

# **Depositional Conditions Associated with Bank-attached Separation Bars, Brushy Canyon Formation\***

**Anjali M. Fernandes<sup>1</sup>, Andrew L. Petter<sup>2,3</sup>, David Mohrig<sup>2</sup>, and Ronald J. Steel<sup>2</sup>**

Search and Discovery Article #50718 (2012)\*\*

Posted September 17, 2012

\*Adapted from oral presentation at AAPG Annual Convention and Exhibition, Long Beach, California, April 22-25, 2012

\*\*AAPG©2012 Serial rights given by author. For all other rights contact author directly.

<sup>1</sup>Dept. of Geological Sciences, University of Texas at Austin, Austin, TX ([anjalifernandes@mail.utexas.edu](mailto:anjalifernandes@mail.utexas.edu))

<sup>2</sup>Dept. of Geological Sciences, University of Texas at Austin, Austin, TX

<sup>3</sup>St. Anthony Falls Laboratory, University of Minnesota, Minneapolis, MN

## **Abstract**

The Brushy Canyon Formation, a predominantly fine-grained turbidite system, was deposited on the slope and basin floor of the late Paleozoic Delaware Basin. Our project focuses on resolving intra-channel sediment sorting within upper-slope channel deposits, and comparing these deposits to channel fills on the proximal basin floor.

The depositional facies on the upper slope fall into two broad classes: A) open-channel facies associated with bypass of sediment to deeper water; and B) channel-filling facies associated with bed aggradation and significant loss of channel relief. Deposits accumulating during bypass are interpreted to be eddy bars located in bank-attached zones of flow separation. These deposits are characterized by packages of steeply inclined beds composed of planar-stratified, trough cross-stratified or sub- to super-critically climbing rippled deposits, with abundant mud drapes ( $D_{50}=110\mu\text{m}$ ). The channel-filling deposits form thick-bedded, sometimes gravel-rich, sandstone bodies which are structureless or which possess stratification associated with migrating dunes and intra-channel barforms ( $D_{50}=156\mu\text{m}$ ). On the proximal basin floor, the channel-filling sandstones ( $D_{50}=110\mu\text{m}$ ) are dominated by stratification associated with trains of dunes climbing at sub- to super-critical angles, indicating high rates of deposition from suspension.

Grain-size analyses show that particles in the 200-400 $\mu\text{m}$  range are common in the channel-filling deposits of upper-slope channels, but are poorly represented in the upper-slope eddy bars and the channel fills on the proximal basin floor. The eddy bars and basin-floor channel fills primarily consist of particles finer than 200 $\mu\text{m}$ , which we interpret as the size fraction that was fully-suspended on the upper slope. This size fraction dominates the eddy-bar deposits because only fully suspended particles can be advected into the bank-attached zones of flow separation in significant volumes. We will synthesize depositional styles and grain-size data in order to:

1) produce a facies model for thick bank-attached bar deposits built in zones of flow separation associated with planform irregularity in submarine channels, 2) estimate flow velocities and current thicknesses; and 3) assess sediment sorting and storage between channels on the upper slope and the proximal basin floor.

### References

- Beaubouef, R.T., C.R. Rossen, F.B. Zelt, M.D. Sullivan, D.C. Mohrig, D.C. Jennette, J.A. Bellian, S.J. Firedman, R.W. Lovell, and D.S. Shannon, 1999, (eds.), Deep-water sandstones, Brushy Canyon Formation, West Texas, unpaginated.
- Dietrich, W.E., 1982, Settling velocity of natural particles: *Water Resources Research*, v. 18/6, p. 1651-1626.
- Dietrich, W.E., and J.D. Smith, 1984, Processes controlling the equilibrium bed morphology in river meanders, *in* C.M. Elliott, (ed.), *River meandering: American Society of Civil Engineers*, New York, p. 759-769.
- Fitchen, W.M., 1997, Lower Permian sequence stratigraphy of the western Delaware Basin margin, Sierra Diablo, West Texas: Ph.D. Dissertation, University of Texas at Austin, 264 p.
- Hickin, E.J., 1979, Concave-bank benches on the Squamish River, British Columbia, Canada: *Canadian Journal of Earth Sciences*, v. 16/1, p. 200-203.
- Rubin, Y., S.S. Hubbard, A. Wilson, and M.A. Cushey, 1998, *Aquifer characterization: Handbook of Groundwater Hydrology*, CRC Press, p. 10.1-10.33.
- Rubin, D.M., J.C. Schmidt, and J.N. Moore, 1990, Origin, structure, and evolution of a reattachment bar, Colorado River, Grand Canyon, Arizona: *Journal of Sedimentary Petrology*, v. 60/6, p. 982-991.
- Smith, J.D., 1977, Modeling of sediment transport on continental shelves, *in* E.D. Goldberg, I.N. McCave, J.J. O'Brien, and J.H. Steele, (eds.), *The sea; ideas and observations on progress in the study of the seas; Volume 6, Marine modeling: John Wiley & Sons*, New York, New York, p. 539-577.
- Smith, D.G., S.M. Hubbard, D.A. Leckie, and M. Fustic, 2009, Counter point bar deposits; lithofacies and reservoir significance in the meandering modern Peace River and ancient McMurray Formation Alberta, Canada: *Sedimentology*, v. 56/6, p. 1655-1669.

Wiberg, P.L., and J. D. Smith, 1987, Calculations of the critical shear stress for motion of uniform and heterogeneous sediments: Water Resources Research, v. 23/8, p. 1471-1480.

Woodyer, K.D., 1975, Concave-bank benches on Barwon River, N.S.W.: Australian Geographer, v. 13/1, p. 36-40.

Wright, F.W., 1962, Abo reef: prime west Texas target: Oil and Gas Journal, v. 60, p. 188-194.

# Depositional Conditions Associated with Bank-attached Separation Bars, Brushy Canyon Formation

Anjali M. Fernandes<sup>1</sup>, Andrew L. Petter<sup>1,2</sup>,  
David Mohrig<sup>1</sup>, Ronald J. Steel<sup>1</sup>

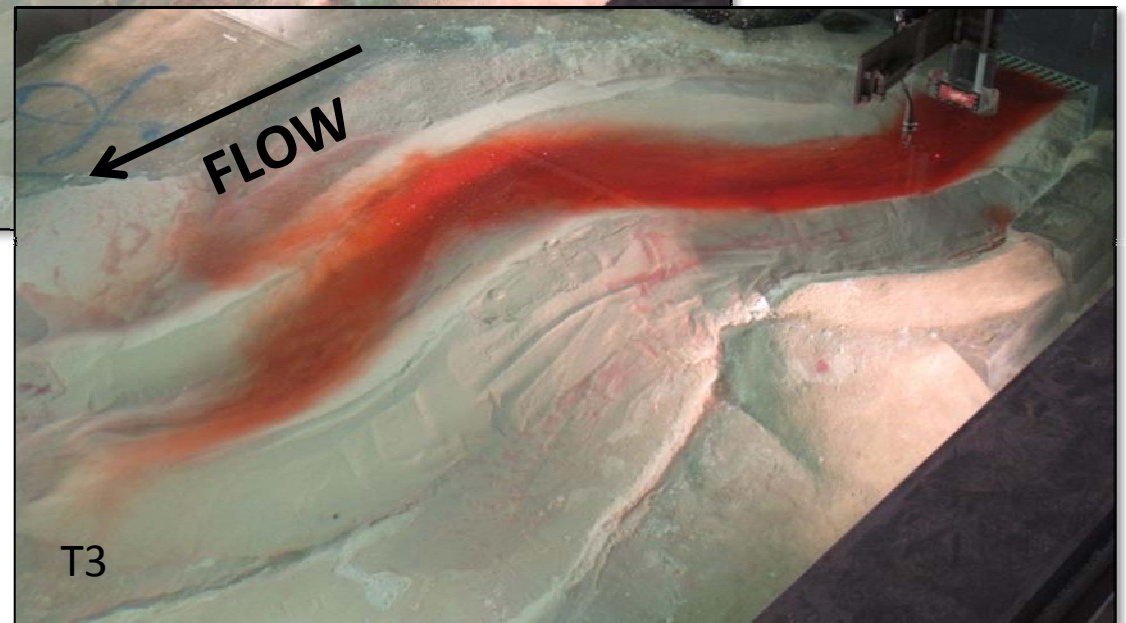
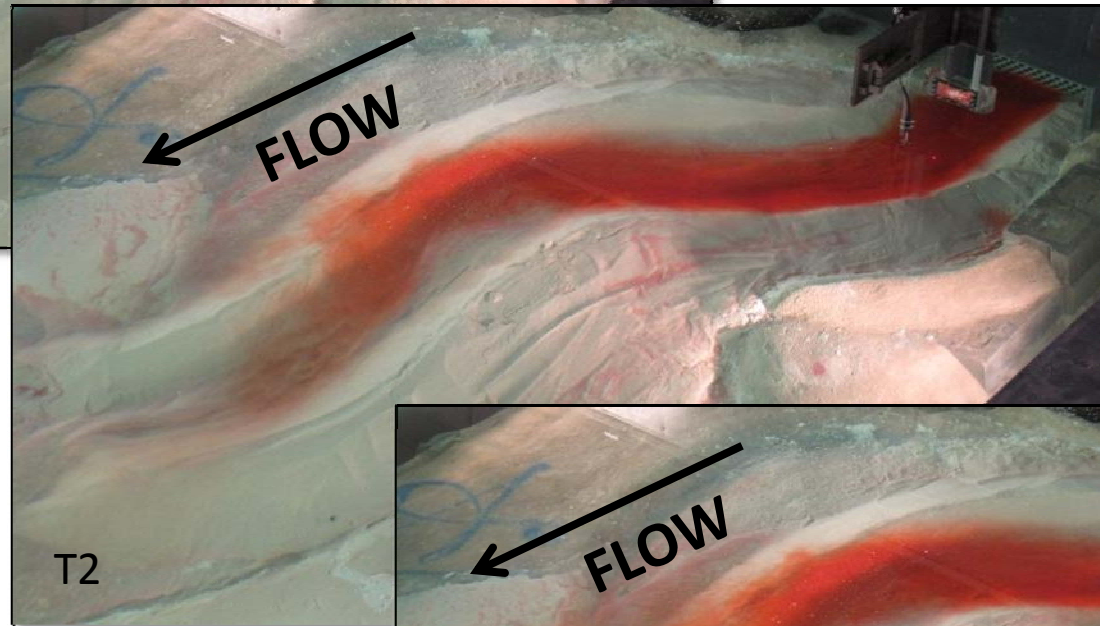
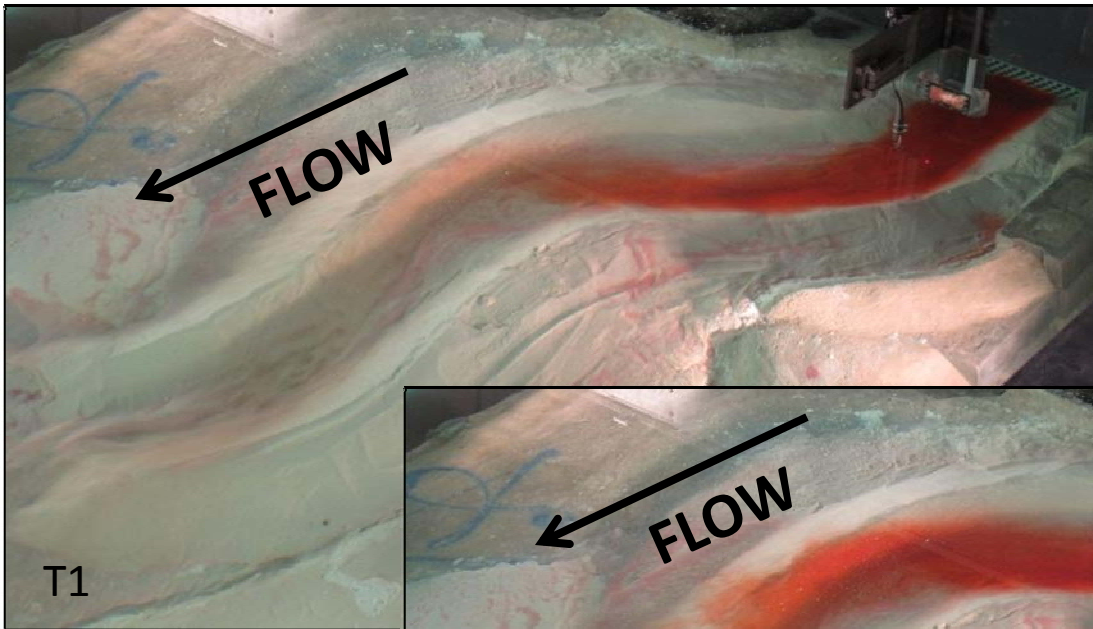
<sup>1</sup> The University of Texas at Austin

<sup>2</sup> St. Anthony Falls Laboratory, University of Minnesota





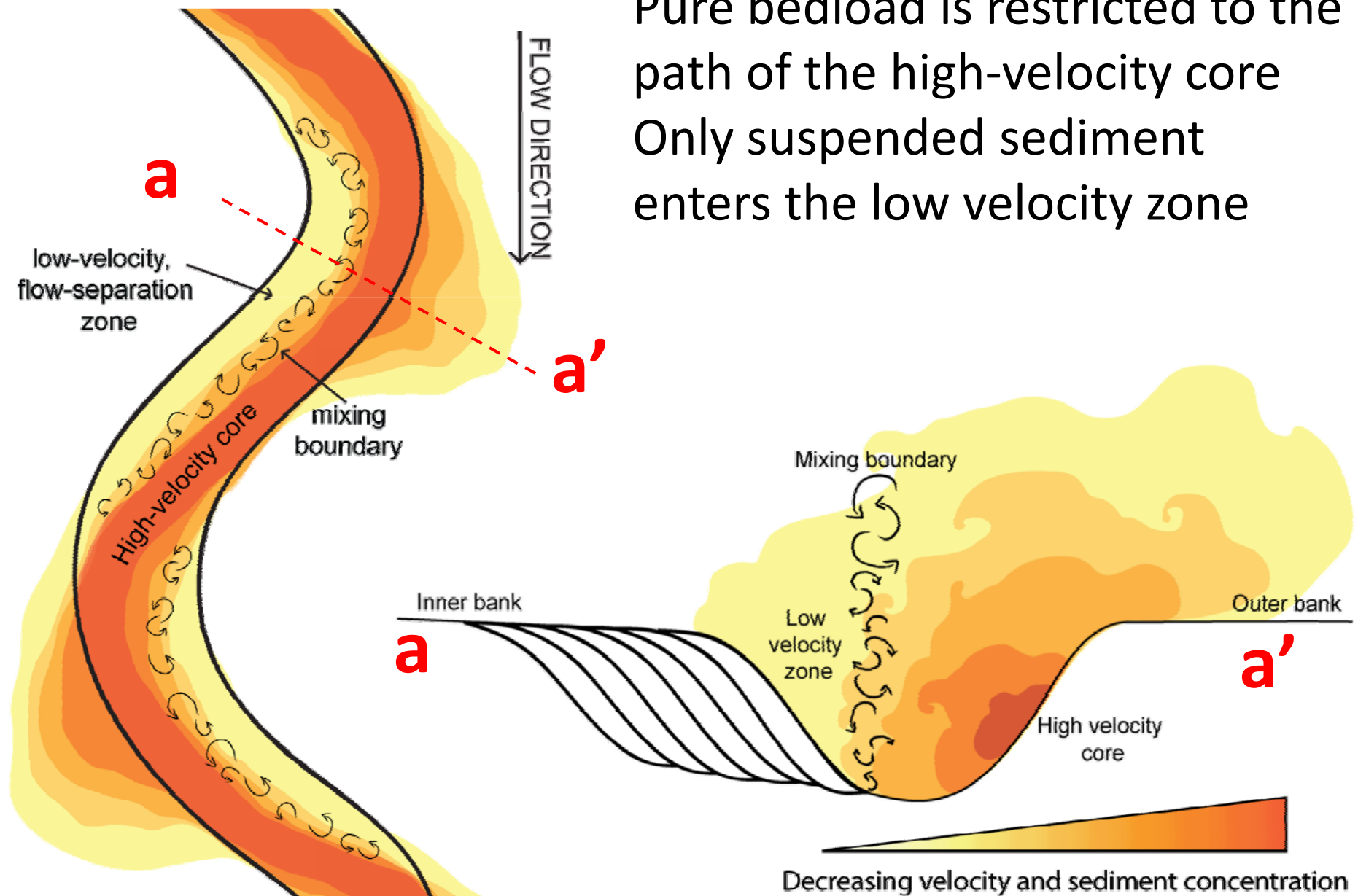
# Turbidity currents in submarine channel bends



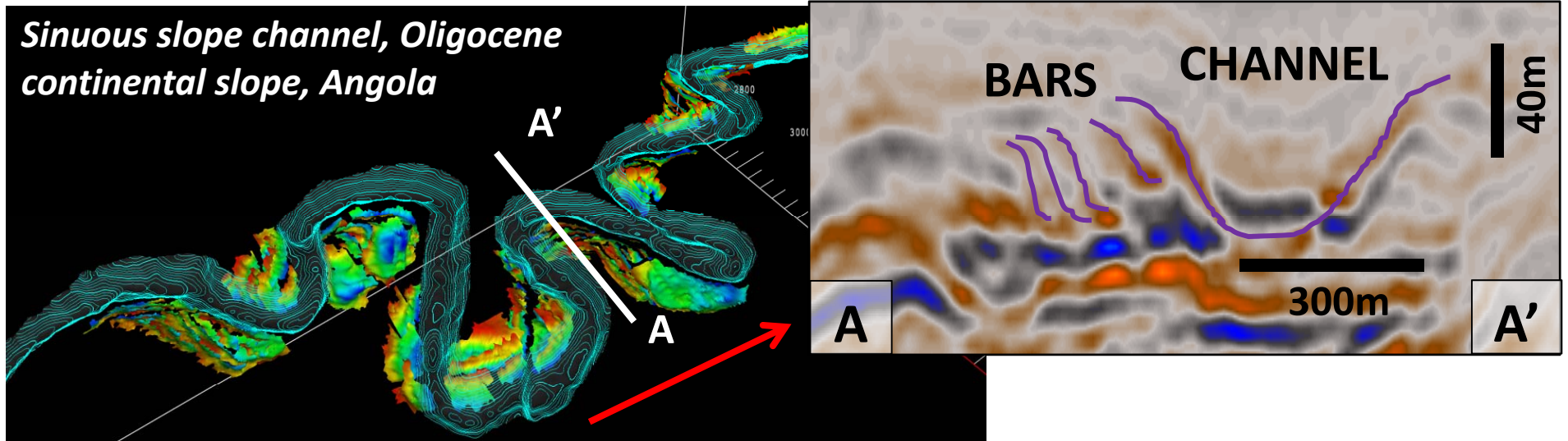
Run-up at the outer banks of bends contributes to extreme flow separation from the inner banks

# Flow separation at channel bends

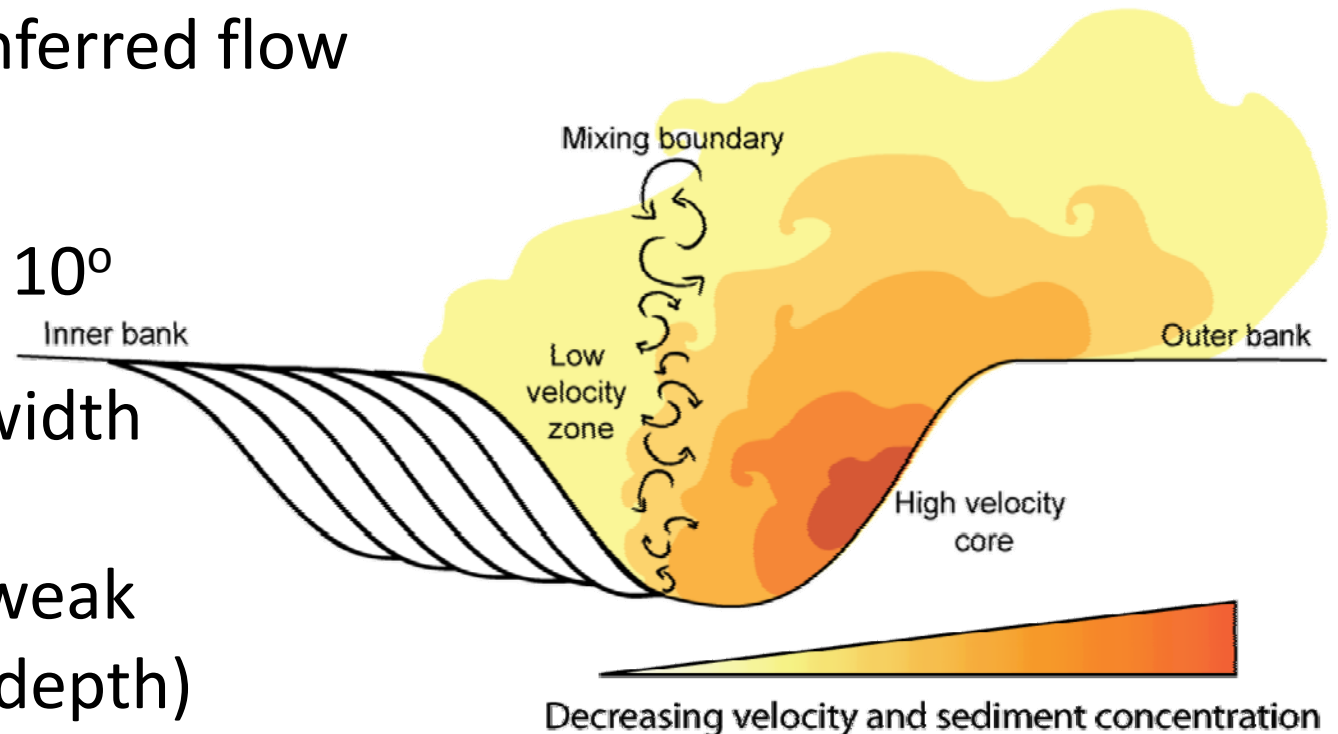
Pure bedload is restricted to the path of the high-velocity core  
Only suspended sediment enters the low velocity zone



# 3-D geometries to processes of bar construction



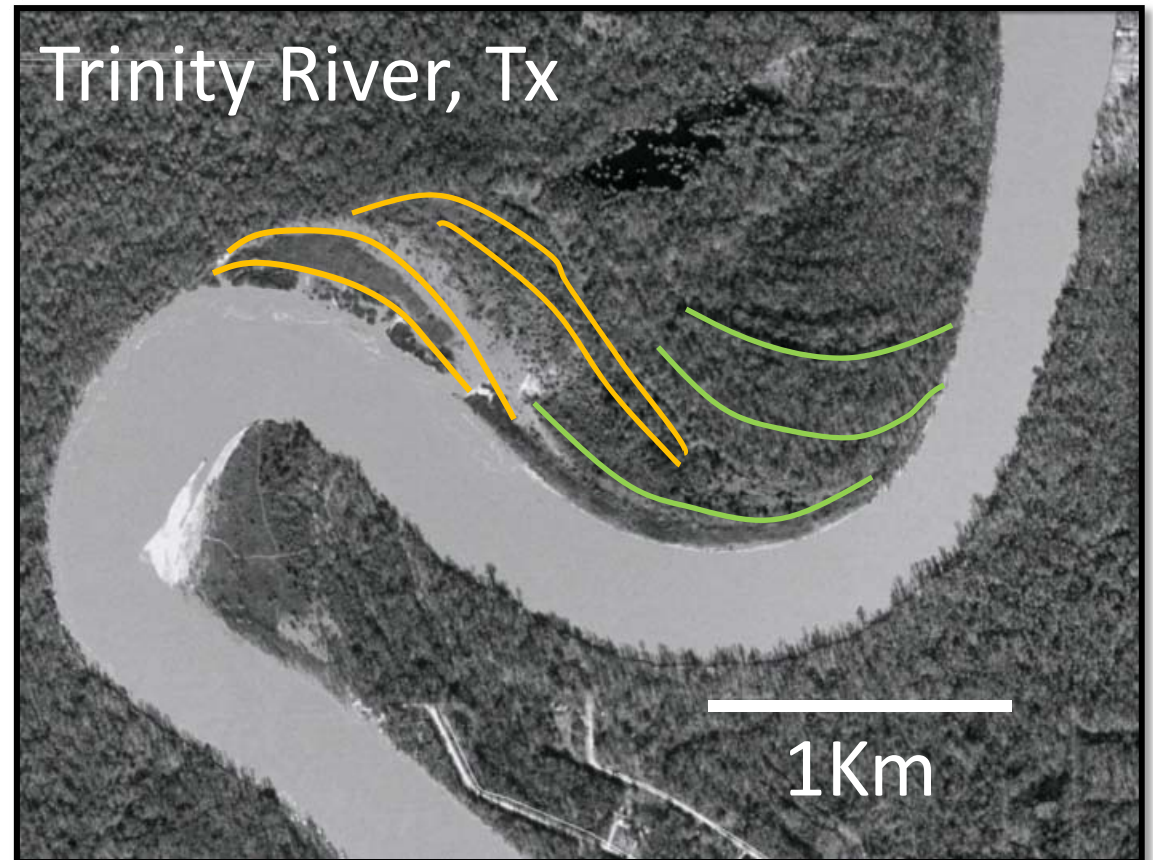
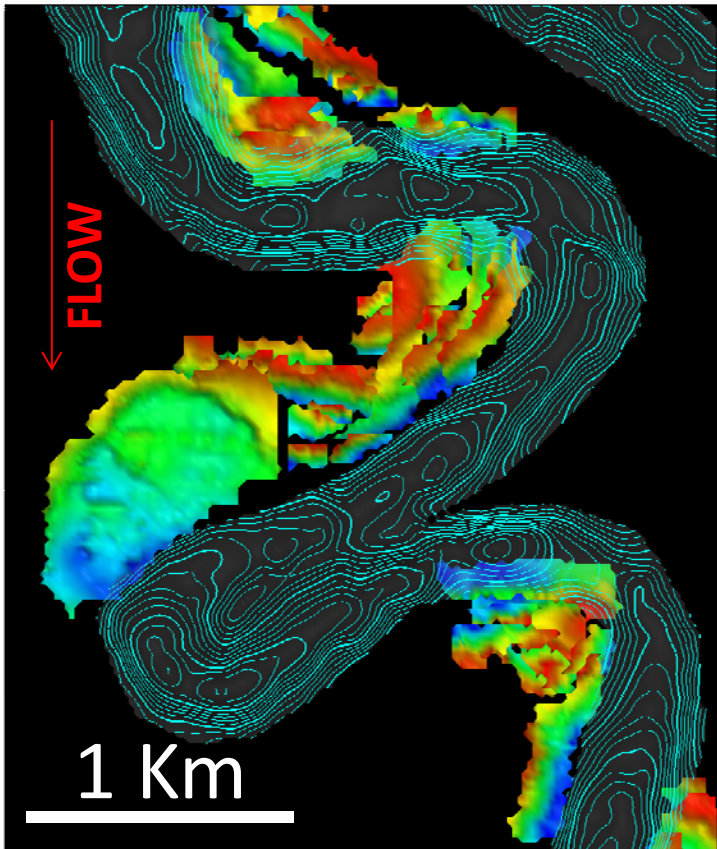
- Constructed in inferred flow separation zones
- Median slopes =  $10^\circ$
- Narrow relative width
- Associated with weak incision (20-40% depth)



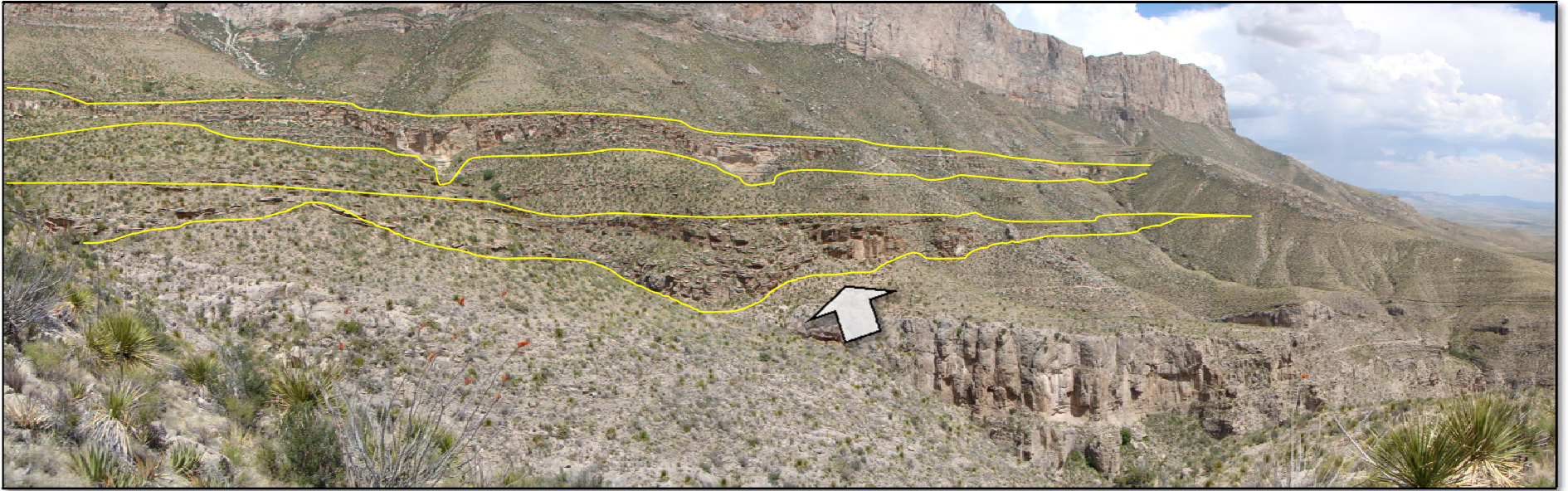


# Concave Benches

- Pure suspension deposition in separation zones  
(*Woodyer, 1975; Hickin, 1979; Smith et al., 2009*)
- Hitherto under-recognized in submarine channels



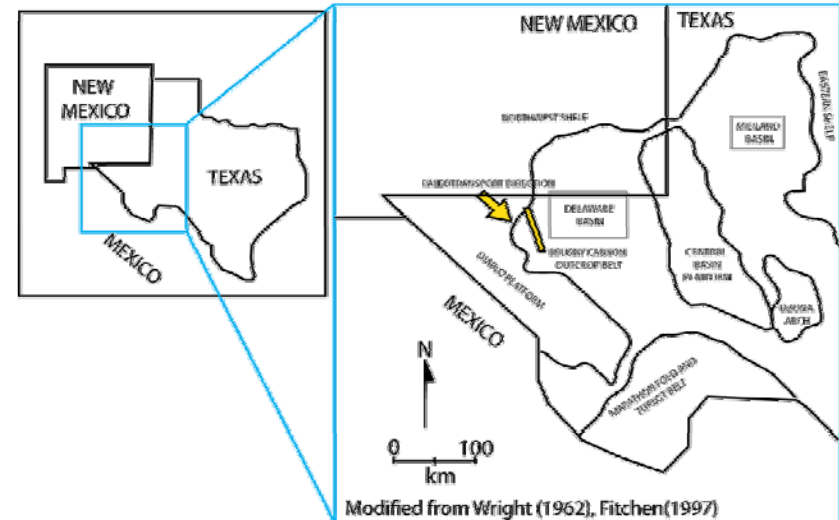
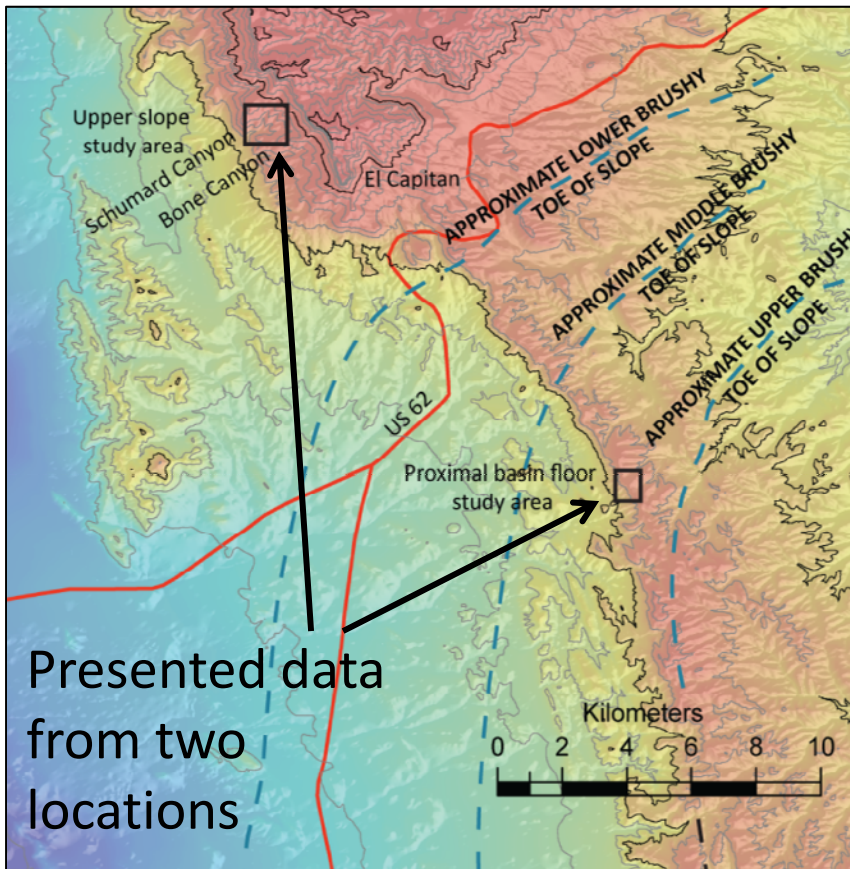
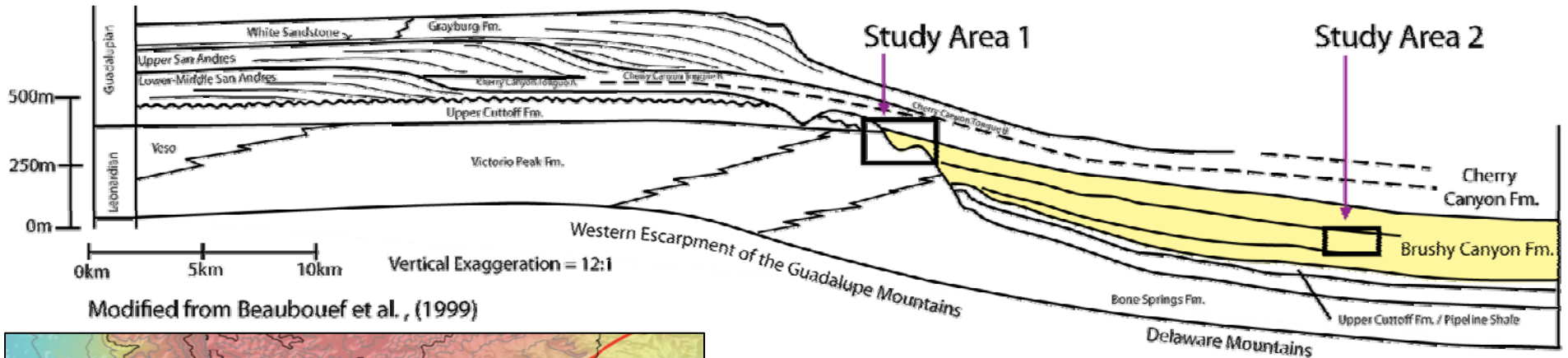
## Upper slope channels in the Brushy Canyon Fm.



Thick submarine slope channel fills exposed in dip-oblique section, in Shumard Canyon in the Guadalupe Mountains



# Study locations on slope and basin floor



Clastic sediments deposited on a carbonate margin during low sea level and sub-aerial exposure of carbonate shelf



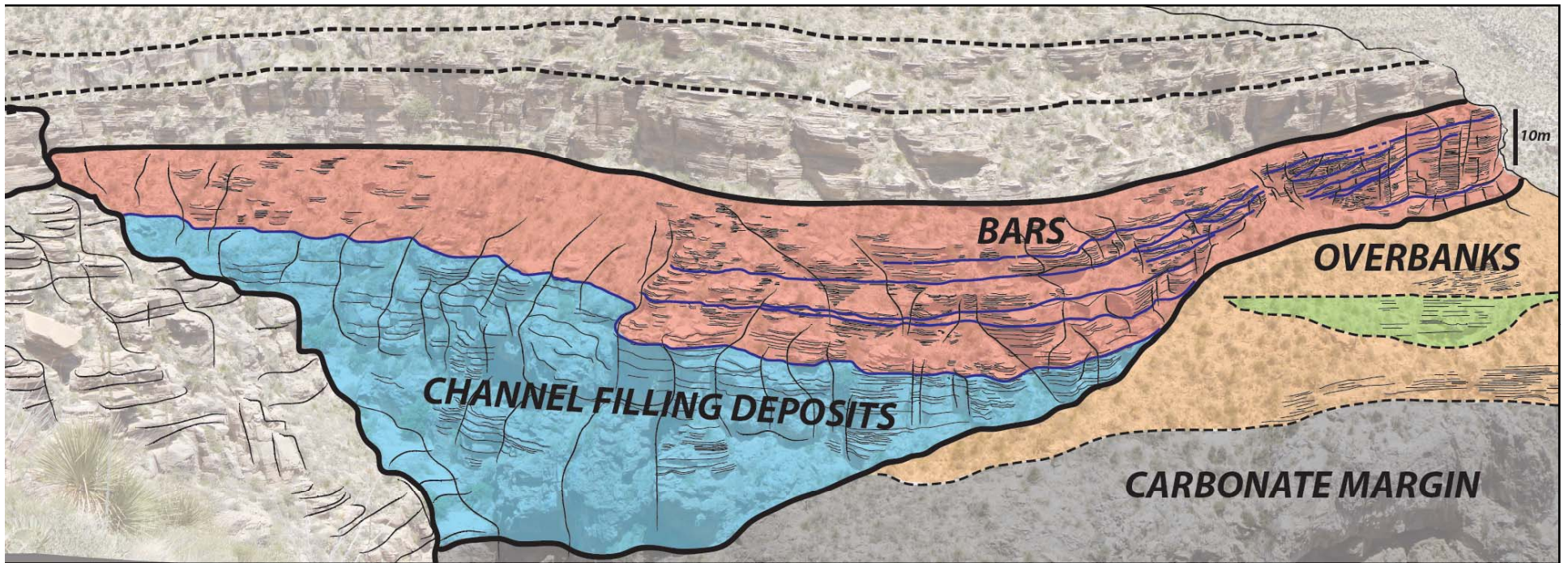
# Deposit styles



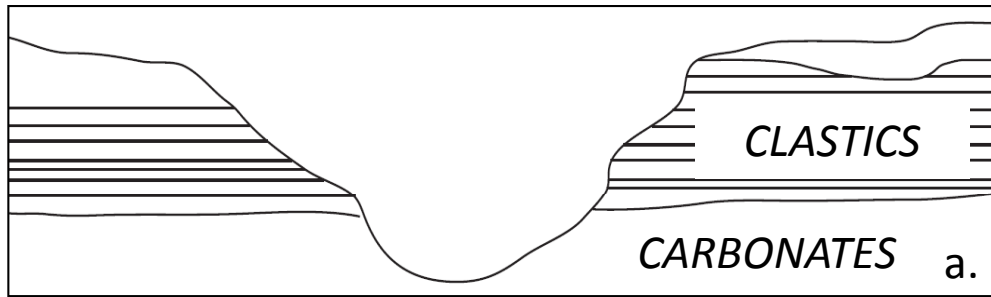
1. **Channel filling deposits:** Thickly-bedded, relatively coarse, bedload filling in topography
2. **A bank-attached bar:** Finer grained sediment settling from suspension in low-velocity zones along channel margins, while channel relief is maintained



# Deposit styles

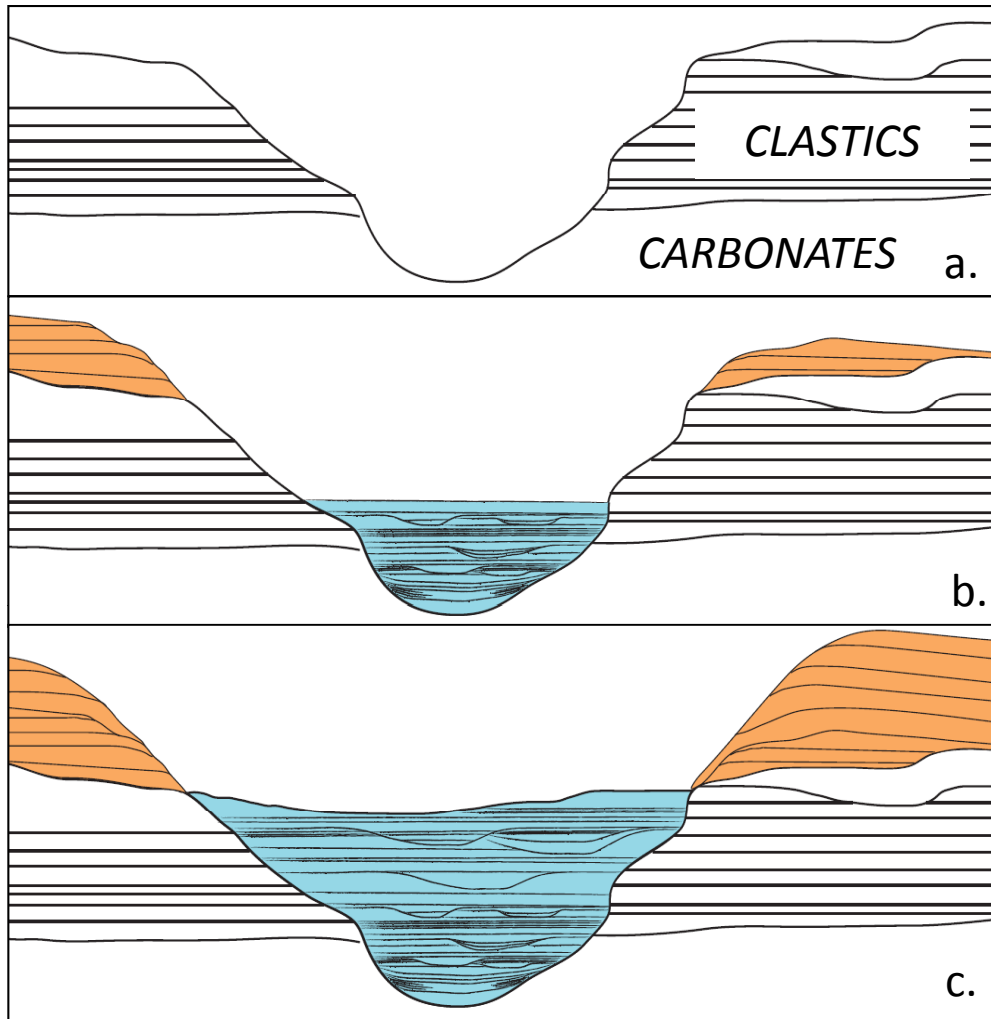


1. **Channel filling deposits:** Thickly-bedded, relatively coarse, bedload filling in topography
2. **A bank-attached bar:** Finer grained sediment settling from suspension in low-velocity zones along channel margins, while channel relief is maintained



Interpreted  
stages of studied  
channel fill

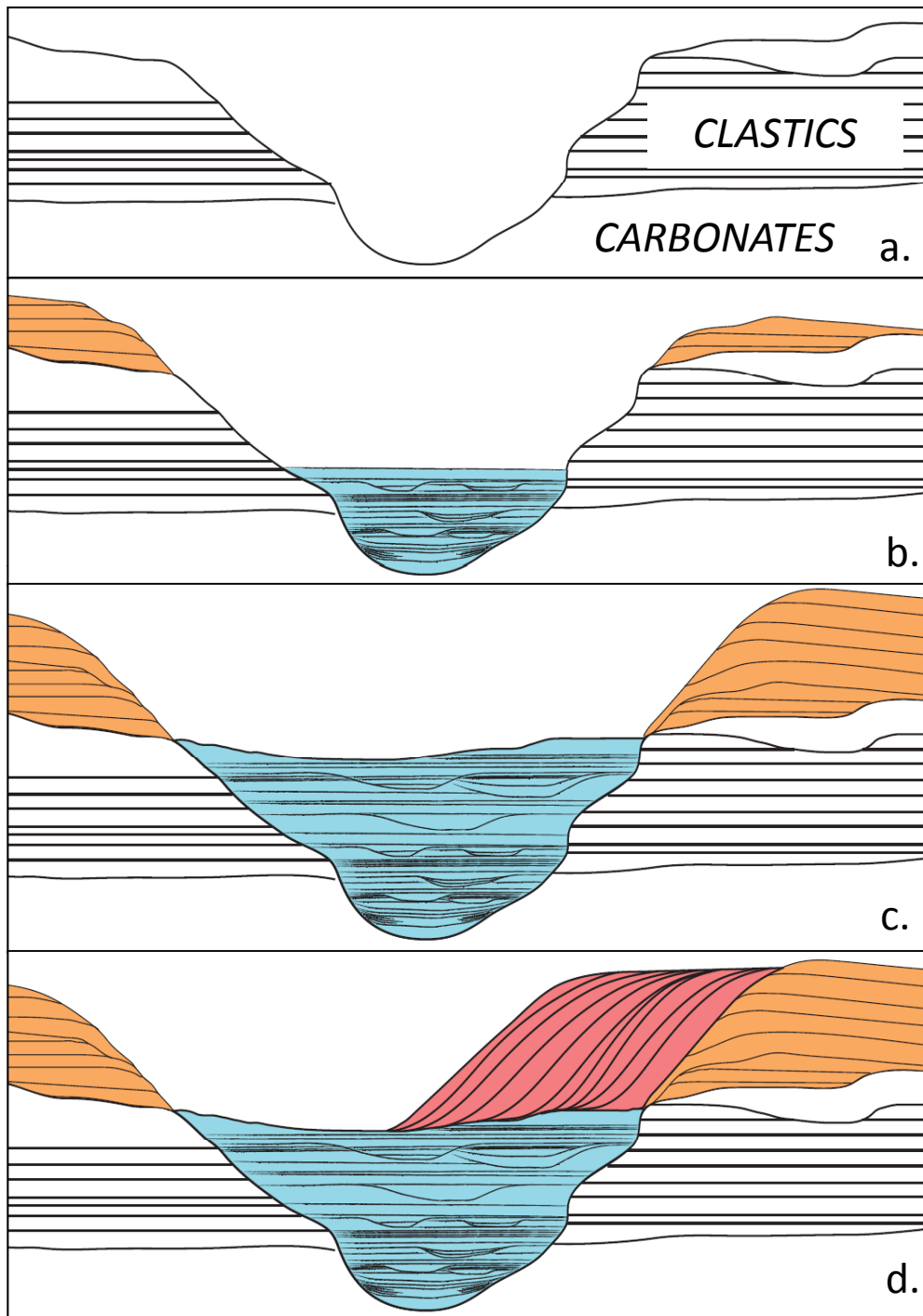
# Interpreted stages of studied channel fill



Stage 1: Channel bed aggradation reduces channel relief



# Interpreted stages of studied channel fill



Stage 1: Channel bed aggradation reduces channel relief

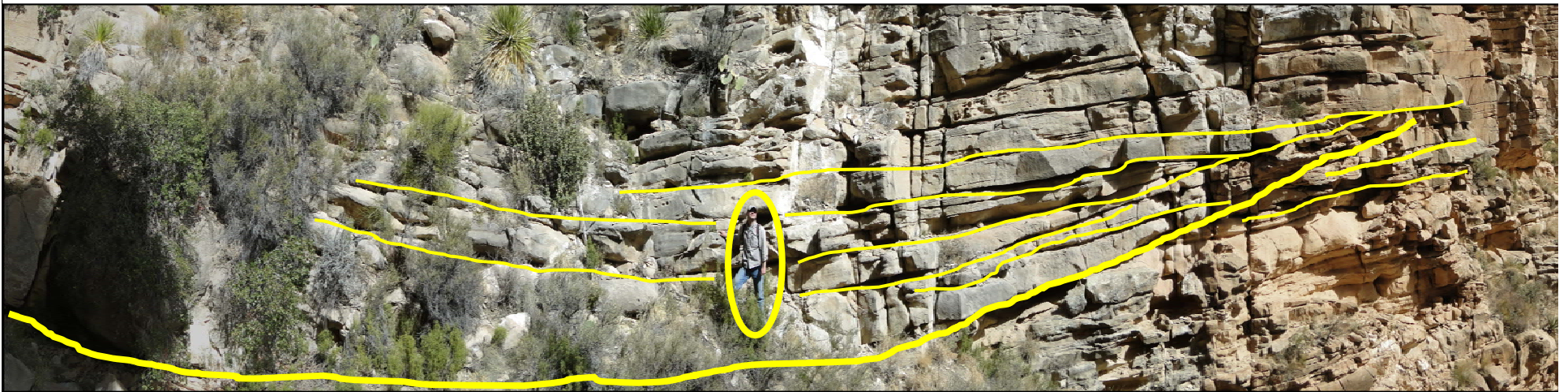
Stage 2: A bar is constructed at the channel margin while channel relief is maintained

# Channel filling deposits

Gravel-rich deposits



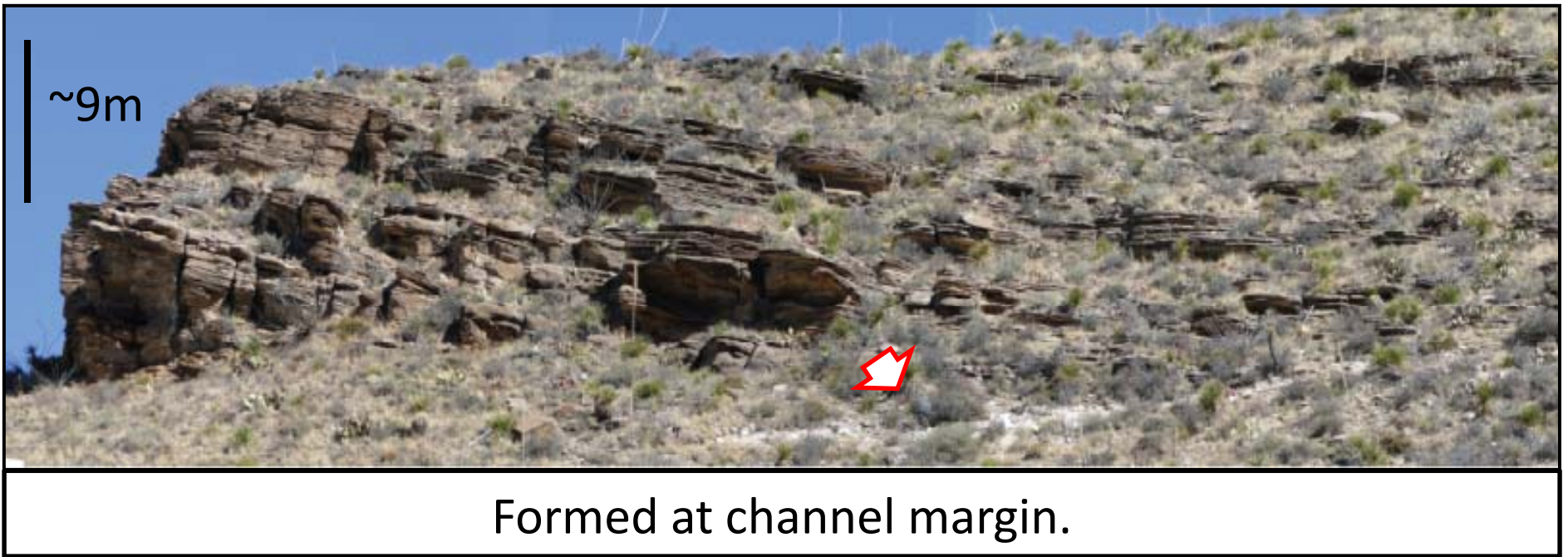
Thickly-bedded, structureless, planar-stratified and dune cross-stratified sands stones. Occasionally gravel-rich.



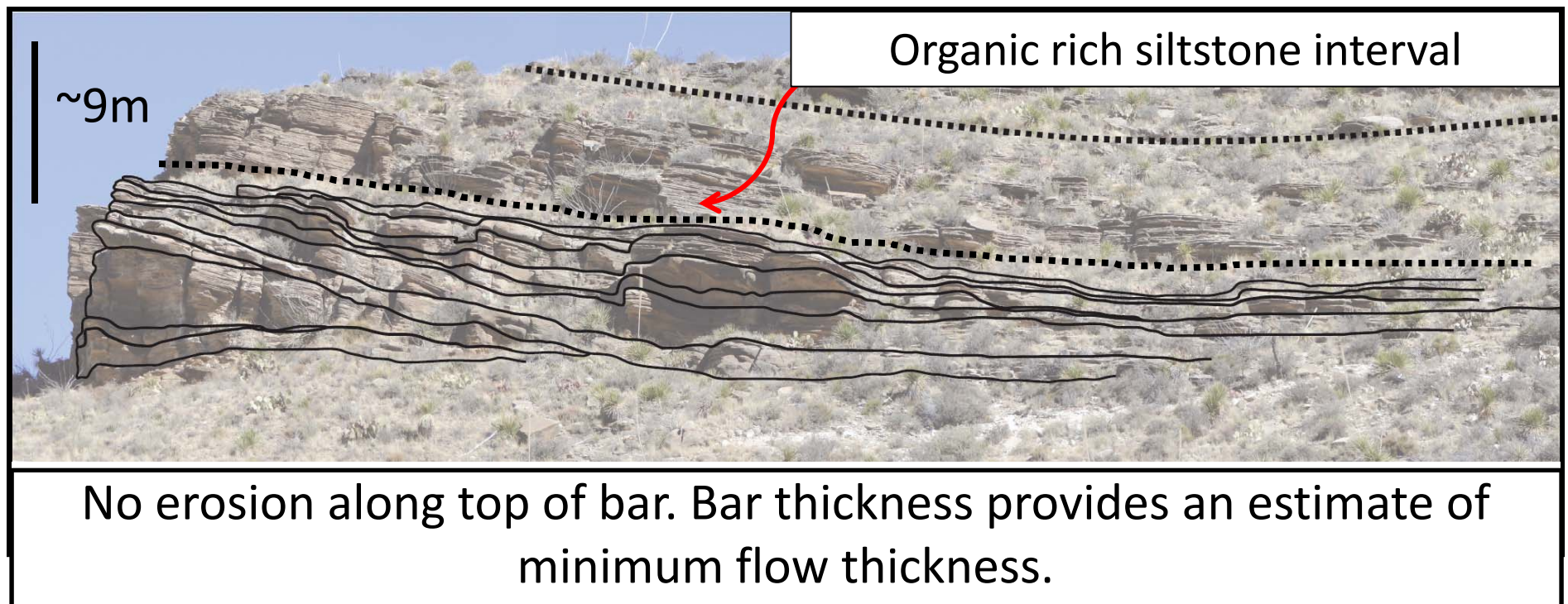
Filling topographic lows



# Bank-attached separation bar

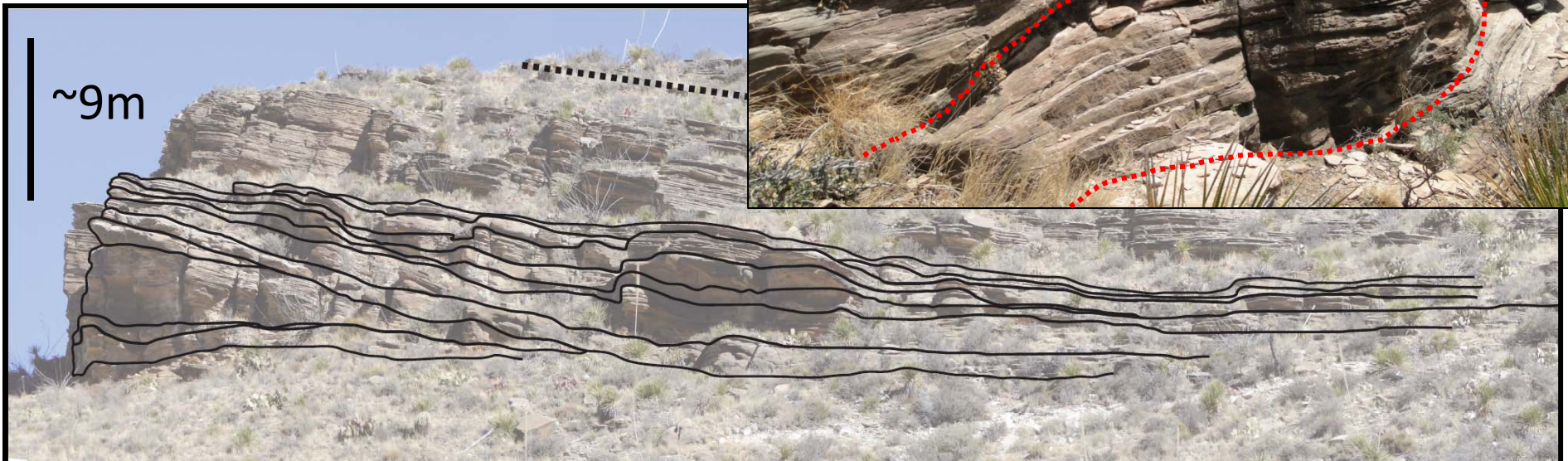
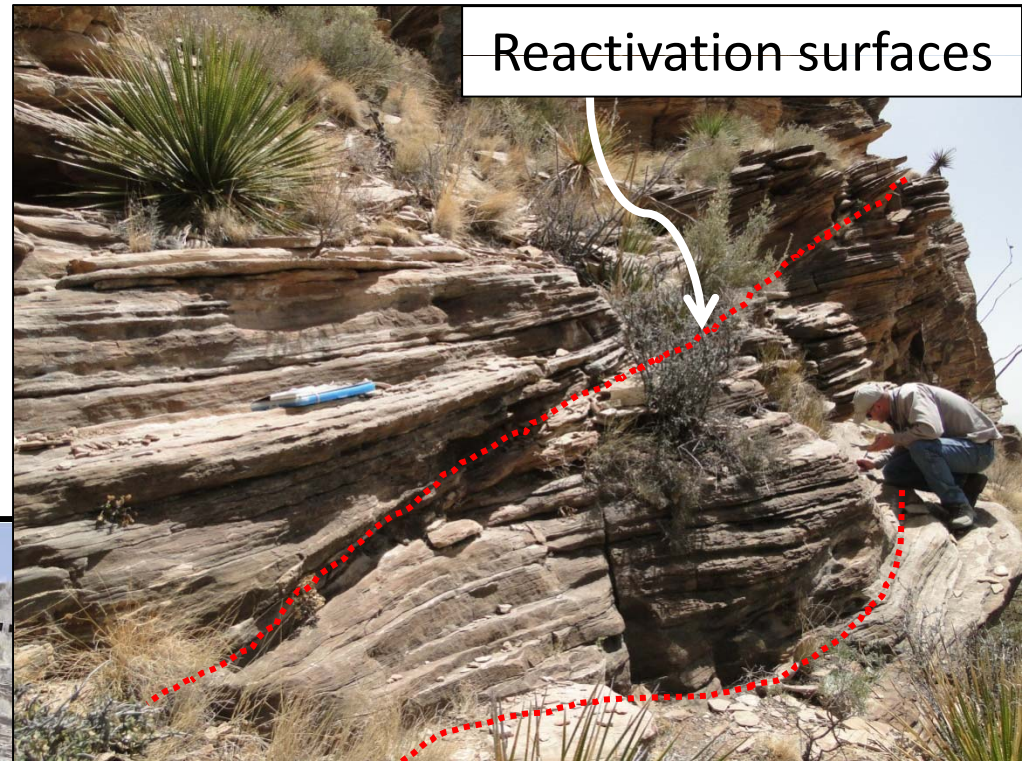
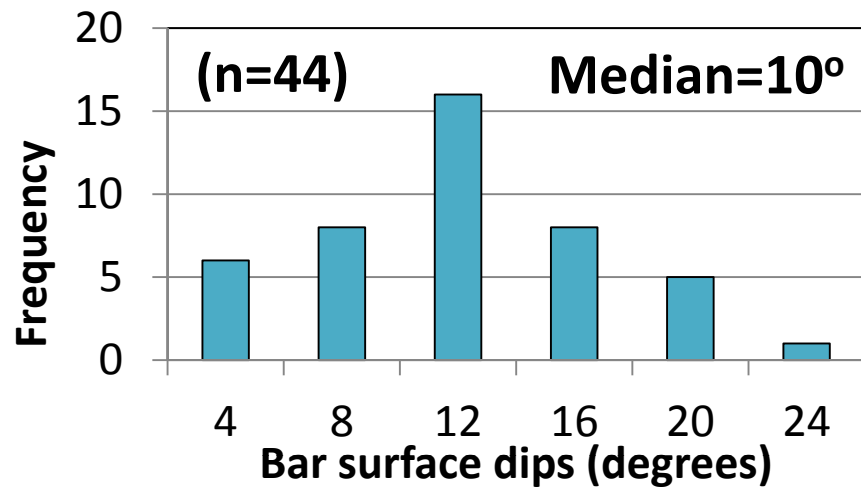


# Bank-attached separation bar





# Bank-attached separation bar

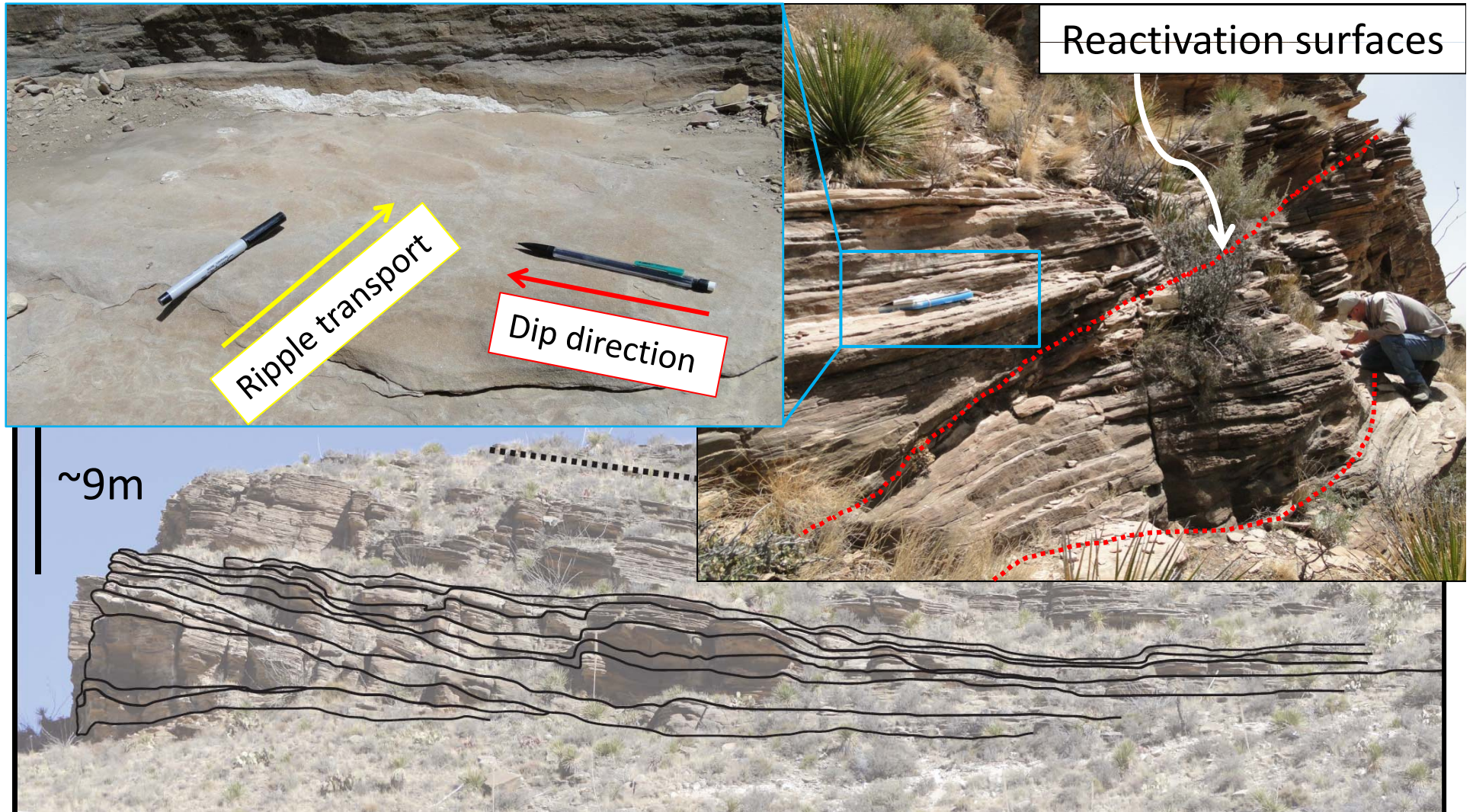


High bedding dips.

Reactivation surfaces associated with the changing shape of the bar.



# Bank-attached separation bar



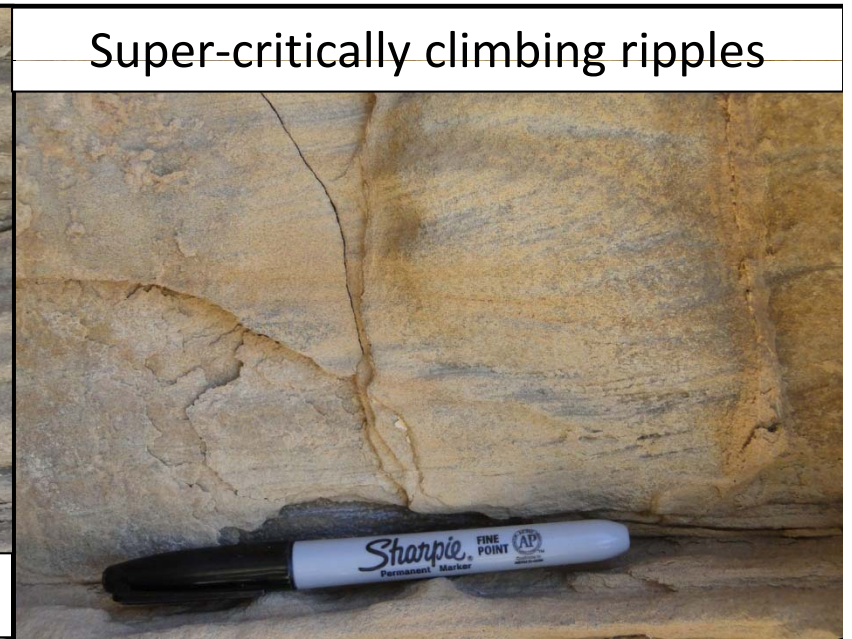
Ripple transport at high angles relative to azimuth of bedding dips.



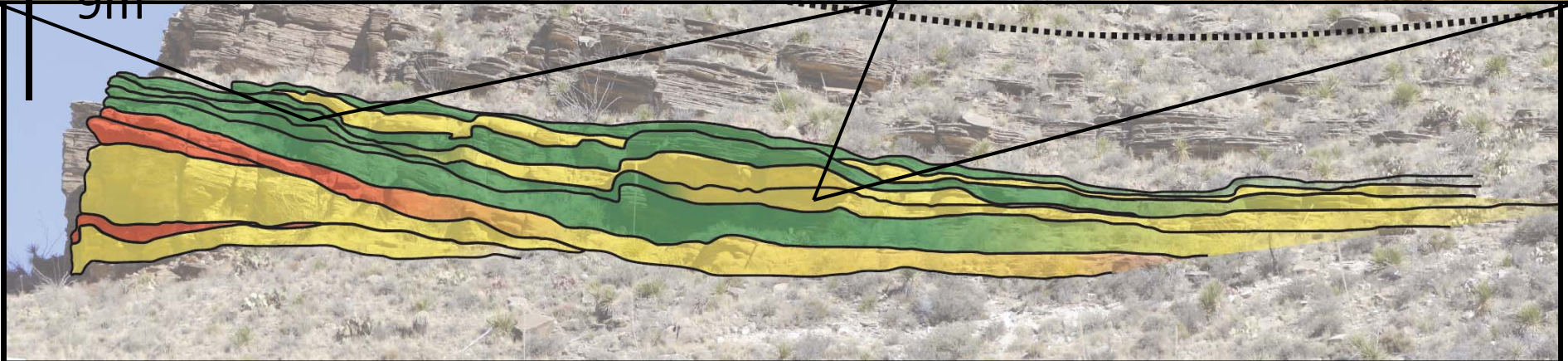
# Bank-attached separation bar






Mud-draped, poorly defined ripples



Super-critically climbing ripples



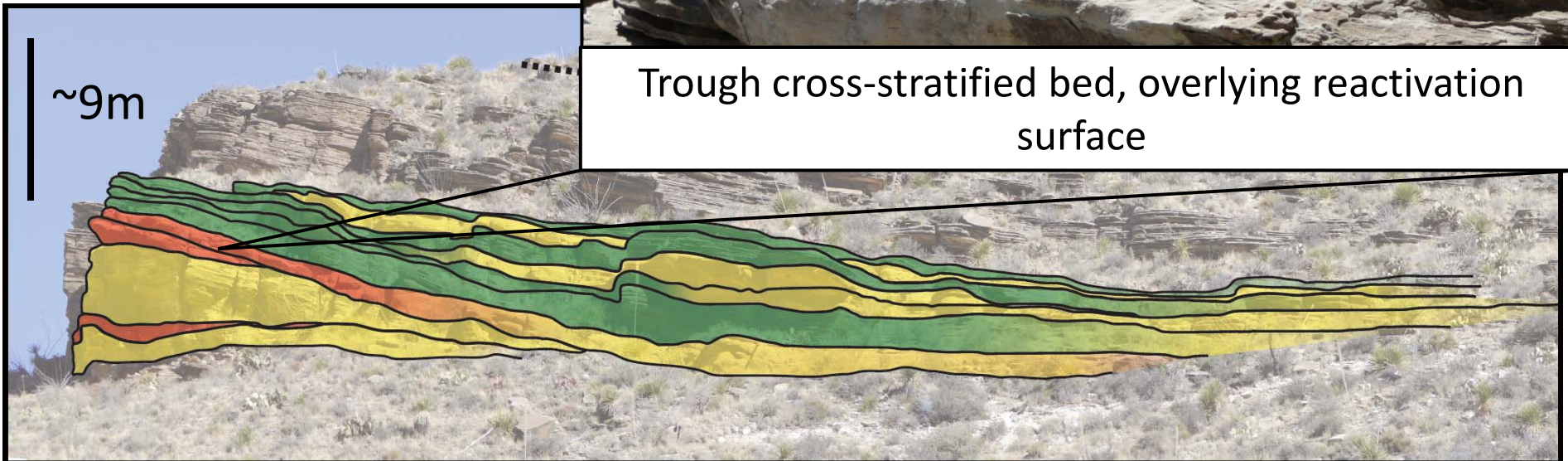
-  RIPPLED SANDS
-  TROUGH CROSS-STRATIFIED SANDS
-  PLANAR STRATIFIED SANDS



# Bank-attached separation bar



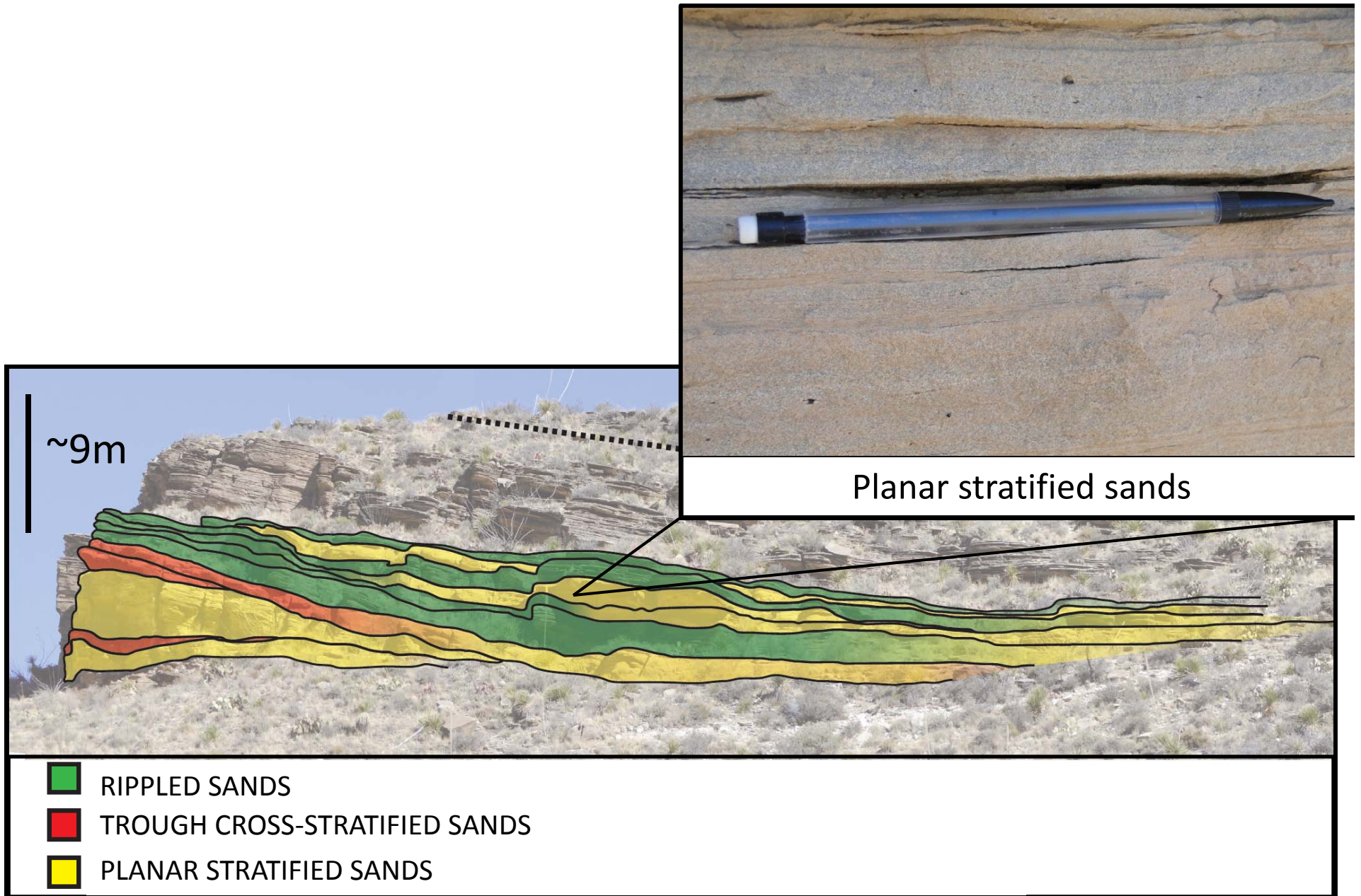
Trough cross-stratified bed, overlying reactivation surface



- RIPPLED SANDS
- TROUGH CROSS-STRATIFIED SANDS
- PLANAR STRATIFIED SANDS

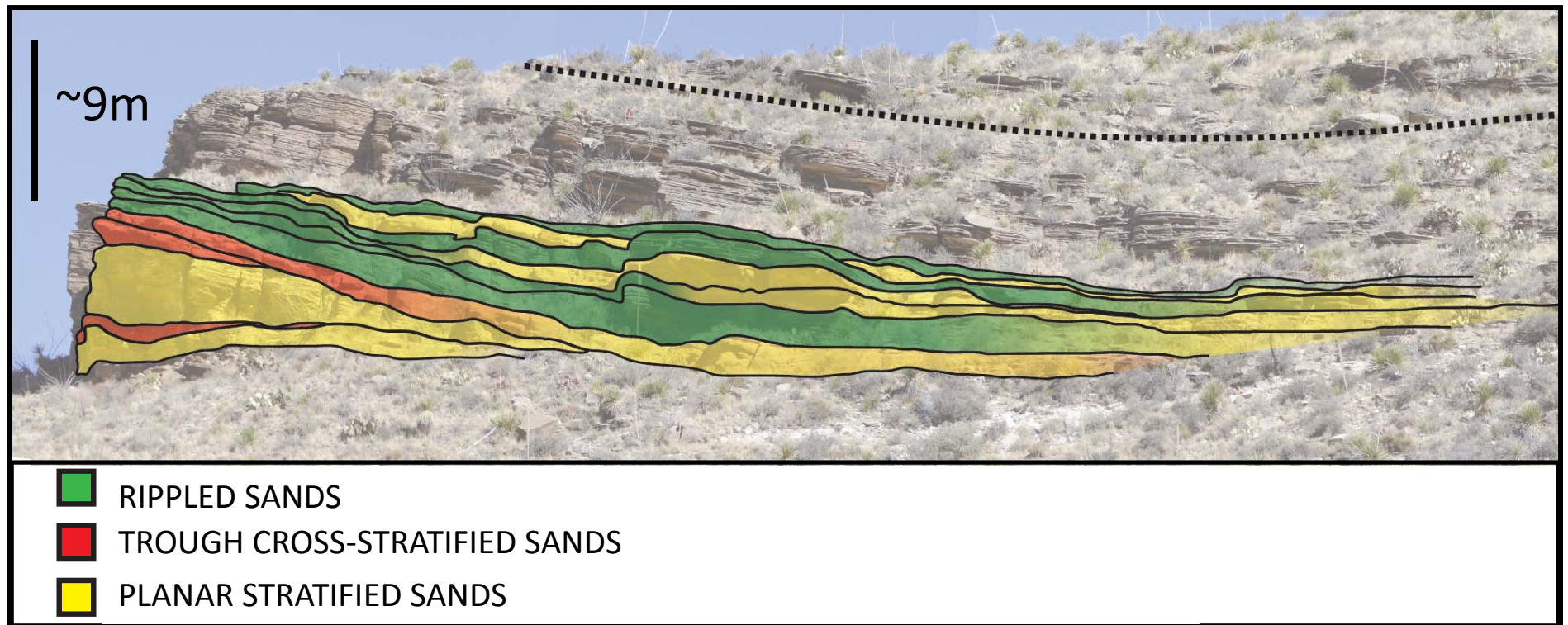


# Bank-attached separation bar



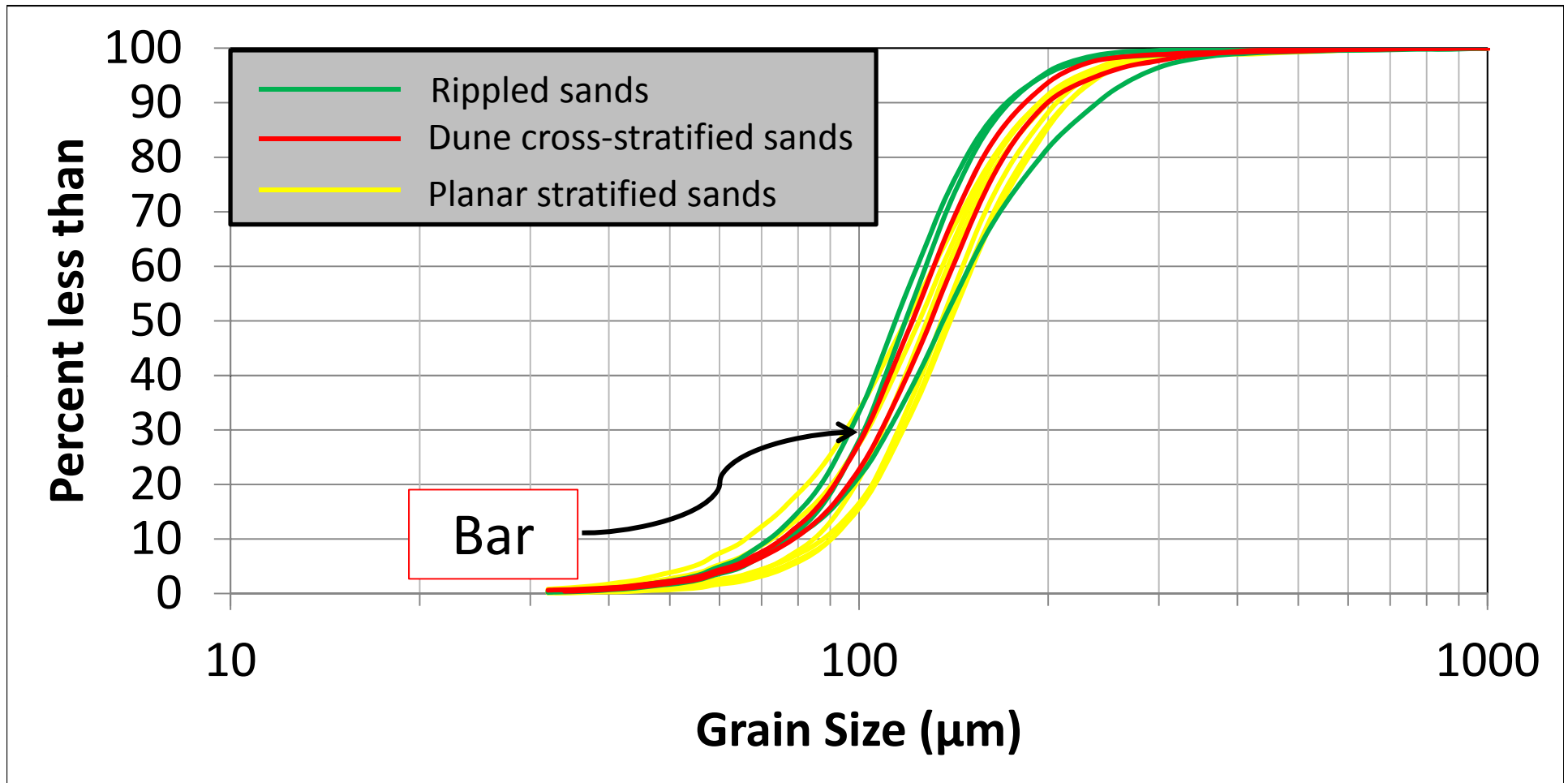
# Bank-attached separation bar

Lateral transitions between facies over small distances





# Bank-attached separation bar

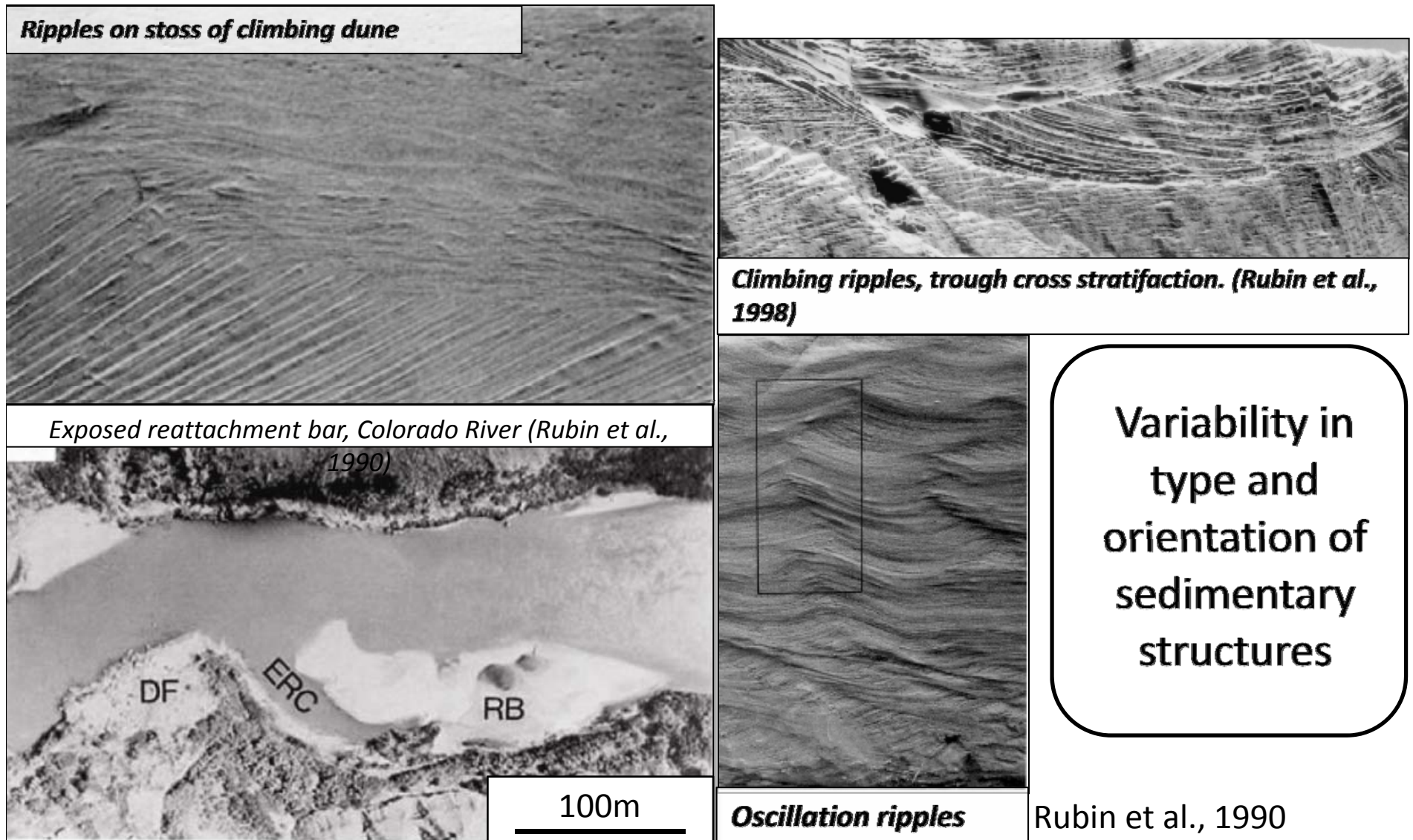


Grain size distributions show **similar composition**.

Abundance of climbing rippled deposits suggest common source from **suspended load**.

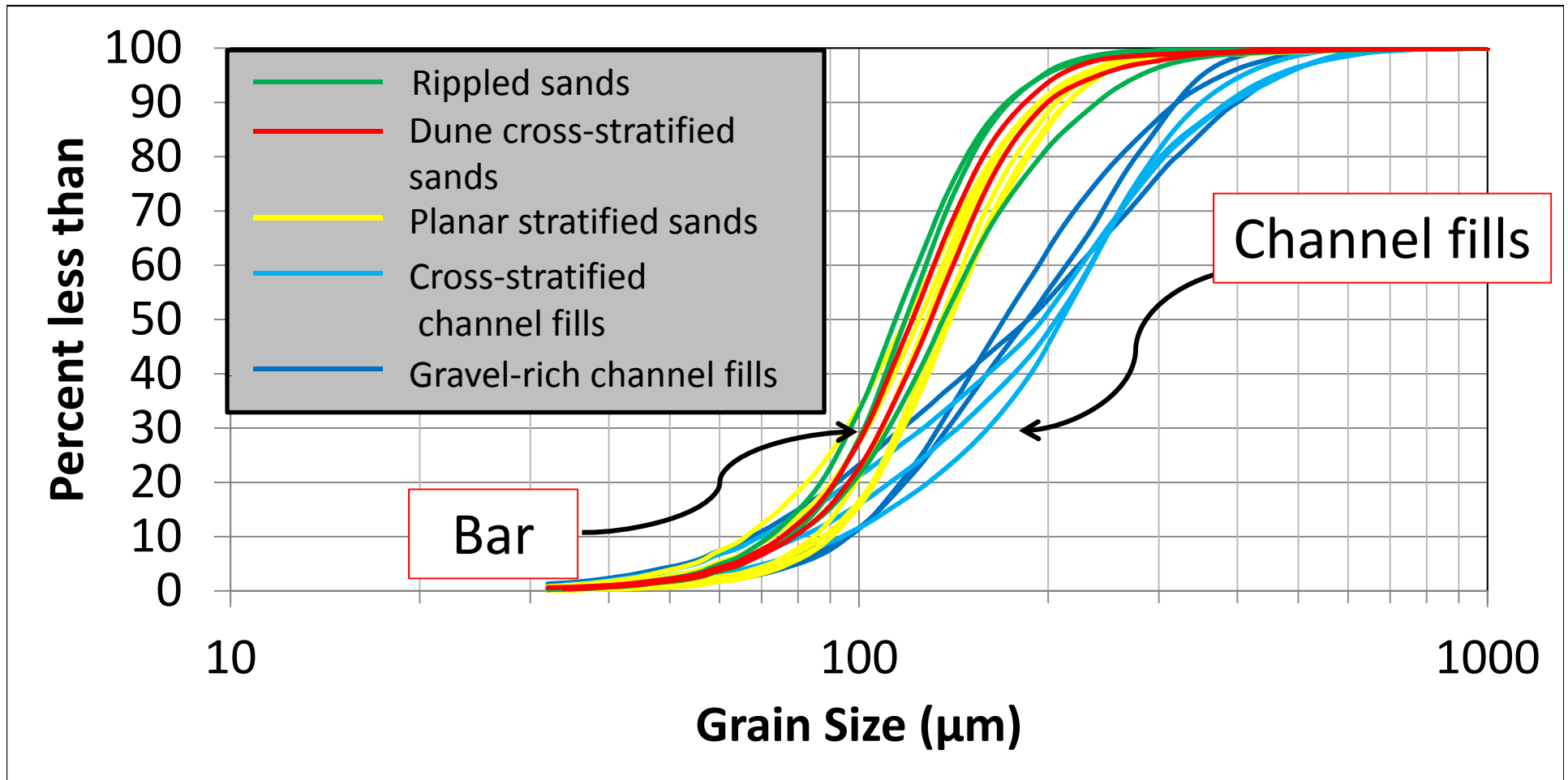
# Re-attachment bar on Colorado river

Similar facies associations in a separation zone bar





# Bars versus channel fills: Grain-sizes of sand



**Coarser** particles in **channel filling** deposits.

Interpreted as bedload or incipiently suspended load  
(travelling near the bed)

# Internal consistency in interpreted transport of sediment

1. Estimate settling velocity  $w_s$  of median clast size of gravels (*Dietrich, 1982*)

2. Estimate shear velocity  $u_*$  of transporting current (*Wiberg & Smith, 1987*)

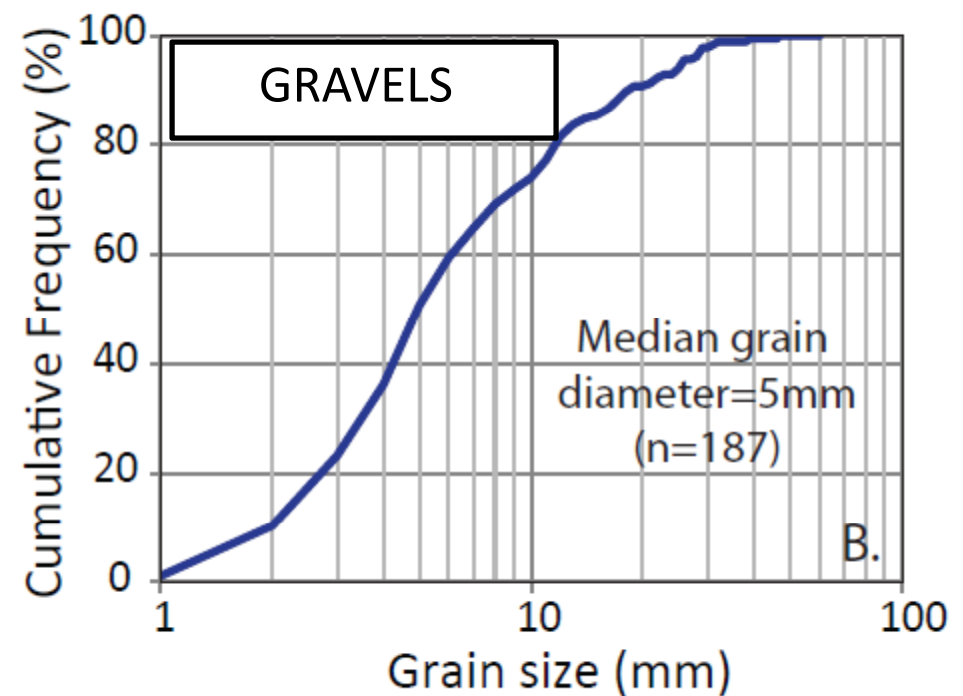
3. Determine fully suspended particle sizes

$$\frac{w_s}{u_*} \leq 0.33 \quad (\text{Smith, 1977})$$

Estimated particle sizes in full suspension  $\leq 231\mu\text{m}$

Actual  $D_{95}$  in bars = **250 $\mu\text{m}$**

Gravel clast clusters



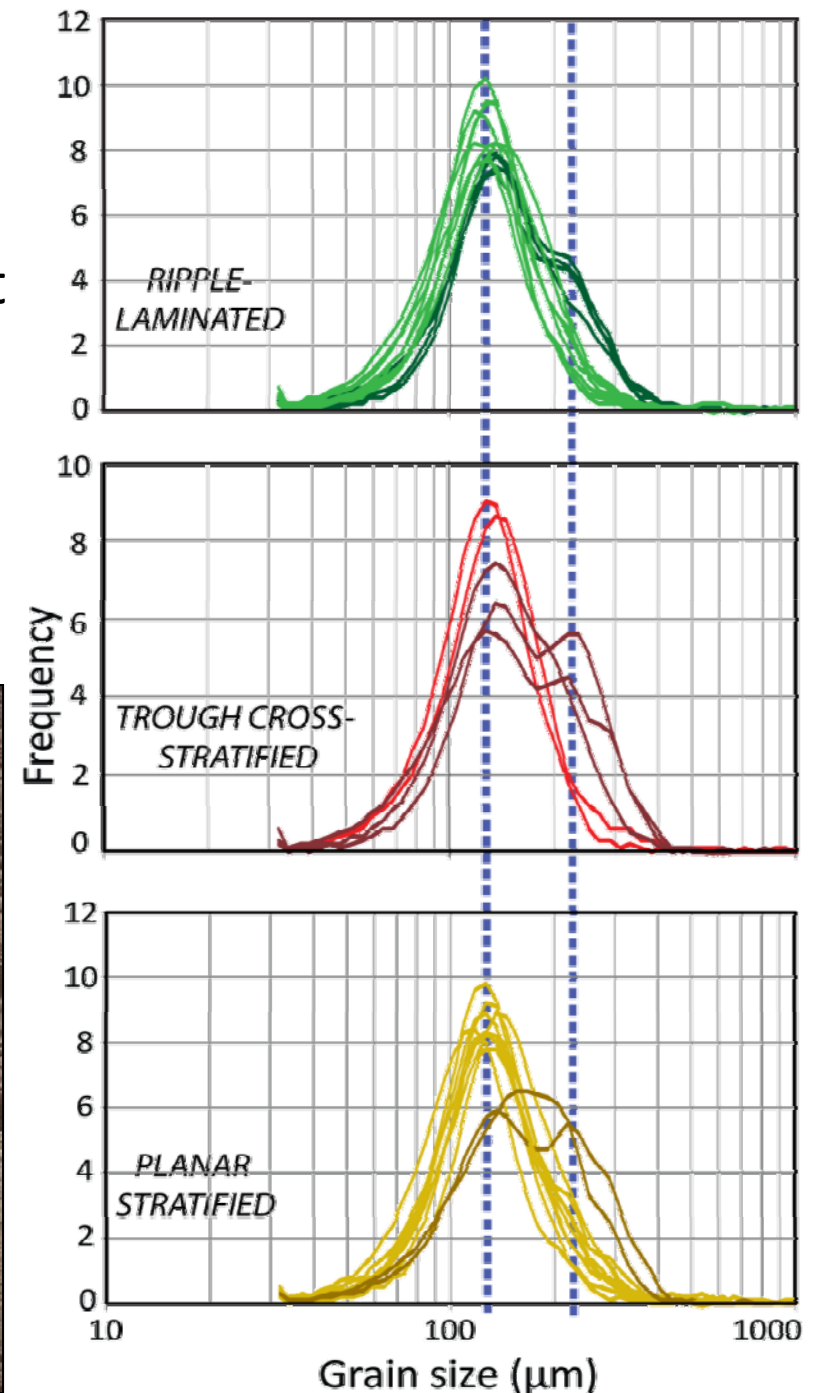


# Bimodality in bar deposits

Distributions usually unimodal (peak at  $120\mu\text{m}$ )

## Occasionally bimodal

Second sediment source from **closer to the thalweg** (second peak at  $210\mu\text{m}$ )



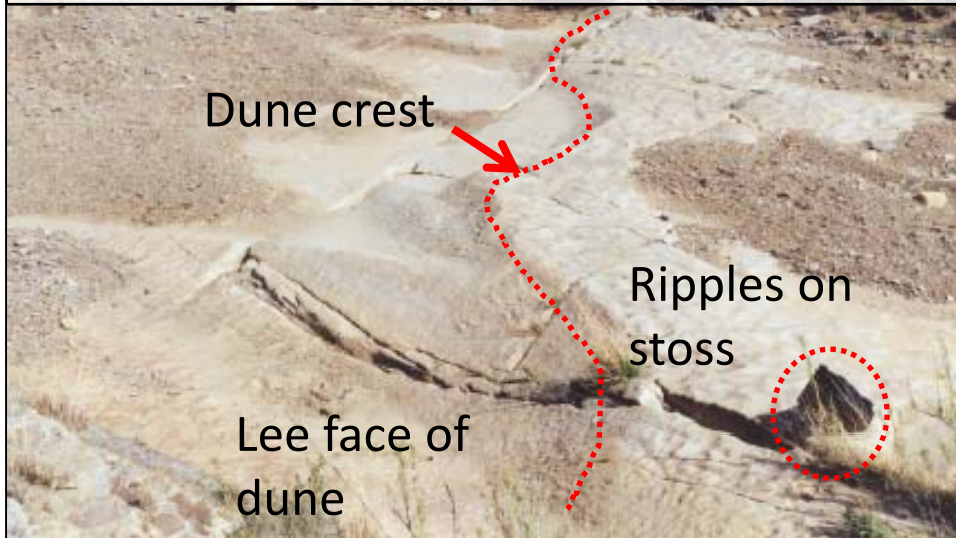


# Basin floor channel fills

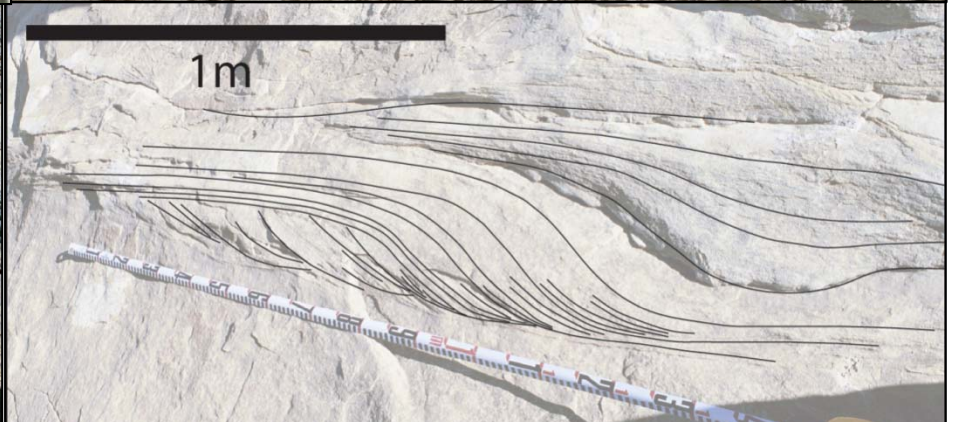
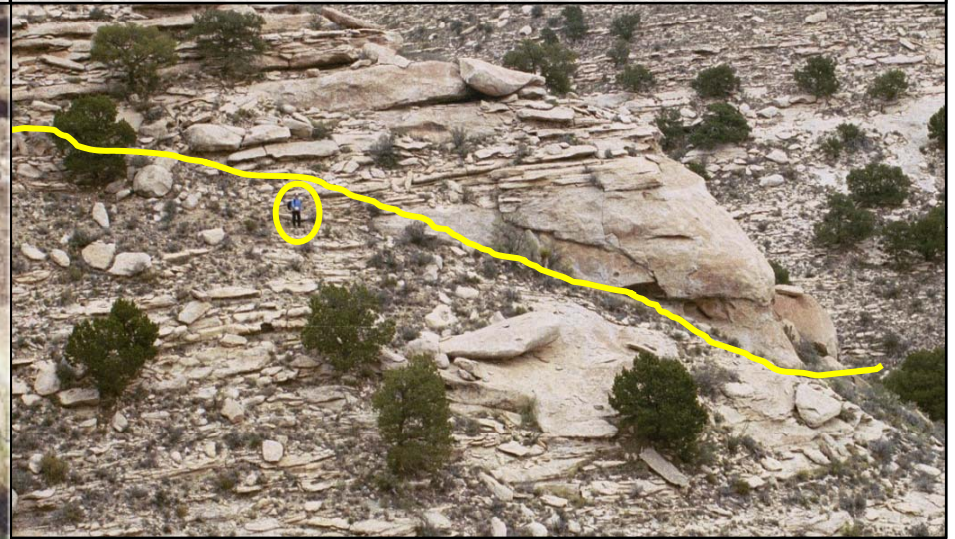
Sub- to super-critically **climbing 3-D dunes** in channel fills.

Requires sediment sourced from **suspended load**.

*Exposed 3-D dune, basin floor channel fill*



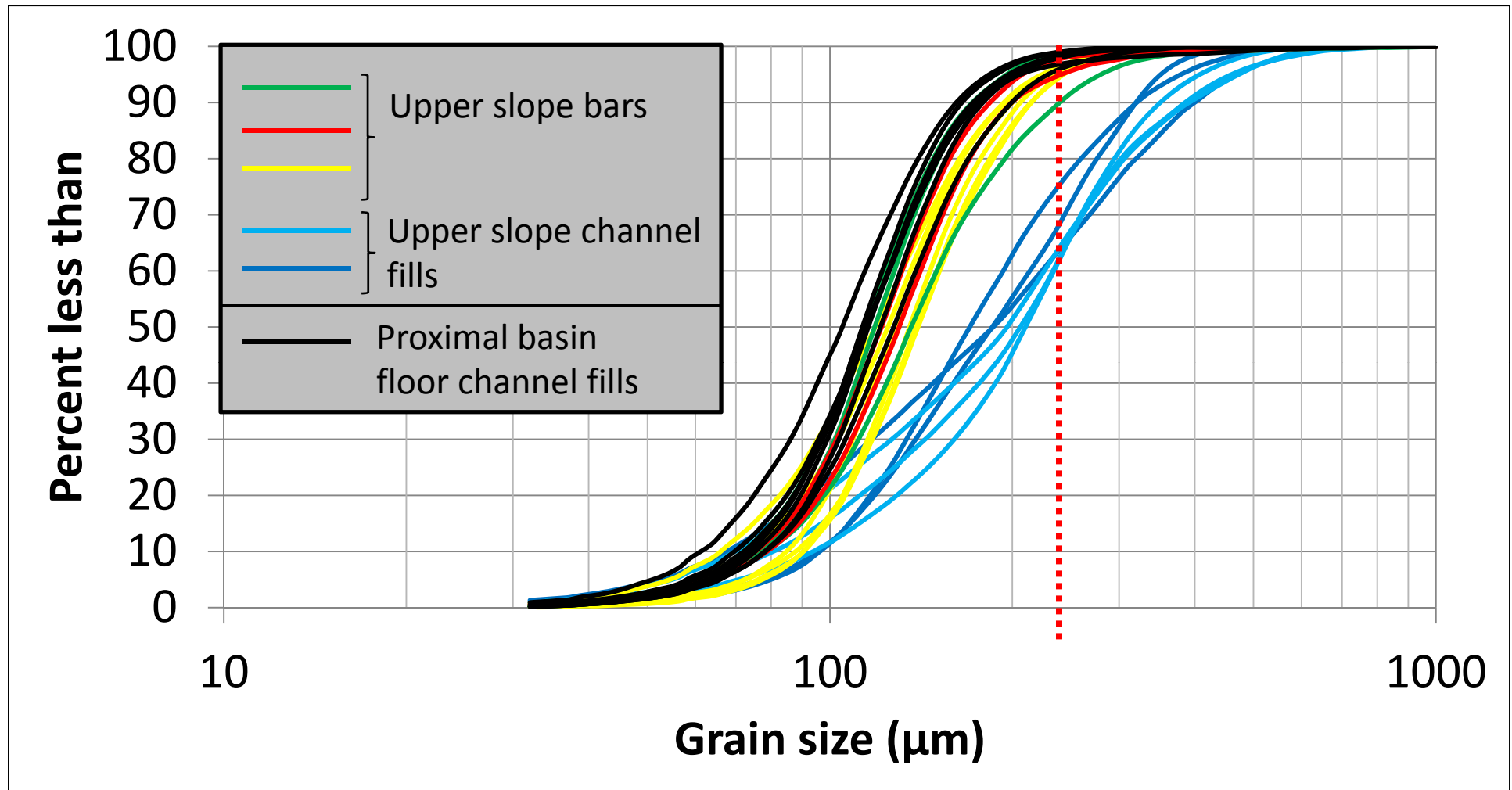
*Channel plug: Proximal basin floor*



*Climbing dunes, basin floor channel fills*



# Linking bars to channel fills on the basin floor

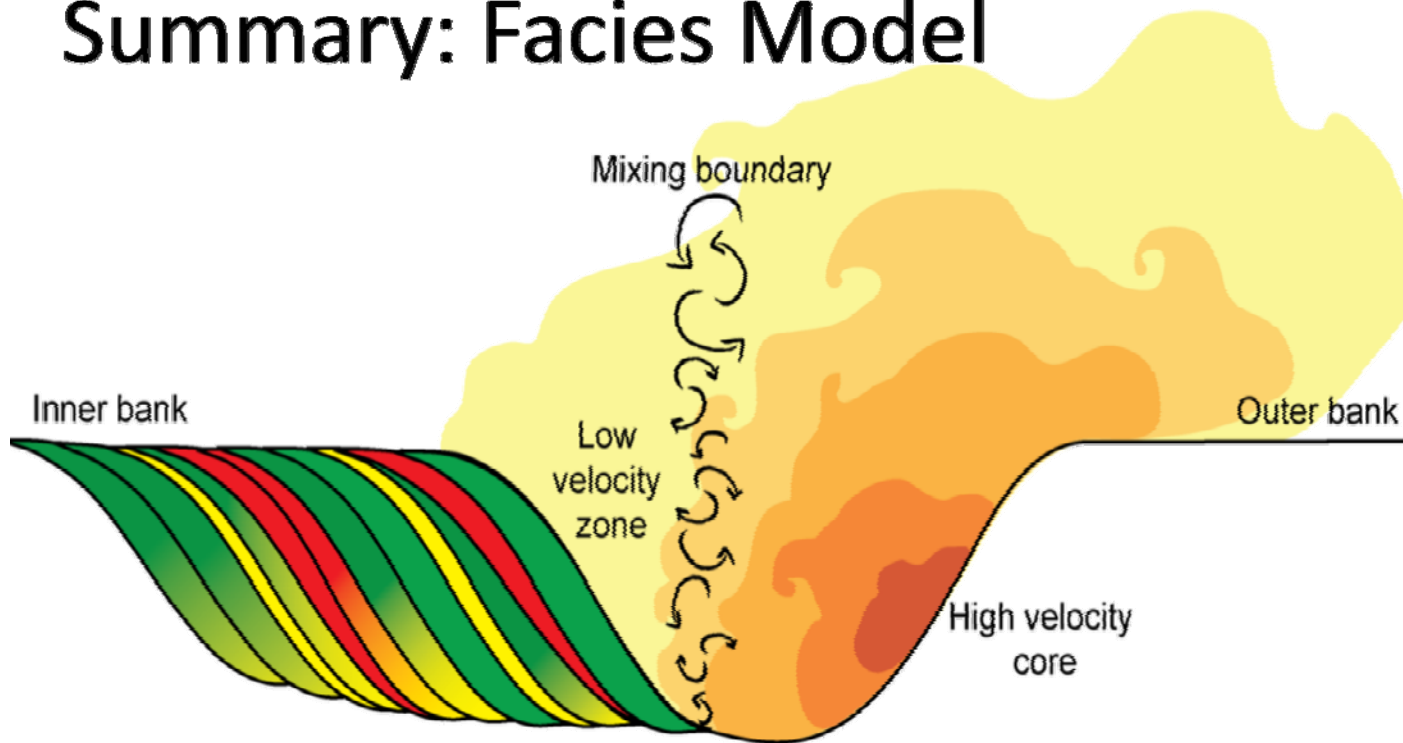


$D_{50}$  (basin floor)=110-150 $\mu\text{m}$ ,  $D_{50}$  (upper slope bar)=130-160  $\mu\text{m}$

Particle sizes <220 $\mu\text{m}$  bypassed to basin floor

Particle sizes >220 $\mu\text{m}$  selectively removed via deposition on slope





# Summary: Facies Model



Constructed  
chiefly from  
**fully suspended**  
sediment

**Reworked** to  
varying degrees  
on the bed.

**Proximity to the high-velocity core** dictates degree of  
reworking and bar composition.

-  Ripple laminations
-  Trough cross-stratification
-  Planar stratification
-  Reactivation surfaces

Small amounts of fully suspended  
sediment are **trapped in low velocity**  
**zones** on the slope, the rest is  
**bypassed to the basin floor.**



# Acknowledgements



AAPG Grants- in-Aid

SEPM Weimer Research Grant

West Texas Geological Society John  
Emery Adams Scholarship