Authigenic Illite in the Sandstone Reservoirs of Taiyuan Formation, Northeast Ordos Basin*

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Abstract

The Upper Palaeozoic sandstones are important hydrocarbon reservoirs in the Northeast Ordos Basin. More extensive illitization is observed in the Upper Carboniferous Taiyuan Formation than those in other reservoir units. In this research, the nature and porosity effects of illite in the Taiyuan Formation of Northeast Ordos Basin have been investigated by means of thin-section observations, scanning electron microscope (SEM), and X-ray diffraction techniques. We have shown different types of illite in the sandstones based on their textural features. Detrital illite, the least abundant component, frequently occurs as tangentially arranged, grain-lining, ragged plates. A detrital origin can be distinguished by its poorly-crystalline morphology with wrinkled surface and blurred boundary, suggesting it was suffering from sedimentary transport. Diagenetic illite, recognized by their delicate, fibrous or lathy morphology, occurs commonly as pore-filling phases formed by feldspar alteration and radially-oriented, grain-rimming cements. It is also observed replacing kaolinite, as laths are extending from pseudohexagonal plates.

It is only in the sandstone of Taiyuan Formation that kaolinite concentration does not increase at the expense of feldspar. This intense illitization, and relative scarcity of kaolinite, may be due to the prolonged exposure to saline depositional pore-waters during Miaogou transgression in the early Taiyuan period. The growth of authigenic K-feldspar, occurring as intergrowths with grain-lining illite, implies that a high concentration of potassium ions were present in the original ground waters. K-feldspar dissolution, inhibited or slowed down by early penetration of K-rich fluids, may provide other potential sources of potassium for illite formation at the expense of kaolinite during burial diagenesis. This process has a greater temperature dependency and burial histories indicate that the Taiyuan Formation entered the illitization window (130° C and 140° C) at 108-135 Ma.

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There is a positive correlation (R^2 = 0.55) between the abundance of authigenic illite and variations in thin-section porosity for Taiyuan sandstone reservoirs. This is interpreted to indicate that the majority of the total sandstone volume is secondary porosity, produced primarily by illitization at the expense of kaolinite and K-feldspar.

References

Liu, D., P. Peng, C. Pan, and L,K. Wang, 2008, Influence of different experiment conditions in the procedure of (⁴⁰) Ar- (³⁹) Ar dating of authigenic illite: Bulletin of Mineralogy Petrology and Geochemistry, v. 27/2, p. 169-174.

Yang, Y., W. Li, and L. Ma, 2005, Tectonic and stratigraphic controls of hydrocarbon systems in the Ordos Basin: a multicycle cratonic basin in Central China: AAPG Bulletin, v. 89/2, p. 255-269.

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Outline

Background information:

Basin location, Study area, Stratigraphy, and Sedimentary facies

What does the petrography of the investigated sandstones tell us?

- Investigated by thin section observations, SEM analysis, XRD measurements
- Rock composition, types of porosity, as well as diagenetic clay minerals associated with feldspar alteration

Authigenic illite in the sandstone reservoirs of Taiyuan Formation

Abundance, Occurance, Morphology

Origin of authigenic illite in the Taiyuan Formation

- Implications for pore fluid chemistry
- Source of potassium for illite precipitation

Basin Location



One of the largest, the oldest, and still an important hydrocarbon Province in China

Contains sediments with a total thickness of 4000–6000m

Covers an area of 320,000km²

Map of Basins in China

	Stratigraphic System			Group/ Formation	Thickness (m)
	Neogene Neogene			, ,	
	[0]	Paleogene			
	Mesozic Cenozoic	Cretaceous	K ₁	Zhidan	100-1200
		Jurassic	J ₃	Fenfanghe	100-1200
			J ₂	Anding	150
				Zhiluo	200-600
			J ₁	Yanan	250-300
		Triassic	T3	Fuxian	0-156
				Yanchang	200-1400
			T ₂	Zhifang	1100
			T ₁	Liujiagou	260-280
	ic	Permian	P ₃	Shiqianfeng	250-280
П			P ₂	Shangshihezhi	140-200
П				Xiashihezhi	100-200
П			P ₁	Shanxi	37-125
П		Carboniferous	C ₂	Taiyuan	22-276
				Benxi	0-40
	Paleozo	Ordovician	O ₃	Beiguoshan	270-800
				Jiangjiawan	230
			O ₂	Pingliang	130-215
			Oı	Majiagou	100-900
		Cambrian	€3	Changshan	90
				Gushan	270
			€₂	Zhangxia	170
				Xuzhuang	120
			€1	Mantou	70
				Houjiashan	100
	Pr	oterozoic			
	Archean				

Contain more than 95% of the total oil resources of the basin

system

Upper
Paleozoic
(C-P)
Sandstone

Gas

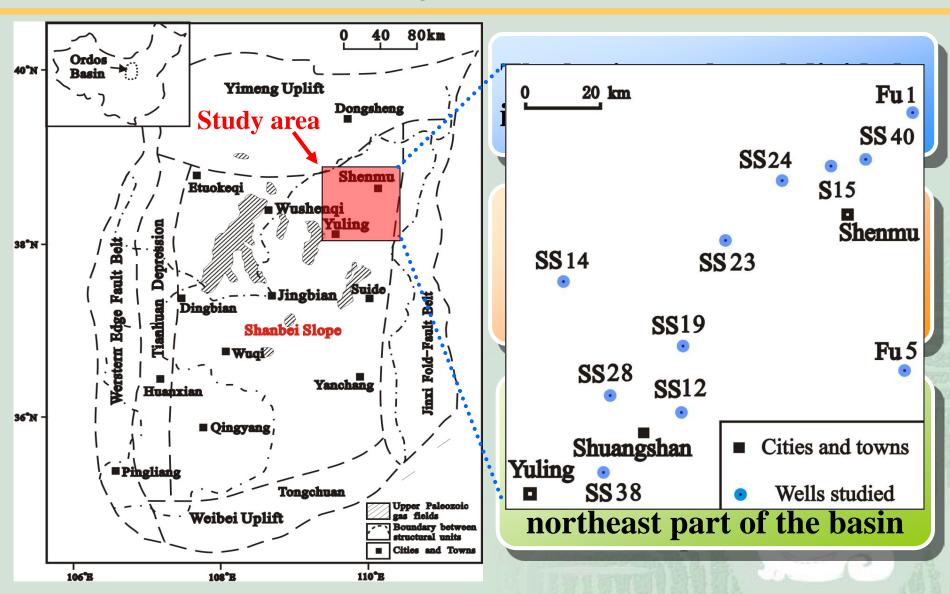
Lower Paleozoic (O) Carbonate

The main gas exploration objective For quite a long period

Hydrocarbon systems in the Ordos Basin

Successions
involved in our
project "Role
of diagenesis
on reservoir
quality in Upper
Paleozoic
sandstones, Ordos
basin"

Study Area



Index map of the Ordos basin structural units. Upper Paleozoic gas fields shaded.

Geologic Properties

The gas system is characterized by widely distributed coal measure source rocks and tight sandstone reservoirs

Source Rocks

—Coals and associated dark mudstones occurring in the Carboniferous–lower Permian coal measures

Cap Rocks

—Thick mudstone in the upper Permian strata

Reservoirs

—Mainly Pennsylvanianearly Permian tight sandstones the focus of this study

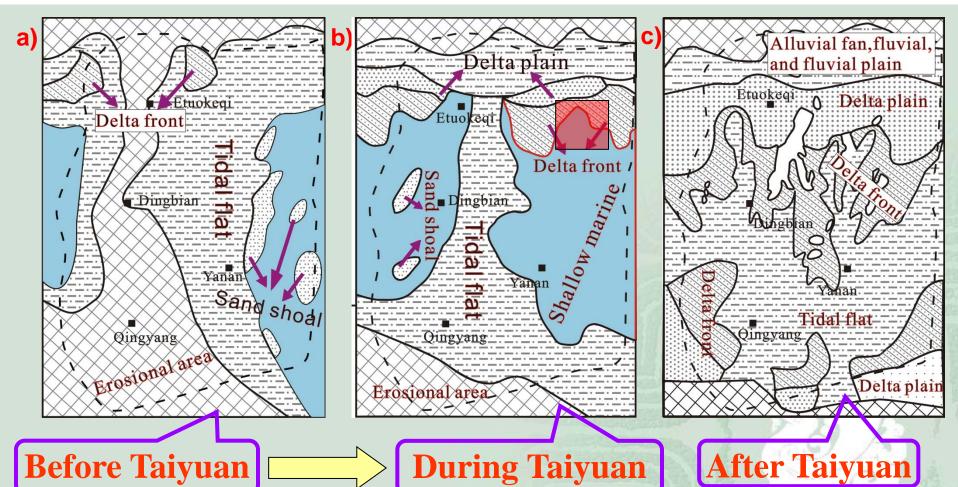
System Series Within the study area, Taiyuan Formation is: The principal hydrocarbon reservoir unit. Mainly near-shore marine deposits being composed of Permian interbedded sandstones and mudstones with some interbedded carbonate levels, Also a coal measure with a 1-3 m thick coal seam Coal delta. and swamp Cap rock nian Marine shore plain Reservoir and swamp rock Penns Bauxitic Benxi Restricted mudstone platform

Stratigraphic columns, depositional environments, and sourcereservoir-seal associations in the Ordos basin for the Upper Paleozoic. after Yang et al., 2005

Sedimentary facies

Maps of sedimentary facies in the Ordos basin during the deposition of

a) Benxi Formation; b) Taiyuan Formation and c) Shanxi& Xiashihezi Formations

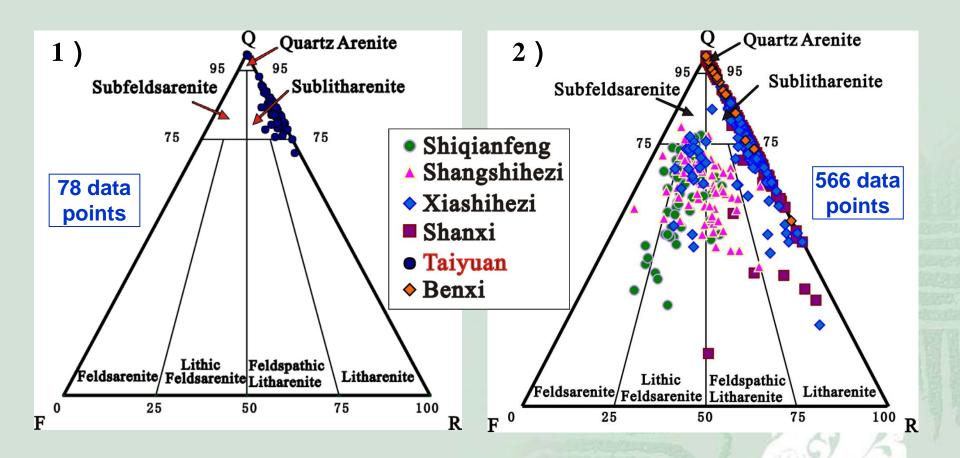


Miaogou transgression

Modified after Yang et al., 2005

Petrologic data

——Framework grain composition of sandstones



Classification of sandstone samples collected from :

1) Taiyuan Formation 2) Other Formations of Upper paleozoic strata

Questions

What created this nearly absence of feldspar in the Taiyuan Formation samples?

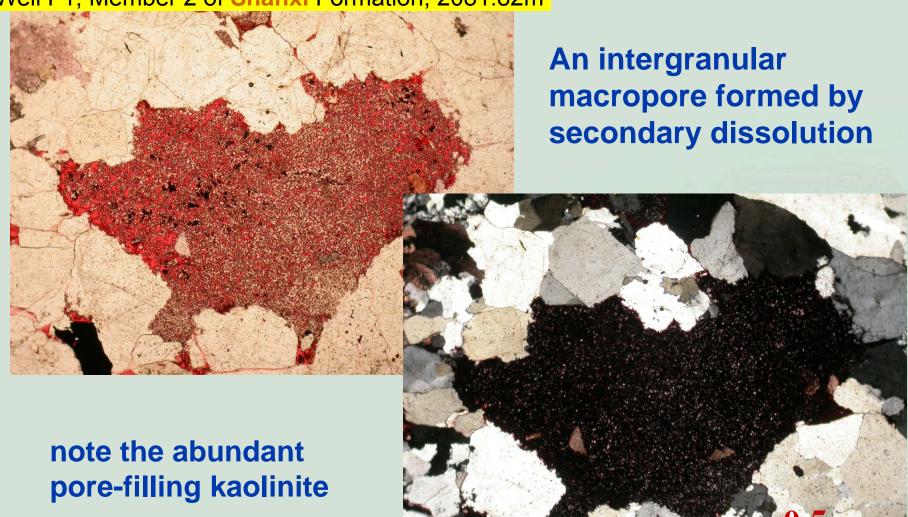
— Depositional or diagenetic?

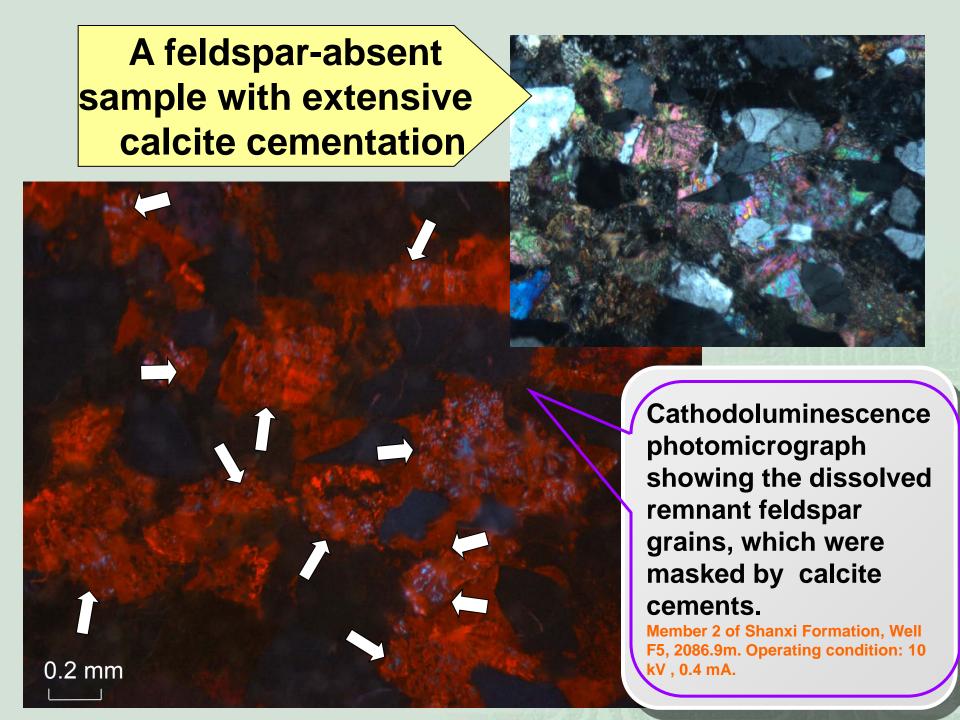
How does it compared to the samples from other Formations?



A feldspar crystal has been partially removed by dissolution, leaving molded pores and remnant portions of feldspar

Well F1, Member 2 of Shanxi Formation, 2081.82m

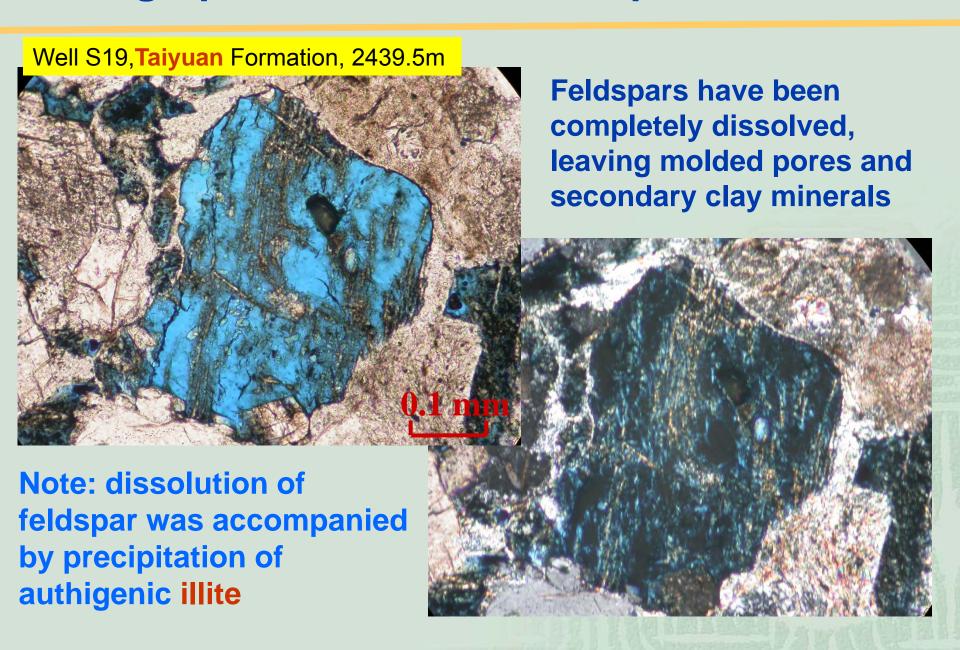




Well SS12, Xiashihezi Formation, 2328.52m

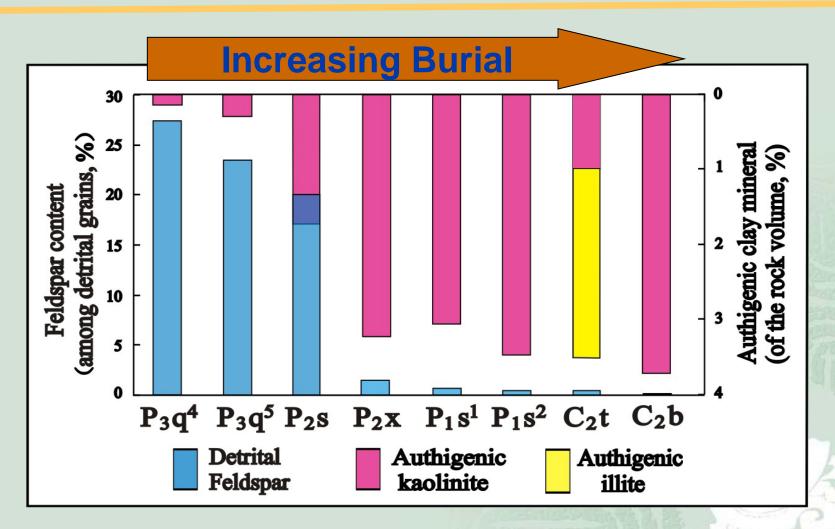
Feldspar grain has been almost completely dissolved, leaving molded pores and secondary clay mineral

Note: dissolution of feldspar was accompanied by precipitation of authigenic kaolinite



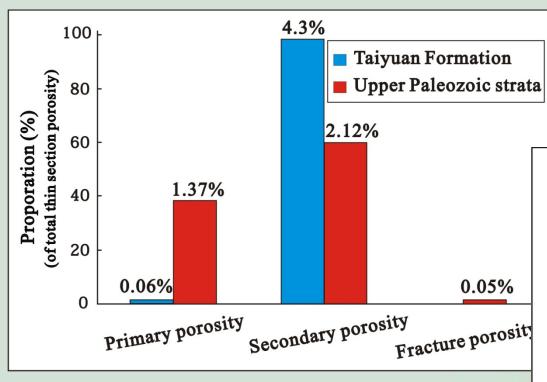
Petrologic data

—— Covariance of feldspar content with secondary clays



Taiyuan Formation dose not match general feldspar-kaolinite content covariance trend represented in other Upper Paleozoic sandstones

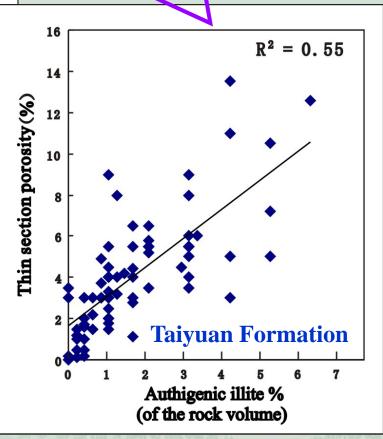
Amounts of main pore types

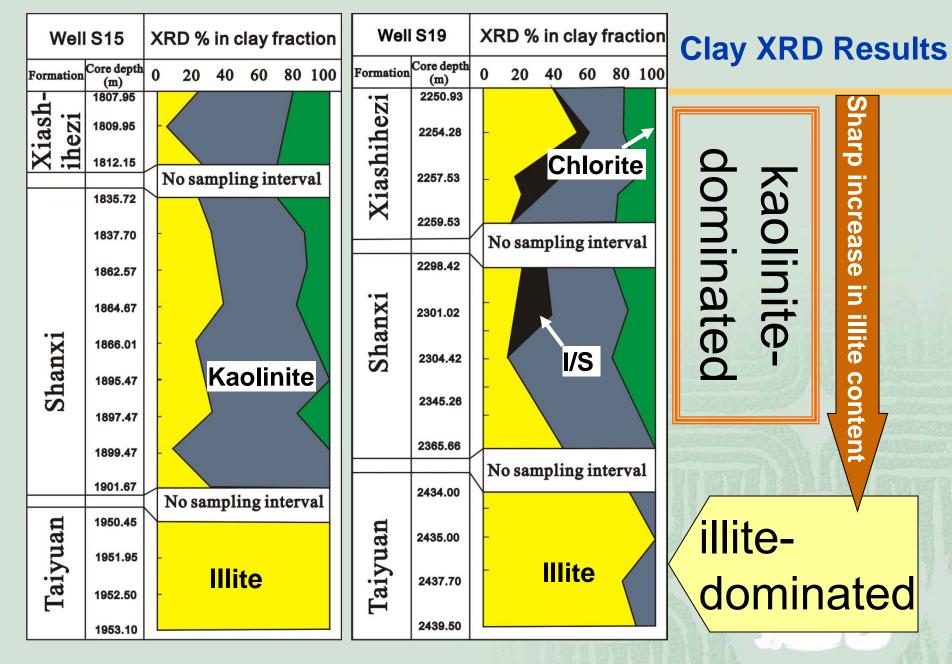


Increased porosity towards the greater degree of illitization

➤ The pores of upper Paleozoic sandstone reservoirs are mainly secondary in origin

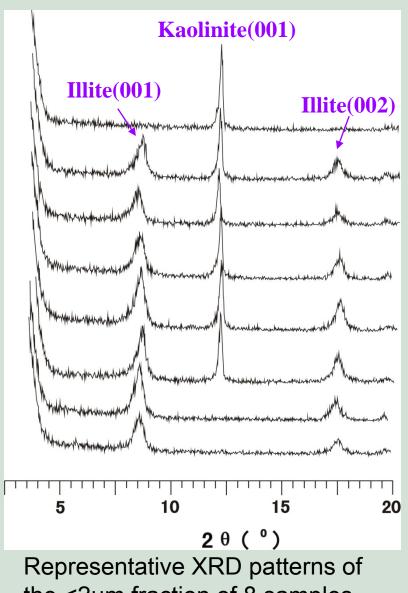
Secondary porosity (predominantly provided by dissolution of feldspar grains) is nearly the only type of porosity observed in Taiyuan Formation



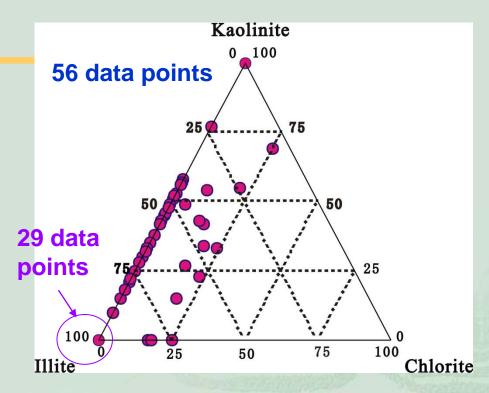


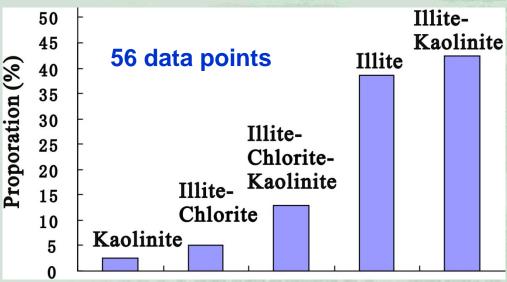
Vertical distribution of clay minerals(<2 µm) in two of the wells studied

Clay XRD Results



the <2µm fraction of 8 samples

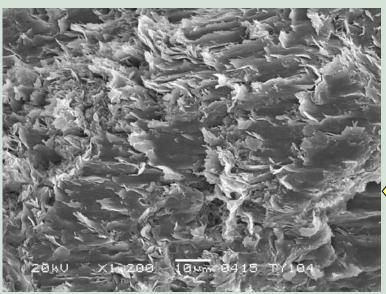




Types of illite

Detrital illite

- Abundance: volumetrically minor
- Morphology: poorly-crystallized with wrinkled surface and blurred boundary
- Cocurrence: as tangentially arranged, grain-lining, ragged plates



Authigenic illite

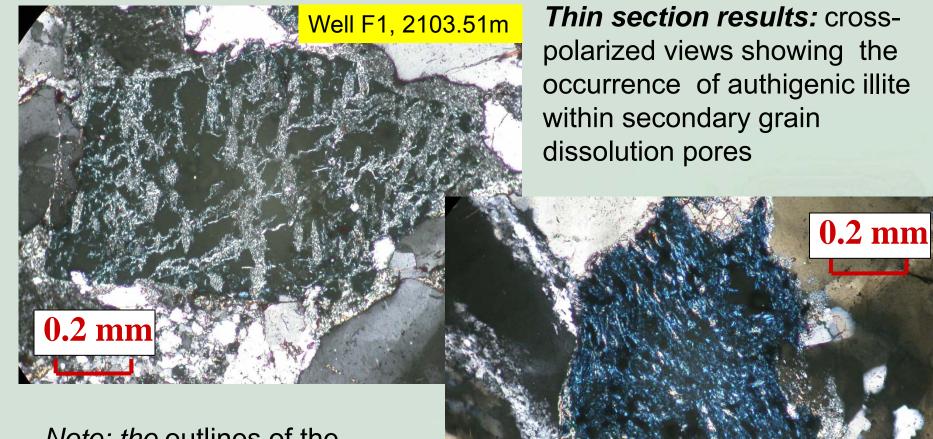
- Abundance: widespread
- Morphology: delicate, fibrous or lathy crystals
- **Occurrence:**
 - Pore-filling/bridging
 - Feldspar-replacing
 - Kaolinite-transitional
 - Grain-coating

Detrital illite: skins of tangentially arranged illite visible on grain surface.

Well:Shuang 28, Depth: 2787.28m

Pore-filling illite

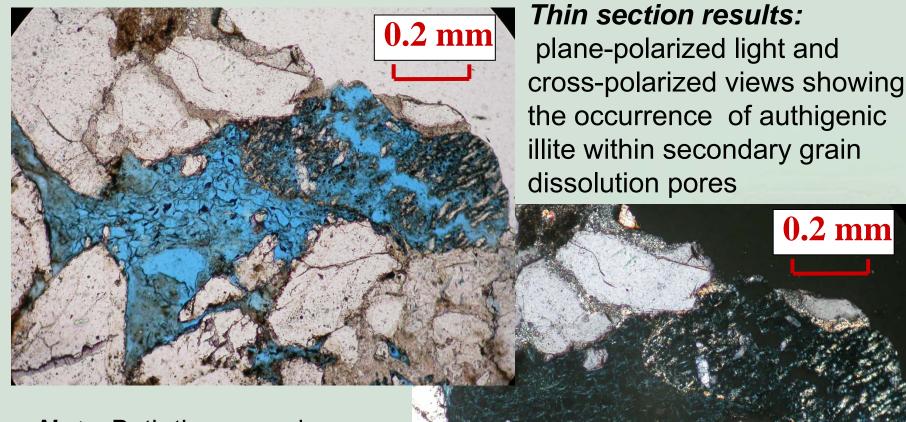
Well S15, 1953.1m



Note: the outlines of the original detrital feldspar shape are preserved

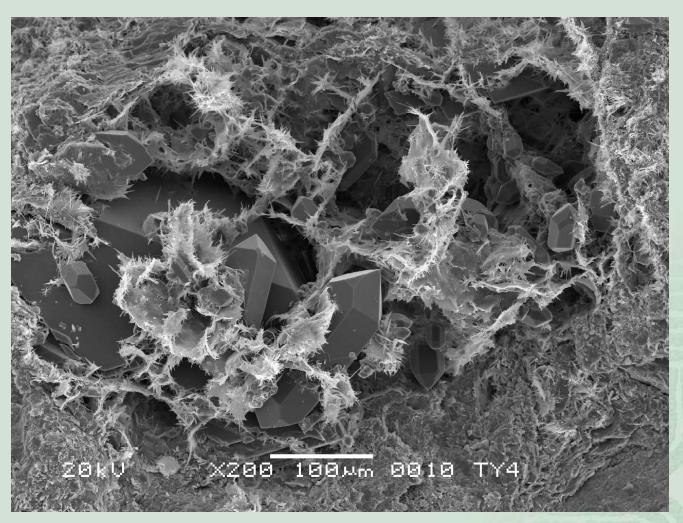
Pore-filling illite

Well SS 23, 2434.68m



Note: Both the secondary dissolution porosity and secondary pore-filling illite cements are uncompacted

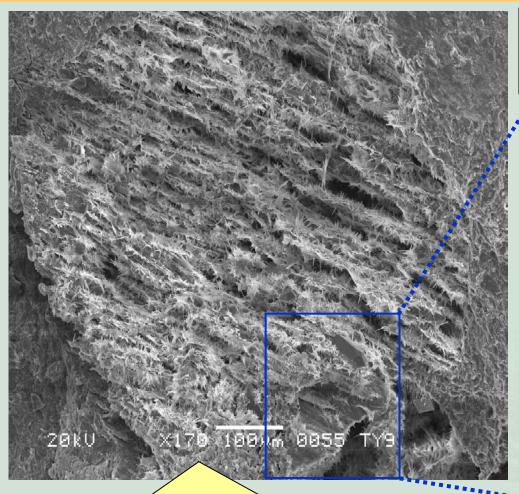
Pore-bridging illite



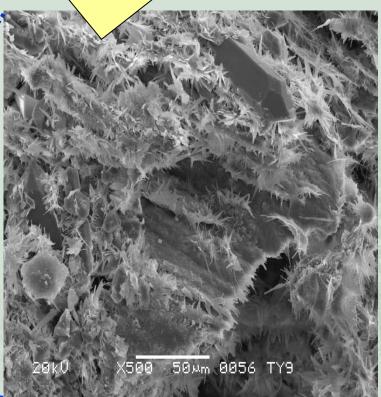
- ➤ Growing along the general orthogonal orientation, possibly indicating feldspar cracks
- ➤ Apparent simultaneous growth of illite and quartz
- ➤ Note the conversion of the dissolution pore scale from macro to micro by this illite

Well SS 14, 2792.2m

Feldspar-replacing



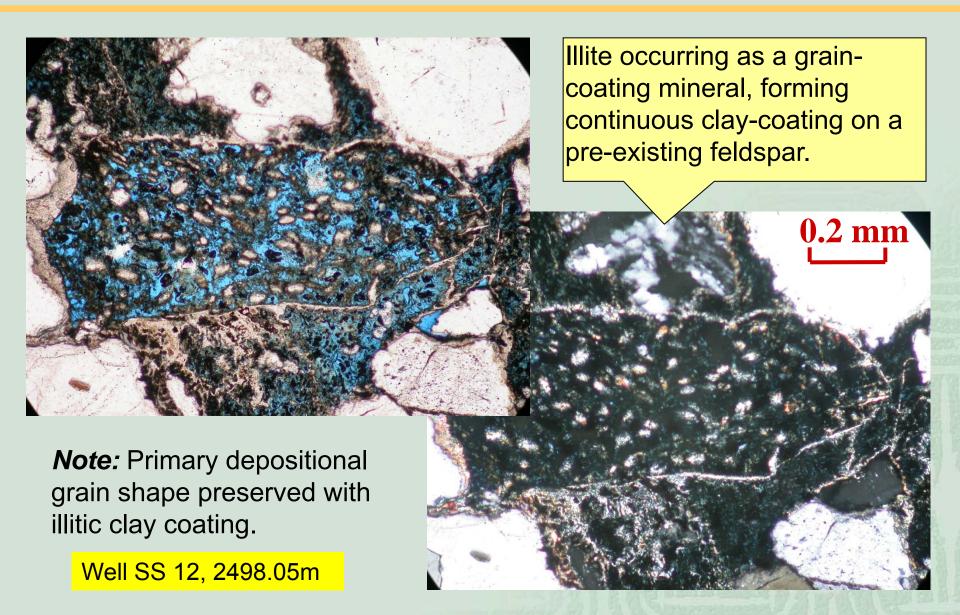
Remnant portions of feldspar and authigenic quartz



SEM photomicrographs showing the replacement of feldspar by illite occurs preferentially along cleavage traces

Well F1, 2120.07m

Grain-coating



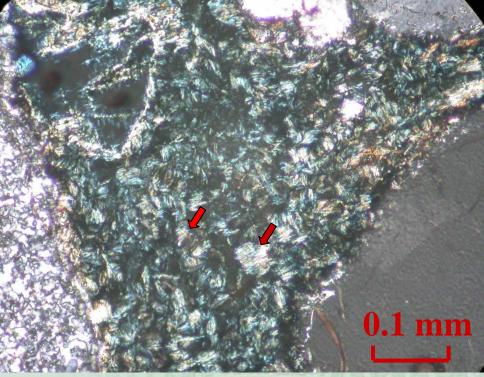
Kaolinite-to-illite transition

Well S19, 2434m

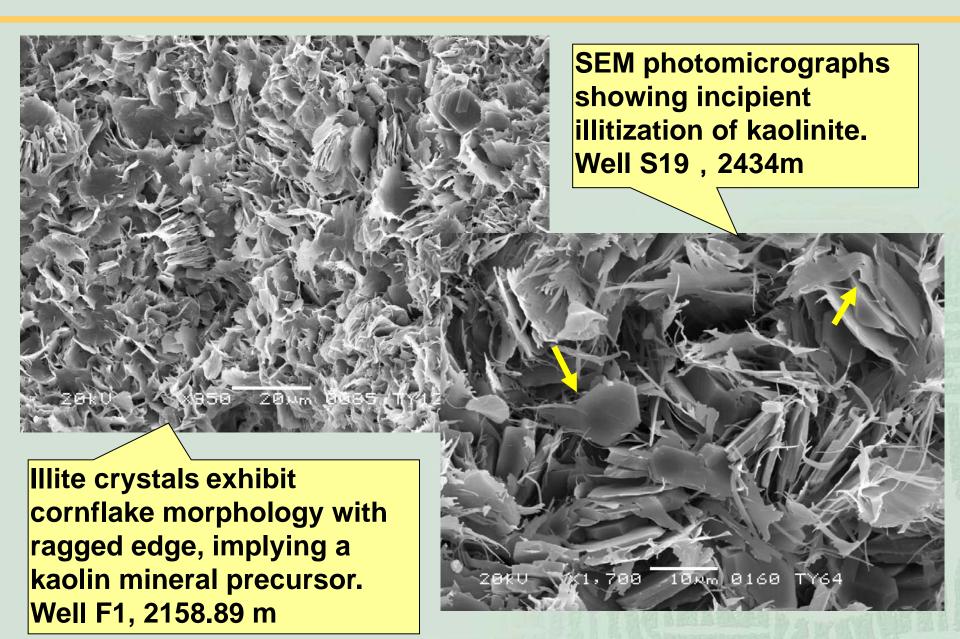
Some of platy illitic clays (red arrows) are arranged into books, probably pseudomorphing earlier formed kaolinite.

booklet: typical crystallite morphology of kaolinite

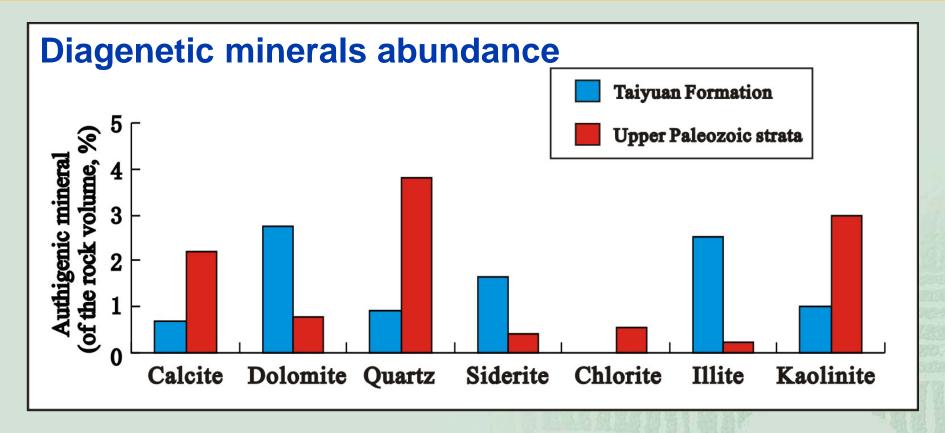
High-order interference colors



Kaolinite-to-illite transition



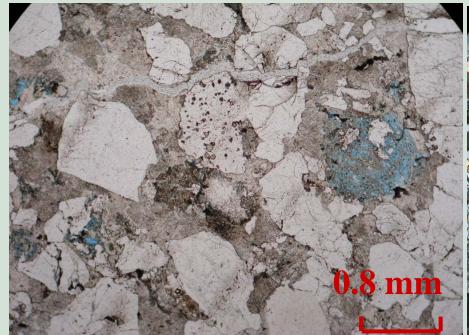
Origin of authigenic illite

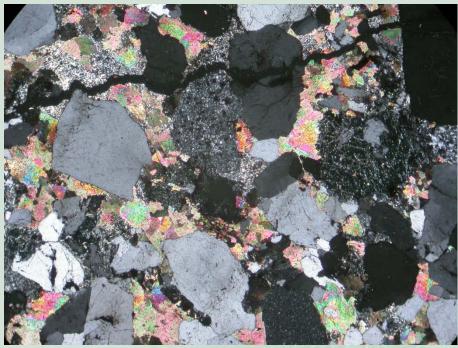


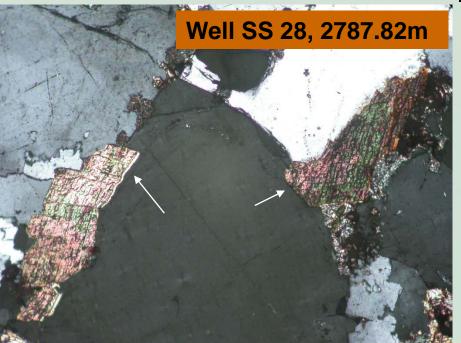
The striking diagenetic feature for Taiyuan Formation is:

- relative scarcity of acid-prefer minerals
- relatively rich in alkali-prefer minerals

-Taiyuan Formation has Pore-water with relatively high alkalinity level







In many sandstones, dolomite cementation pre-dates burial diagenetic features

—— dolomite was an early cement

—— large amounts of alkalinity fluids moving through the Taiyuan Formation rocks at shallow depth

Origin of authigenic illite

Two stages of feldspar dissolution are inferred:

Early-stage, Open system: Feldspar-Kaolinite

$$\sim 2KAlSi_3O_8 + 2H^+ + H_2O = Al_2Si_2O_5(OH)_4 + 4SiO_2 + 2K^+$$

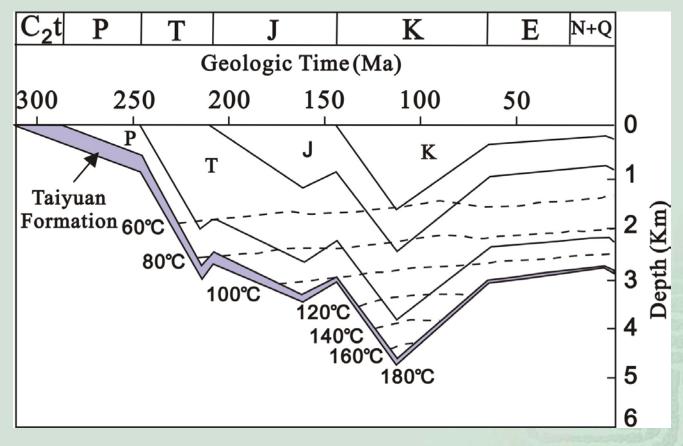
- —— Possibly in response to organic acid-rich fluids derived from the coal measures
 - ——Slowed down due to large amounts of alkalinity fluids moving through the Taiyuan Formation rocks.
- Late-stage, Closed system:

main mechanism responsible for illite formation

$$\bowtie$$
 $|KAlSi_3O_8 + Al_2Si_2O_5(OH)_4 = 2SiO_2 + KAl_3Si_3O_{10}(OH)_2 + H_2O_1$

- —— Possibly in response to increased temperature
- Much of the K-feldspar reaction occurred during burial diagenesis, which is suggested by the presence of uncompacted secondary dissolution porosity and secondary pore-filling authigenic illites

Origin of authigenic illite



Burial and thermal history model for northeast Ordos basin depicting incremental burial history for the Late Carboniferous through to the present (after Liu, 2008)

The reaction of K-feldspar + kaolinite to form illite

- has a greater temperature dependency
- require minimum temperature of ~120°C, which corresponds to thermodynamical destabilization of kaolinite + K-feldspar assemblage
- Taiyuan Formation had reached this temperature threshold



Thank you!

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