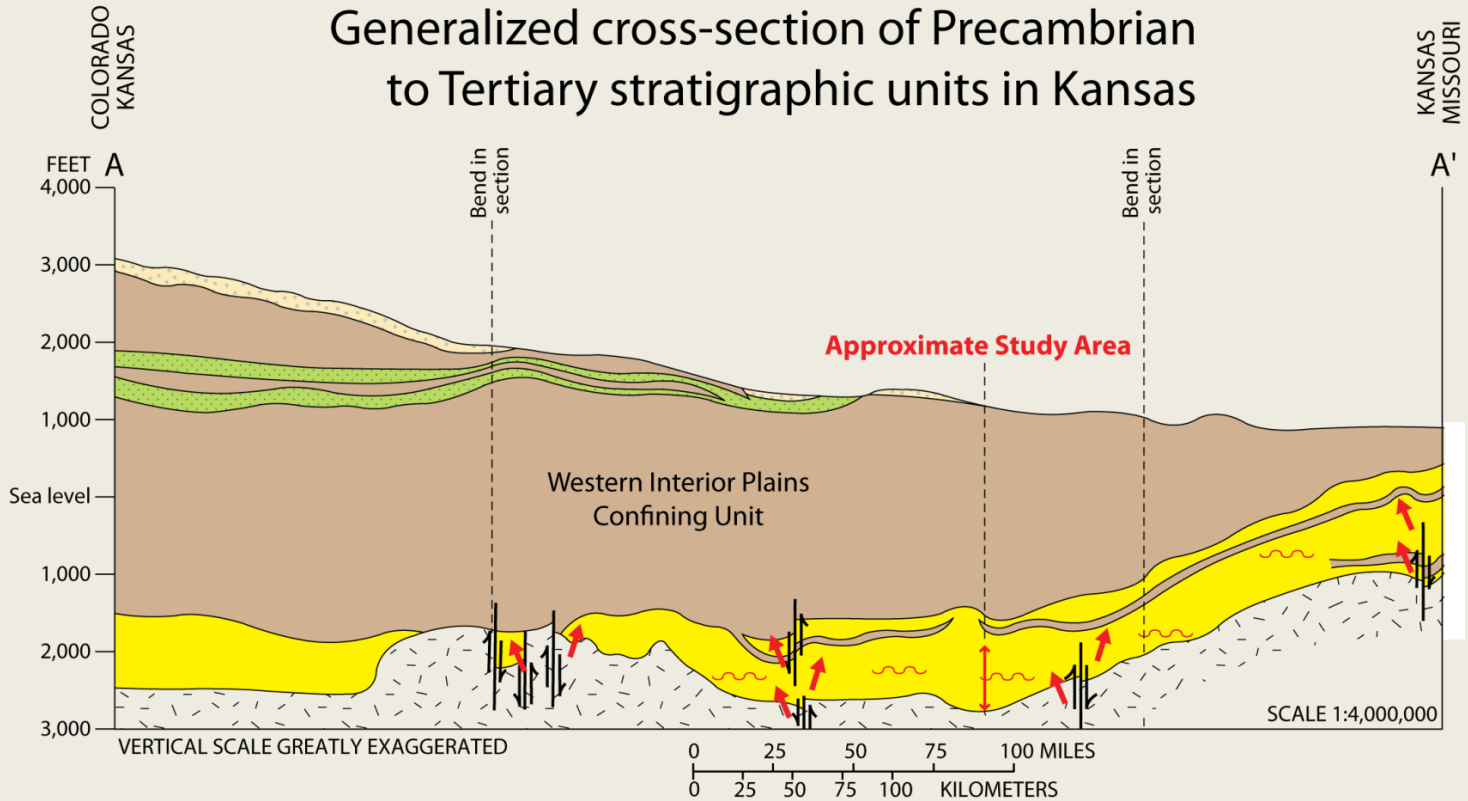
Stage 3 - Hydrothermal Flow



Stage 3 - Calcite $^{87}\text{Sr}/^{86}\text{Sr}$

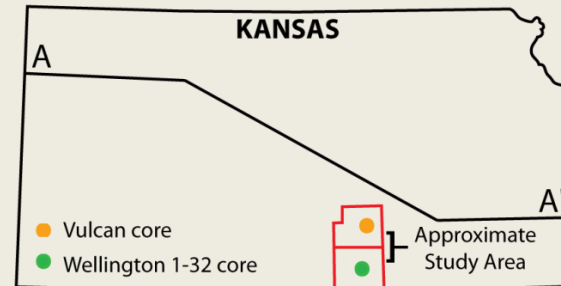


Stage 3 - Fracture-Controlled Hydrothermal Fluid Flow and Calcite-Laramide?



EXPLANATION

- Tertiary-Quaternary sand and gravel
- Cretaceous sandstones
- Cambrian-Mississippian Strata
- Confining unit / Pattern indicates crystalline rocks
- Fault—Arrows show direction of relative movement
- Fluids associated with vertical migration of fluids from units below the Arbuckle Group



Conclusions

- Three-stage evolution of hydrothermal systems
- Two stages of regional advective flow
 - Mississippian to Cambrian-Ordovician section acted as a regional aquifer and Pennsylvanian acted as a leaky confining unit.
 - In the regional aquifer cross-formational connections allowed lower density, warmer fluids to concentrate at the top of the aquifer, something predicted on the basis of density
- Better porosity is related to late structure and hydrothermal flow immediately below unconformity
- A third stage of hydrothermal fluid flow was localized by later (possibly Laramide) faults and fractures and led to localized hydrothermal systems