

Viscoplastic Deformation of Shale Gas Reservoir Rocks and its Relation to the In-Situ Stress Variations Observed in a Well from Barnett Shale*

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Abstract

We have studied the viscoelastic/viscoplastic properties of various shale gas reservoir rocks through laboratory triaxial creep experiments under confining pressures representative of in situ conditions. We find that the viscoelastic stress relaxation behavior of these rocks vary considerably and is generally larger for rocks with more clay and organic content. The strain response of the rock is adequately described by a power-law function of time, and its magnitude is approximately linear against the magnitude of the applied differential pressure but insensitive to the confining pressure. Motivated by these observations, we described the rheology of the shales in the framework of linear viscoelasticity in order to calculate the differential stress accumulation/relaxation that would occur in these reservoirs over geological time scales. Variation in viscoelastic properties within the Barnett Shale can create differential stress variations on the order of one to tens of MPa, consistent with fluctuations in stress difference inferred from observations of wellbore failures. Although time-dependent deformational behaviors of intact reservoir rocks are not studied routinely in the lab, we suggest that it can have a significant impact in determining the current in-situ state of stress when reservoir deformation takes place over geological time scales.

Reference Cited

Hagin, P.N., and M.D. Zoback, 2004, viscous deformation of unconsolidated reservoir sands (Part II): Linear Viscoelastic models: Geophysics, v. 69/3, p. 742-751.

VISCOPLASTIC DEFORMATION OF SHALE GAS RESERVOIR ROCKS AND ITS LONG-TERM EFFECT ON THE IN-SITU STATE OF STRESS

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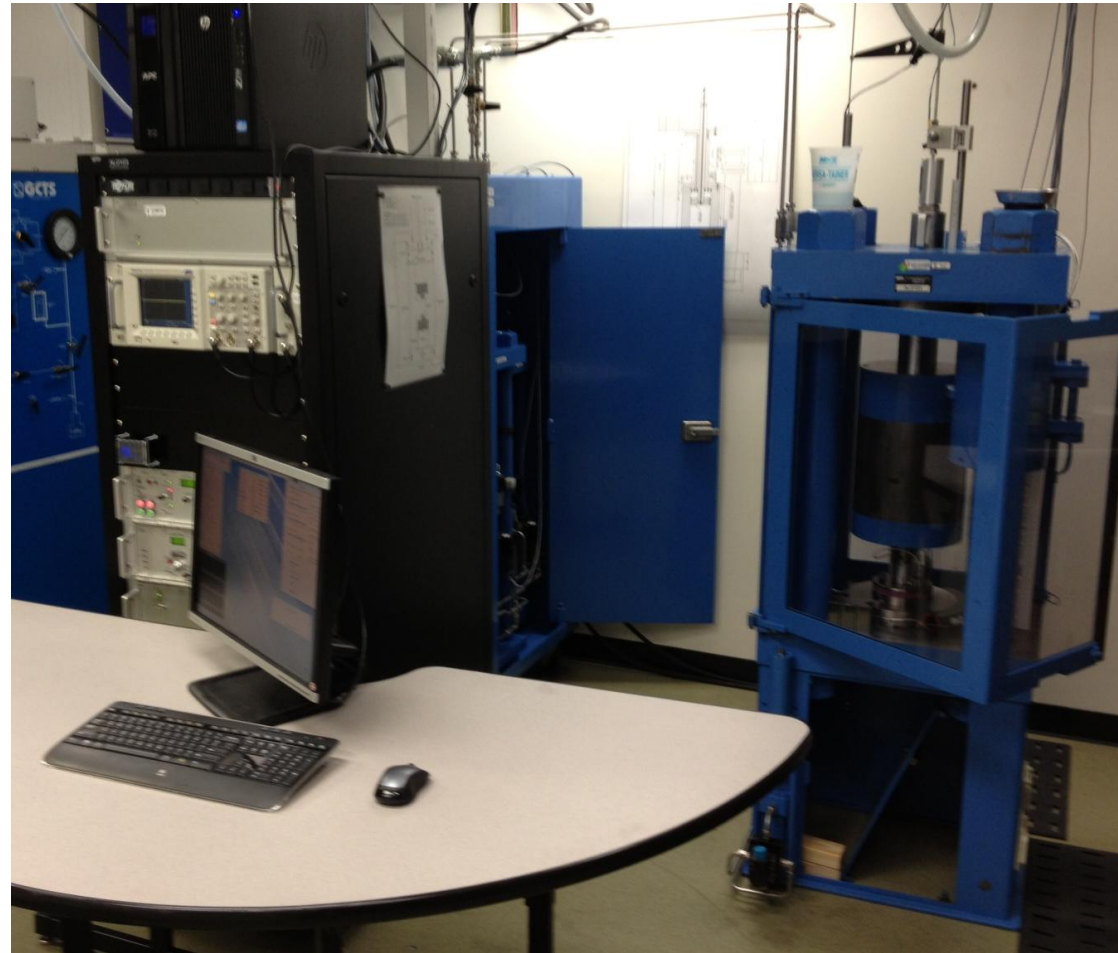
* Now at GeoForschungZentrum, Potsdam



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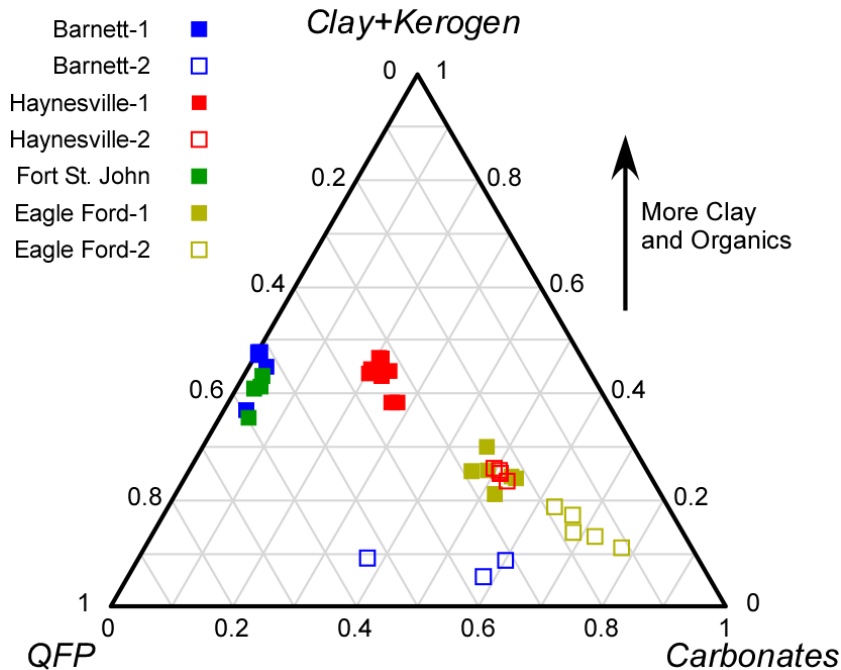


Creep Studies of Shale





Gas Shale Samples



Barnett-1



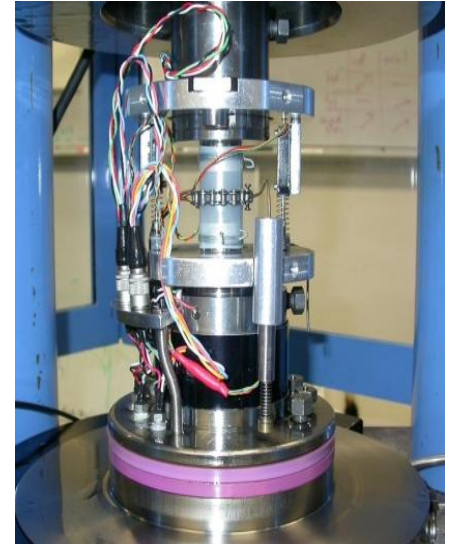
Barnett-2

- 4 different shale gas reservoirs
- Barnett, Haynesville, Eagle Ford divided into subgroups
subgroup-1 more clay-rich and organic-rich than subgroup-2
- 1" diameter cylindrical samples
- Cylinder axis vertical and horizontal to bedding plane



Organic Rich Shales

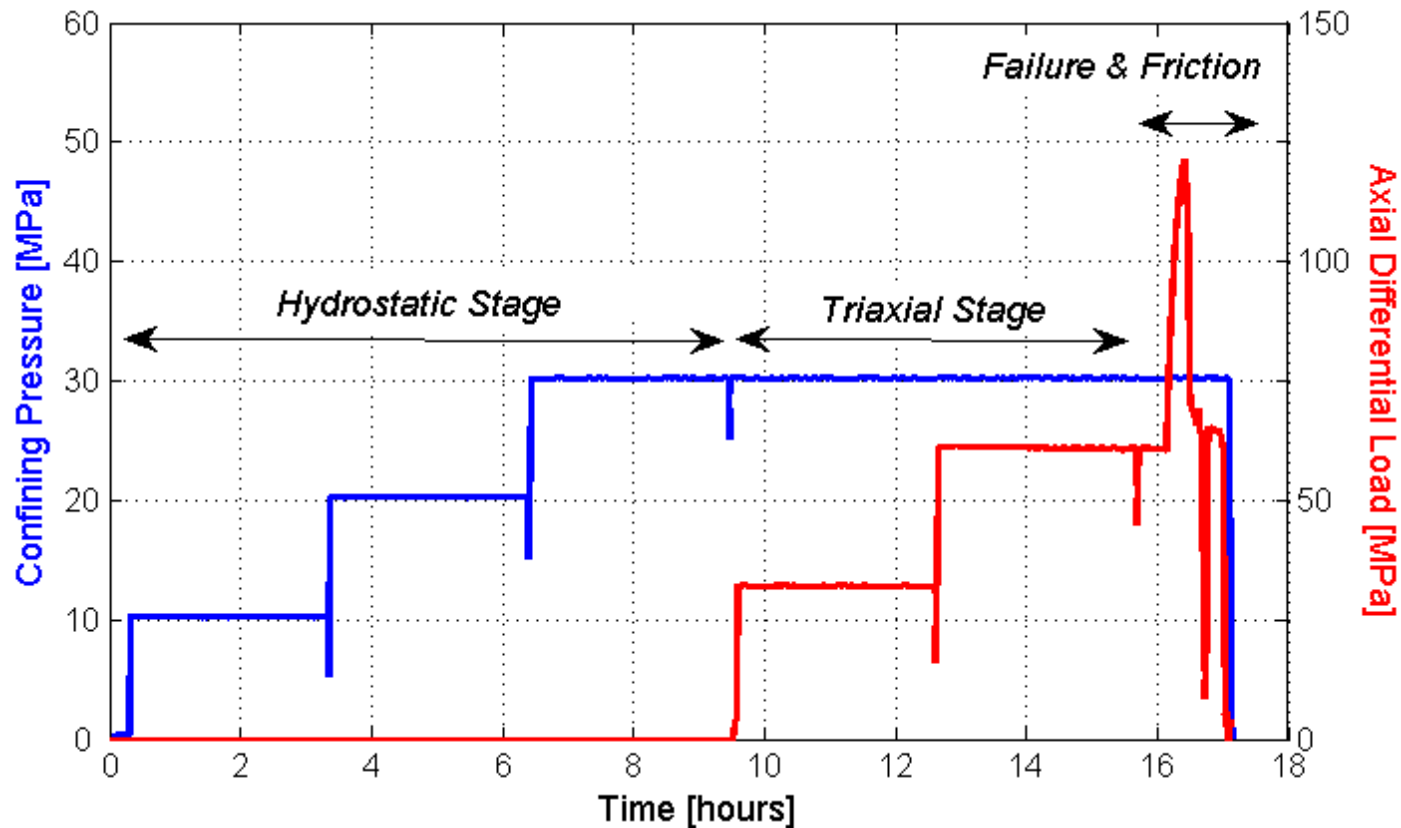
Sample group	Clay	Carbonate	QFP	TOC (wt%)
Barnett-dark	29-43	0-6	48-59	4.1-5.8
Barnett-light	2-7	37-81	16-53	0.4-1.3
Haynesville-dark	36-39	20-23	31-35	3.7-4.1
Haynesville-light	20-22	49-53	23-24	1.7-1.8
Fort St. John	32-39	3-5	54-60	1.6-2.2
Eagle Ford-dark	12-21	46-54	22-29	4.4-5.7
Eagle Ford-light	6-14	63-78	11-18	1.9-2.5



- 3-10 % porosity
- All room dry, room temperature experiments
- In-situ (and lab) effective stress between 15-30 MPa



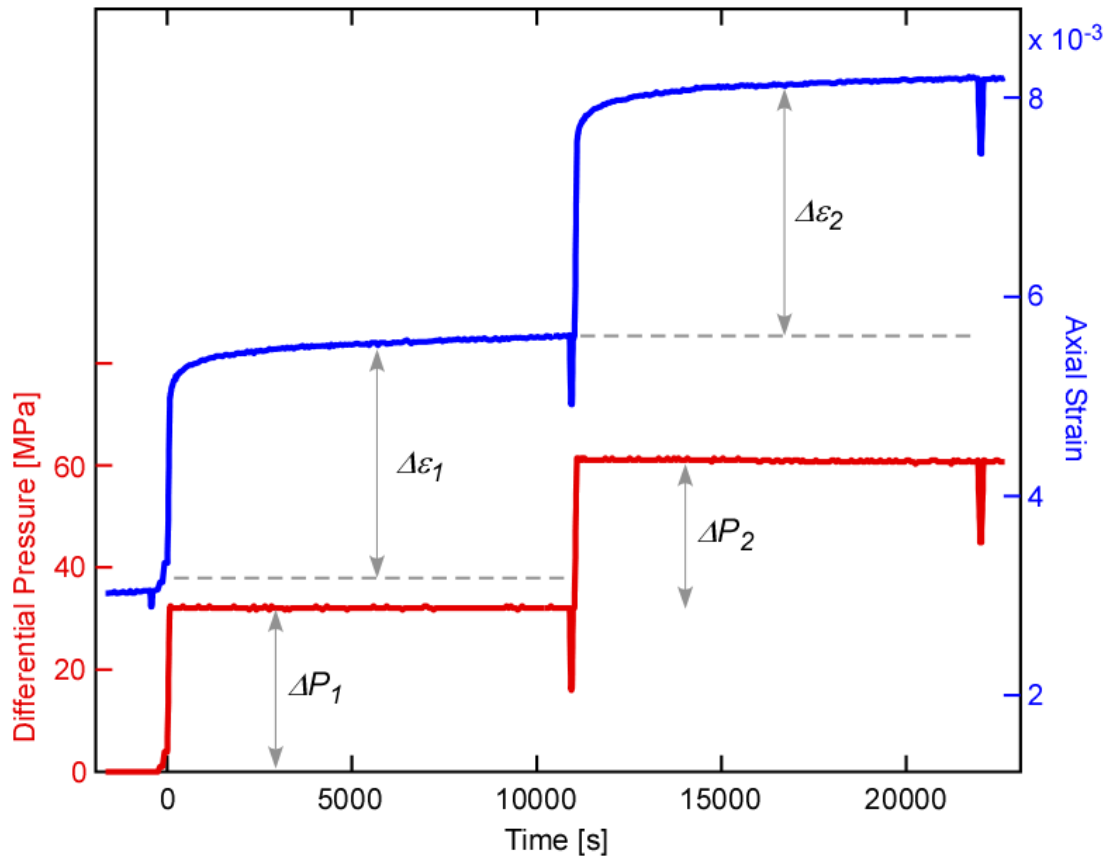
Typical Experimental Procedure



- Hydrostatic Stage: Bulk modulus, hydrostatic creep
- Triaxial State: Young's modulus, Poisson ration, triaxial creep
- Failure&Friction: Onset of dilatancy, intact/frictional strength



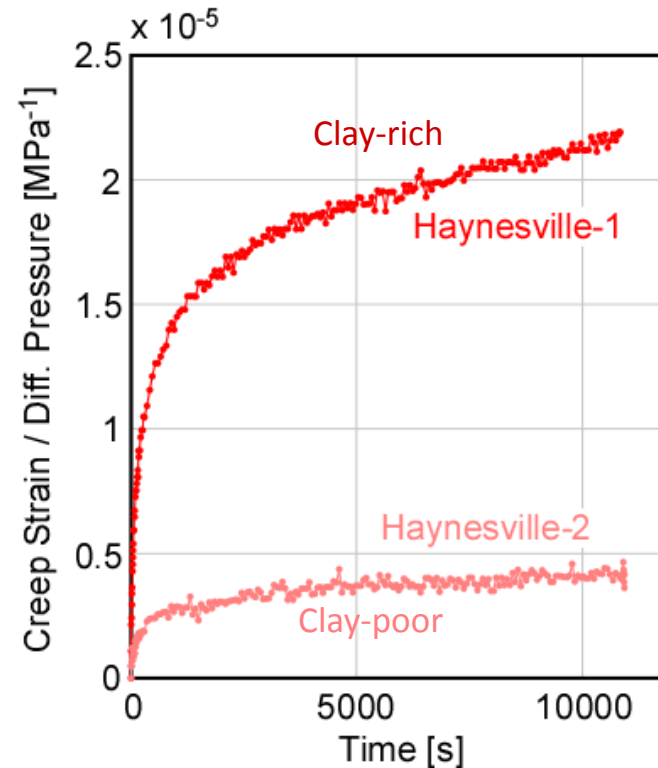
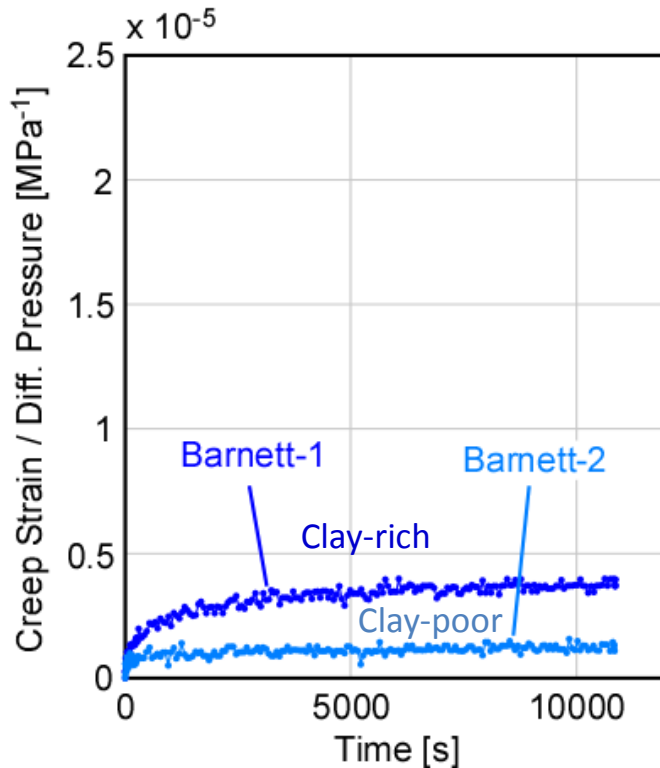
Experimental Procedure



- Confining pressure 10-60 MPa
- Differential pressure applied in several steps
- Held differential pressure constant for 3 hrs ~ 2 weeks
- Observed creep behavior to constrain $J(t)$



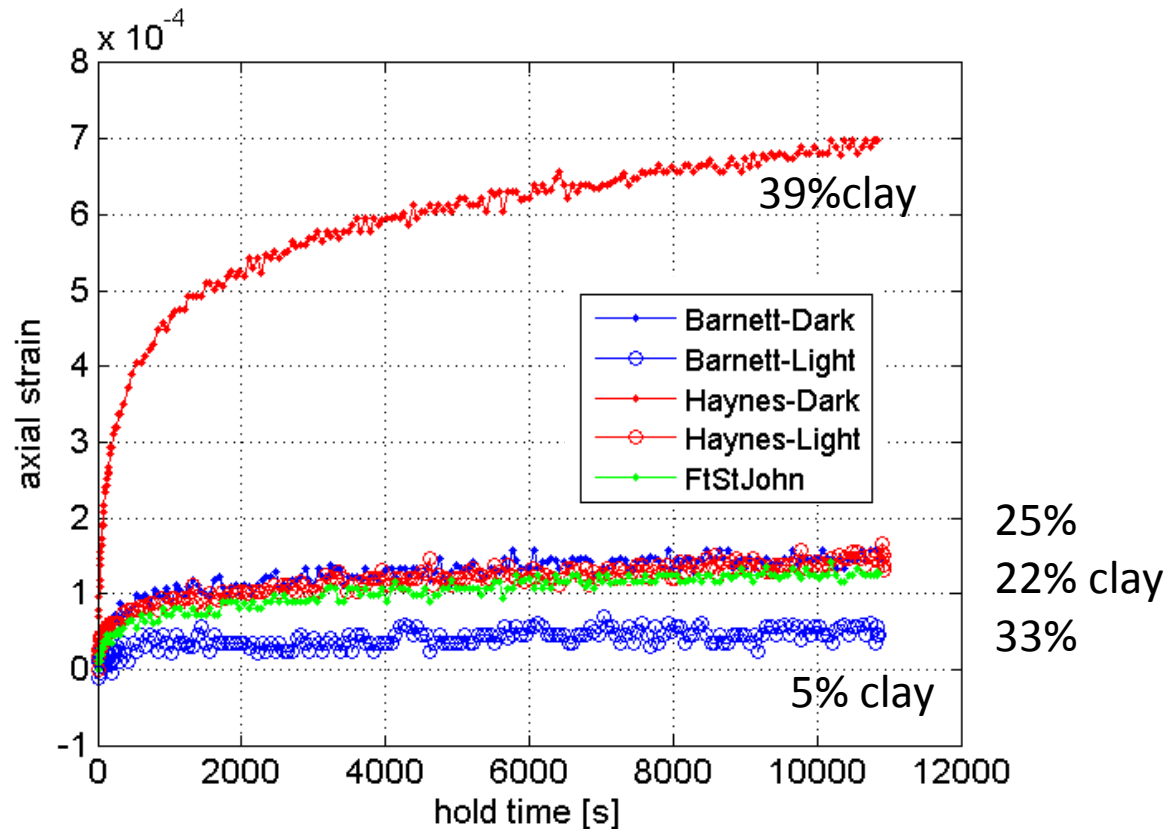
Clay/Organic Content Promotes Ductility



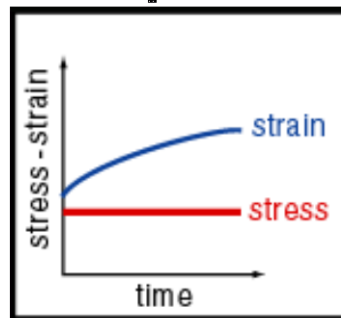
- Creep (ductility) is more pronounced in clay-rich, organic-rich rocks in all reservoirs
- Volume reduction and velocity increase during creep suggests creep is accommodated by compaction in clays and organics



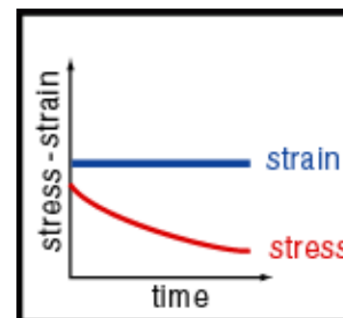
Clay/Organic Content Promotes Ductility



Creep Strain

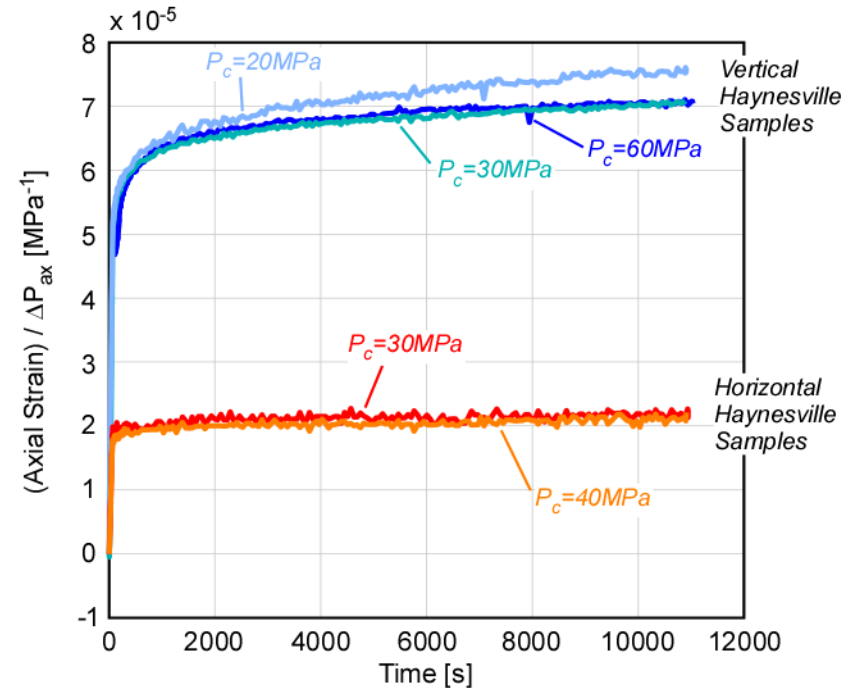
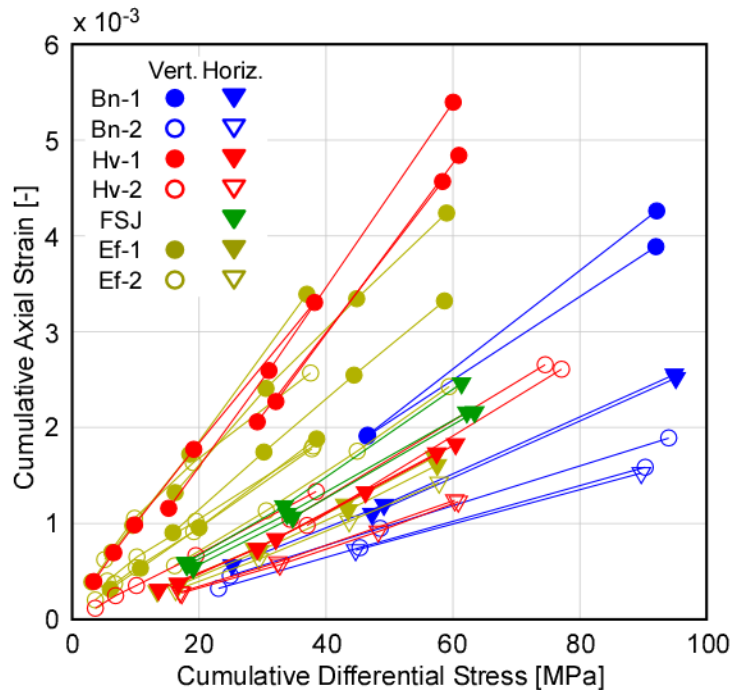


Stress Relaxation





Pressure Dependence



- Relation between cumulative strain and axial pressure is roughly proportional
- The amount of deformation is not dependent on the confining pressure (isotropic stress)

→ Strain is linear against differential pressure

→ Justifies the use of linear viscoelasticity



Quantitative Analysis via Linear Viscoelasticity

- Addition of time parameter to linear elasticity**

modulus $C \rightarrow E(t)$,
compliance $S \rightarrow J(t)$

$E(t)$: Relaxation Modulus

$J(t)$: Creep Compliance

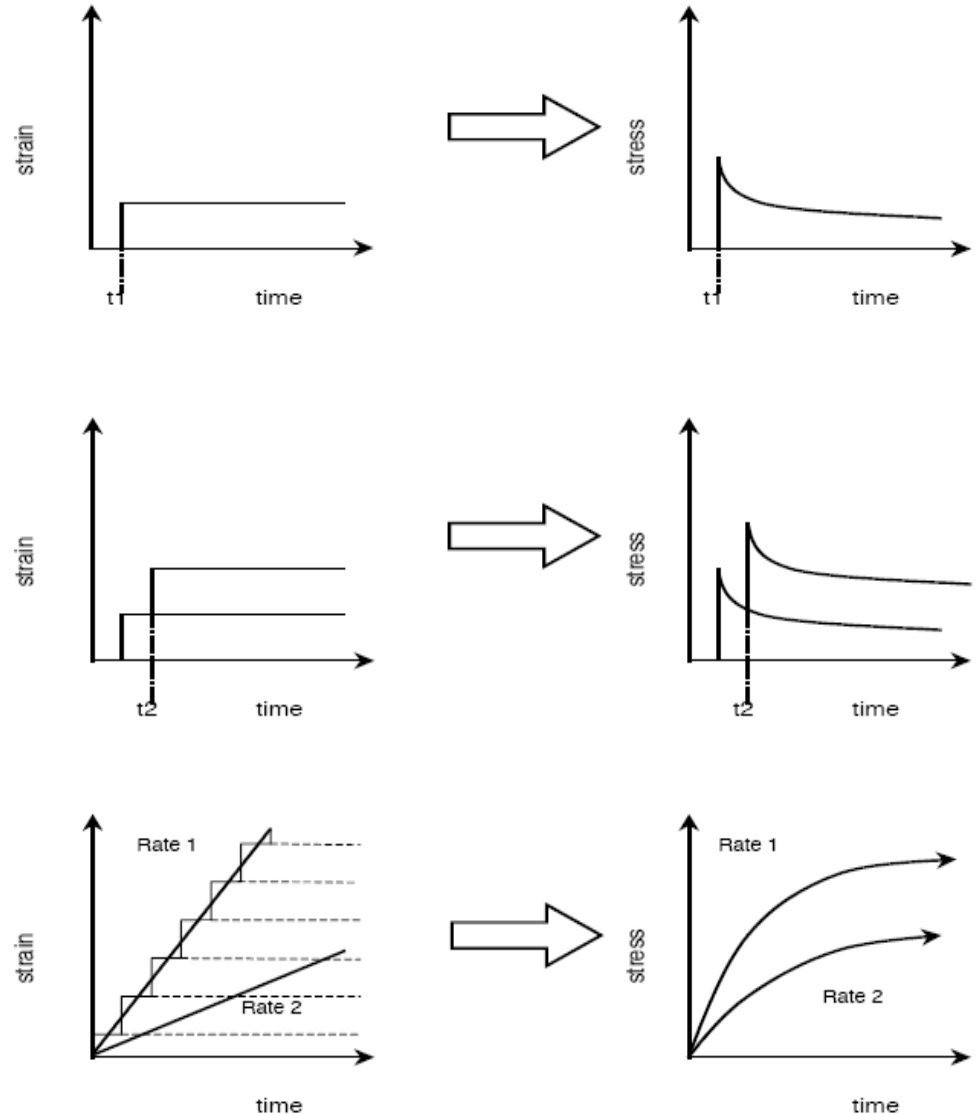
- Linear superposition**

$$\sigma(t) = \int_0^t E(t-\tau) \frac{d\varepsilon}{d\tau} d\tau$$

$$\varepsilon(t) = \int_0^t J(t-\tau) \frac{d\sigma}{d\tau} d\tau$$

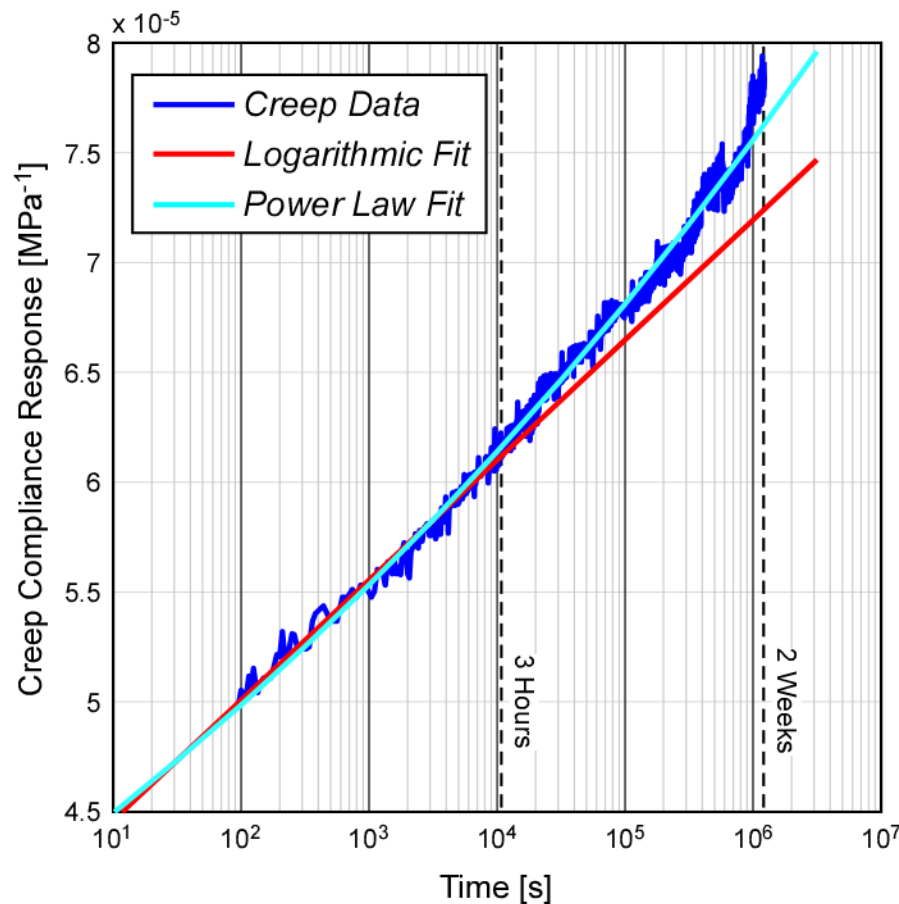
In the Laplace domain,

$$E(s)J(s) = \frac{1}{s^2}$$





Power-law Creep Compliance, $J(t)=Bt^n$



- Characteristics of creep
 1. Creep does not reach asymptote
 2. Creep rate continues to decrease
- Power-law expression has better long-term predictability of creep behavior
- Power-law expression also known to be useful for concrete, asphalt



Convenience of the Power-law

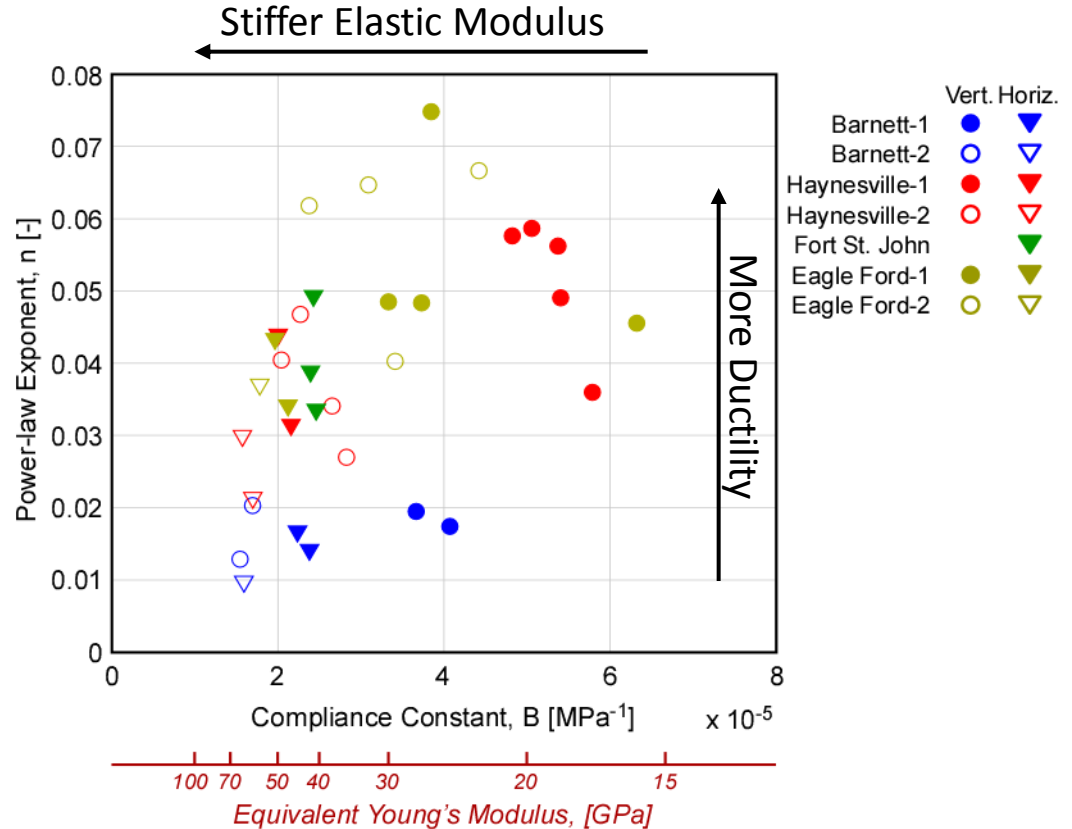
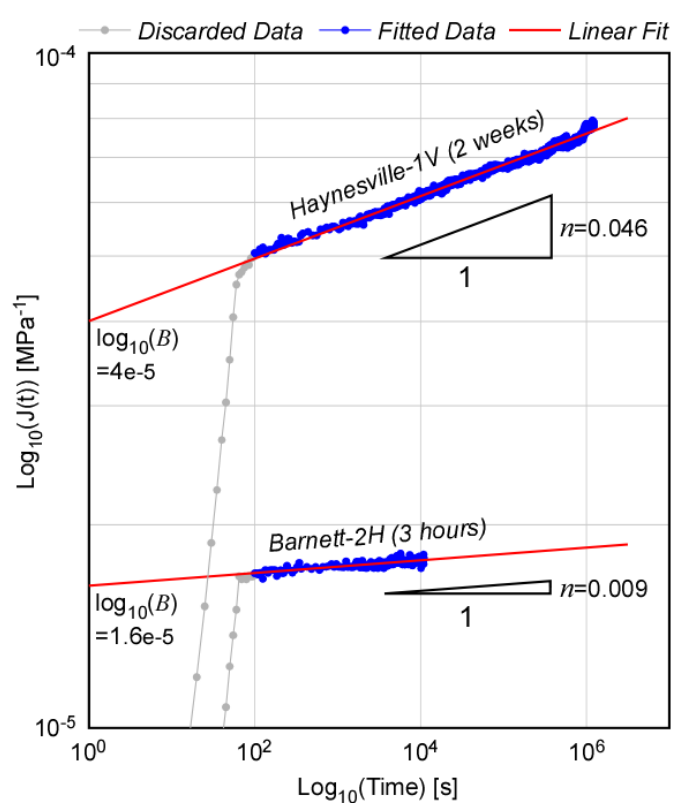
- Relaxation Modulus, $E(t)$, is easily obtained

$$J(t) = Bt^n$$
$$J(s) = \frac{B\Gamma(n+1)}{s^{1+n}}$$
$$E(s) = \frac{1}{s^2 J(s)} = \frac{1}{B\Gamma(n+1)} \frac{1}{s^{1-n}}$$
$$E(t) = \frac{1}{B\Gamma(1+n)\Gamma(1-n)} t^{-n} \approx \frac{1}{B} t^{-n} \quad (n \ll 1)$$
$$\sigma(t) = \dot{\epsilon} \frac{1}{B(1-n)} t^{1-n}$$

- $1/B$ is roughly equal to the elastic Young's modulus
- n is the power-law exponent, the degree of ductility, or how much time-dependent deformation you get



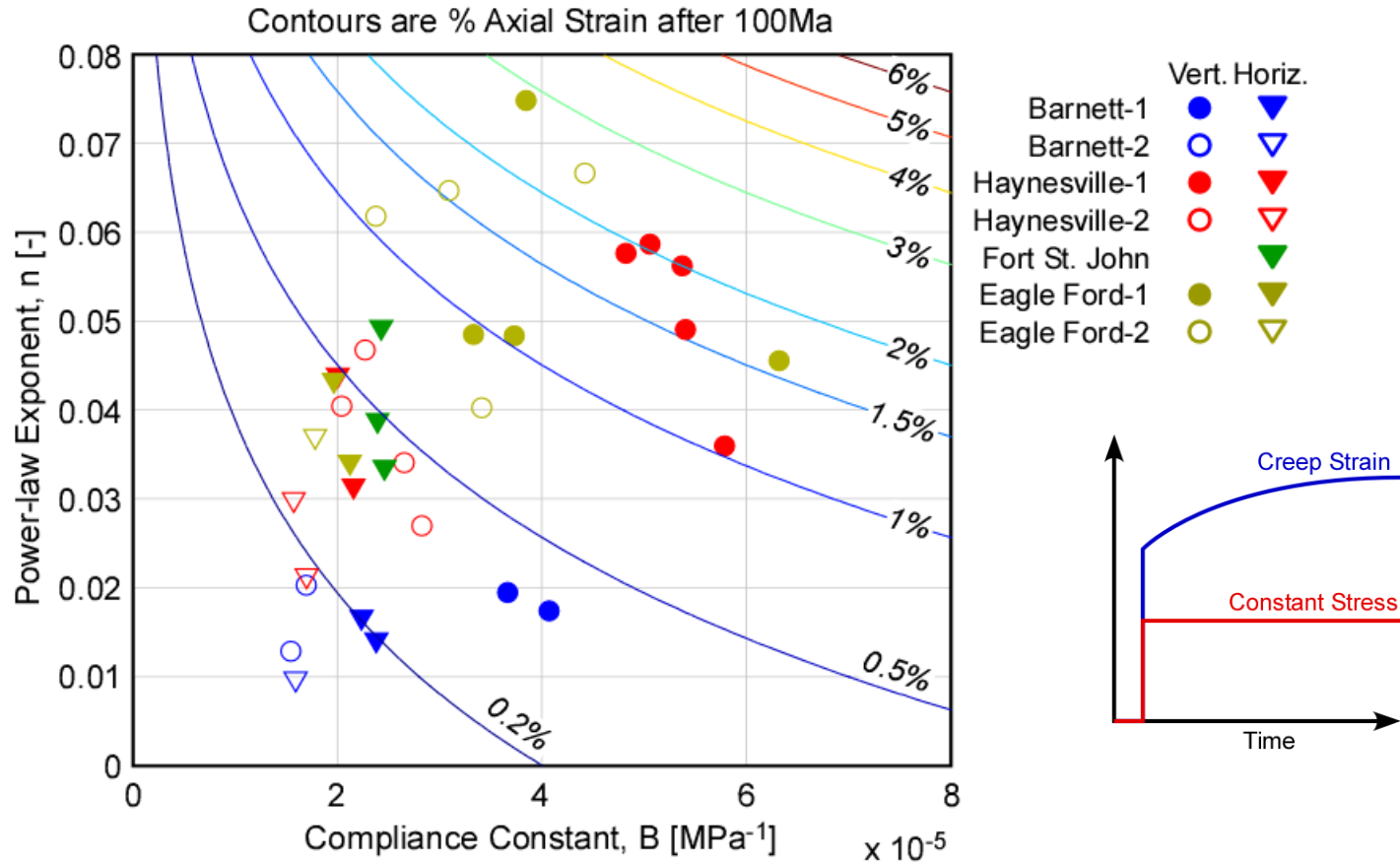
Constitutive Parameters in $J(t)=Bt^n$



- Creep data fit by a line in log-log space
→ B and n are obtained for each sample
- Samples exhibit wide variety of behavior



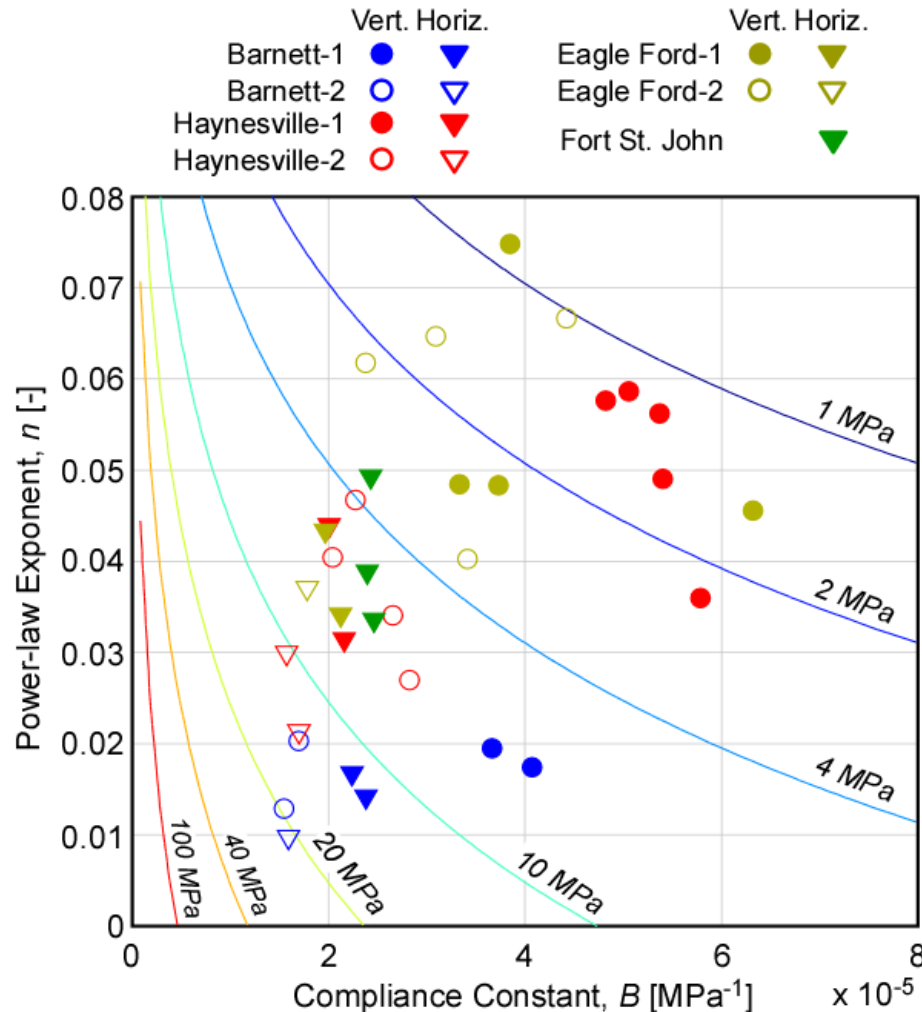
Creep Strain over Geological Time



- Contours are % strain under 50 MPa differential load
- Reasonable axial strain magnitudes of 0.1~3%



Predicting Stress Anisotropy over Geological Time



□ Stress Accumulation under constant strain rate

□ 150 Ma - Half of age of Barnett shale

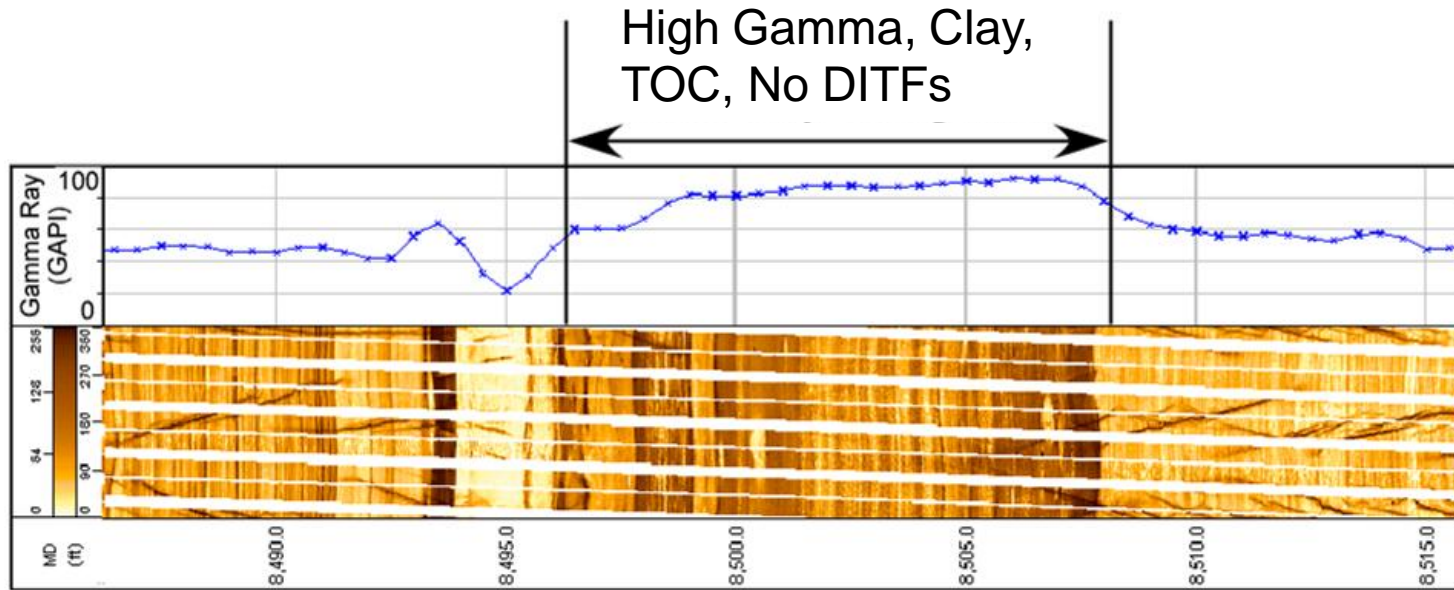
□ 10^{-19} s^{-1} - Stable intraplate

$$\sigma(t) = \dot{\epsilon} \frac{1}{B(1-n)} t^{1-n}$$

□ Significant stress relaxation observed for high n



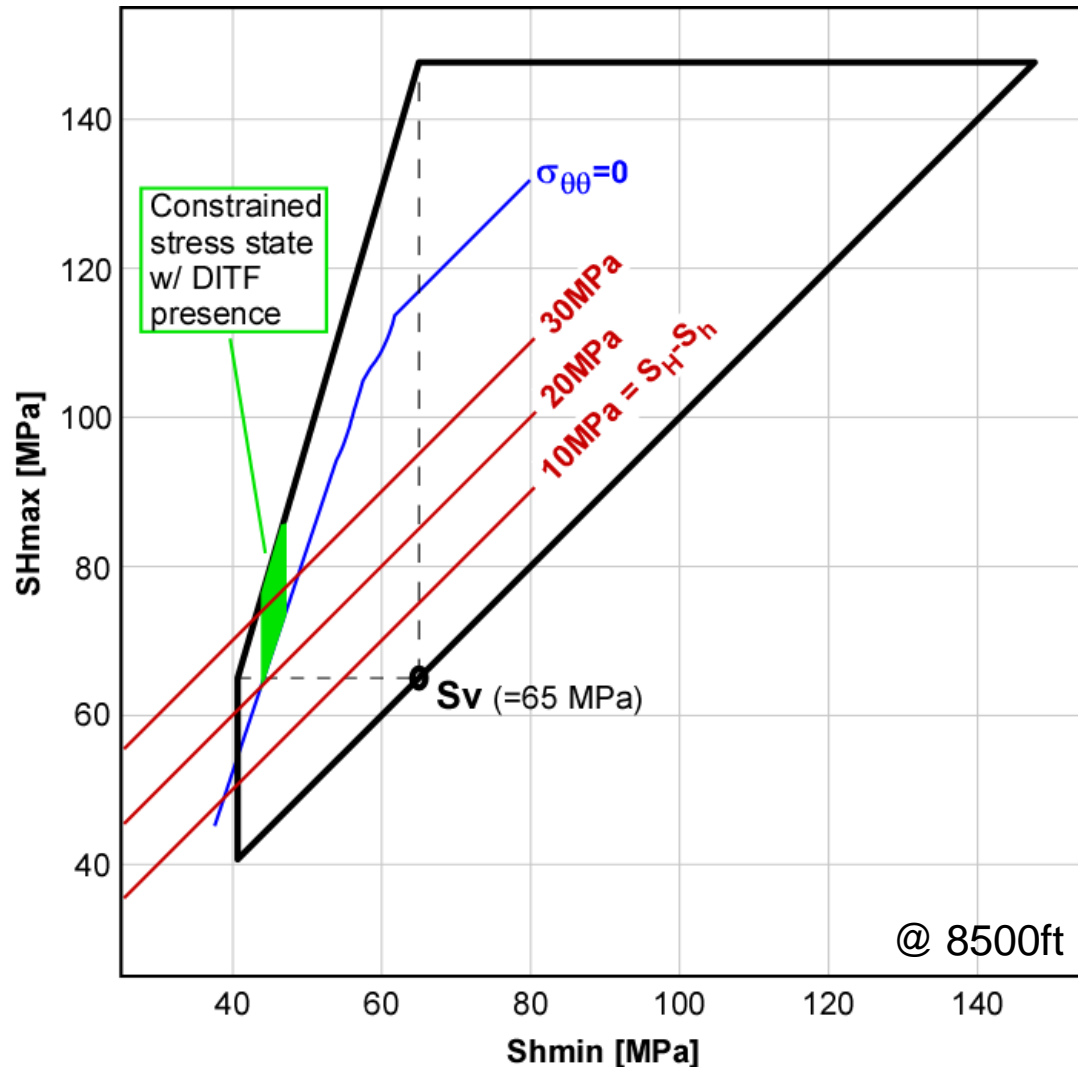
Stress Variation in the Barnett Shale



- DITF's disappears frequently in clay-/organic-rich layers
- **Horizontal Stress Differences** ($S_{Hmax} - S_{hmin}$) is relatively small in these clay-/organic-rich layers



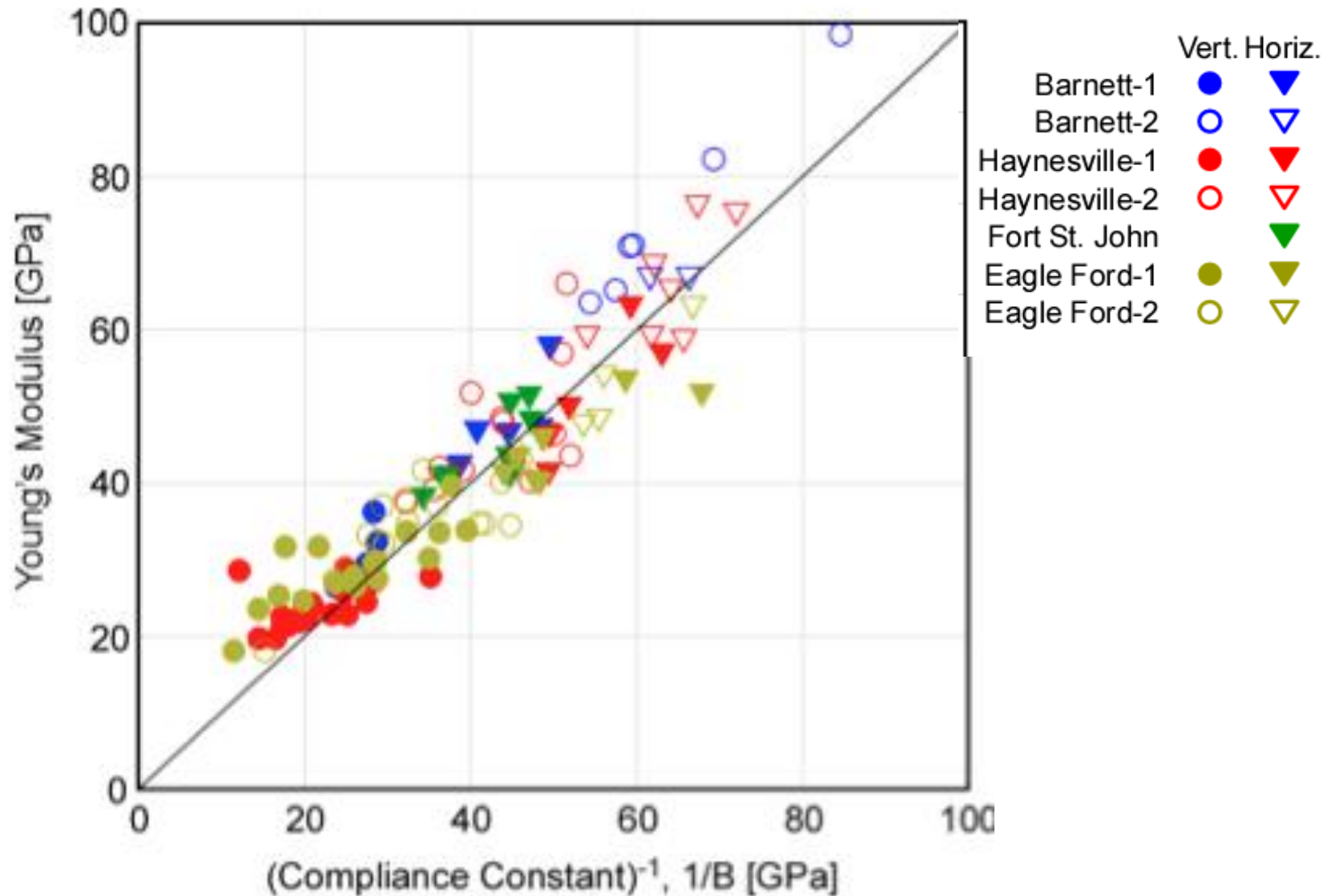
Independent Estimate of Stress Magnitude



- **Green region:** stress state constrained from DITF presence
- **Red lines:** contours of constant horizontal stress difference
- DITF occurrence changes at around 20-25 MPa horizontal stress difference ($S_{Hmax} - S_{hmin}$)

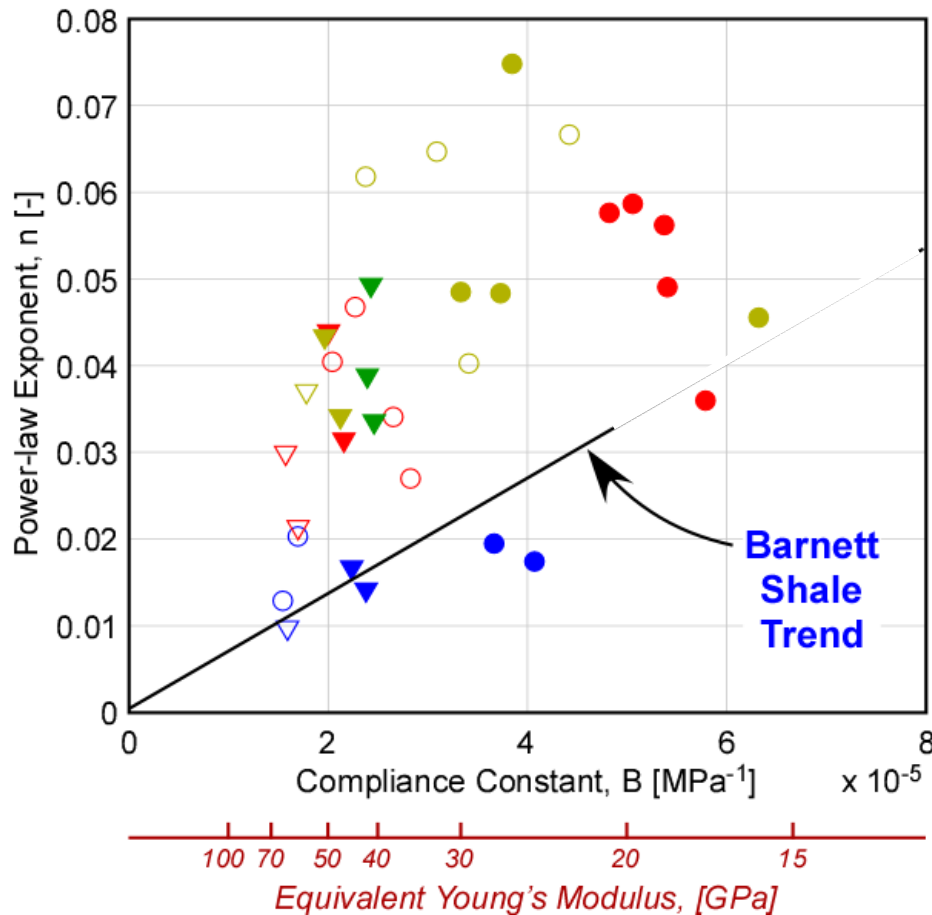


A Useful Correlation





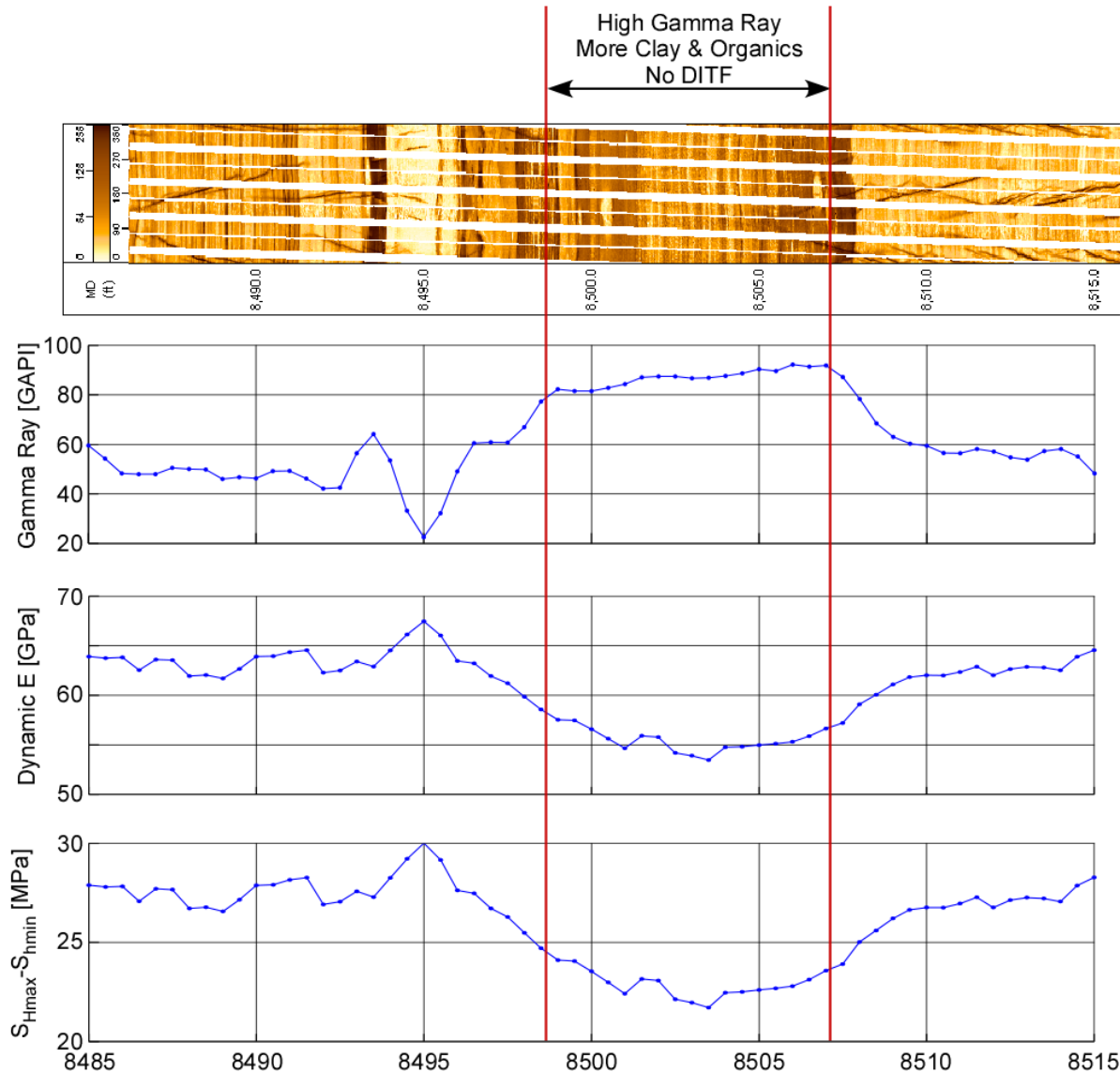
Lab Calibration of “ n ” for the Barnett Shale



- From sonic log data,
→ Dynamic E
(\approx Static E)
- From empirical trend
 $n \approx 2000/3 \cdot B$,
→ constant B
→ constant n
- From viscoelastic theory
→ stress difference



Stress Variations in the Barnett Shale



- Stress analysis predicts lowered differential stress at where DITF is absent
- Transition of DITF presence/absence occurs at somewhere below 25 MPa horiz. stress diff.



Summary

- Organic Rich Shales Creep at Room Temperature
- Clay + Kerogen Content Affects How Much Creep Occurs
- A Power Law Constitutive Law Seems to Describe Creep Behavior
- The Constitutive Law (and Lab-Determined Parameters) Predict Reasonable Values for Stress Anisotropy
- Our Current Research is Focusing on the Creep Behavior of Specific Shales with Mineralogy and Other Factors