

Cenozoic Evolution of Carbonate Shelf and Ramp Habitats: Insights from Paleoceanography*

Pamela Hallock¹ and Luis Pomar²

Search and Discovery Article #50663 (2012)**

Posted July 31, 2012

*Adapted from oral presentation at AAPG Annual Convention and Exhibition, Long Beach, California, April 22-25, 2012

**AAPG©2012 Serial rights given by author. For all other rights contact author directly.

¹University of South Florida. College of Marine Science, 140 7th Avenue South. St. Petersburg, FL 33701 (pmuller@usf.edu)

²Universitat de les Illes Balears. Departament de Ciències de la Terra, Ctra. Valldemossa km 7.5; E-07122 Palma de Mallorca. Spain

Abstract

Understanding biological, geochemical, and oceanographic processes influencing modern shelf and ramp carbonate deposition is essential when interpreting carbonate sedimentation in the geologic record. Yet recognizing the limitations of uniformitarianism is equally important. Cenozoic carbonate-producing ecosystems emerged from the remnants of Cretaceous biotas, evolving in the warm alkaline oceans of a Greenhouse world, then modifying in response to emerging Icehouse conditions. The latter included stronger latitudinal and bathymetric temperature gradients, declining carbon dioxide concentrations in the atmosphere, and declining calcium concentrations and alkalinity in the oceans. Paleocene-Eocene photic-dependent carbonates tended to be dominated by calcitic coralline red algae and larger benthic foraminifers (LBF), with aragonitic corals and calcareous green algae restricted temporally and spatially. Conceptual models suggest that episodic changes in ocean circulation and thermocline stratification that accompanied high latitude cooling during the Cenozoic provided impetus for turnover in light-dependent (chlorozoan) biotas. For example, comparison of Eocene through Miocene paleotemperature data for surface to thermocline gradients with the history of LBF assemblages indicates that the latter were most diverse and productive when deeper waters were warmest and gradients were weakest. Higher extinction rates corresponded with times when surface to thermocline gradients increased. In contrast, zooxanthellate corals, while relatively diverse in the Eocene, were restricted as reef builders. As Icehouse conditions emerged, aragonite production by corals and calcareous algae became more widespread, with a setback in the early and middle Miocene when coralline algae again dominated. Moreover, the proliferation of reef-building coralline algal taxa into shallow-water habitats in the late Miocene paralleled the emergence of shallow-water corals and new clades of zooxanthellae, indicating co-evolution of these critical reef taxa. Implications of these observations indicate that enhanced understanding of deeper photic-zone (30-100 m) carbonate systems can enhance interpretations especially of Paleogene shelf and ramp carbonate depositional environments, and likely older carbonate systems, especially those deposited during Greenhouse conditions.

References

- Beavington-Penney, S.J., and A. Racey, 2004, Ecology of extant nummulitids and other larger benthic Foraminifera; applications in paleoenvironmental analysis: *Earth Science Reviews*, v. 67/3-4, p. 219-265.
- Dutton, A., K.C. Lohmann, and R.M. Leckie, 2005, Insights from the Paleogene tropical Pacific: Foraminiferal stable isotope and elemental results from Site 1209, Shatsky Rise. *Paleoceanography* v. 20:PA3004, doi:10.1029/2004PA001098
- Hallock, P., 1981, An analysis of the energetics of algal symbiosis: *GSA Abstracts with Programs*, v. 13/7, p. 466.
- Hallock, P., 1987, Fluctuations in the trophic resource continuum; a factor in global diversity cycles?: *Paleoceanography*, v. 2/5, p.457-471.
- Hallock, P., 1988, Diversification in algal symbiont-bearing foraminifera; a response to oligotrophy?: *Revue de Paleobiologie, Special Issue* v. 2, p. 789-797.
- Hallock, P., and W. Schlager, 1986, Nutrient excess and the demise of coral reefs and carbonate platforms: *Palaaios*, v. 1-4, p. 389-398.
- Hallock, P., I. Premoli Silva, and A. Boersma, 1991, Similarities between planktonic and larger foraminiferal evolutionary trends through Paleogene paleoceanographic changes, *in* I. Premoli Silva, and A. Boersma, (eds.), *The oceans of the Paleogene: Palaeogeography Palaeoclimatology Palaeoecology*, v. 83/1-3, p. 49-64.
- Leichter, J.J., G. Shellenbarger, S.J. Genovese, and S.R. Wing, 1998, Breaking internal waves on a Florida (USA) coral reef: a plankton pump at work?: *Marine Ecology Progress Series*, v. 166, p. 83-97.
- Lear, C.H., H. Elderfield, and P.A. Wilson, 2000, Cenozoic Deep-Sea Temperatures and Global Ice Volumes from Mg/Ca in Benthic Foraminiferal Calcite: *Science*, v. 287, p. 269-272.
- Pomar, L., and P. Hallock, 2008, Carbonate factories: A conundrum in sedimentary geology. *Earth- Science Reviews*, v. 87, p. 134-169
- Stanley, S.M., and L.A. Hardie, 1998, Secular oscillations in the carbonate mineralogy of reef-building and sediment-producing organisms driven by tectonically forced shifts in seawater chemistry: *Palaeogeography Palaeoclimatology Palaeoecology*, v. 144/1-2, p. 3-19.
- Susanto, R.D., and A.L. Gordon, 2005, Velocity and transport of the Makassar Strait throughflow: *Journal of Geophysical Research*, v. 110/C1, 10 p.
- Wolanski, E., P. Colin, J. Naithani, E. Deleersnijder, and Y. Golbuu, Large amplitude, leaky, island-generated internal waves around Palau, Micronesia: *Estuarine, Coastal and Shelf Science*, v. 60, p. 705-716.

Zachos, J.C., and G.R. Dickens, 2000, The assessment of the biogeochemical feedback response to the climatic and chemical perturbations of the LPTM, in B. Schmitz, B. Sundquist, and F.P. Andreasson, (eds.), Early Paleogene warm climates and biosphere dynamics; short papers and extended abstracts: GFF, v. 122/1, p. 188-189.

Website

Internal Waves in the Indonesian Seas. Web accessed 18 July 2012. <http://www.ldeo.columbia.edu/~dwi/iwaves.html>

Cenozoic Evolution of Carbonate Shelf and Ramp Habitats: Insights from Paleooceanography

Pamela Hallock

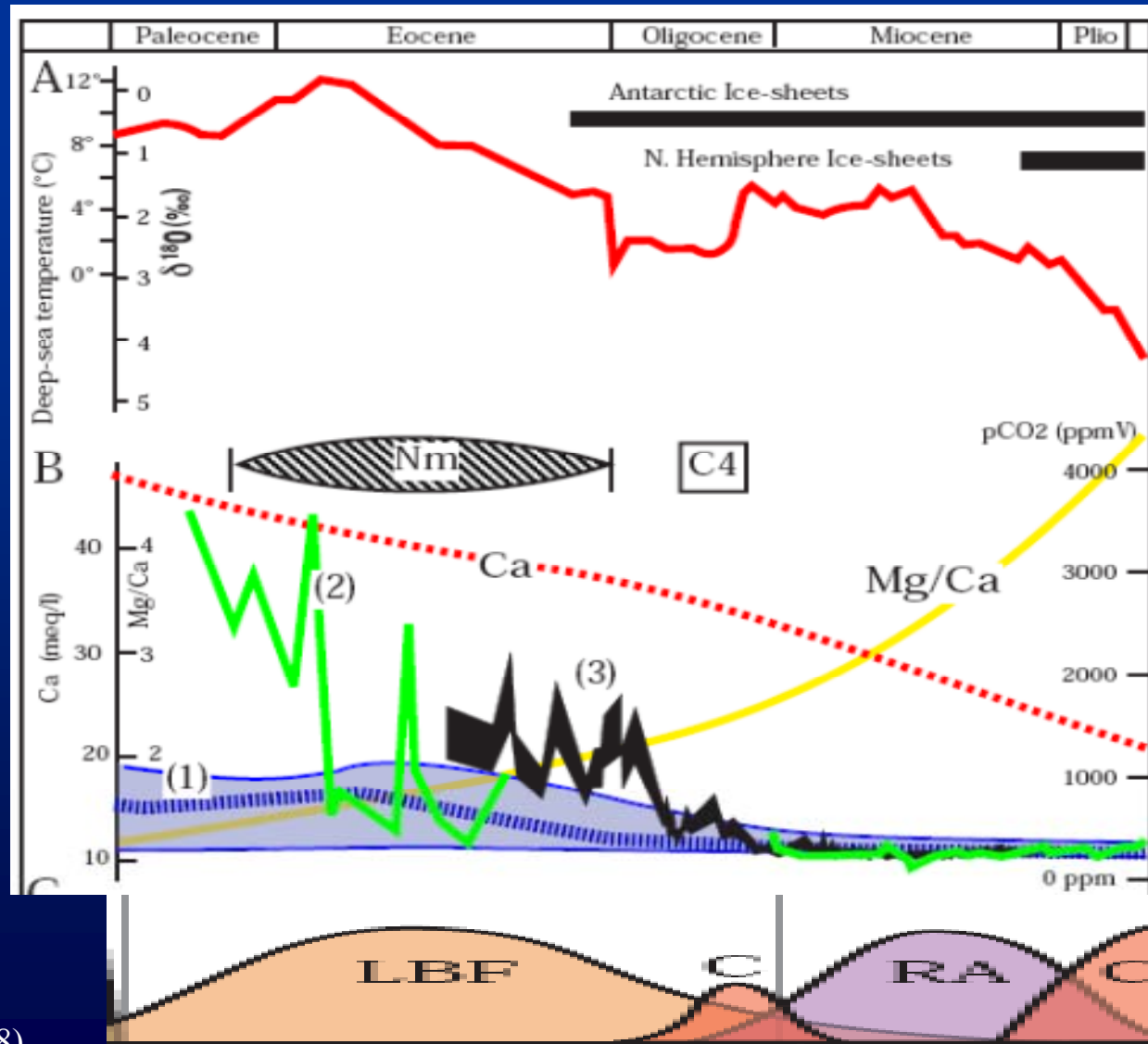
*University of South Florida. College of Marine Science
140 7th Avenue South. St. Petersburg, FL 33701. USA*

Luis Pomar

*Universitat de les Illes Balears. Departament de Ciències de la Terra
Ctra. Valldemossa km 7.5; E-07122 Palma de Mallorca. Spain*

The Paradox of the Cenozoic

Why did larger forams thrive in the Greenhouse World?
Why do coral reefs thrive in the Icehouse World?

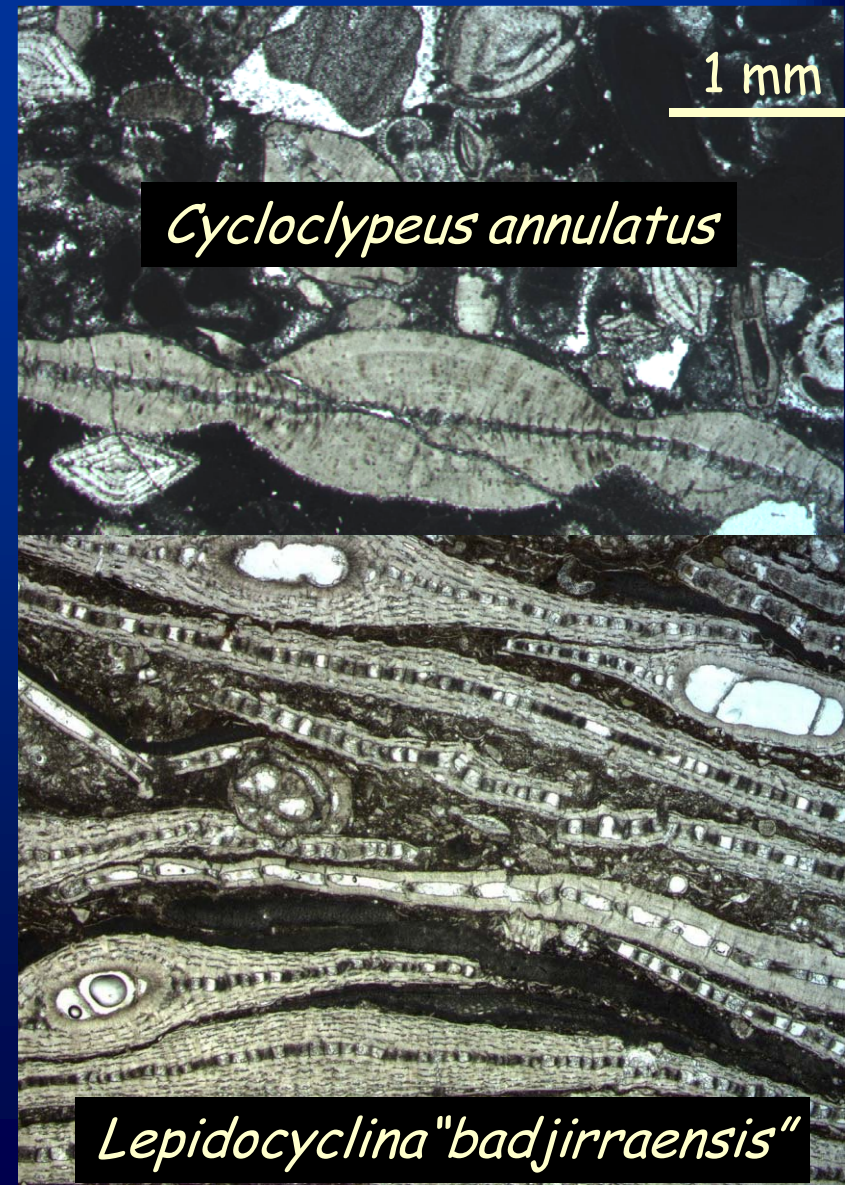


Overview

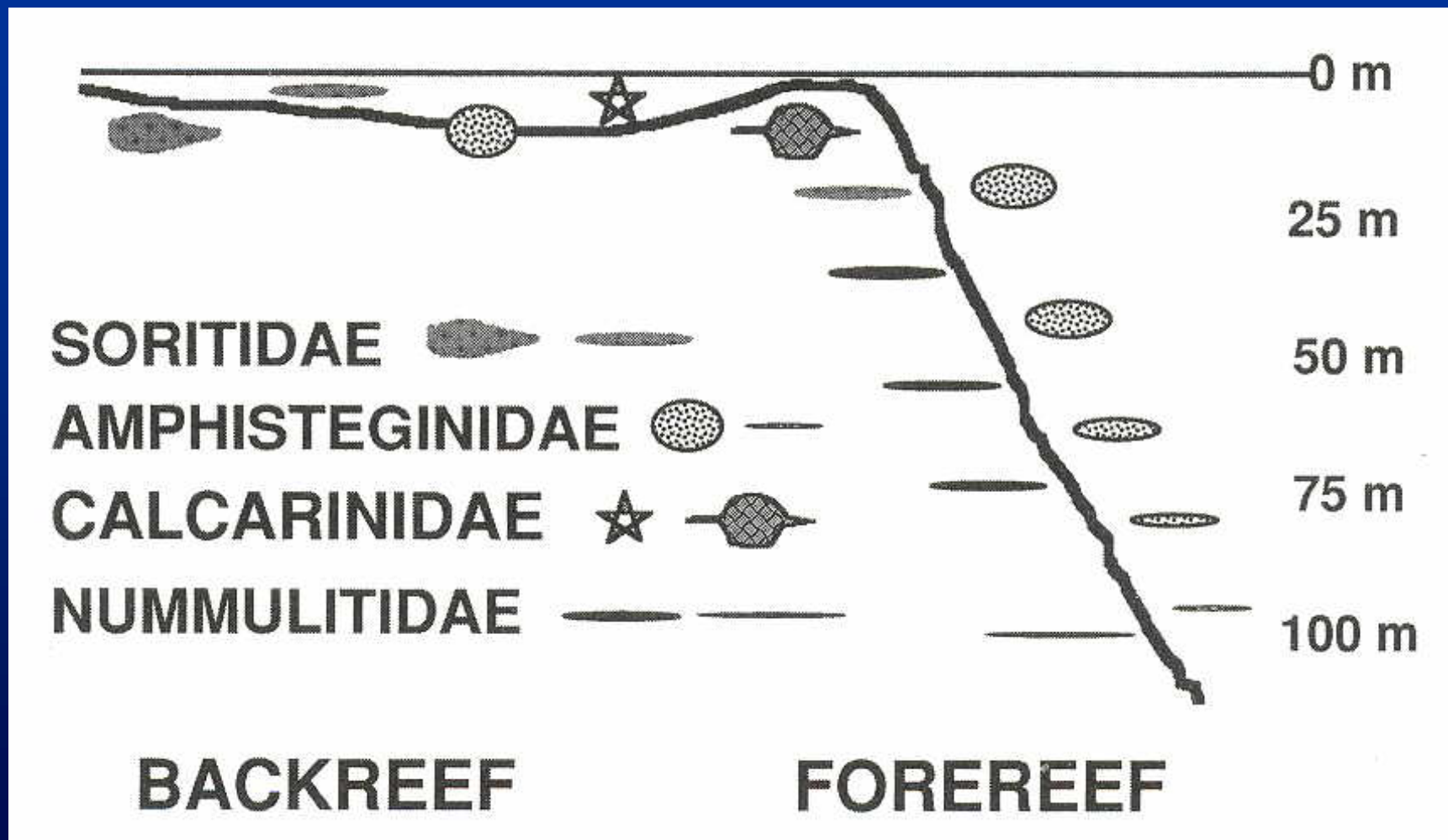
- Paleoenvironmental indicators for carbonate shelves and ramps
 - Larger benthic foraminifers (LBF)
- Paleooceanographic influences on LBF evolution
- Risk-benefit considerations of algal symbiosis
- How do changes in modes and rates of ocean circulation influence LBF?
- How do deep-sea paleotemperatures compare with diversification and extinction rates of Cenozoic LBF?
- New interpretations
- Concluding remarks

Larger Benthic Foraminifera (LBF)

- Important benthic carbonate producers in Cenozoic
 - Especially prevalent in Paleogene
- Morphologies adapted to algal symbiosis
 - Complex interior morphologies
 - Overall shape trends



Distributions of modern LBF: strongly dependent on light and hydrodynamics



Shape as a function of environment

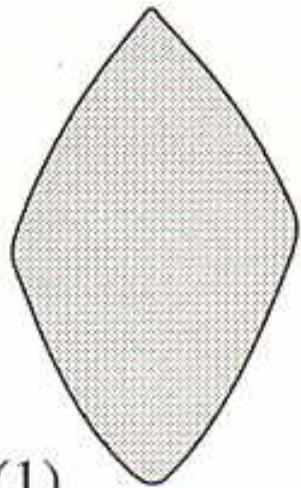
Interspecific differences



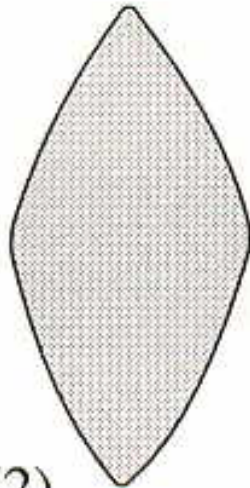
A. lobifera

Amphistegina lessonii

A. papillosa



(1)



(2)



(3)



(4)



(5)

High light
High energy

High light
High energy

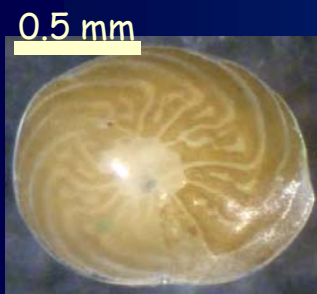
High light
Low energy

Low light
Low energy

Low light
Low energy



Intraspecific variability

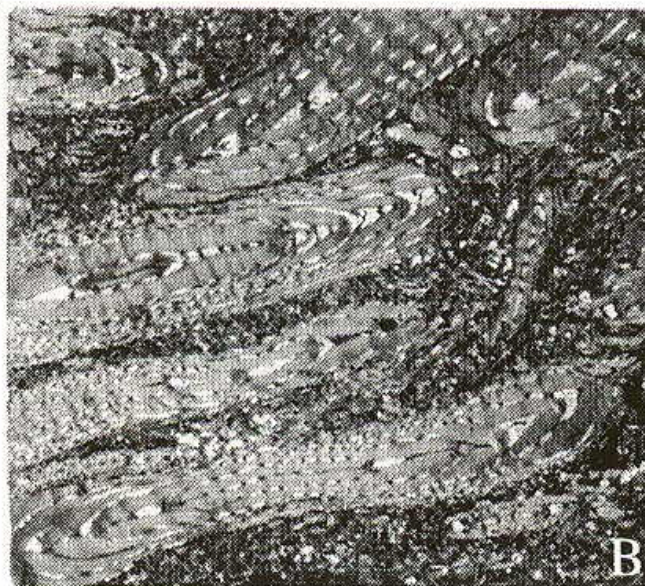


A-forms with
'robust', ovate
tests dominate



Field of view 4 mm

B-forms dominate
in intermediate
parts of depth range



Field of view 14 mm

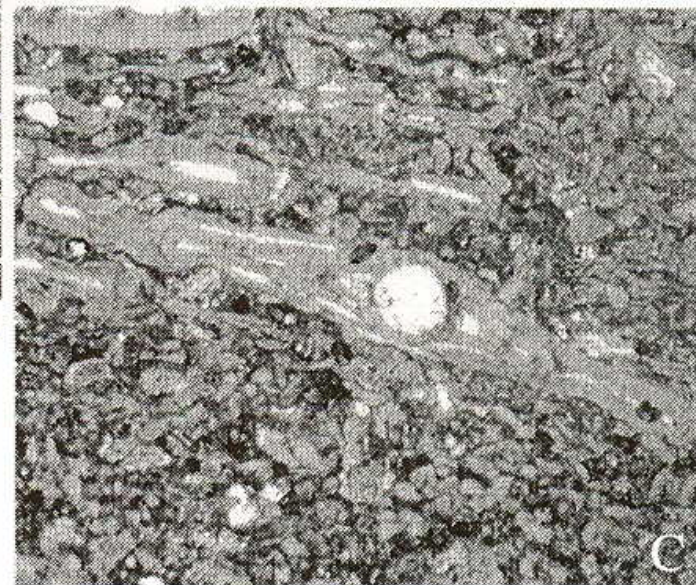
**Cenozoic LBF are widely
used to interpret
paleoenvironments**

Shallow

Depth Range

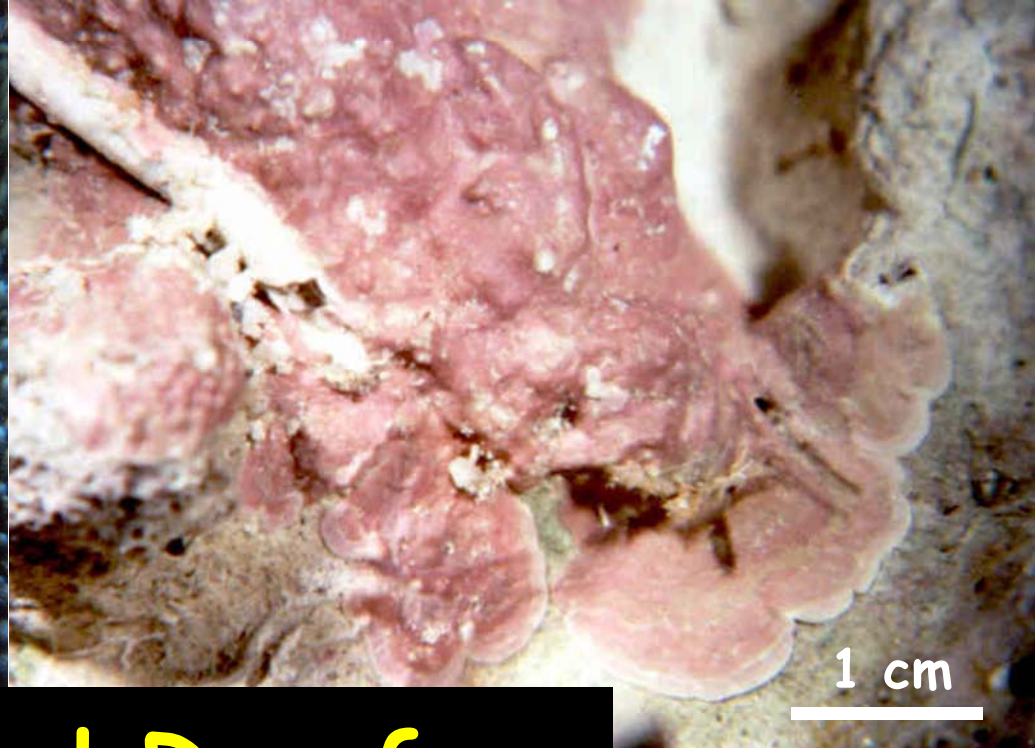
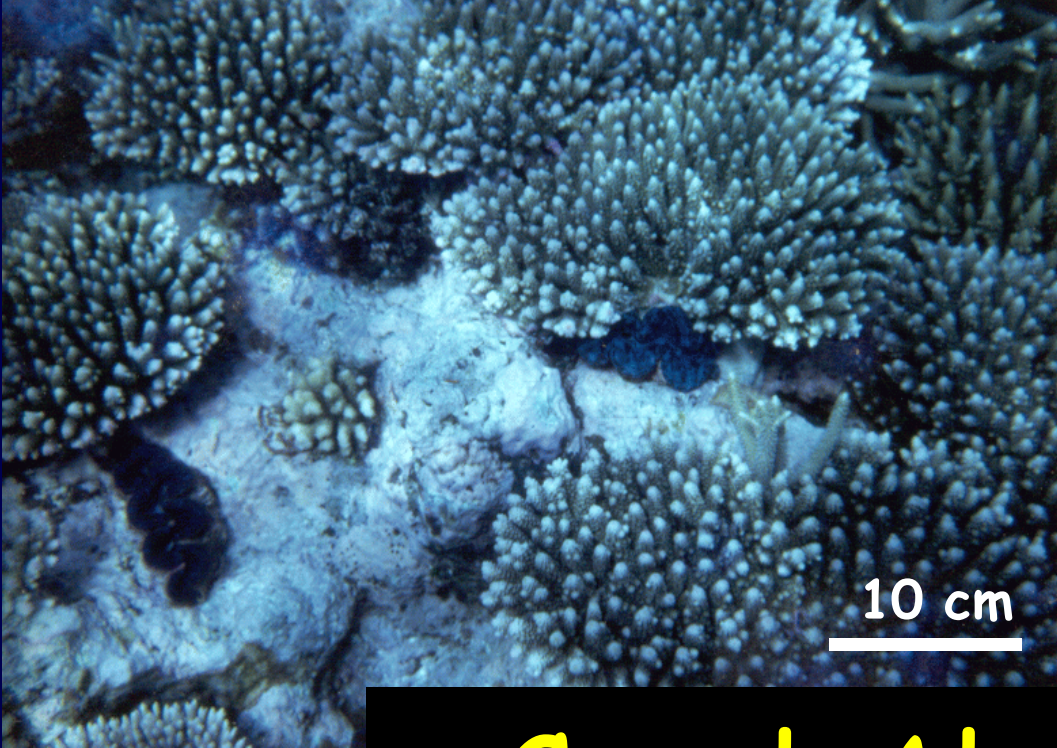
Deep

Elongate A-forms
dominate

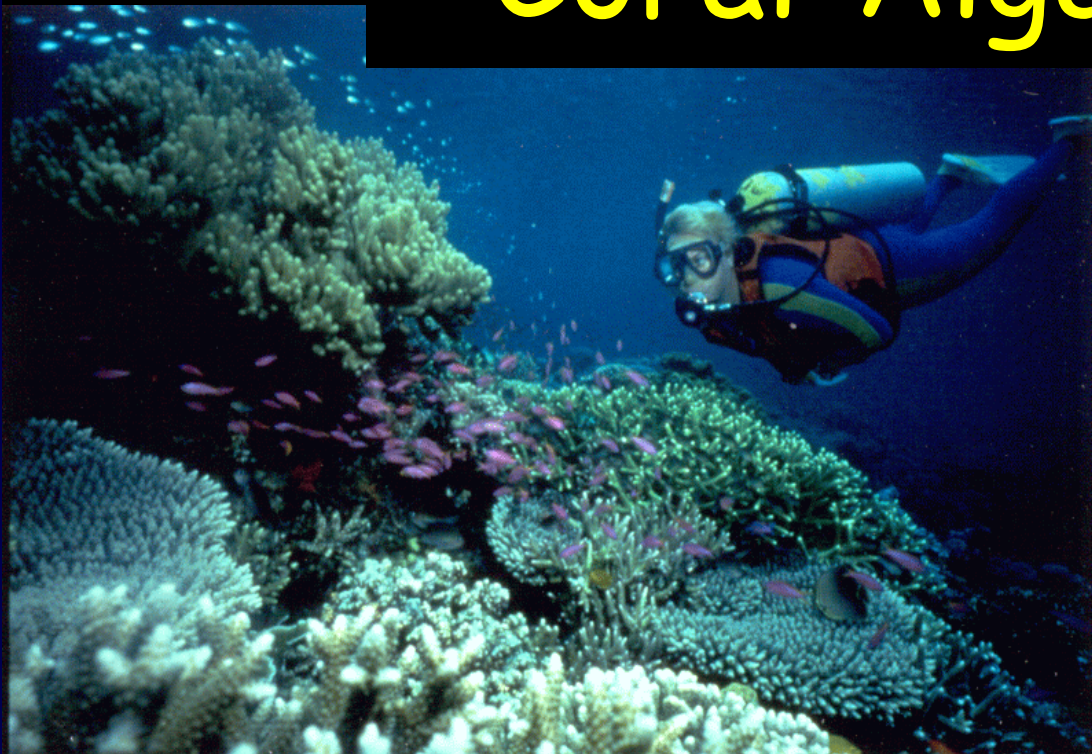


Field of view 3.0 mm

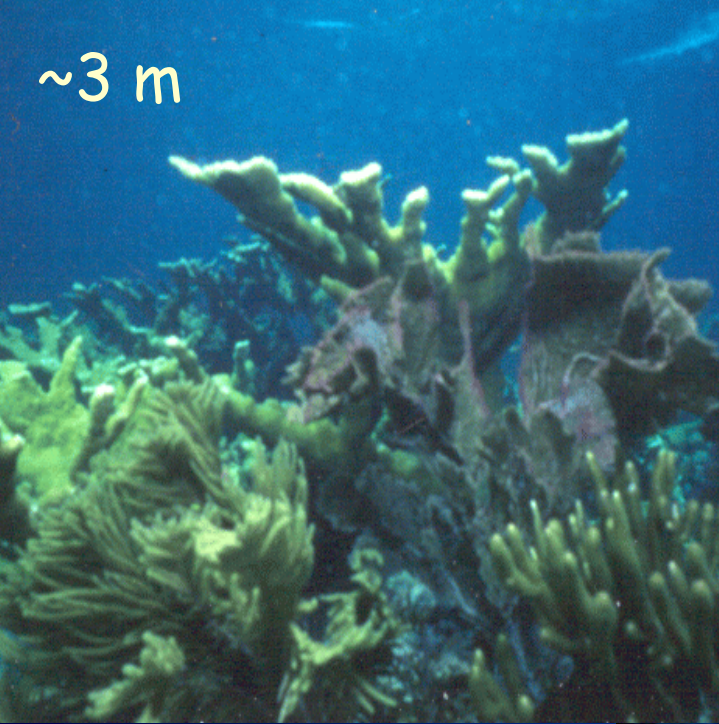
Decreasing light levels and water energy



Coral-Algal Reefs

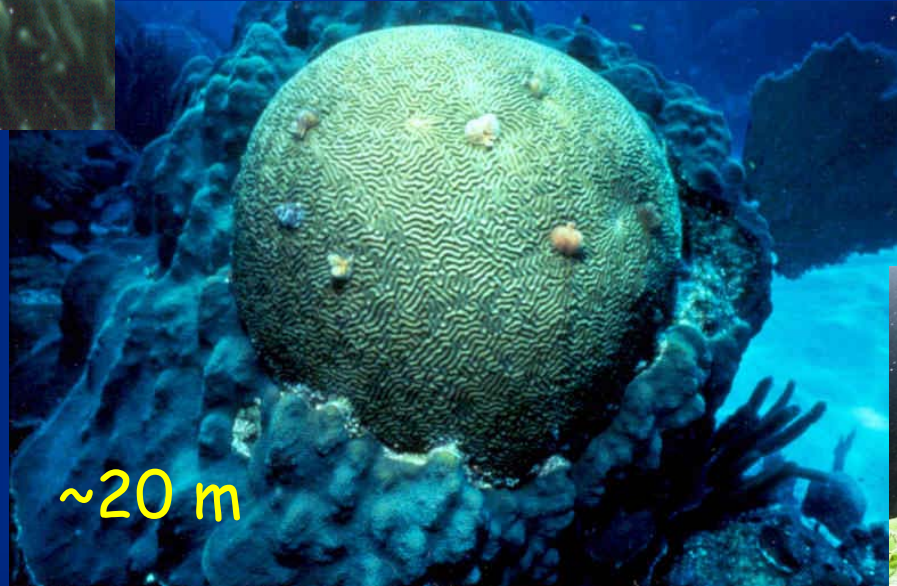


~3 m



Coral morphologies also vary predictably with light and hydrodynamics

~20 m



~60 m



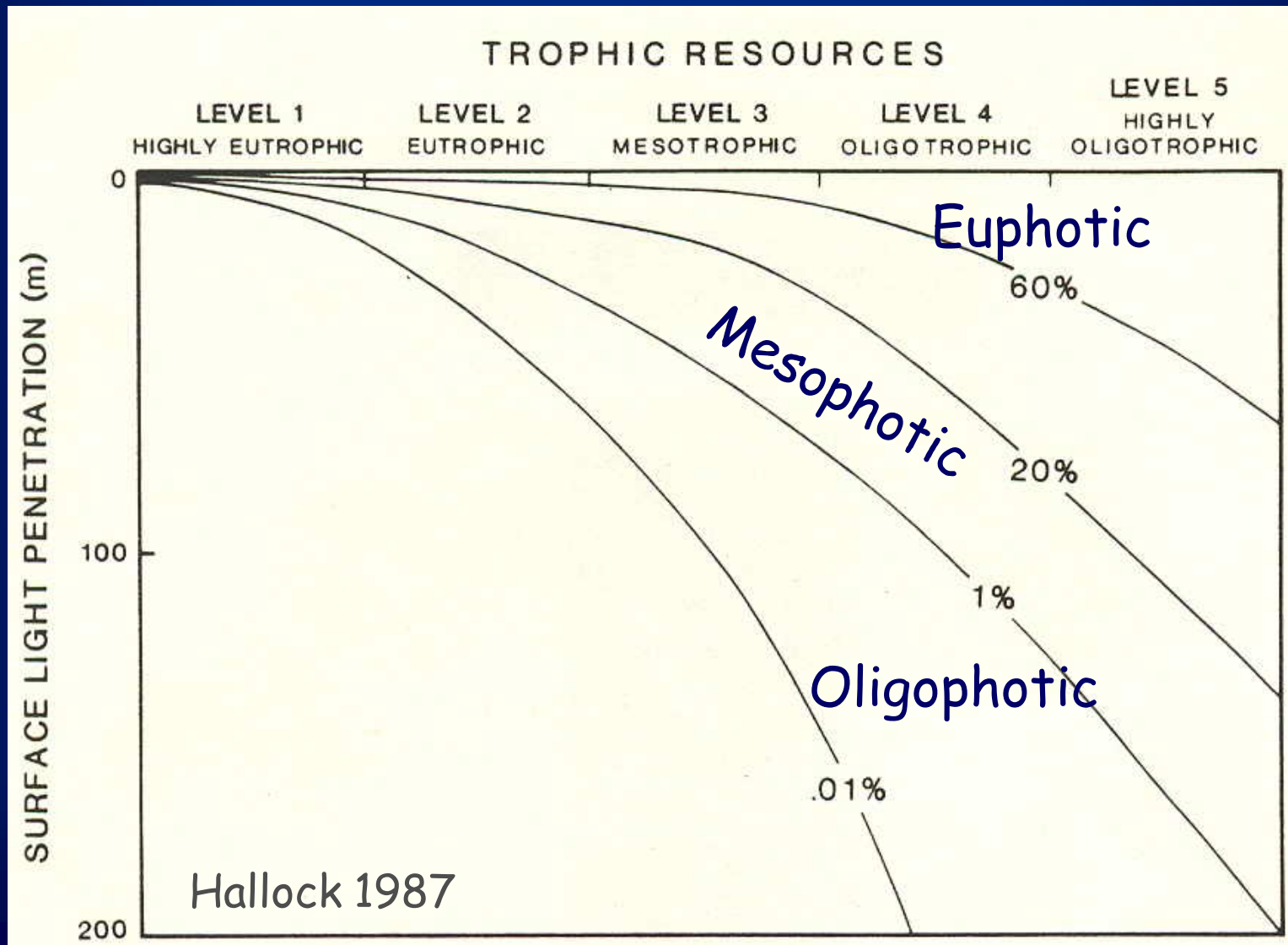
Photo courtesy of A.C. Hine.

Trophic (nutritional) resources

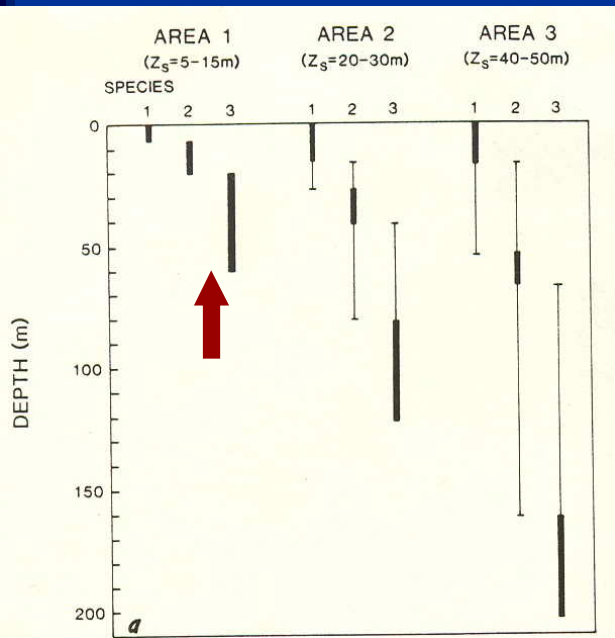
- In sunlight, trophic resources are functions of temperature and nutrient flux (N, P, trace nutrients)
- Below the photic zone, organic carbon = trophic resources
- Algal symbiosis is energetically most advantageous in resource-poor, warm, sunlit environments
- Diversities of habitats are strongly influenced by trophic resources
- Changes in trophic resources are linked to changes in ocean circulation
- Paleogene planktic and LBF evolutionary trends linked to paleoceanographic events

[Hallock 1981, 1987, 1988; Hallock & Schlager 1986; Hallock et al. 1991]

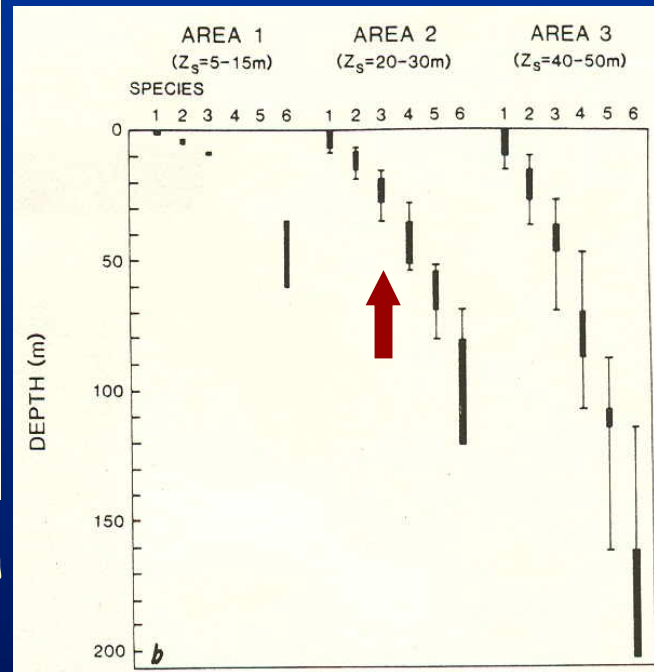
Trophic Resources and Light Penetration: inversely related



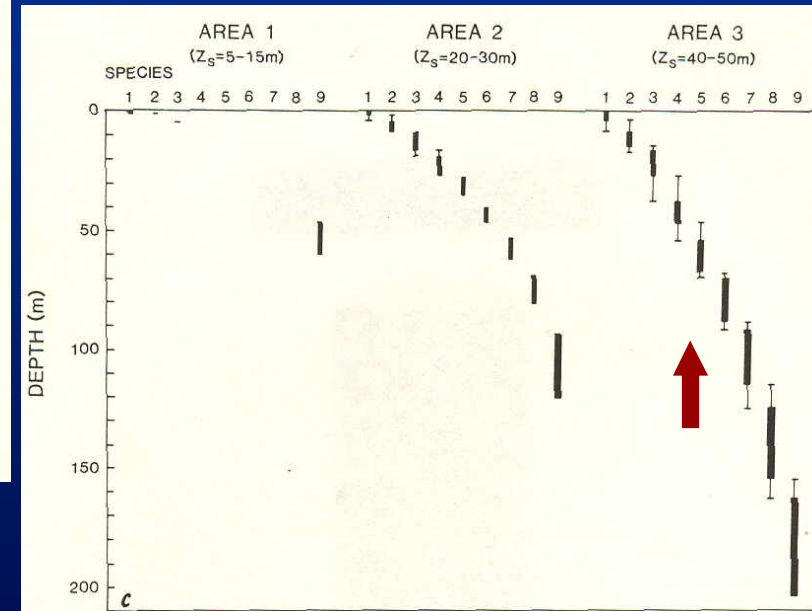
Predictably transparent seawater provides maximum potential for niche specialization in taxa dependent upon photosynthesis



Limited light penetration accommodates 3 species












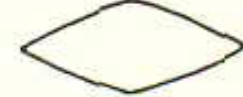



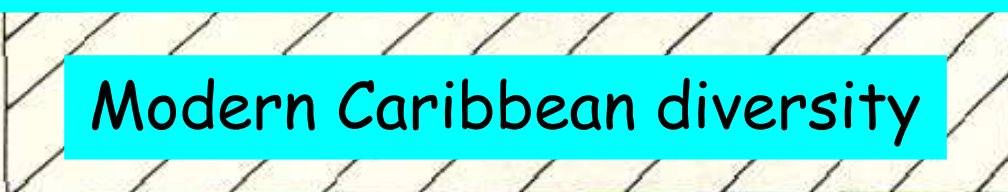




Medium light penetration accommodates 6+ species



Very clear water accommodates 9+ very specialized species

Simple model by Hallock (1987)





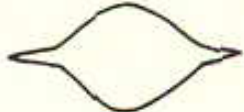
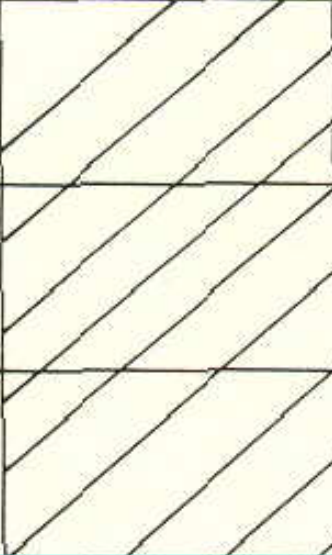



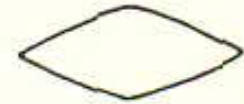





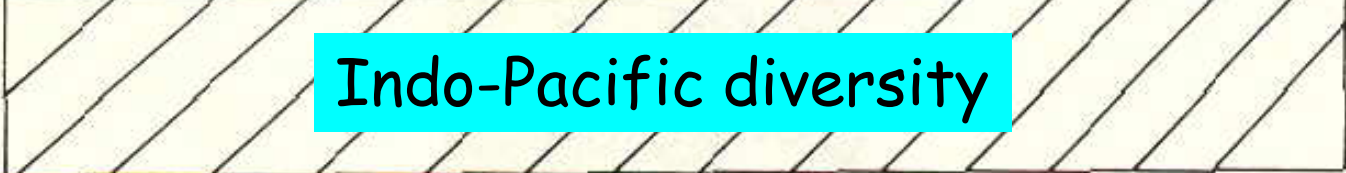
← T R C - HOLOCENE OCEANS →
 ← T R C - MAJOR MARINE TRANSGRESSION →

TROPIC RESOURCES PHOTIC HABITAT	LEVEL 1 HIGHLY EUTROPHIC	LEVEL 2 EUTROPHIC	LEVEL 3 MESOTROPHIC	LEVEL 4 OLIGOTROPHIC	LEVEL 5 HIGHLY OLIGOTROPHIC
SHALLOW ($<5\text{m}$)					
INTERMEDIATE ($5\text{--}30\text{m}$)					
MOD. DEEP ($30\text{--}70\text{m}$)					
DEEP ($70\text{--}130\text{m}$)					
VERY DEEP ($>130\text{m}$)					

Modern Caribbean diversity





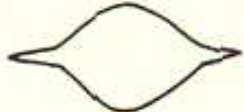
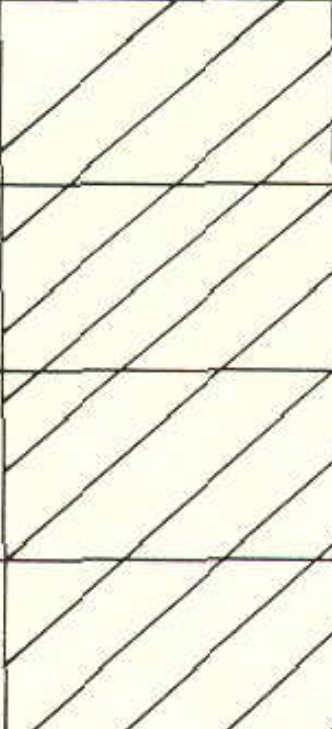



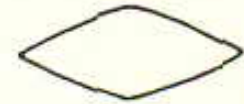







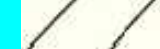

Modern Caribbean diversity

← T R C - HOLOCENE OCEANS →
 ← T R C - MAJOR MARINE TRANSGRESSION →

TROPIC RESOURCES EUPHOTIC HABITAT	LEVEL 1 HIGHLY EUTROPHIC	LEVEL 2 EUTROPHIC	LEVEL 3 MESOTROPHIC	LEVEL 4 OLIGOTROPHIC	LEVEL 5 HIGHLY OLIGOTROPHIC
SHALLOW (<5m)					
INTERMEDIATE (5-30m)					
MOD. DEEP (30-70m)					
DEEP (70-130m)					
VERY DEEP (>130m)					

Indo-Pacific diversity

← T R C - HOLOCENE OCEANS →
 ← T R C - MAJOR MARINE TRANSGRESSION →

TROPIC RESOURCES EUPHOTIC HABITAT	LEVEL 1 HIGHLY EUTROPHIC	LEVEL 2 EUTROPHIC	LEVEL 3 MESOTROPHIC	LEVEL 4 OLIGOTROPHIC	LEVEL 5 HIGHLY OLIGOTROPHIC
SHALLOW (<5m)					
INTERMEDIATE (5-30m)					
MOD. DEEP (30-70m)					
DEEP (70-130m)					
VERY DEEP (>130m)					

Mid-Eocene diversity



Calcareous green algae

Aragonite Producers



Most scleractinian corals



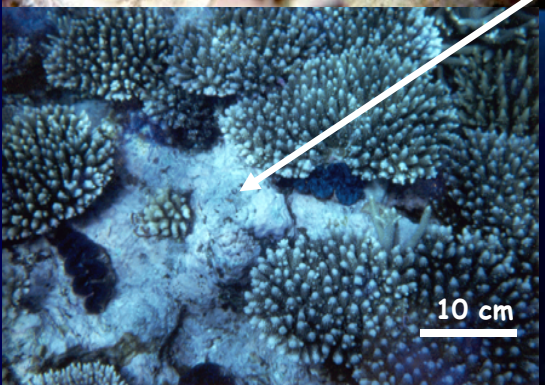
~10 cm

**HYPERCALCIFICATION
TYPICALLY
INVOLVES
PHOTOSYNTHESIS**
[Stanley and Hardie 1998]



**Mg-Calcite
Producers**
(> 8 mole% MgCaCO_3)

Coralline red algae



10 cm



10 m

**Calcite
Producers**

Many benthic foraminifers

Planktic foraminifers



~0.2 mm

Photo by: David A. Caron, Univ. Southern California



~1 mm

What are the costs of algal symbiosis?



**Aragonite
Producers**



**Calcite
Producers**

Many benthic foraminifers



Planktic foraminifers



~0.2 mm

Photo by: David A. Caron, Univ. Southern California

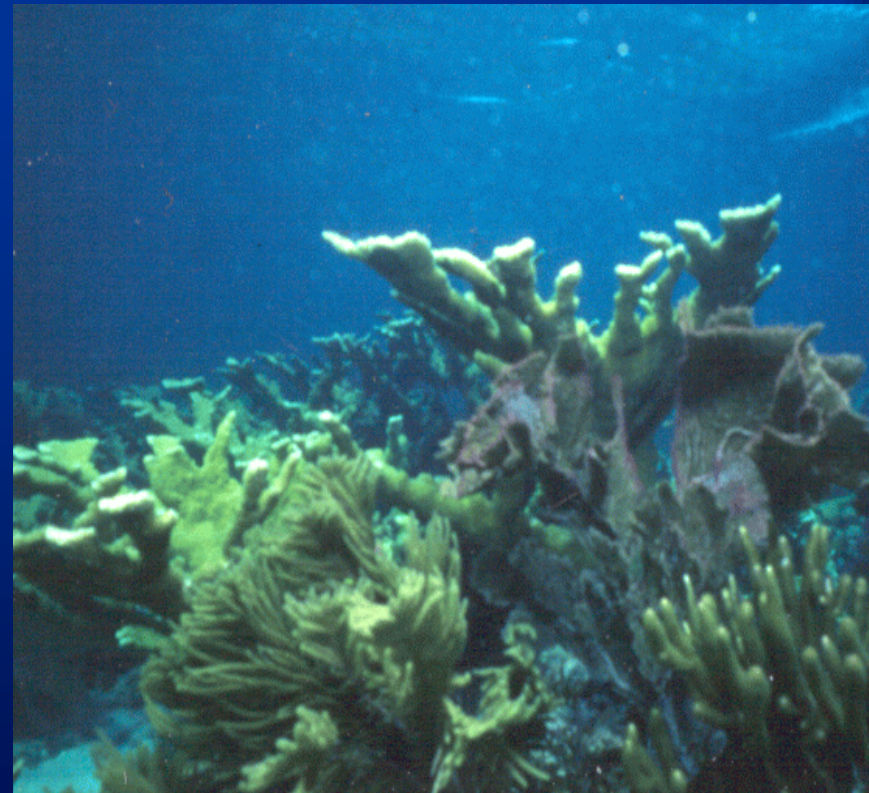


~1 mm

Risk-benefit considerations in high light environments

Euphotic – high energy/ high risk

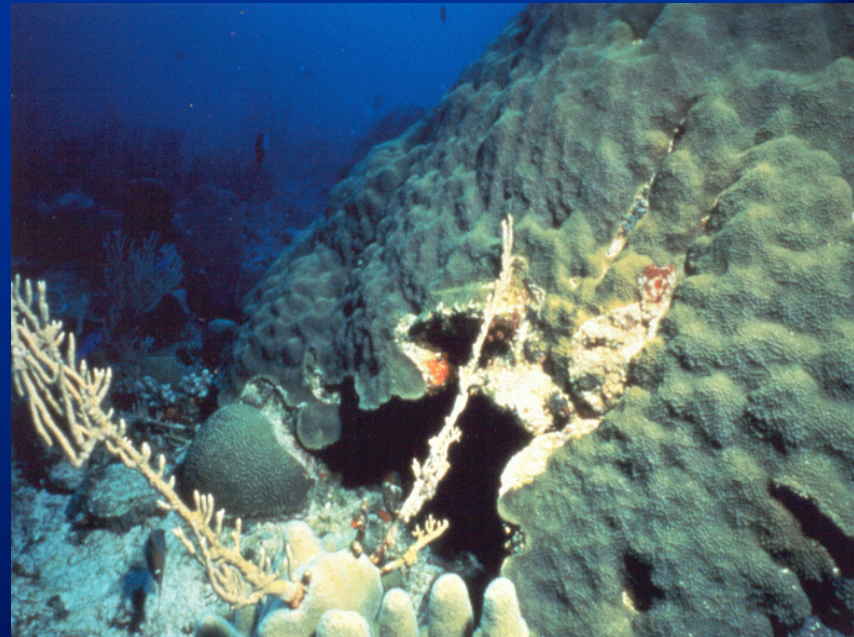
- + abundant energy for/from photosynthesis
- high potential for photo-oxidative stress (bleaching)
- high potential for physical damage
- limited POC resources in very clear waters



Risk-benefit considerations in intermediate light environments

Mesophotic – moderate energy/low risk

- + substantial energy for/from photosynthesis
- + low potential for photo-oxidative stress
- + limited risk of physical damage
- potential for insufficient solar energy with turbidity events
- + potential for energy supplementation with particulates



Risk-benefit considerations in very low light environments?

Oligophotic – low energy/high risk

- limited energy for/from photosynthesis
- + relatively abundant nutrients and particulates
- + low potential for photo-oxidative stress
- + low potential for physical damage
- predictably clear water is essential for symbiosis to be worth the “infrastructural” costs

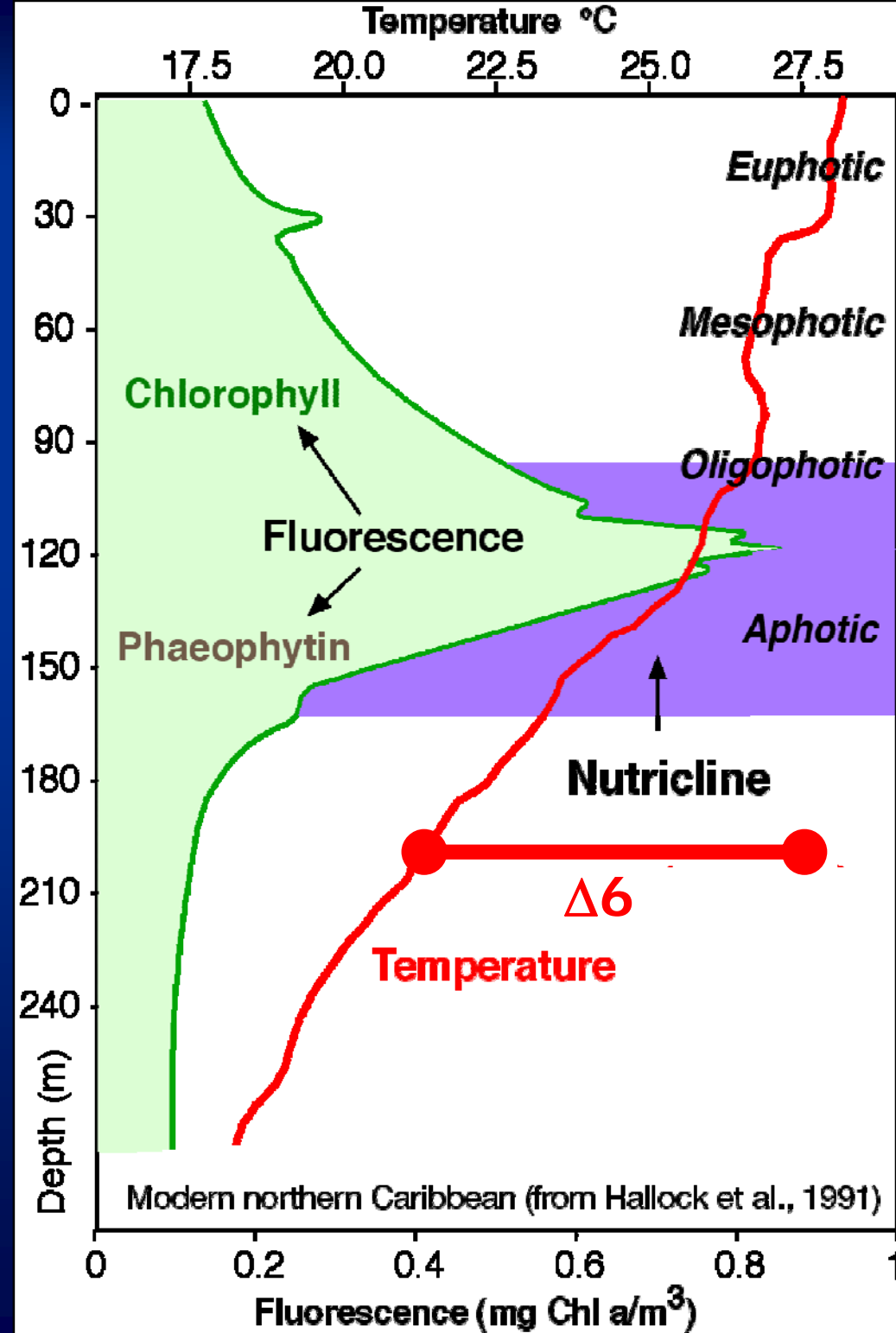


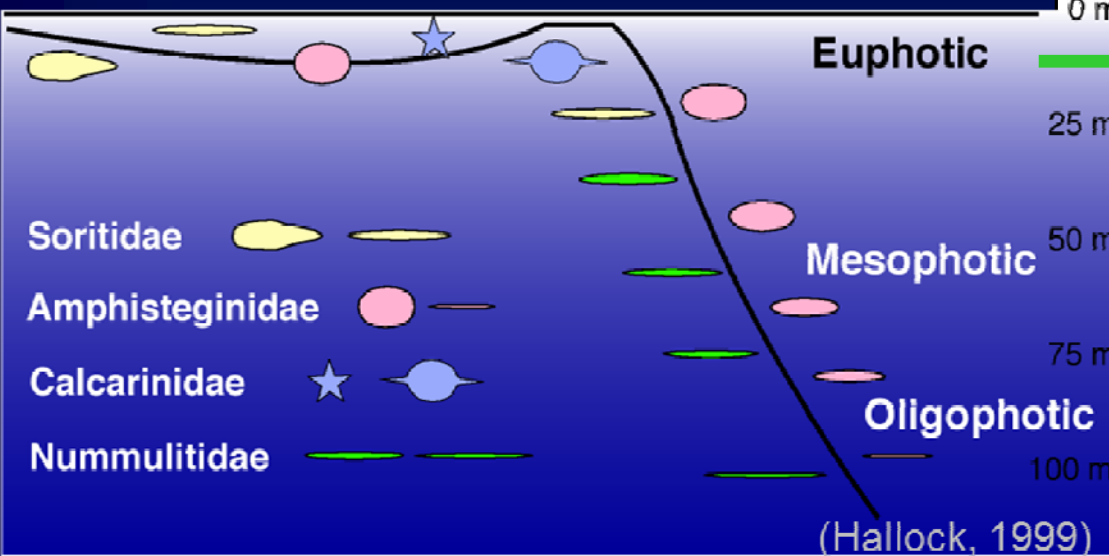
Photo courtesy of A.C. Hine.

Temperature and chlorophyll profiles for modern Caribbean

- 4X change in food supply in 120 m
- 6 degree change in temperature in 200 m

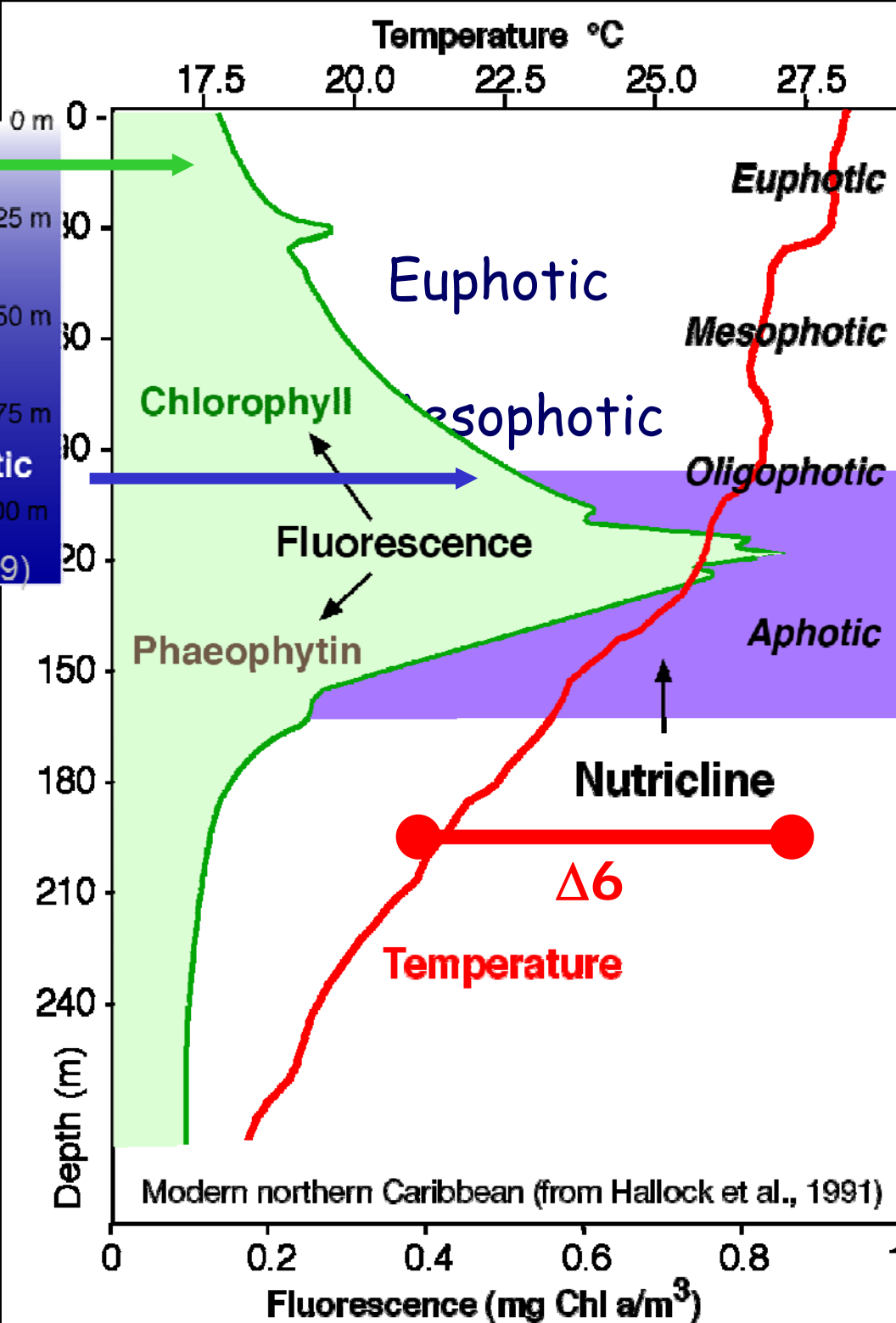
Hallock et al. 1991



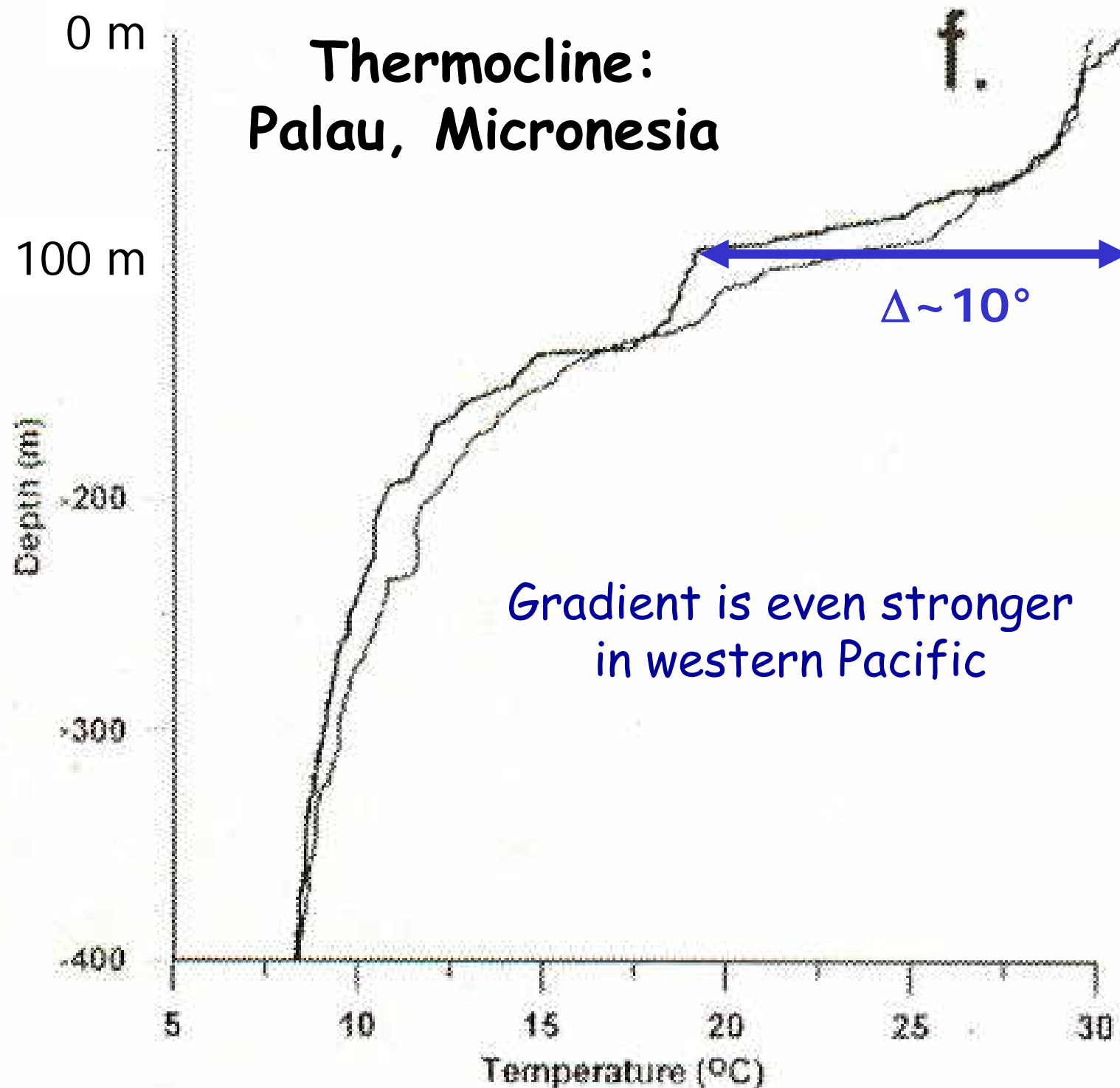


Shallow environments very different than mesophotic & oligophotic environments

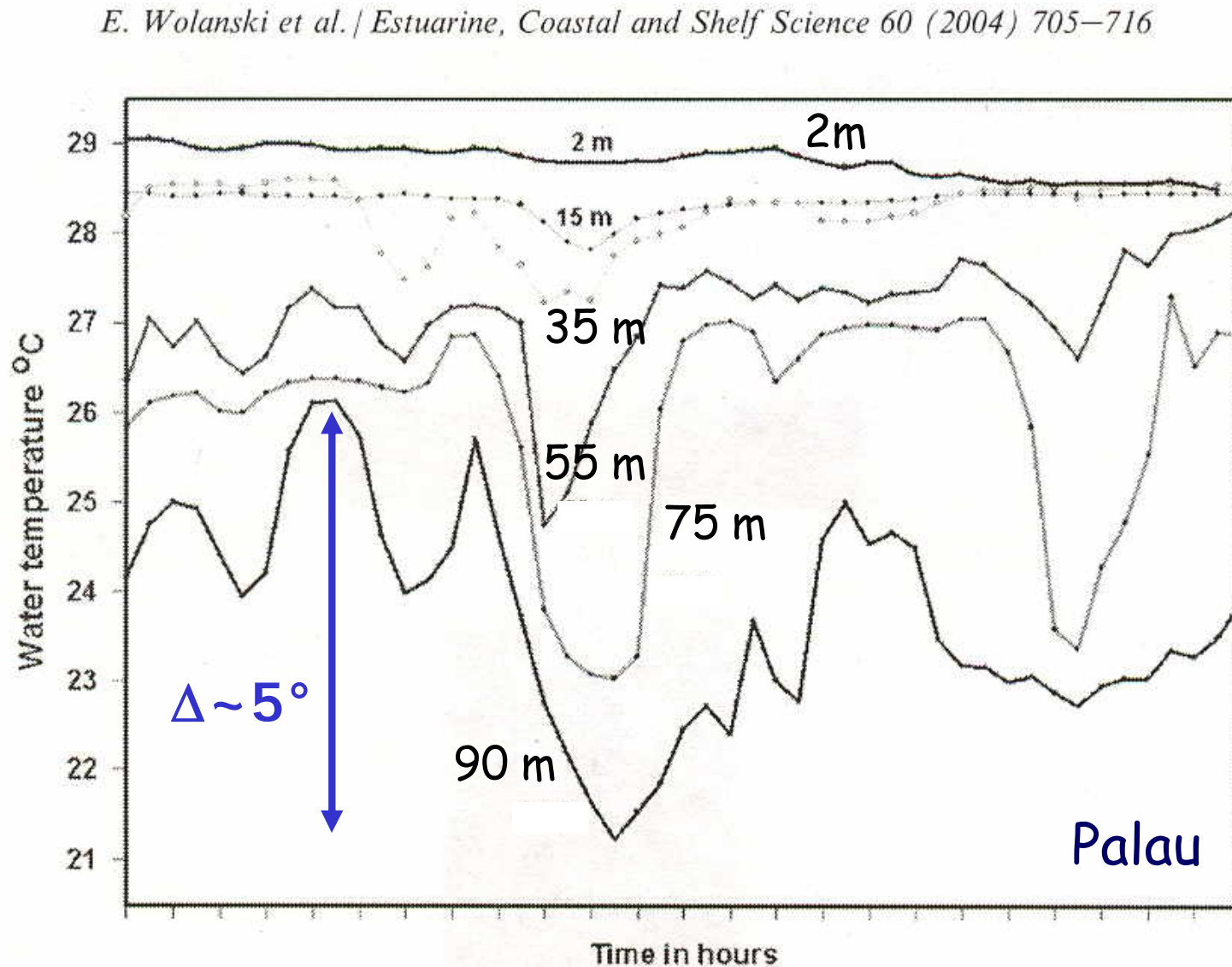
- Solar energy
- Hydrodynamics
- Food supply
- Temperature



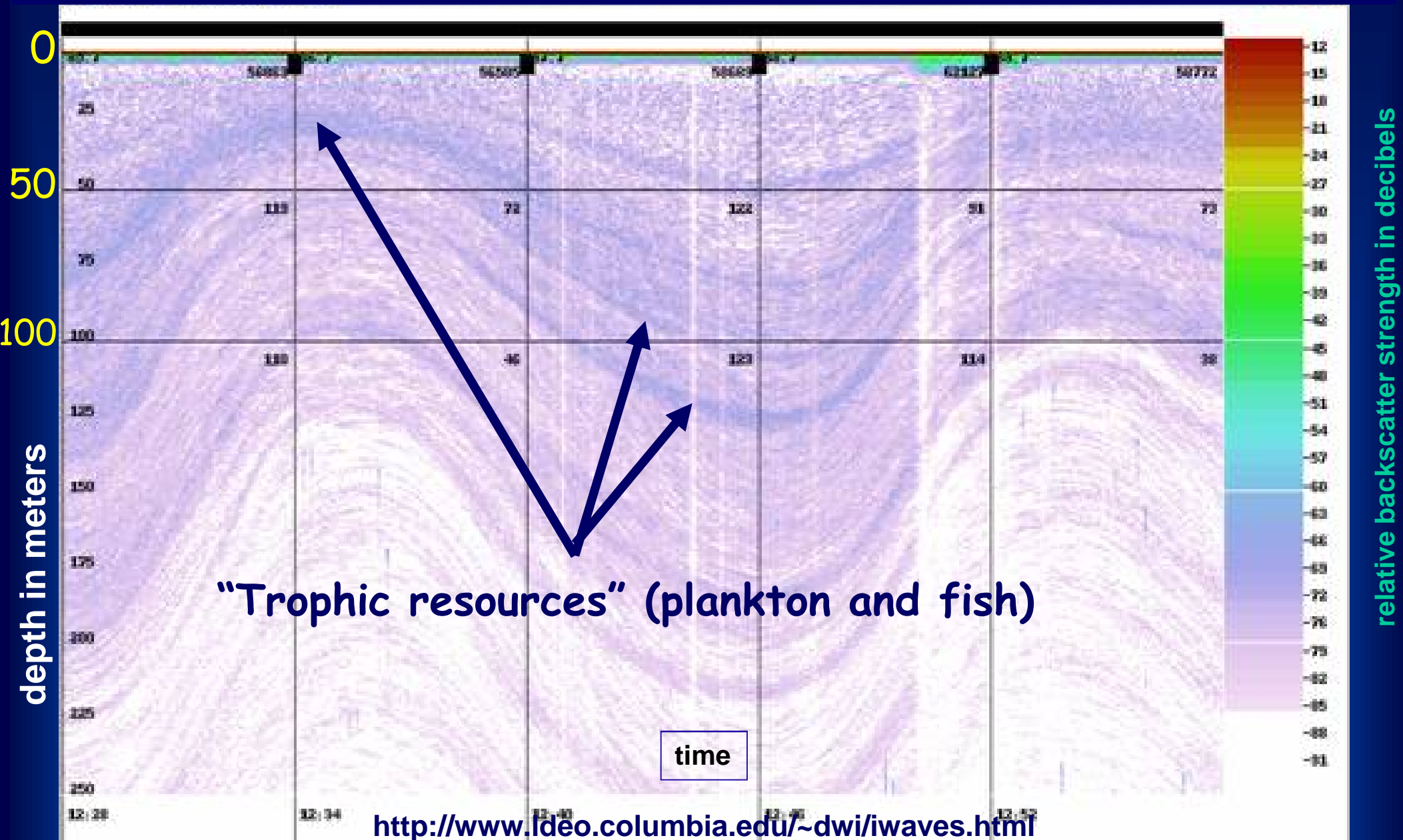
Thermocline: Palau, Micronesia

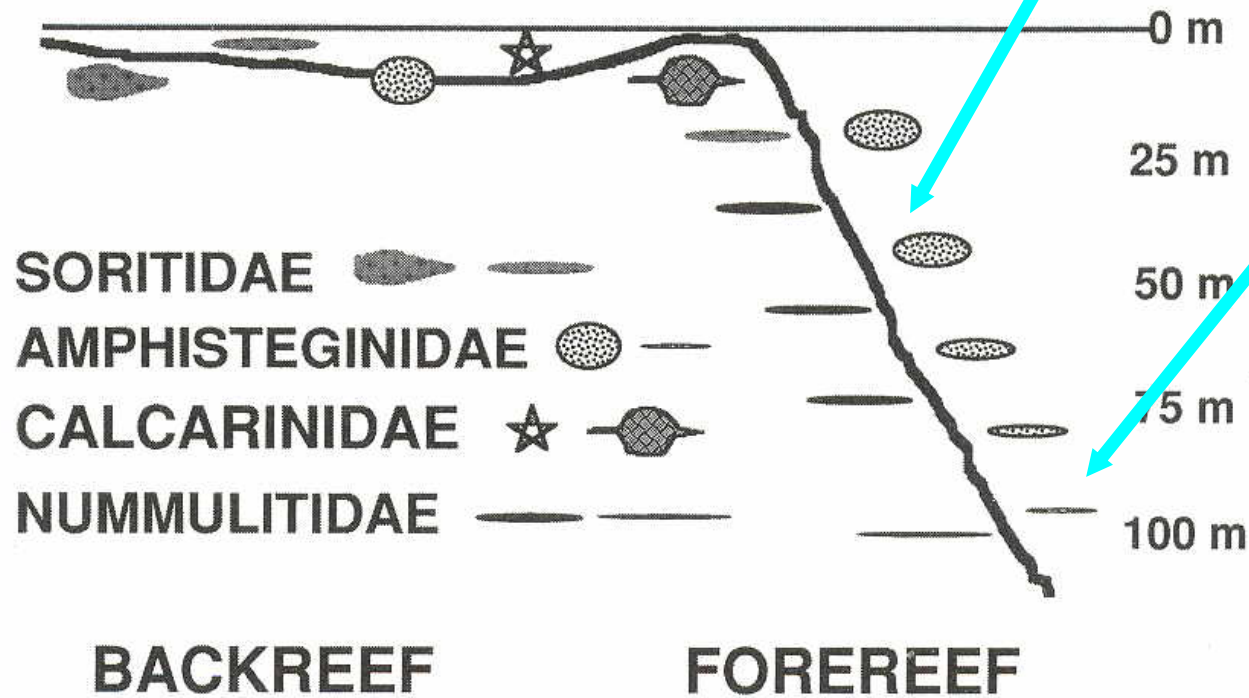
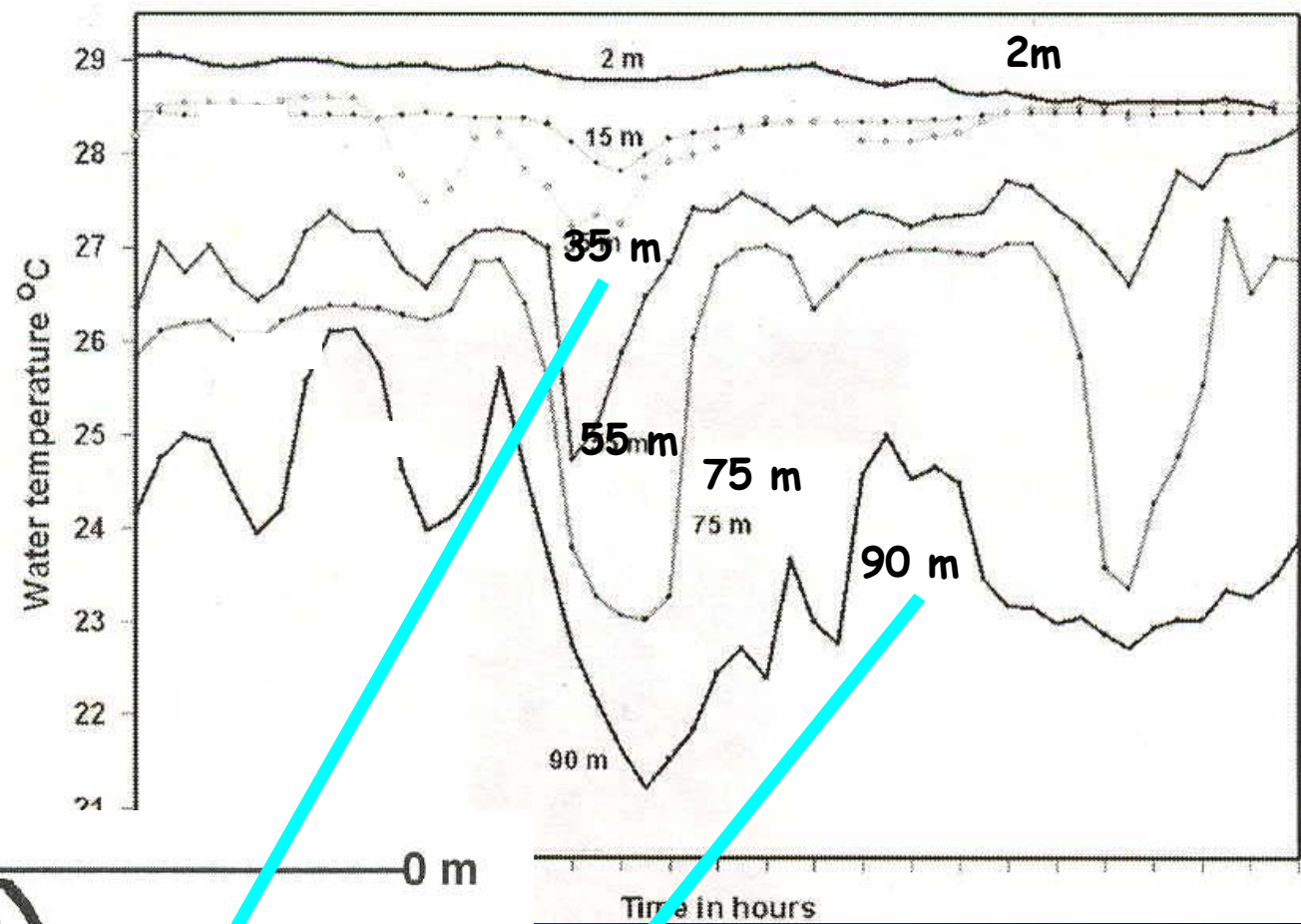


Where depth gradients are strong, internal waves also impart variability on tidal cycles, especially between 30 and 100 m



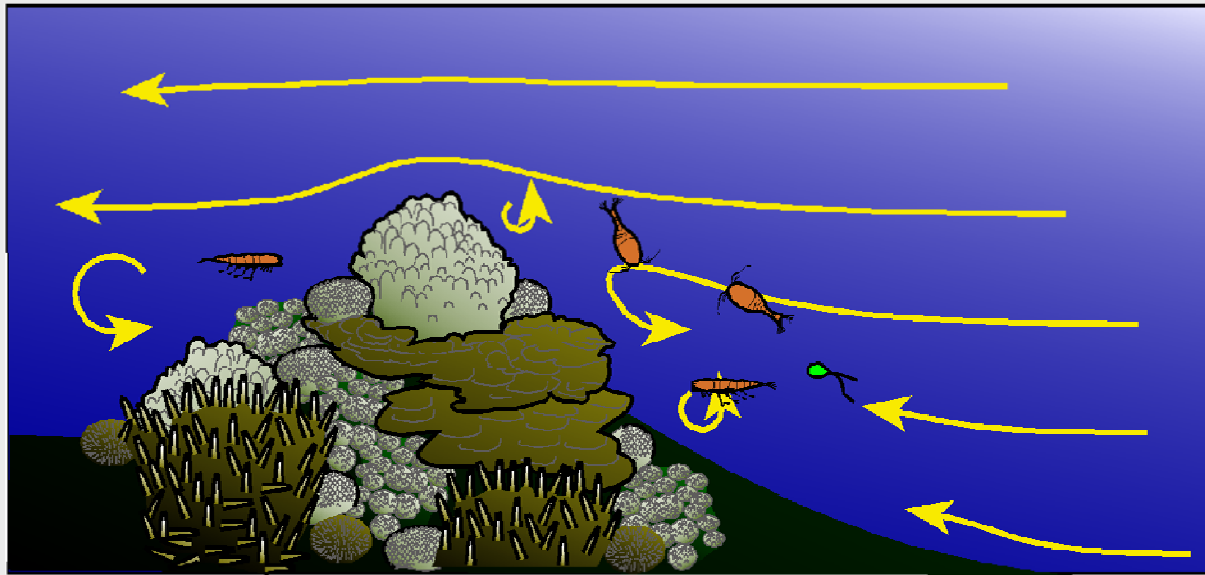
38 kHz echogram of internal waves in the Lombok Strait, Indonesia:
wavelength ~ 1.8 km; speed ~ 1.5 m/s; wave amplitude (peak to trough) > 100 m.
Higher backscatter values indicate higher plankton concentration or large schools of fish (Susanto et al, 2005)



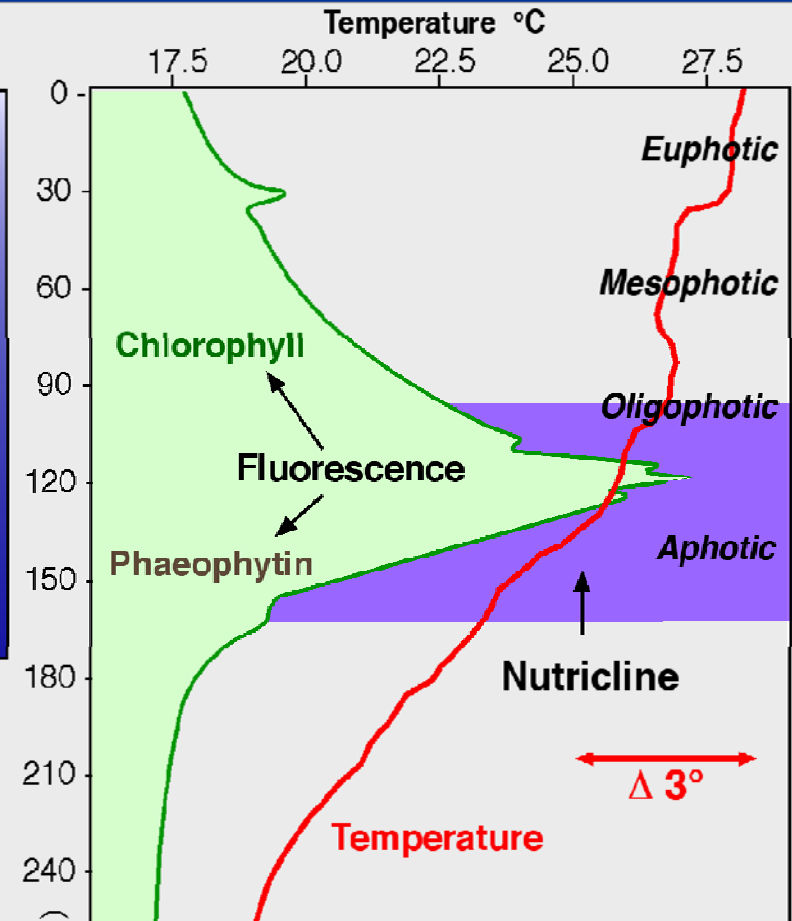


Modern mesophotic
and oligophotic
environments are
eurythermal

Internal waves and metazoan buildups



Pomar et al. (2012)

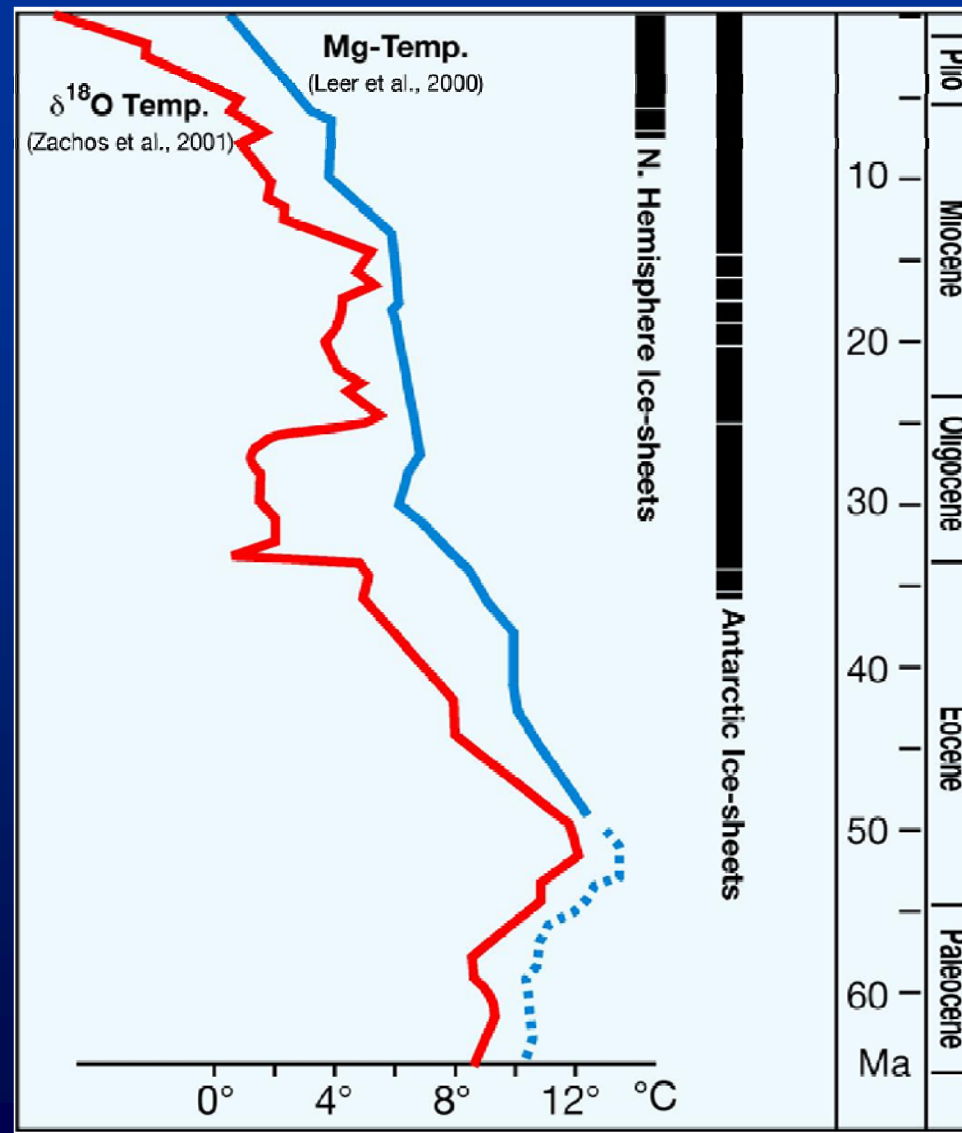


Internal waves (Leichter et al. 1998)

- a predictable, periodic source of transport for shelf-margin reefs
- an important influence on the spatial and temporal heterogeneity of suspended food particles and larval delivery to the benthos

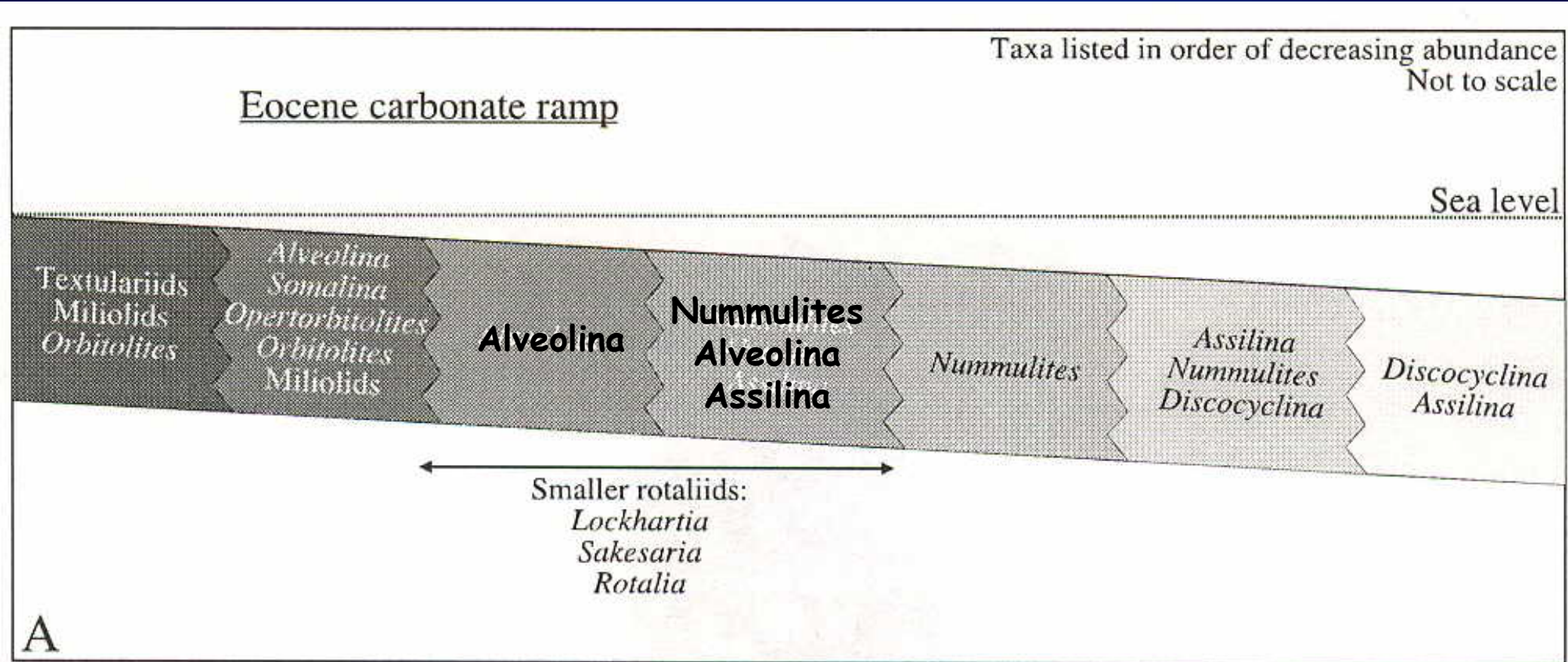
Internal waves provide food resources without limiting light for photosynthesis

Can deep sea paleotemperatures provide clues for shelf and ramp habitats?



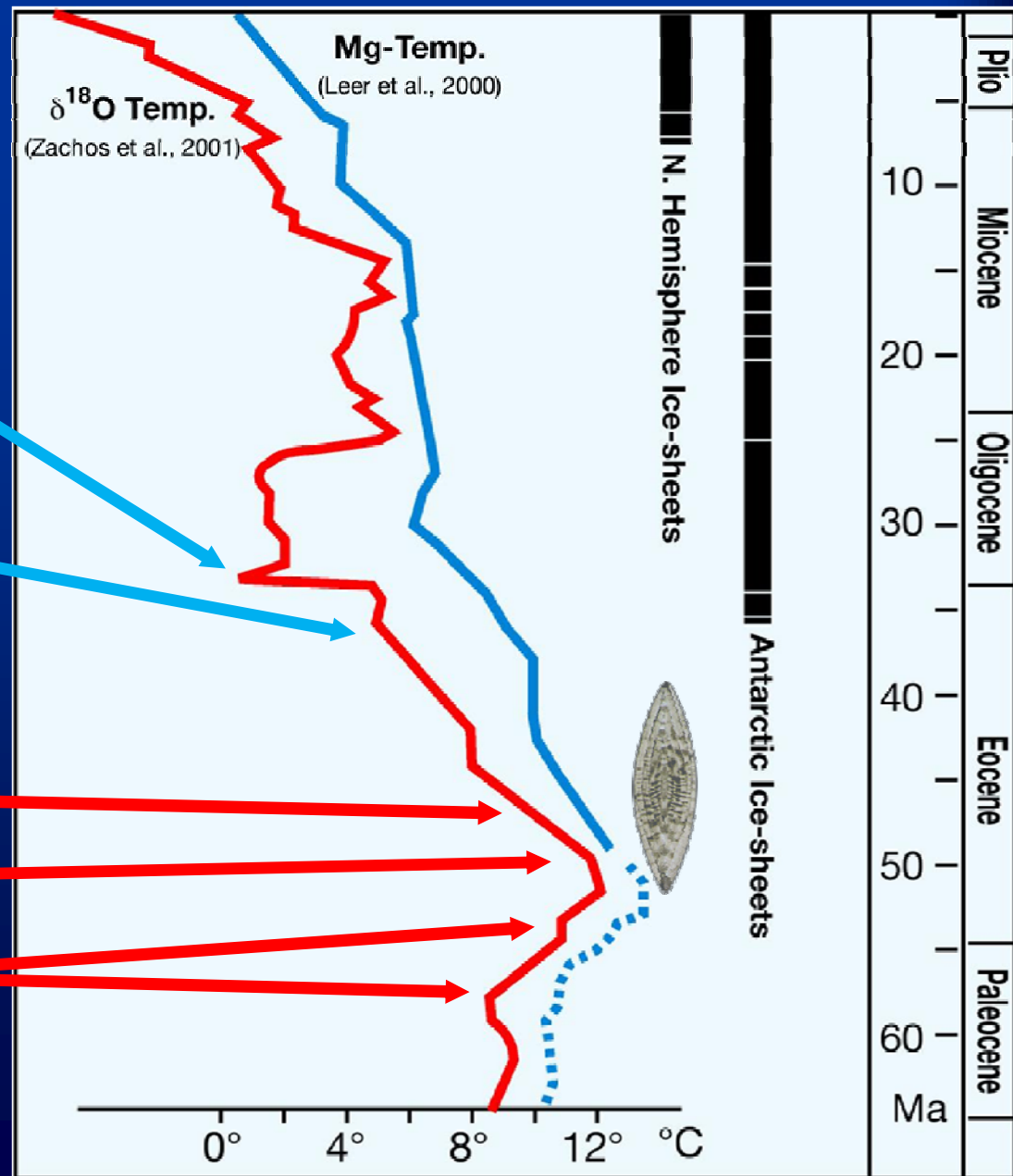
Eocene larger benthic foraminifera

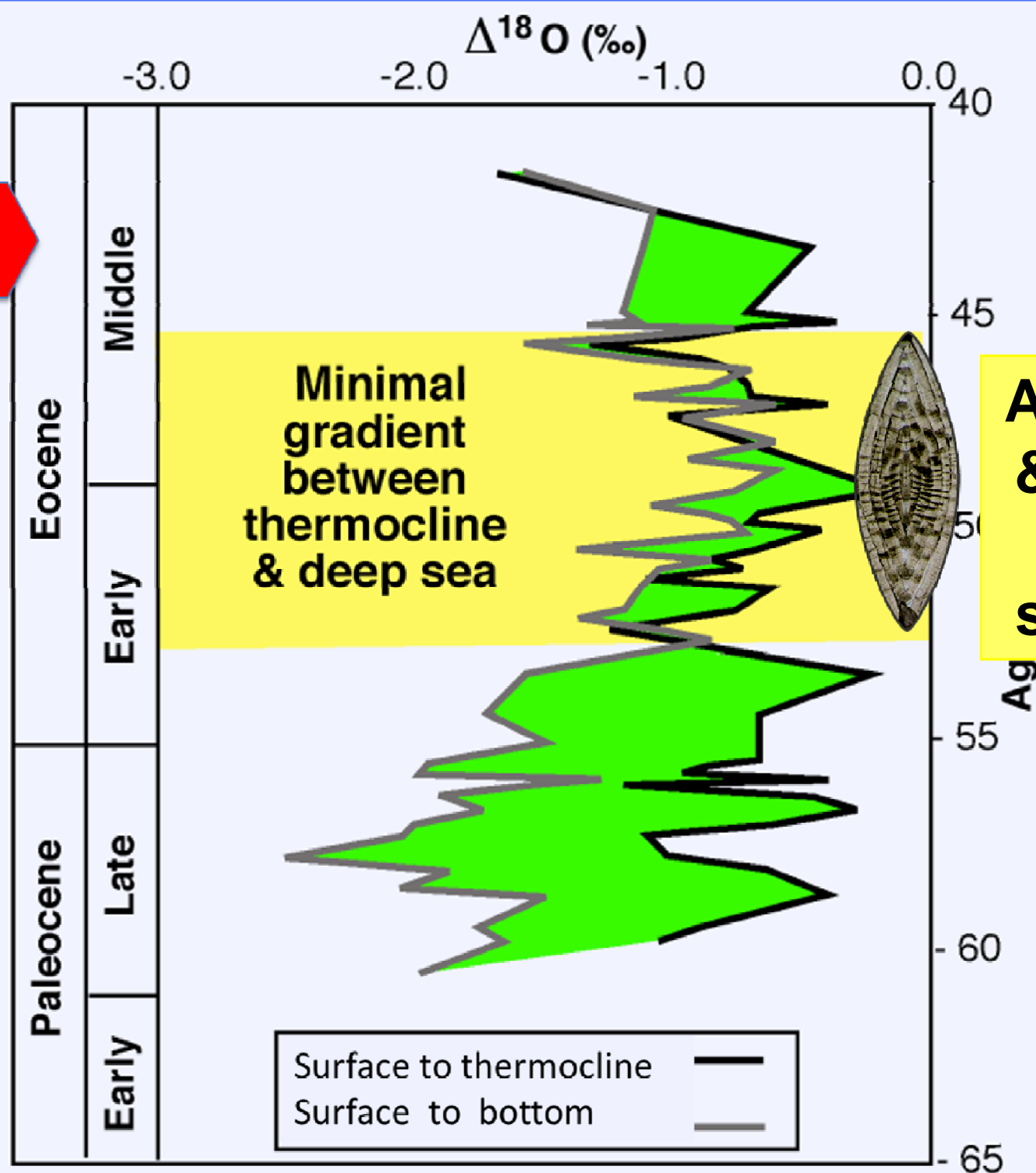
- Highly specialized horizontally and vertically
- What stratification and variability did they encounter?



Deep sea paleotemperatures and Paleocene-Eocene LBF lineages

- Extinctions when bottom temperatures $< \sim 10^{\circ}$
- Extinctions of orthophragmines
- Extinctions of large nummulites and some orthophragmines
- Diversification and peak at bottom temperatures $> \sim 10^{\circ}$
- Large nummulites peak
- 1st large nummulites
- Diversification of forams with complex morphologies



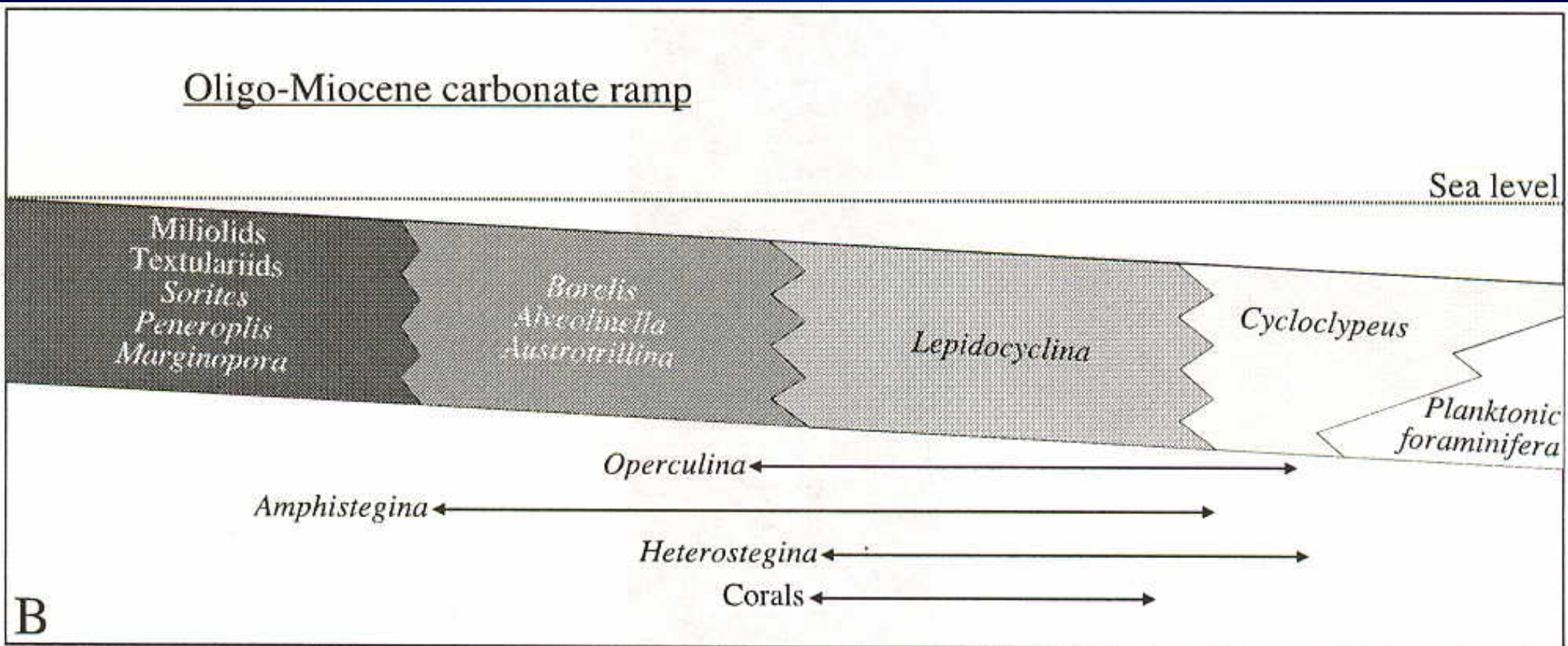


**Apex of nummulites
& orthophragmines
living in outer
shelf/ramp habitats**

Surface to thermocline
vs. surface to bottom
temperature gradients
in subtropical North
Pacific (Shatzky Rise)

Oligo-Miocene LBF Assemblages

- Lower diversity horizontally and vertically
- What stratification and variability did they encounter?

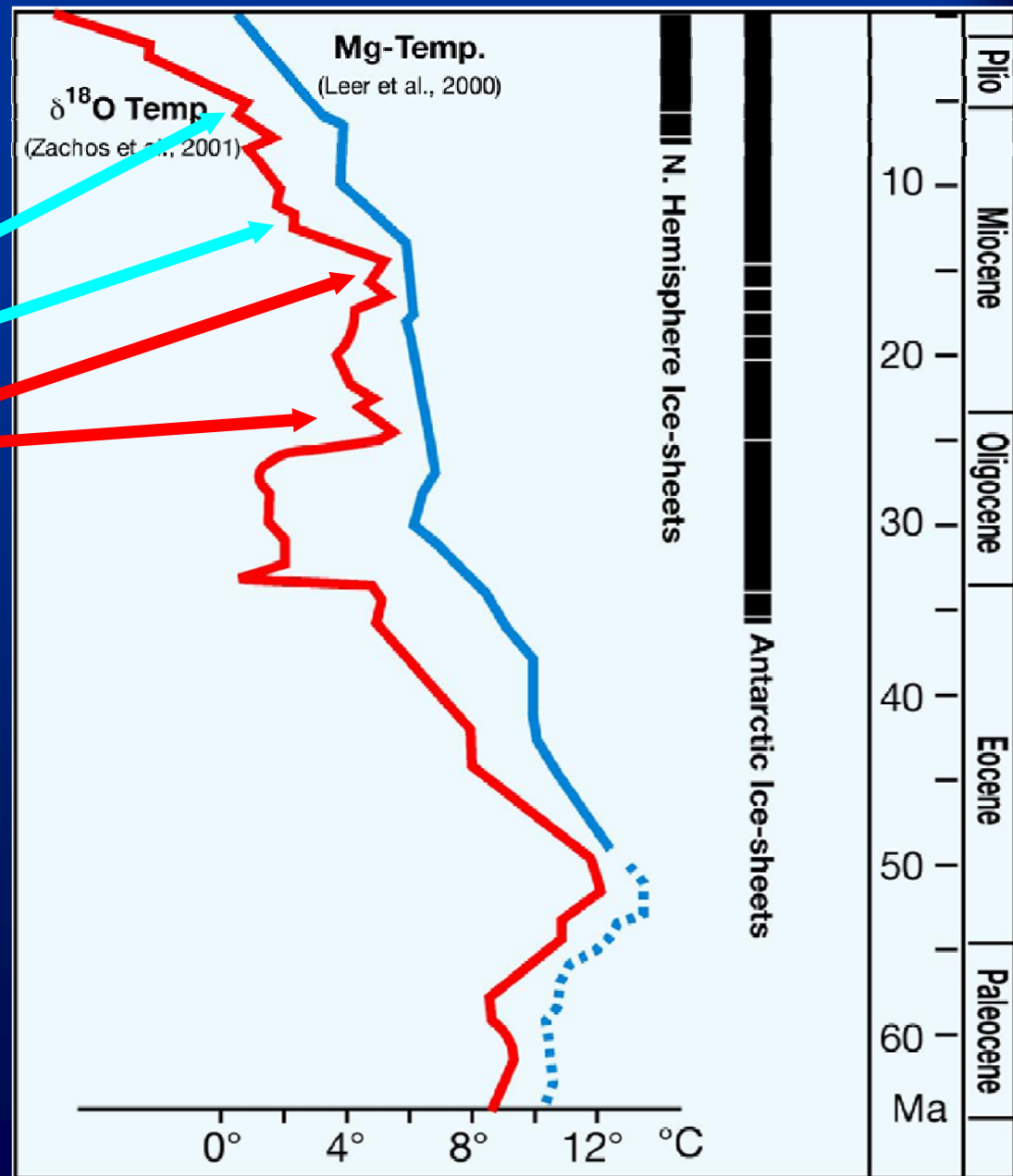


Deep sea paleotemperatures and Oligo-Miocene LBF lineages

Extinction of lepidocyclinids

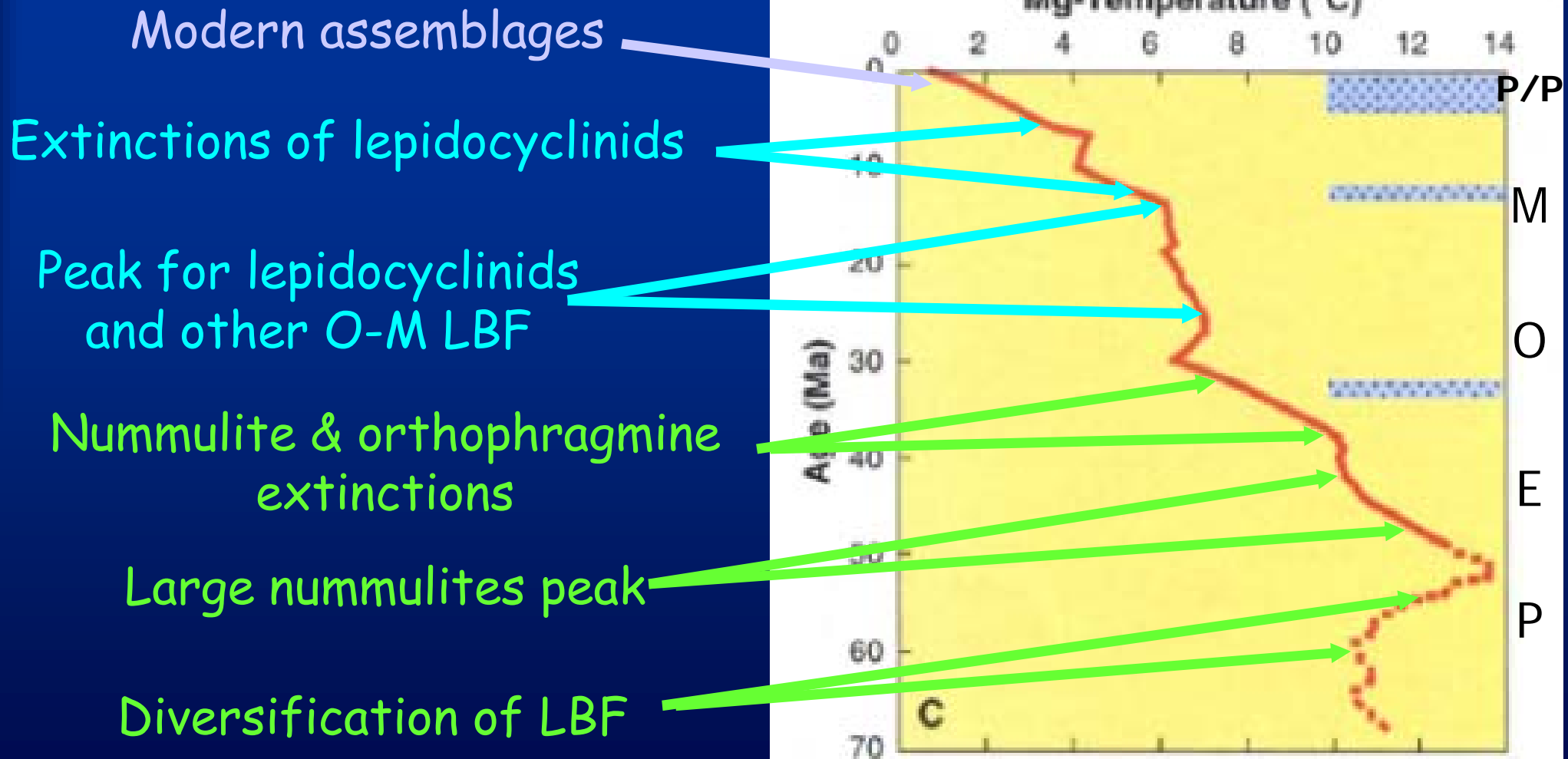
Decline in lepidocyclinids

Peak for lepidocyclinids
and other O-M LBF



From various sources

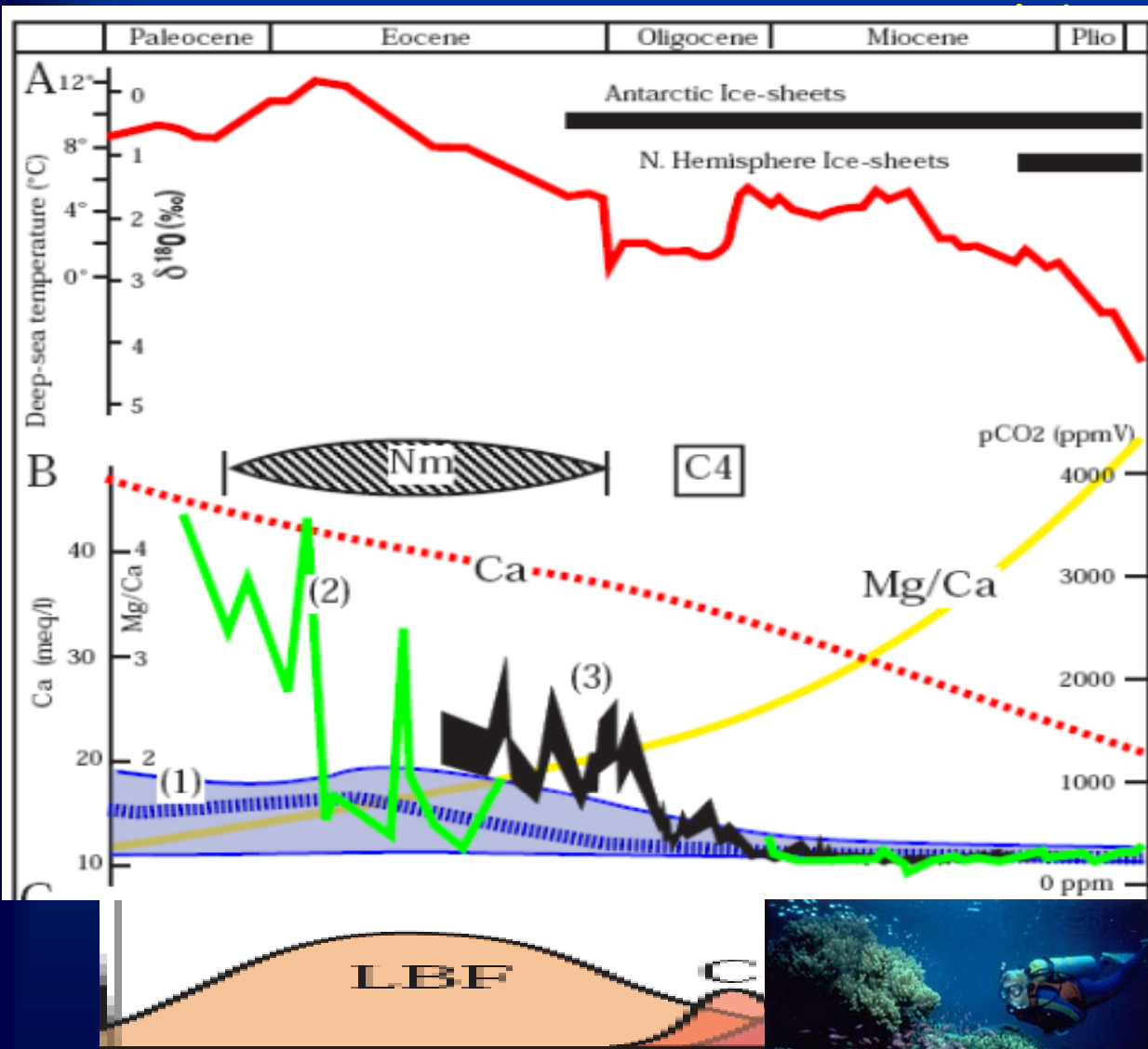
Summary: deep sea temperatures and LBF lineages



Interpretations

1. Cenozoic changes in ocean circulation increased bathymetric as well as latitudinal temperature gradients
 - Also likely influenced bathymetric nutrient gradients
2. Increasing bathymetric temperature gradients strongly influenced mesophotic to oligophotic environments
3. Largest LBF taxa live at oligophotic depths
4. Turnovers in LBF lineages coincide with strong declines in deepwater paleotemperatures
5. Declining diversities in "peak" LBF biotas are consistent with increasing temperature variability in mesophotic to oligophotic environments through the Cenozoic

The Paradox of the Cenozoic: So why do coral reefs like the Icehouse



- Greenhouse → Icehouse
 - $[Ca^{2+}]$ ↓
 - $[Mg^{2+}]$ ↑
 - pCO_2 ↓
- Succession
 - Calcite/Mg-calcite → aragonite/Mg-calcite
 - LBF/RA → corals/RA
- Neogene co-evolution of red algae, corals & light-tolerant zooxanthellae into shallow, high light environments
 - Provided energy for hypercalcification



Concluding Remarks

- Synthesis builds upon previous interpretations of LBF paleoecology and influence of algal symbiosis
- Presents new understanding of temporal variability of shelf environments at 30-100+ m
- Bathymetric gradients have changed dramatically on mid-low latitude ramps and shelf margins through the Cenozoic
- Provides new context to utilize planktic paleotemperature data to better understand evolutionary trends
 - in Cenozoic LBF assemblages
 - in Cenozoic carbonate ramp/shelf communities

Previous and Current Sources of Funding for LBF Studies

Funded by EPA's Science
To Achieve Results (STAR)
Program Grant # R82 -5869



Florida Hurricane
Mitigation Alliance

Note: Funding does not
imply endorsement of
views presented





Thank you