Scaling Relationships in Fluvial Depostional Systems*

Kristy Milliken¹, Mike Blum², and John Martin¹

Search and Discovery Article #30245 (2012)**
Posted August 20, 2012

*Adapted from oral presentation at AAPG Annual Convention and Exhibition, Long Beach, California, April 22-25, 2012

Abstract

Fluvial systems possess a range of scaling relationships that reflect drainage-basin controls on water and sediment flux. In hydrocarbon exploration and production, scaling relationships for fluvial deposits can be utilized to constrain environmental and sequence-stratigraphic interpretations, as well as predict the lateral extent of fundamental reservoir flow units. This study documents the scales of channel fills, point and channel bars, channel belts, and coastal-plain incised valleys from well-constrained Quaternary fluvial systems.

Data on channel-fill and point-bar to channel-belt scales were compiled from published thicknesses for sinuous single-channel systems, with spatial dimensions measured from GoogleEarth. Fluvial systems included in this database span 3 orders of magnitude in drainage area, from continental-scale systems to small tributaries, and span tropical to sub-polar climatic regimes. Channel-fill and channel-belt scales were measured upstream from backwater effects, so as to minimize inclusion of distributive, highly avulsive systems. Scales of incised valleys were derived from well-constrained published examples that are known to have formed during the last 100 kyr glacio-eustatic cycle.

All scaling relationships are represented by statistically-significant power laws, with absolute dimensions that scale to drainage area, but distinct clustering occurs between channel fills, point bars and channel belts, and incised valleys. Mean width-to-thickness ratios for channel fills are ~10:1, whereas point bars commonly range from 70-250:1. Coastal-plain incised valleys from the last glacio-

^{**}AAPG © 2012 Serial rights given by author. For all other rights contact author directly.

¹ExxonMobil Exploration Company, Houston, TX

²ExxonMobil Upstream Research, Houston, TX (mike.blum@exxonmobil.com)

eustatic cycle range from 25-150 m in thickness, and a few kilometers to more than 80 km in width, with width-to-thickness ratios of \sim 600-800.

Scales of Quaternary examples compare well with previous compilations of channel-belt scales interpreted in the ancient record, and with theory. However, the smallest Quaternary incised valleys in our database reside in the uppermost part of the domain of published compilations of ancient incised valleys, with ancient examples overlapping significantly with both interpreted channel fills and channel belts. When interpreted within the context of this database from modern systems, we suggest many ancient examples may have been overinterpreted, which in turn suggests a persistent lack of objective criteria for differentiating channel fills, channel belts, and incised valleys.

References

Blum, M.D., M.J. Guccione, D.A. Wysocki, P.C. Robnett, and E.M. Rutledge, 2000, Late Pleistocene evolution of the lower Mississippi River valley, southern Missouri to Arkansas: GSA Bulletin, v. 112/2, p. 221-235.

Blum, M.D., J.H. Tomkin, A. Purcell, and R.R. Lancaster, 2008, Ups and downs of the Mississippi Delta: Geology, v. 36/9, p. 675-678.

Jackson, R.G., II, 1976a, Depositional model of point bars in the lower Wabash River: Journal of Sedimentary Petroleum, v. 46, p. 579-594.

Jackson, R.G., II, 1976b, Large scale ripples of the lower Wabash River: Sedimentology, v. 23, p. 593-632.

Kolla, V., H.W. Posamentier, and L.J. Wood, 2007, Deep-water and fluvial sinuous channels; characteristics, similarities and dissimilarities, and modes of formation, in R.B. Wynn, and B.T. Cronin, (eds.), Sinuous deep-water channels; genesis, geometry and architecture: Marine and Petroleum Geology, v. 24/6-9, p. 388-405.

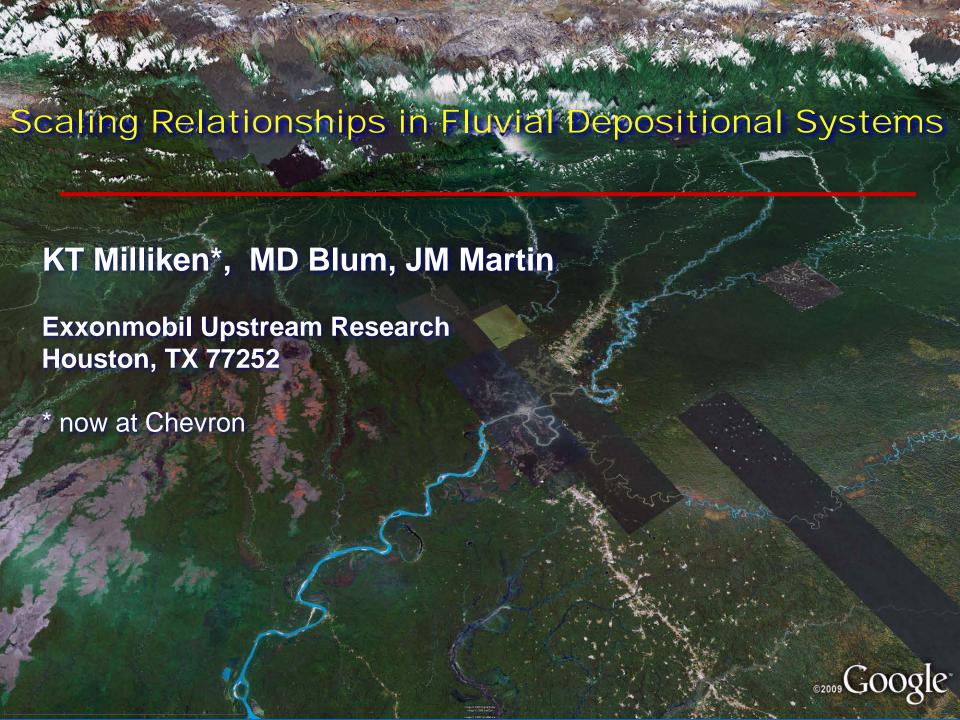
Miall, A.D., 2002, Architecture and sequence stratigraphy of Pleistocene fluvial systems in the Malay Basin, based on seismic time-slice analysis: AAPG Bulletin, v. 86/7, p. 1201-1216.

Reijenstein, H.M., H.W. Posamentier, and J.P. Bhattacharya, 2011, Seismic geomorphology and high-resolution seismic stratigraphy of inner-shelf fluvial, estuarine, deltaic, and marine sequences, Gulf of Thailand: AAPG Bulletin, v. 95/11, p. 1959-1990.

Rittenour, T.M., R.J. Goble, and M.D. Blum, 2005, Development of an OSL chronology for late Pleistocene channel belts in the Lower Mississippi Valley, USA: Quaternary Science Reviews, v. 24/23-24, p. 2539-2554.

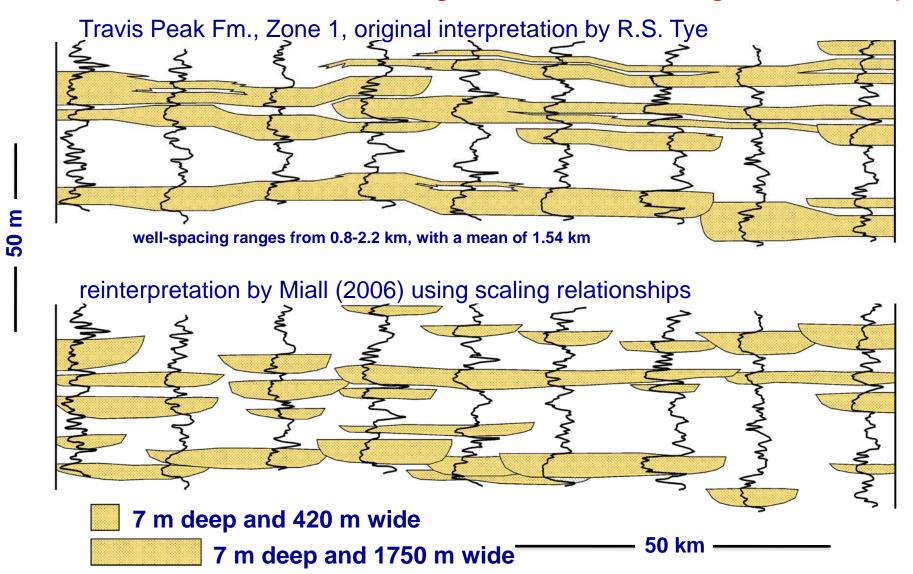
Rittenour, T.M., M.D. Blum, and R.J. Goble, 2007, Fluvial evolution of the lower Mississippi River valley during the last 100 k.y. glacial cycle; response to glaciation and sea-level change: GSA Bulletin, v. 119/5-6, p. 586-608.

Willis, B.J., S.L. Dorobek, and Y. Darmadi, 2007, Three-dimensional seismic architecture of fluvial sequences on the low-gradient Sunda Shelf, offshore Indonesia: JSR, v. 77/3-4, p. 225-238.



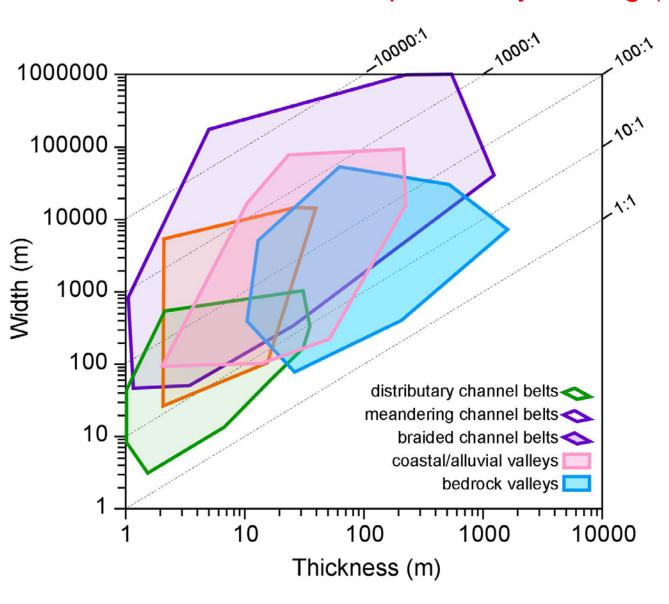


Channel-Belt Sand Bodies – Significance of Scaling Relationships





Motivations – Recent Compilation by Gibling (2006)



Compilation of scales in published literature

Scale domains are very different from modern systems, much larger range

Raises issue of how original data were interpreted, and criteria that were used

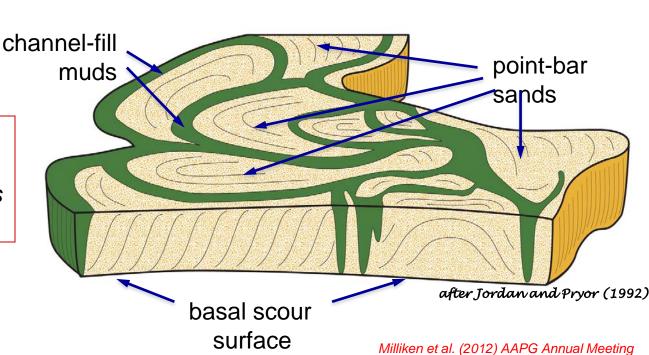
Need to cross-check with modern systems, where dimensions are an observation, not interpretation



Goals – Quantify Scales of 1st Order Fluvial Elements

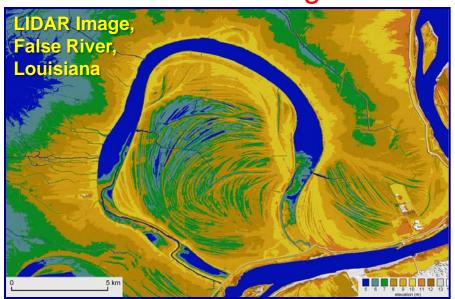
- Construct a database of modern channel-belt and channel-fill scales
 - Google Earth Approach: Scour literature and global landsat / DEM
 - mature point bars, recently cutoff or close to cutoff
 - point bars with thickness measurement, i.e. core through deposit, measure width, thickness, asymetry
- Construct dataset of Late Quaternary incised-valley scales
 - coastal-plain valleys with robust geochronological control

channel fills and channel-belt margins are fundamental barriers to flow

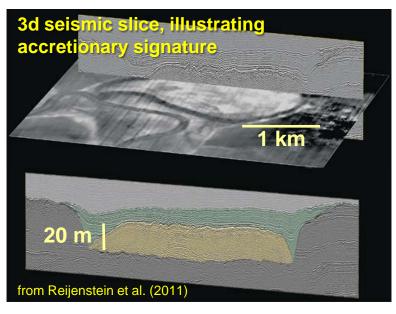




Bar Accretion and Large-Scale Inclined-Strata Sets



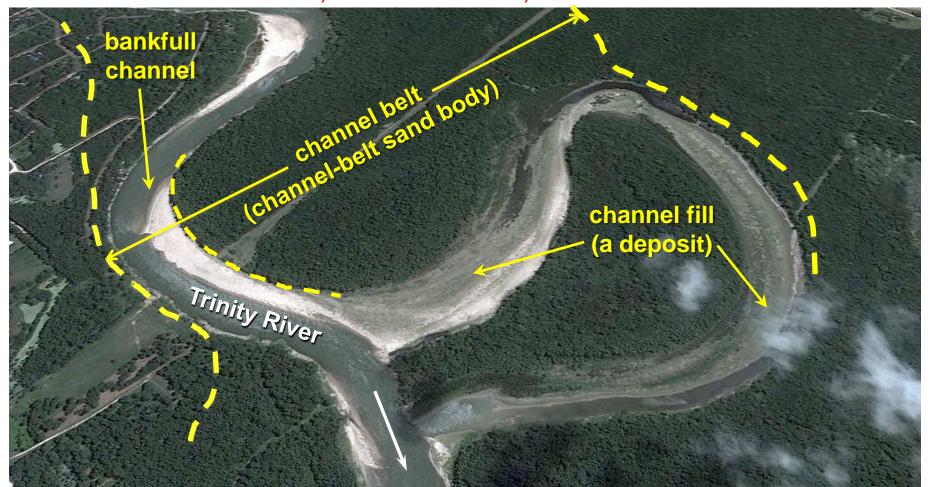








Definitions - Channels, Channel Fills, and Channel Belts

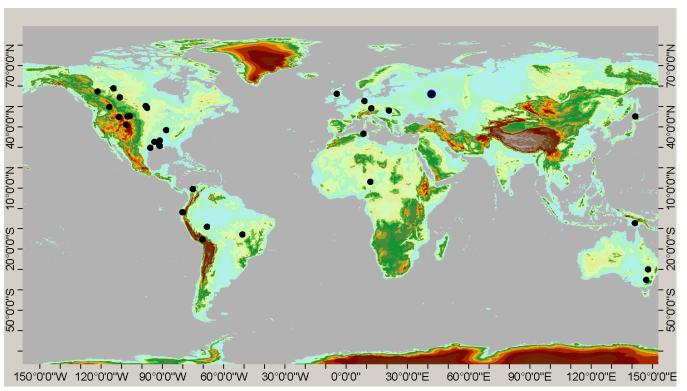






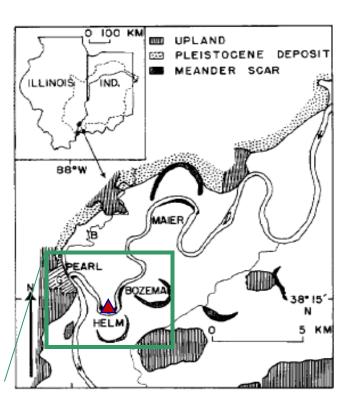
Dataset

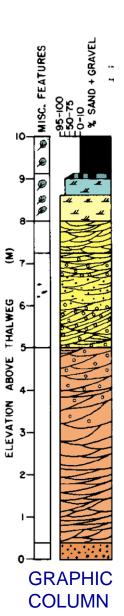
- Channel-belt and channel-fill scales from 38 modern rivers
 - periglacial to tropical
 - drainage areas = 250 to 3,000,000 km²
 - 124 measured meanders
- Incised-valley Scales from 10 Late Quaternary Systems
 - drainage areas = $50,000 \text{ to } 3,000,000 \text{ km}^2$



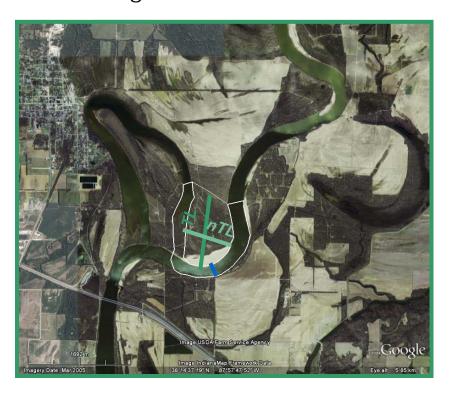


Methodology: Wabash River Example



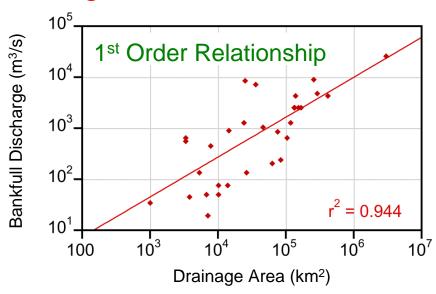


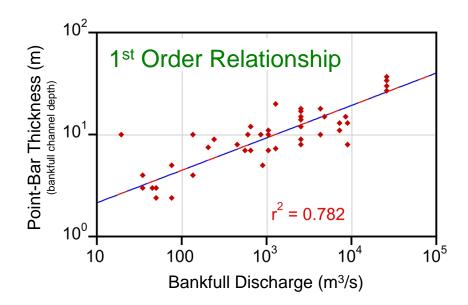
- Thickness: 9 m
- Point bar areal extent: 0.8 km²
- Abandoned Channel Width: 190 m
- Translation Length: 1450 m
- Non-translation length: 650 m
- Drainage Basin: 75,700 km²

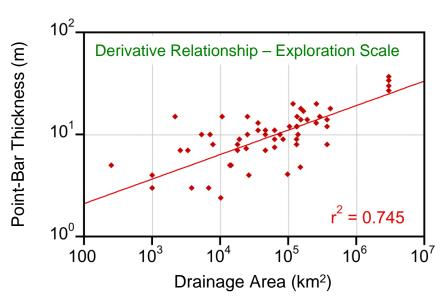


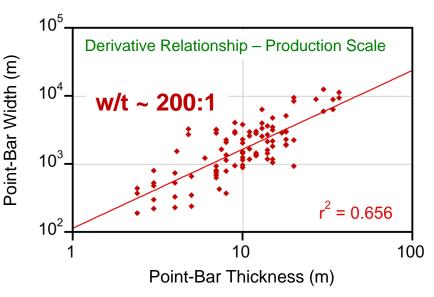


Scaling of Channel Belts



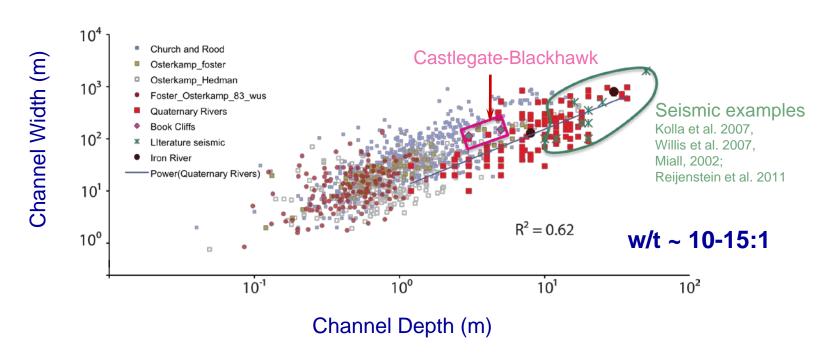


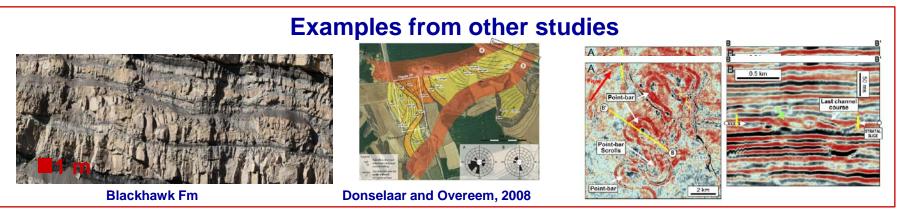






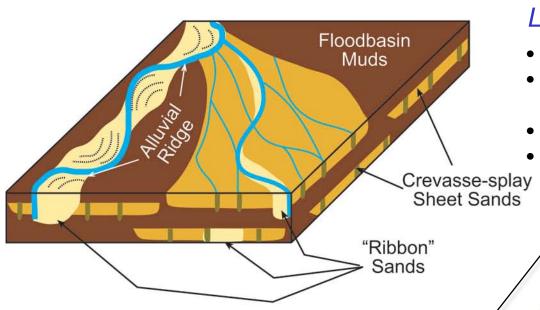
Heterolithic Abandoned Channel Fills







Fundamental Contrasts in Channel-Belt Scales: Why??

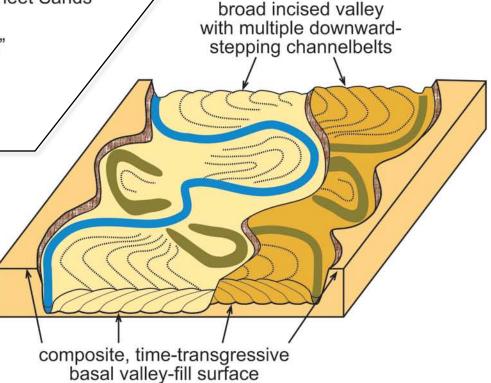


Low-Net End Member

- Highly heterolithic
- Ribbon sands and thin sheet sands encased in muds
- Dominated by avulsion processes
- Highly aggradational stacking

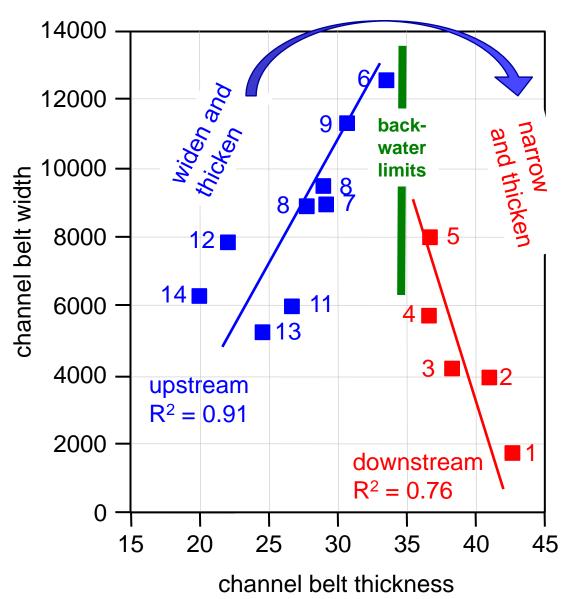
High-Net End Member

- Amalgamated sand bodies
- Channel fills and channel-belt margins are dominant barriers to flow
- Dominated by channel-belt migration
- Fluvial processes confined to discrete valley
- Degradational or low aggradational stacking





Fundamental Contrasts in Channel-Belt Scales: Backwater Effects

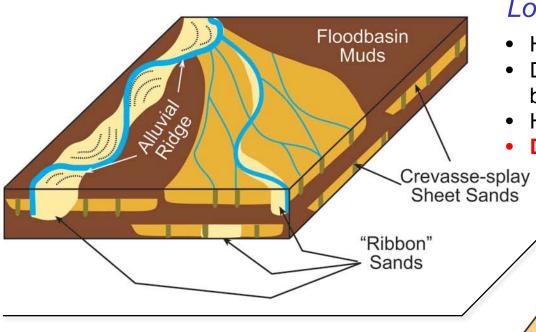








Fundamental Contrasts in Channel-Belt Scales: Why??

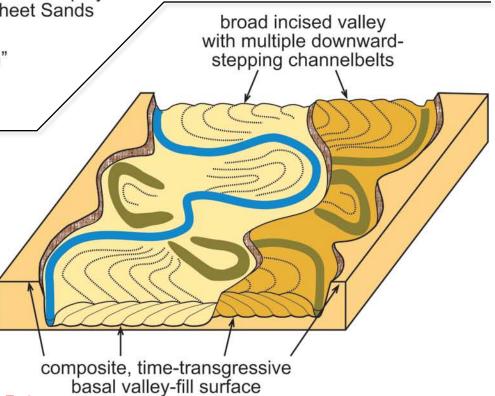


Low-Net End Member

- Highly heterolithic
- Dominated by avulsion processes within backwater length
- Highly aggradational stacking
- Distributary channel belts w/t = 20-50:1

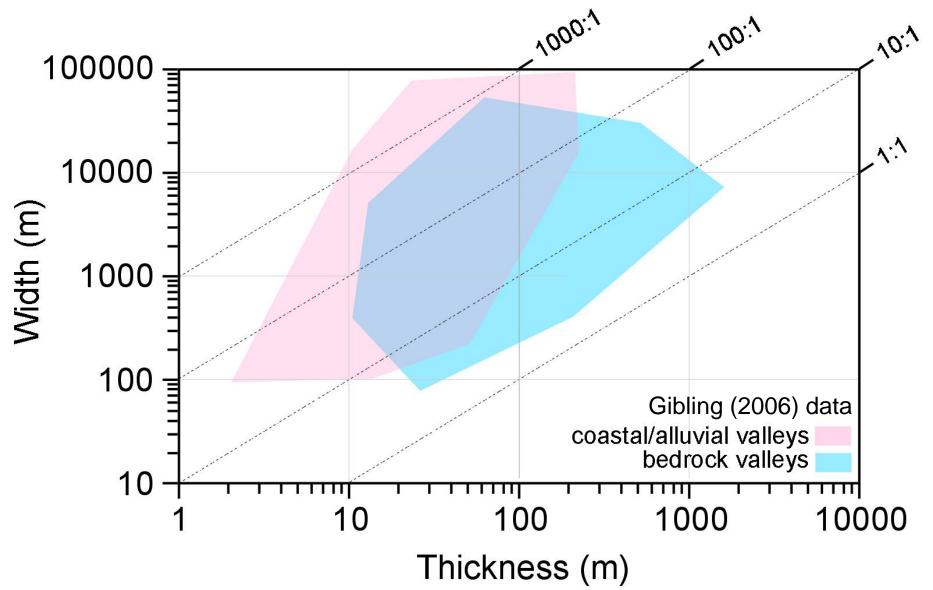
High-Net End Member

- · Amalgamated sand bodies
- Channel fills and channel-belt margins are dominant barriers to flow
- Dominated by channel-belt migration
- Fluvial processes confined to discrete valley
- Degradational or low aggradational stacking
- Channel belts w/t = 200:1, channel fills 15:1



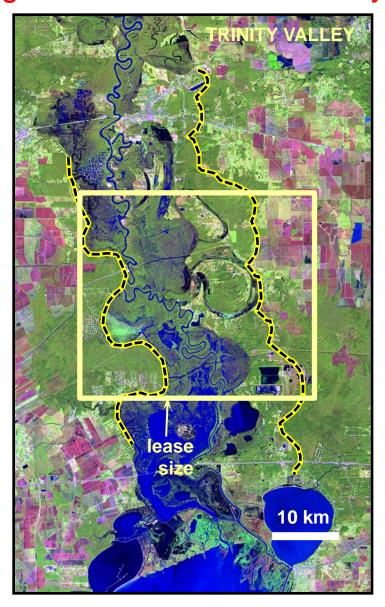


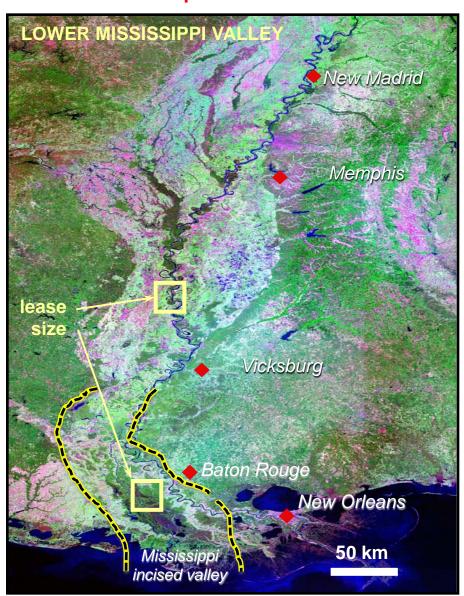
Scales of Incised Valleys - Published Ancient





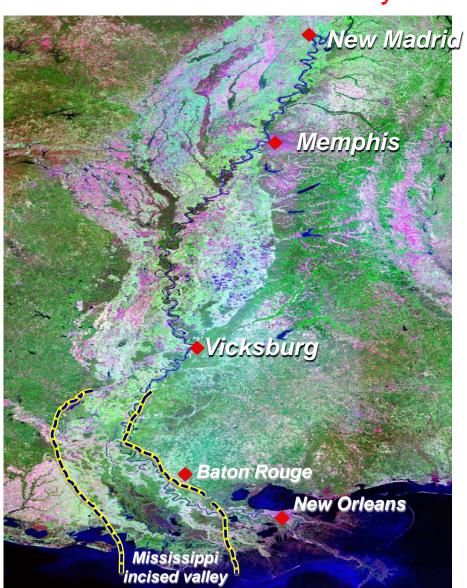
Significance of Incised-Valley Scales to Interpretation

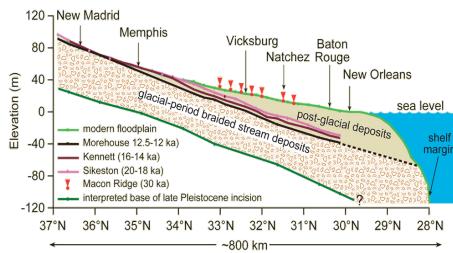


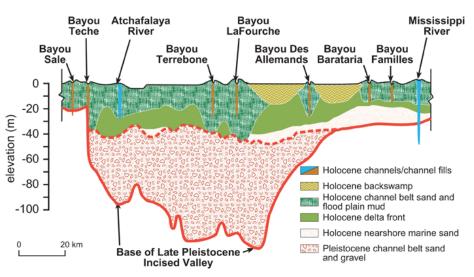




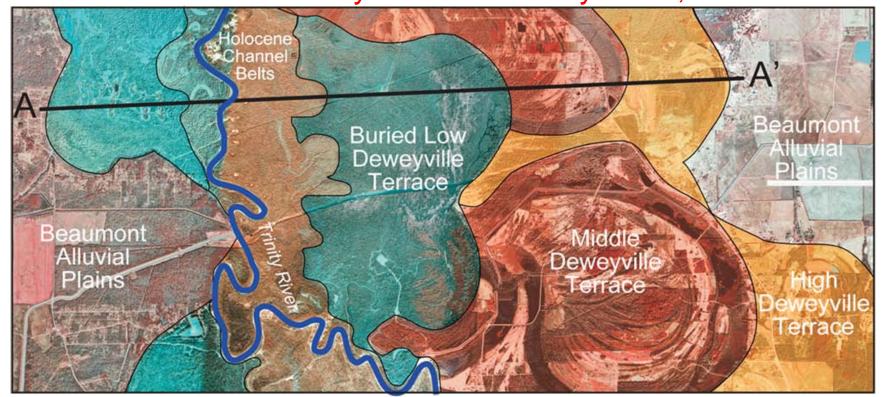
Coastal-Plain Incised Valleys - Lower Mississippi River

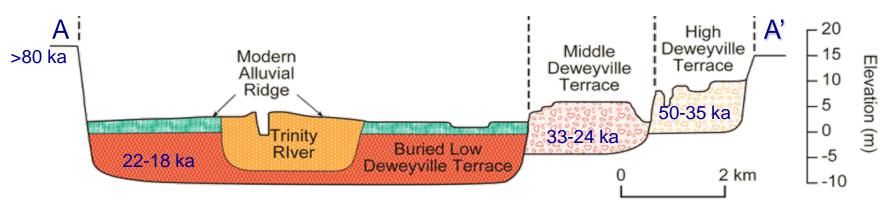






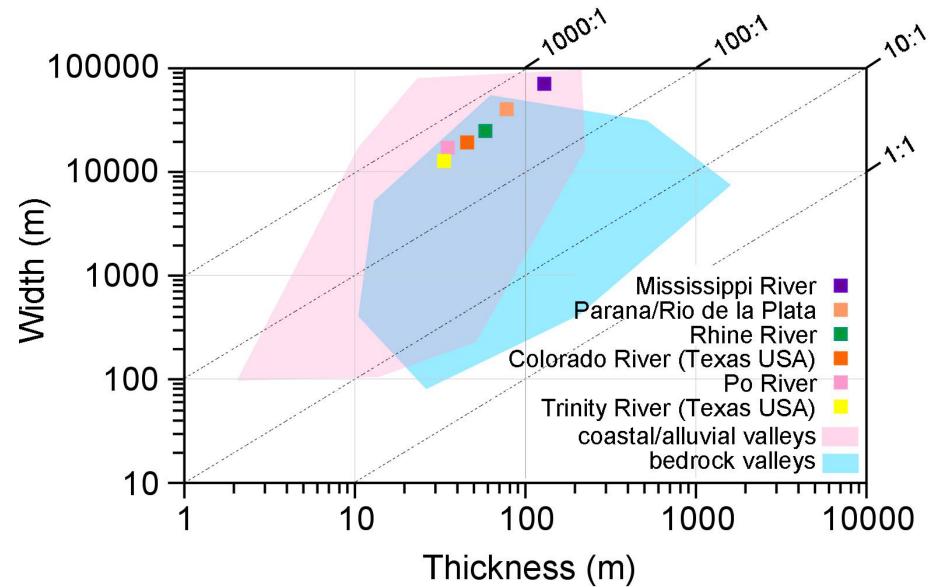
Coastal-Plain Incised Valleys - Lower Trinity River, Texas





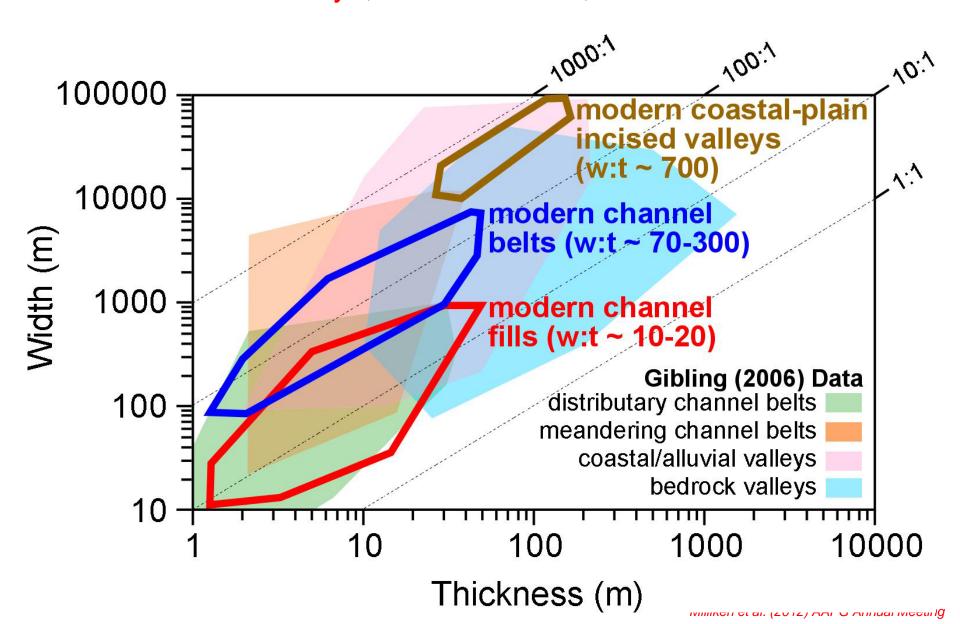


Scales of Incised Valleys – Published Ancient vs. Late Quaternary





Scales of Incised Valleys, Channel Belts, and Channel FIlls





Summary

Fluvial stratigraphic elements exhibit scaling relations that are consistent across a range of river system scales

- Channel-belt sand bodies range from ~100-300:1, with a mean of ~200
- Distributary channel-belts range from ~20-50:1
- Abandoned channel-fills range from 10-30:1
- Each element occupies the same thickness domain, because of scaling to bankfull discharge, but differ in width domains due to lateral migration
- Coastal-plain incised valleys range from 500-800:1

Scaling relations defined from modern systems commonly differ from scales interpreted in the stratigraphic record

- Scaling relationships from modern systems are observations, not interpretations, and tied to process-based understanding of system parameters (discharge, sediment flux)
- Can be used as an additional set of recognition criteria to guide, crosscheck, or calibrate interpretations in outcrop or subsurface data