

Sedimentary Structure Distribution and Modification on the Continental Shelf: Relative Roles of River Input, Sediment Transport and Oceanographic Setting*

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Abstract

Studies of sedimentary structures in modern shallow marine environments influenced by rivers have been conducted in active and passive margin settings from numerous areas worldwide (e.g., shelves adjacent to the Amazon, Mississippi, Eel, Waipaoa Rivers). Despite the large differences in scale between such systems, the offshore progression of sedimentary structures has many similarities. Factors indicated in the control on fine-scale sedimentary structures and their post-depositional modification include episodicity of river inputs, sediment transport mode, water depth and wave base, biological activity and the sediment accumulation rate. Many previous studies have suggested sediment accumulation rate as a dominant control on the preservation of primary physical structures on the continental shelf. However, results from recent studies suggest that, within the normal range of accumulation rates observed in shelf environments, other factors such as water depth, flood input history, and proximity to sediment source are the dominant controls on the occurrence and preservation of physically emplaced sedimentary structures. The timing and history of river flood and storm events is one factor in determining the distribution of event layers on the shelf, and their ultimate preservation. For example, concomitant river flooding and storm conditions favor the generation of wave- and current-supported gravity flows capable of broadcasting flood sediments across the shelf. Out of phase flooding would favor rapid deposition in nearshore and shallow shelf environments. The resultant flood layers have a higher preservation potential if they are buried quickly by deposition during subsequent large floods. Surface gravity waves cause physical reworking of the seabed in water depths shallower than wave base, obliterating original structures and winnowing the seabed of fines, but creating layers and laminations which may be similar (albeit coarser) than originally emplaced flood layers. In deeper waters, reworking of primary sedimentary structures arises from biological activity in the near-surface seabed, and the preservation of physically emplaced structures depends on the relative importance of biological mixing depth and intensity, and the sediment burial rate and history. These studies indicate that factors other than long-term accumulation rates primarily influence the formation and preservation of fine scale sedimentary structures on the continental shelf.

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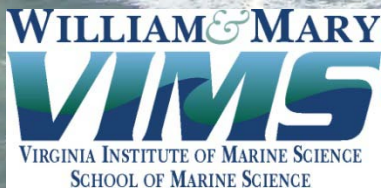
Traykovski, P., W.R. Geyer, J.D. Irish, and J.F. Lynch, 2000, The role of wave-induced density-driven fluid mud flows for cross-shelf transport on the Eel River continental shelf: *Continental Shelf Research*, v. 20, p. 2113-2140.

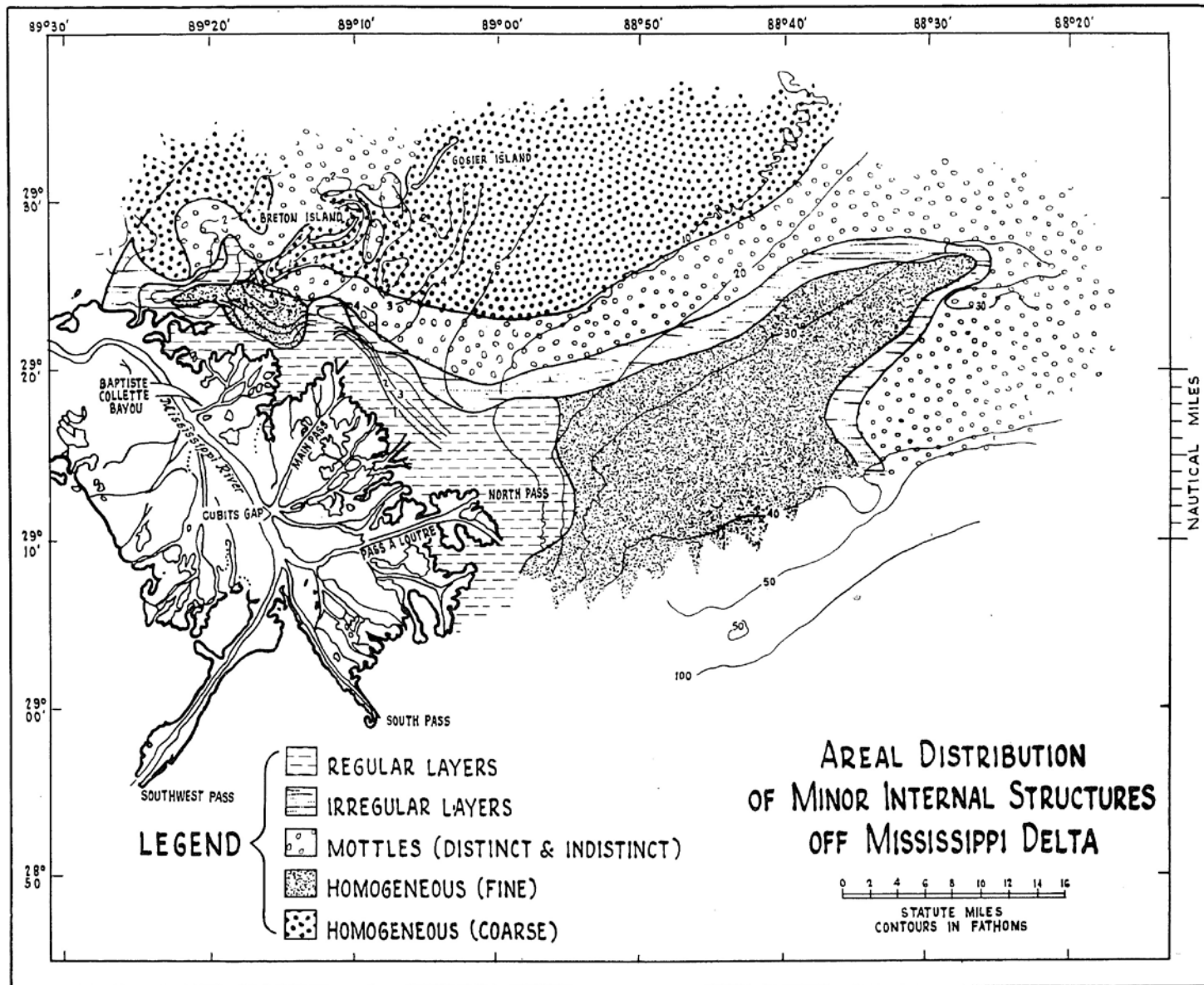
Wheatcroft, R.A., P.L. Wiberg, C.R. Alexander, S.J. Bentley, D.E. Drake, C.K. Harris, and A.S. Ogston, 2007, Post-depositional alteration and preservation of sedimentary strata: *International Association of Sedimentologists, Special Publication* v. 37, p. 101-155.

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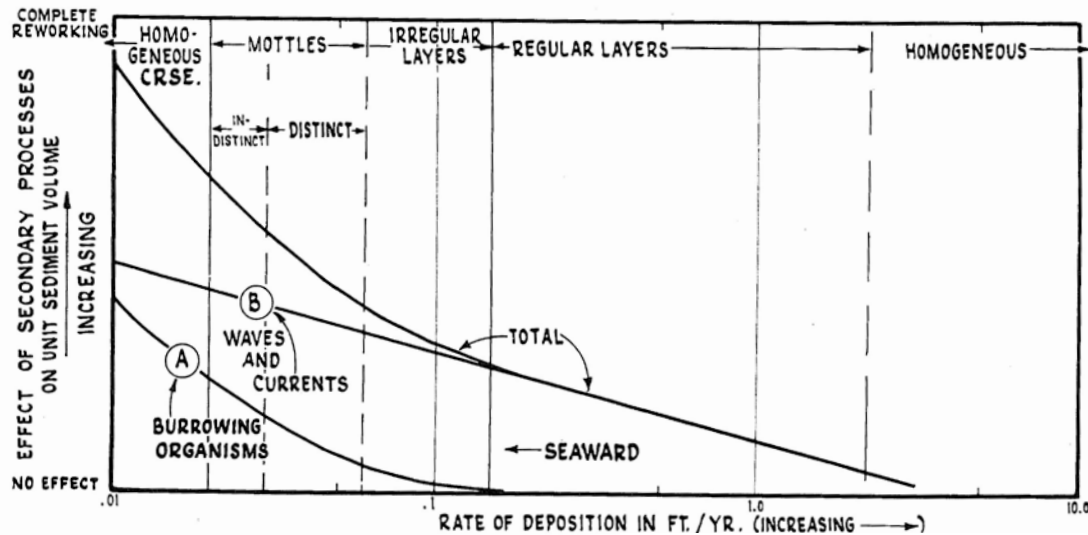
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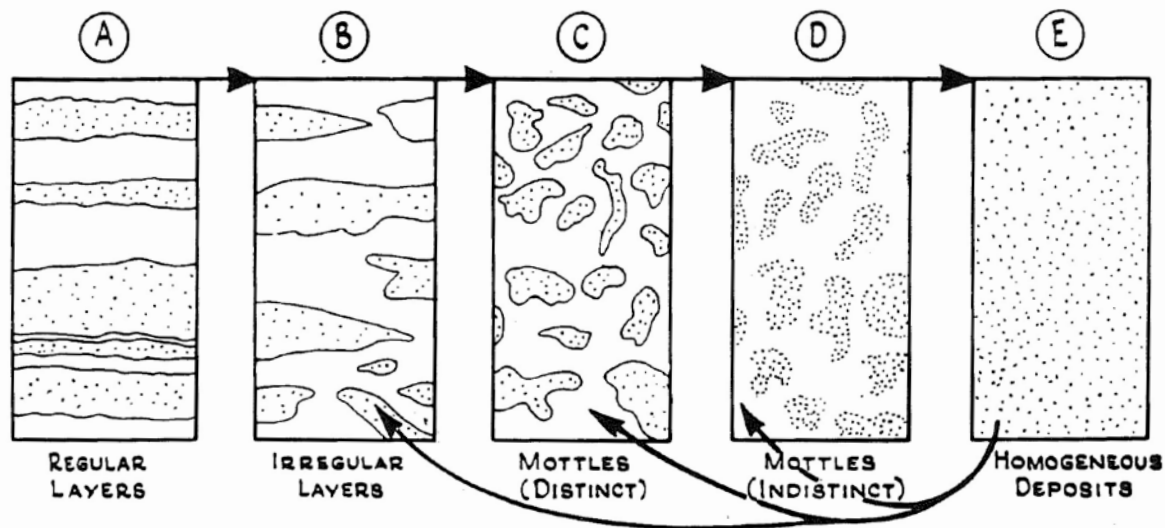




Sediment Accumulation Rate as the Master Variable?



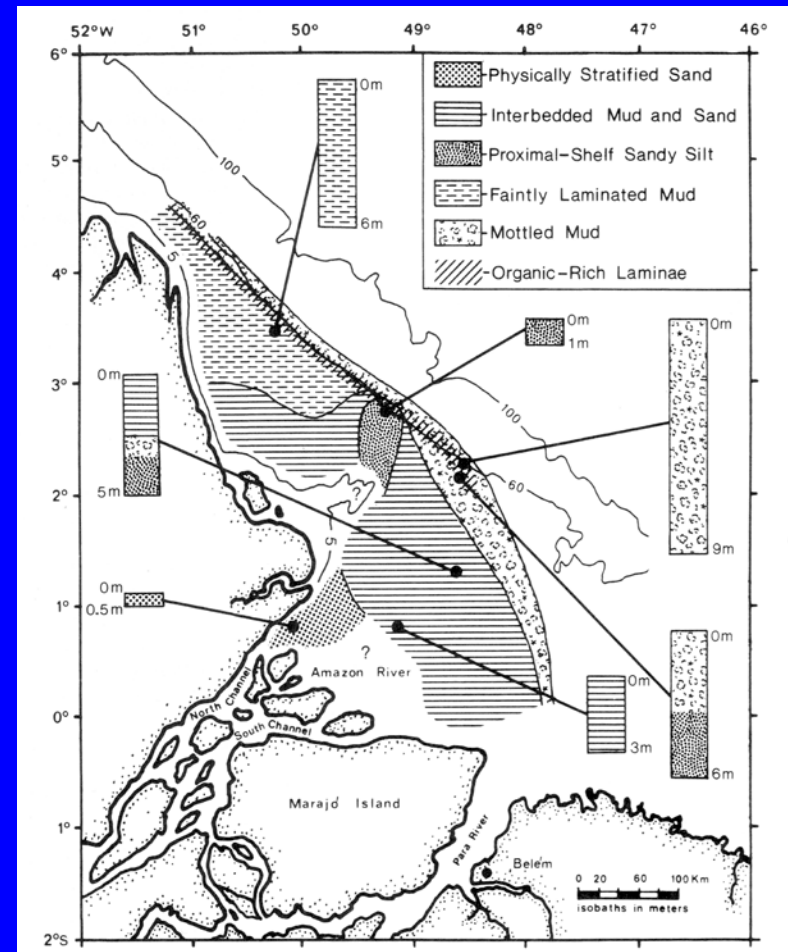
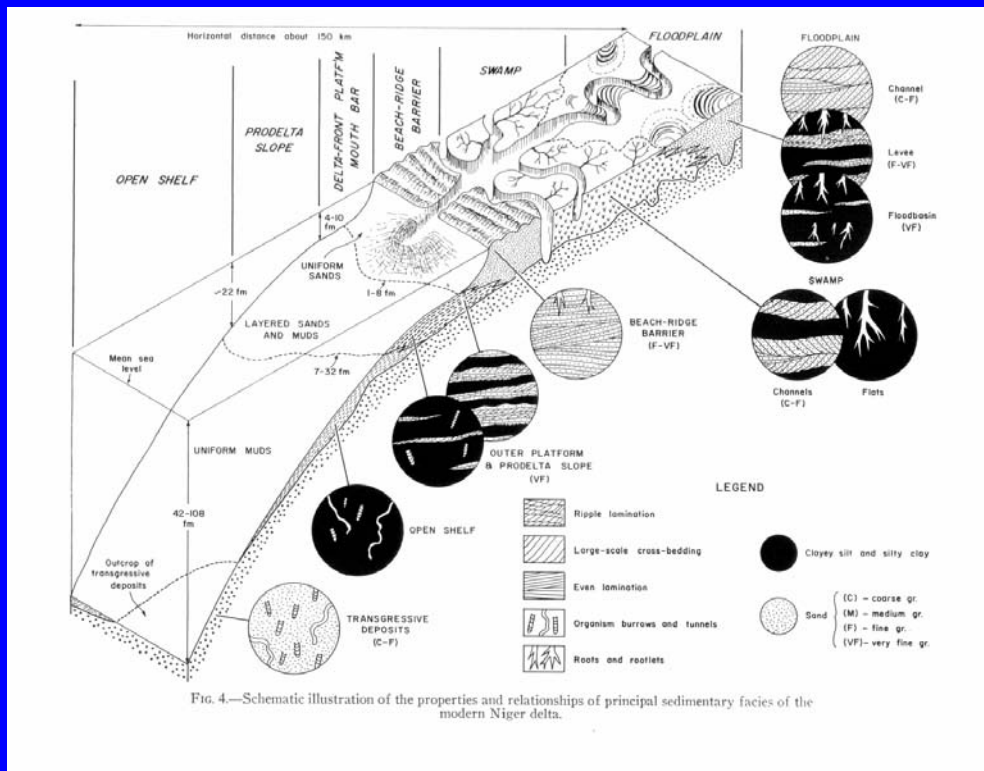
Moore and Scruton, 1957



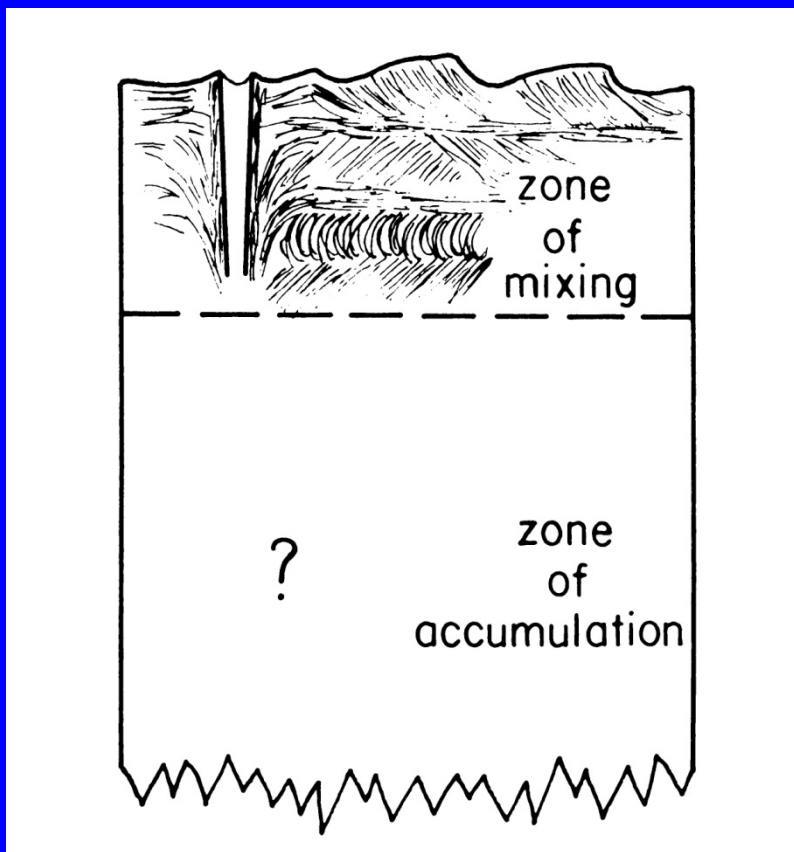
Examples from Other Passive Margins

Amazon Delta (Kuehl et al. 1986)

Niger Delta (Allen, 1965)



Introducing the Relative Roles of Mixing and Accumulation Rate



Biological Mixing

$$\begin{array}{l}
 D_b = \text{mixing coefficient (cm}^2/\text{yr)} \\
 A = \text{accumulation rate (cm/yr)}
 \end{array}
 \left. \vphantom{\begin{array}{l} D_b \\ A \end{array}} \right\} L_b = \text{biological mixing depth (cm)}$$

$$G = \frac{D_b / L_b}{A} = \frac{\text{volume of sediment processed per unit area of seabed per time}}{\text{net accumulation rate}}$$

Physical Mixing

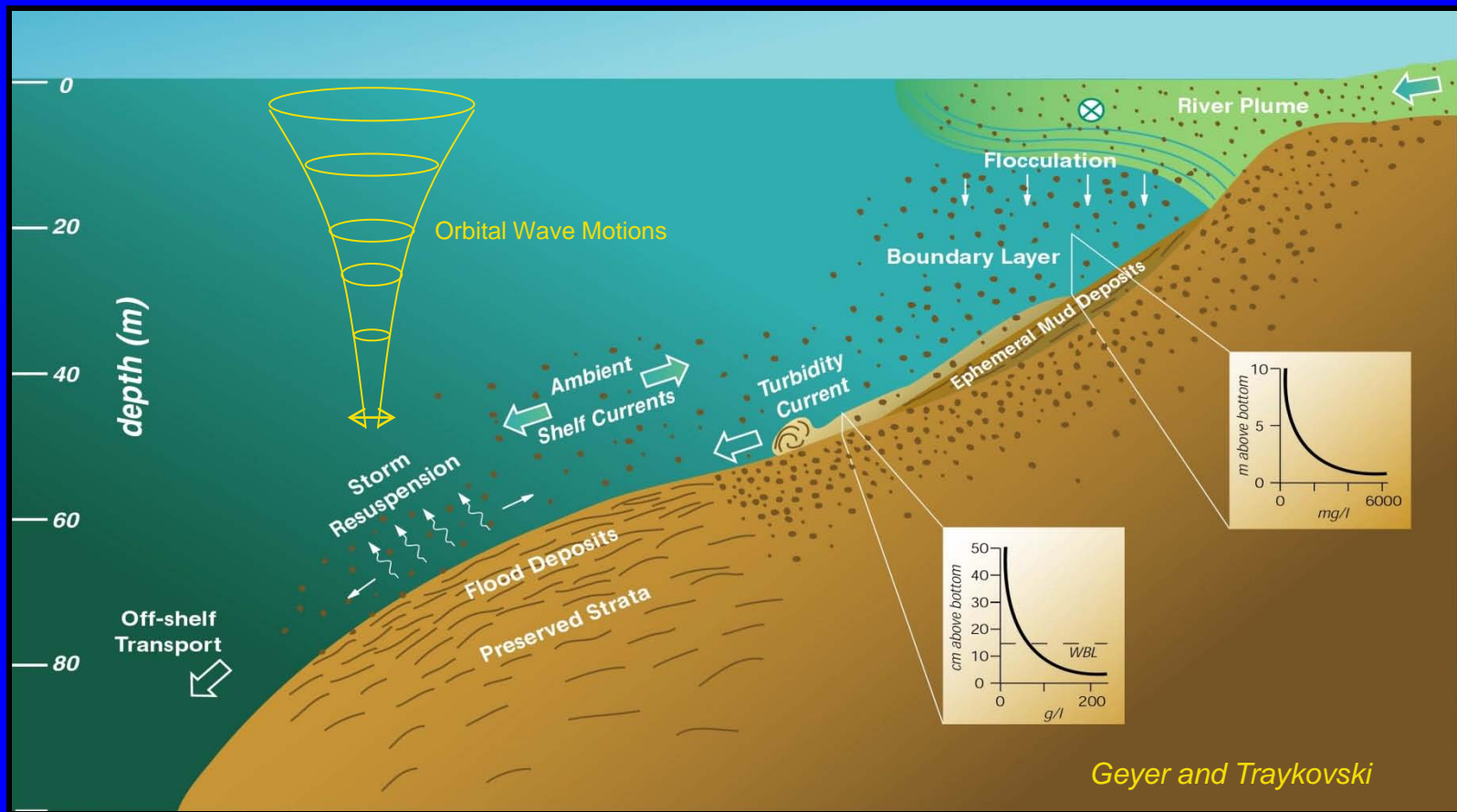
$$\begin{array}{l}
 f_p = \text{frequency of erosion (yr}^{-1}\text{)} \\
 A = \text{accumulation rate (cm/yr)}
 \end{array}
 \left. \vphantom{\begin{array}{l} f_p \\ A \end{array}} \right\} L_p = \text{physical erosion depth (cm)}$$

$$H = \frac{f_p L_p}{A} = \frac{\text{volume of sediment reworked per unit area of seabed per time}}{\text{net accumulation rate}}$$

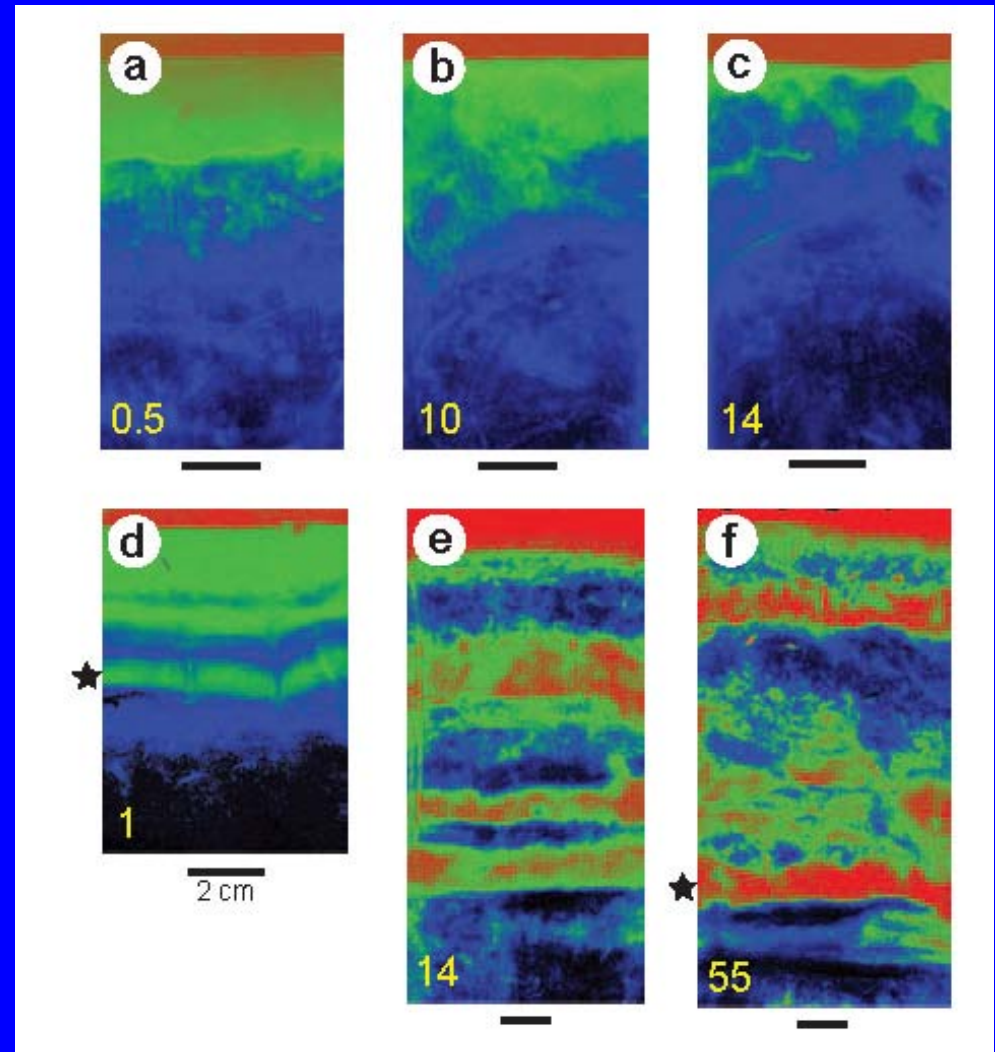
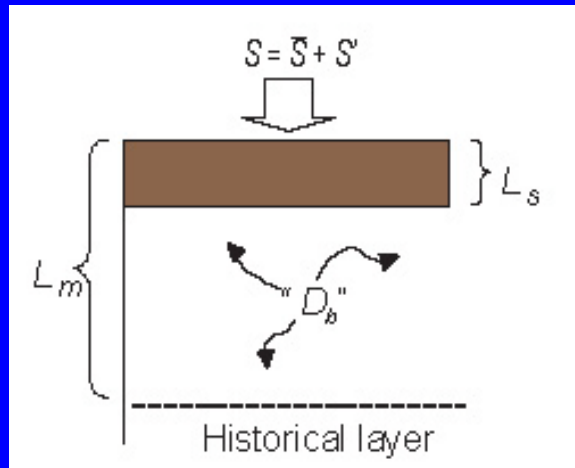
Nittrouer and Sternberg, 1981

Assessing Importance of Sediment Transport Mechanisms

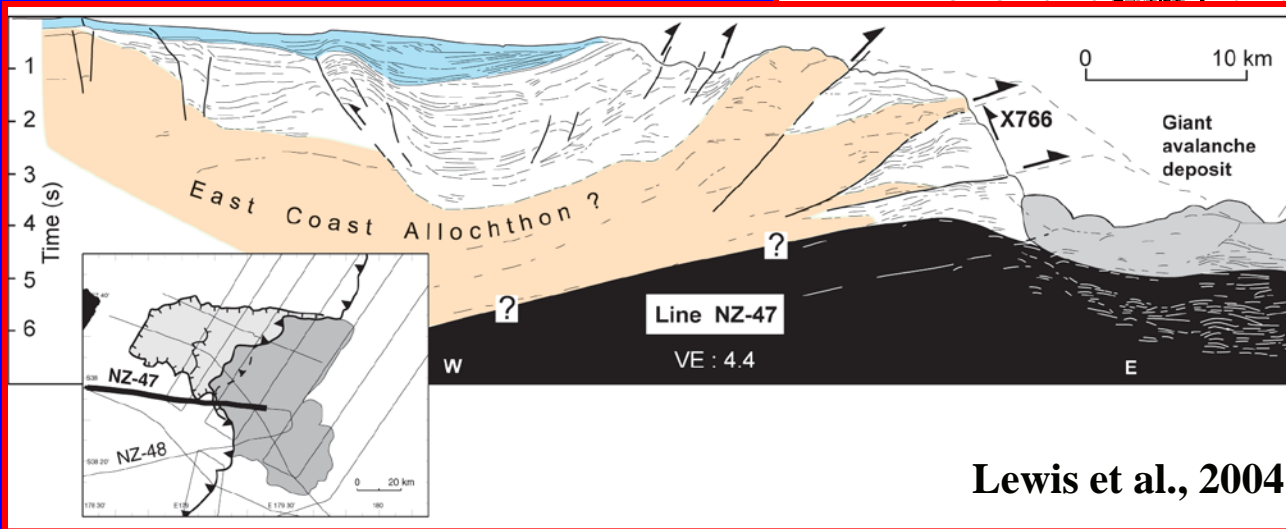
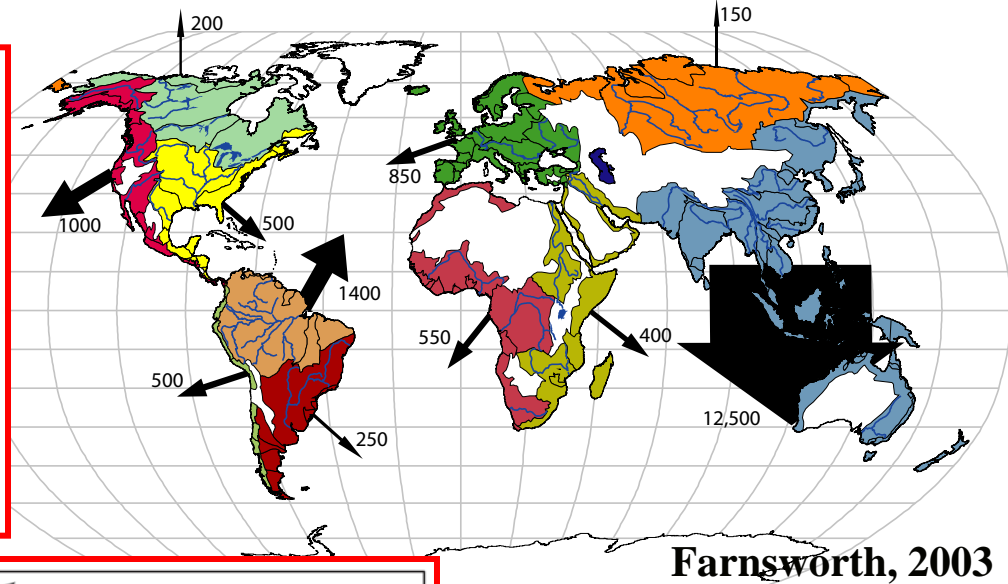
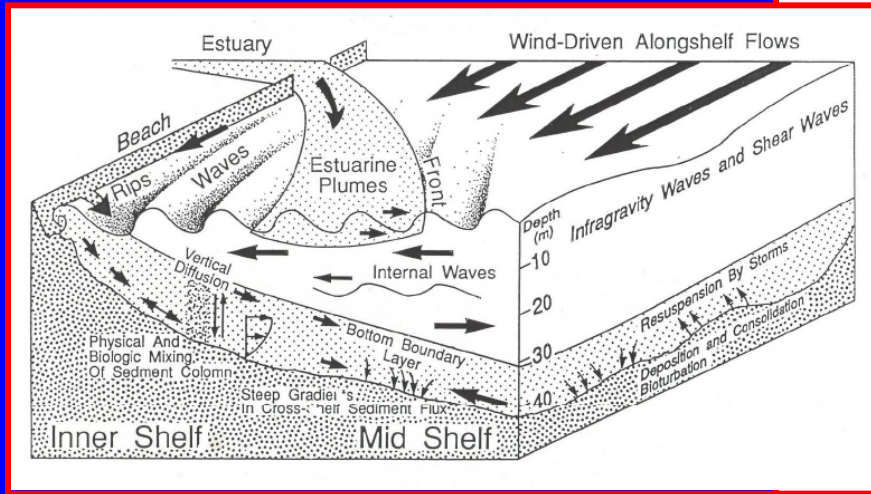
- sediment gravity flows
- dilute suspension



Episodic Burial as the Master Variable?



High Episodic Inputs Dominate Active Margins



Lewis et al., 2004

- high yields
- various transport pathways
- tectonic control

Insights from New Zealand East Coast: Waipaoa and Waiapu Rivers

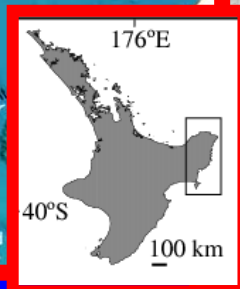
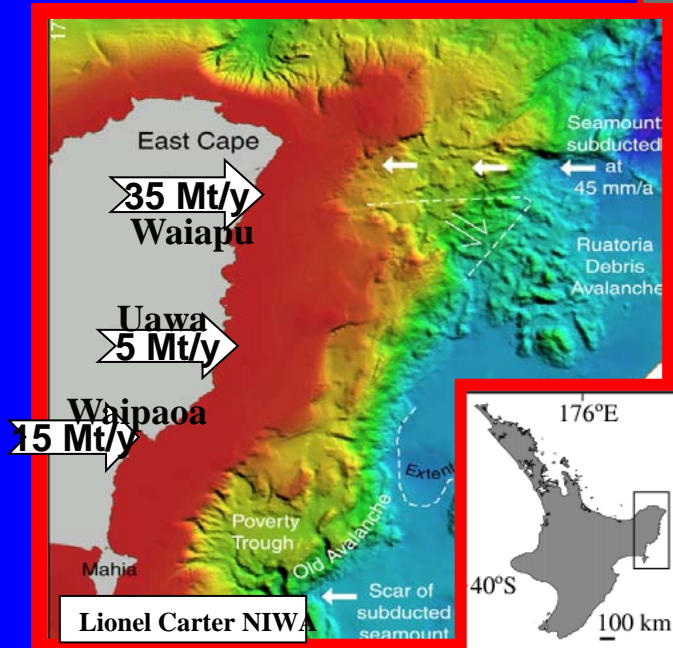
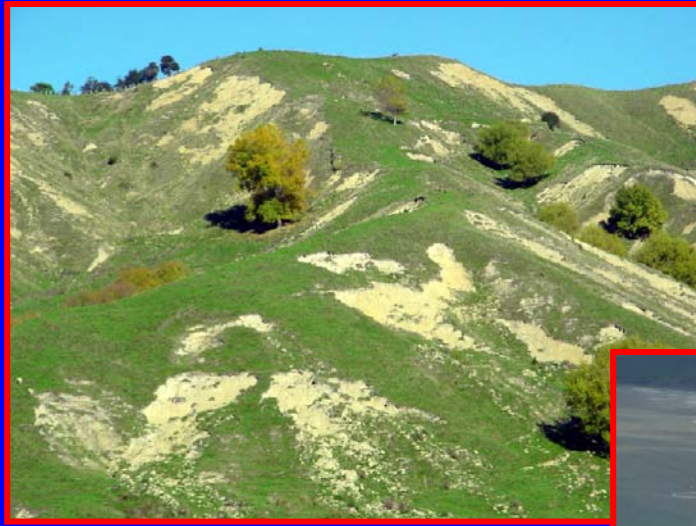
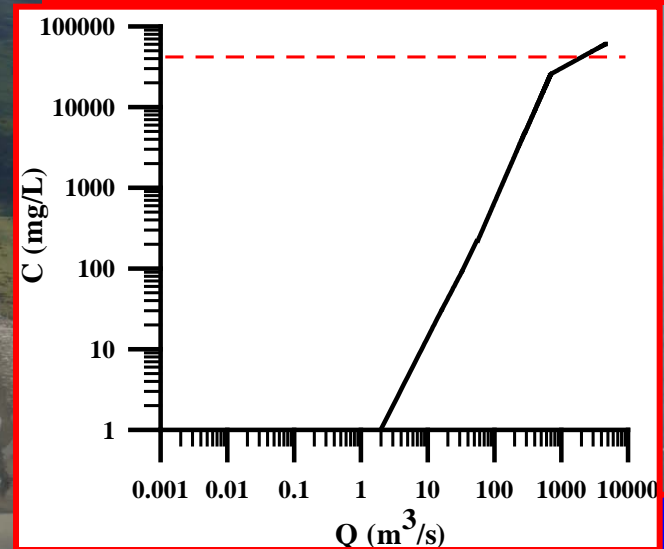
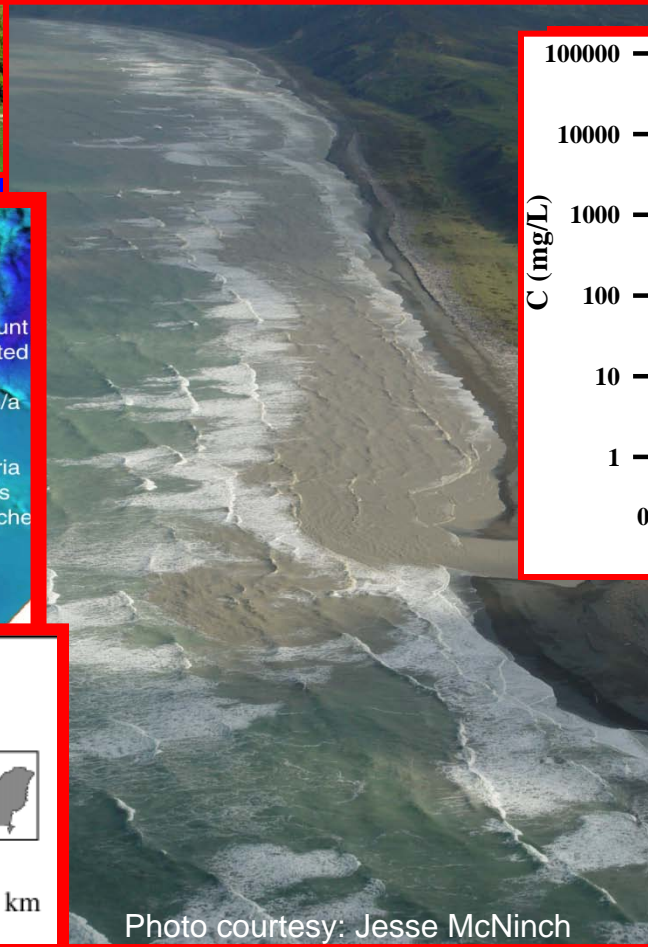
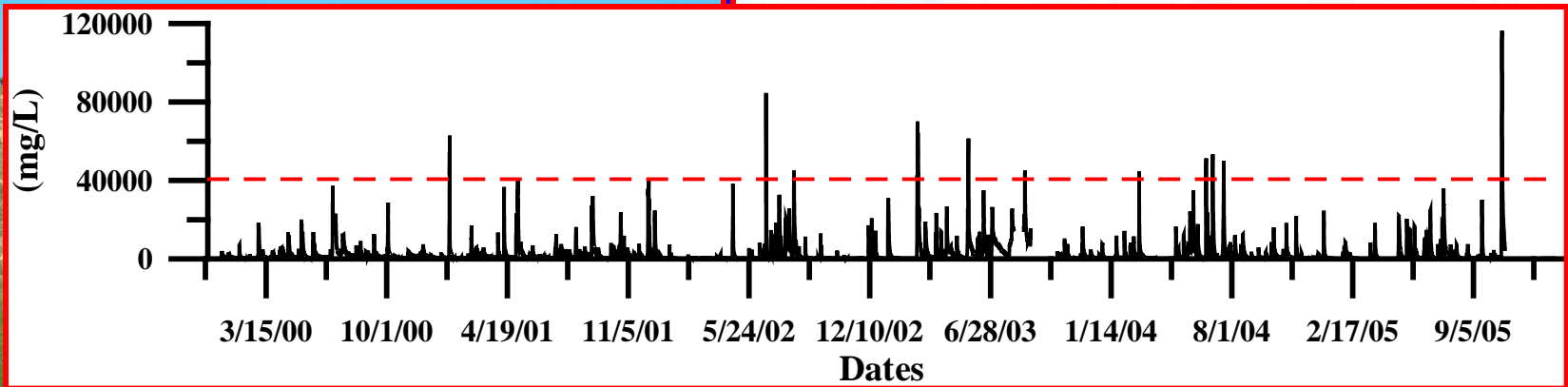


Photo courtesy: Jesse McNinch



Hicks et al. (2004)

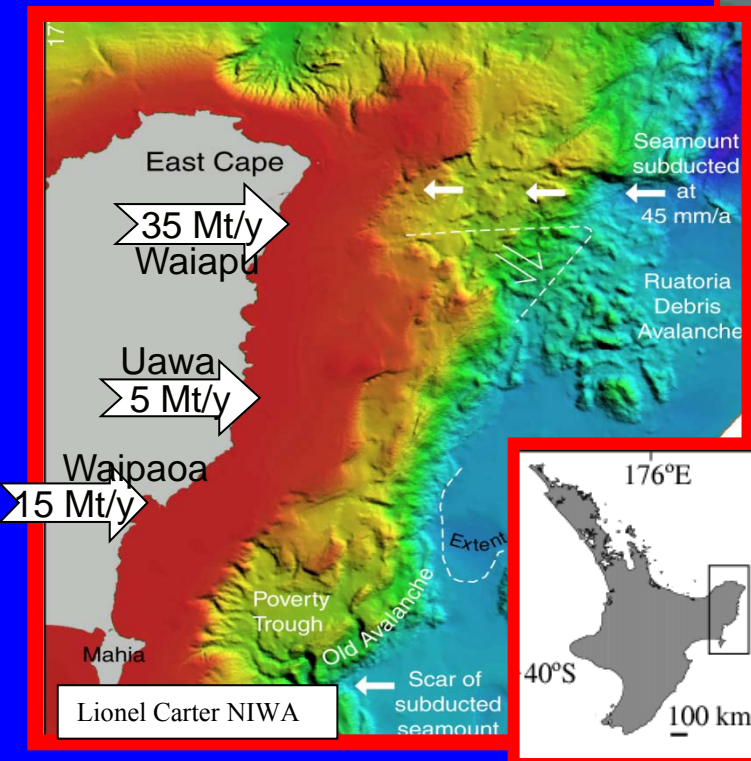
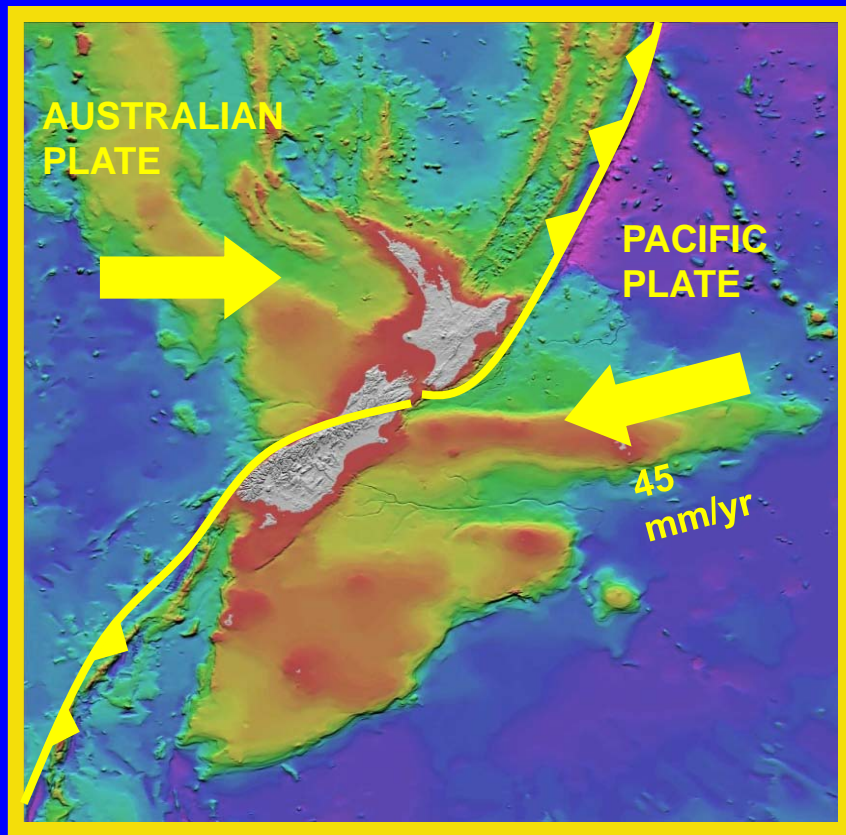


Photo courtesy: Jesse McNinch

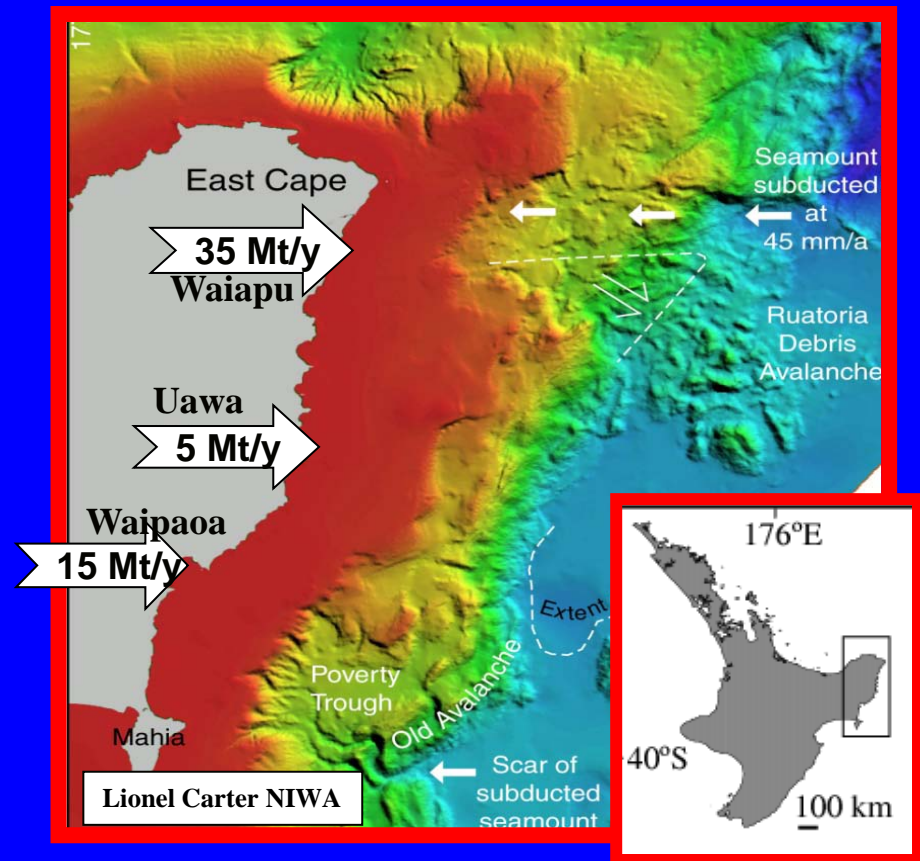
Glorious Mud

Study Area

Active Tectonics



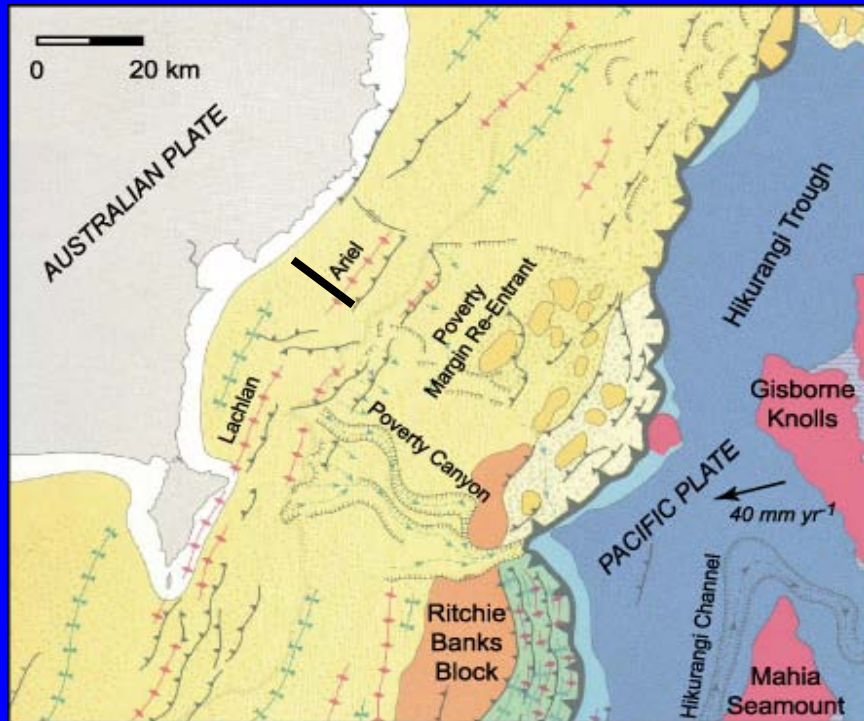
High Sediment Yield



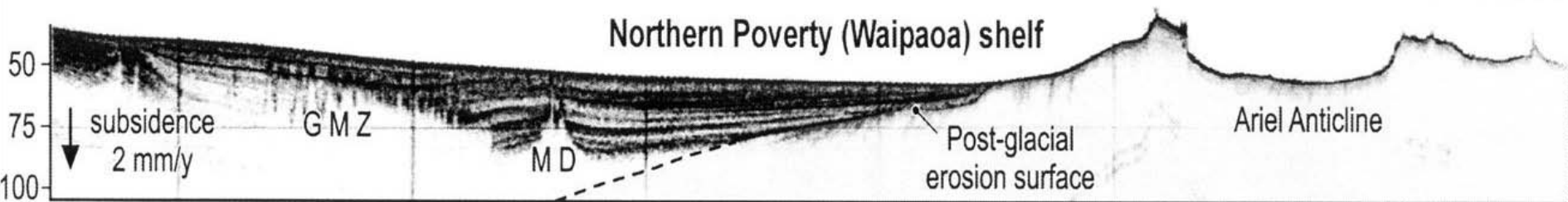
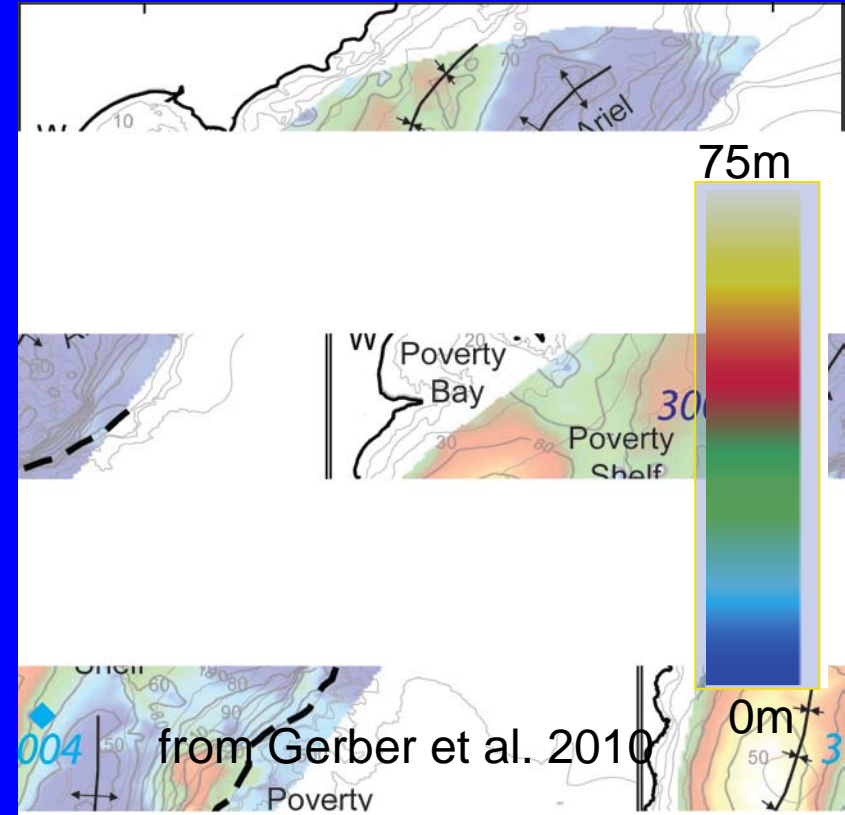
Notes by Presenter (for previous slide):

The catchment of the Waipaoa River originates in the axial ranges of eastern North Island and sediments are emptied into Poverty Bay via the coastal plains of Gisborne. Waipaoa can be classified as a small mountainous river. These rivers often have large sediment yields and are incredibly important components in delivering sediment to the ocean. Milliman and Meade (1983) estimated that 70% of the sediment reaching the ocean is derived from rivers draining southern Asia and islands in the Pacific and Indian Ocean of which most of these rivers are classified as small mountainous rivers.

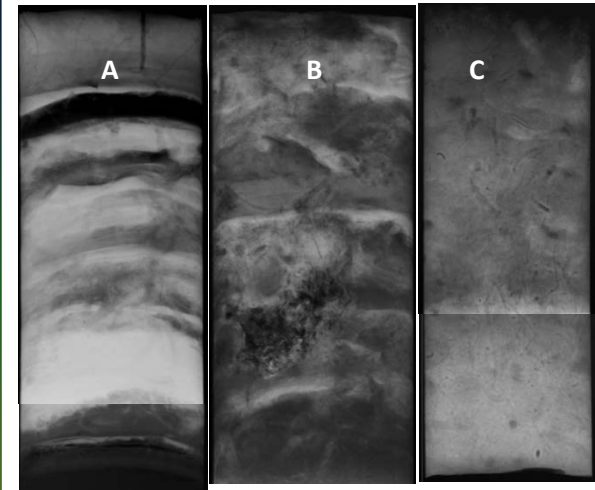
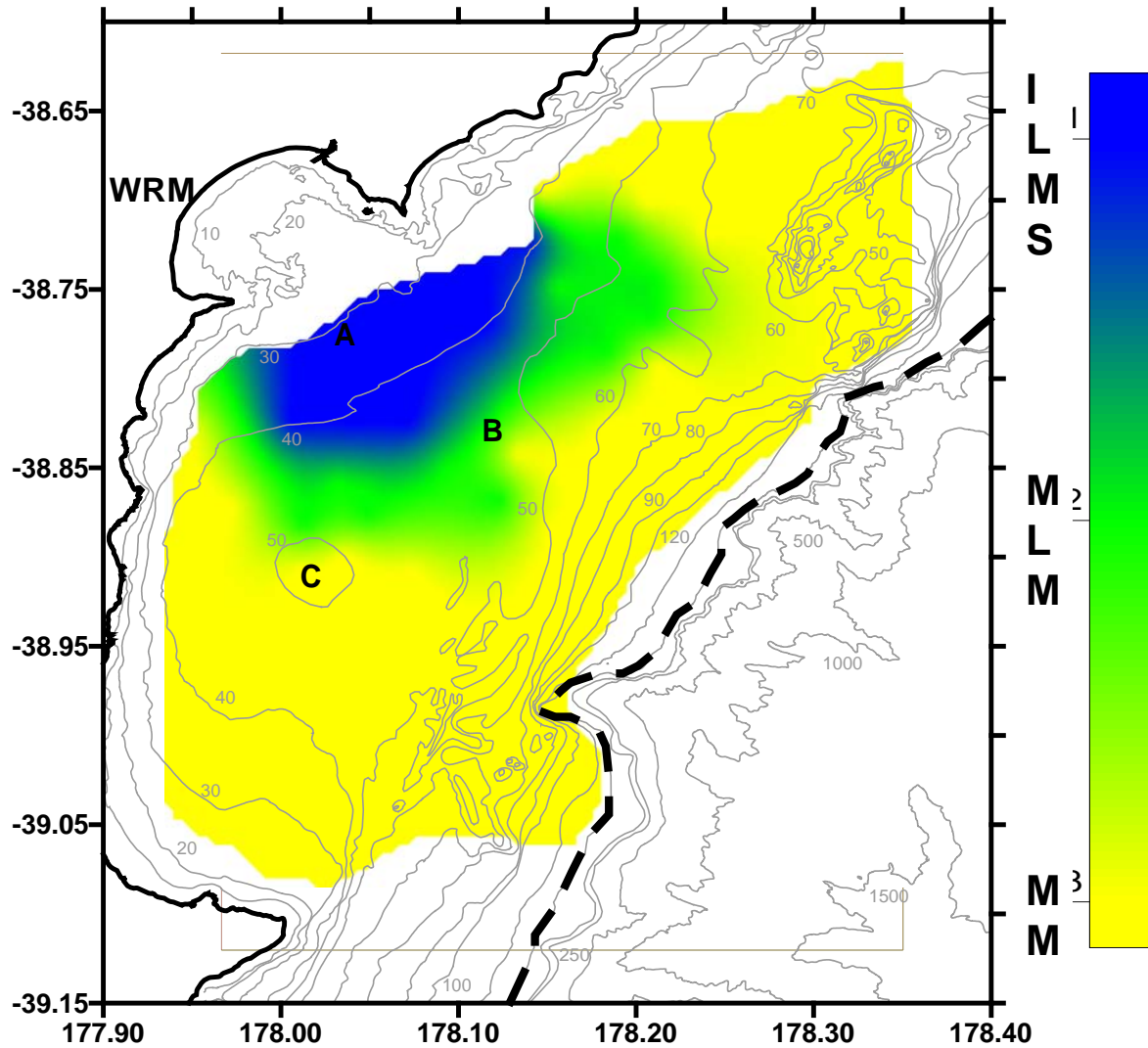
Waipaoa Shelf Setting



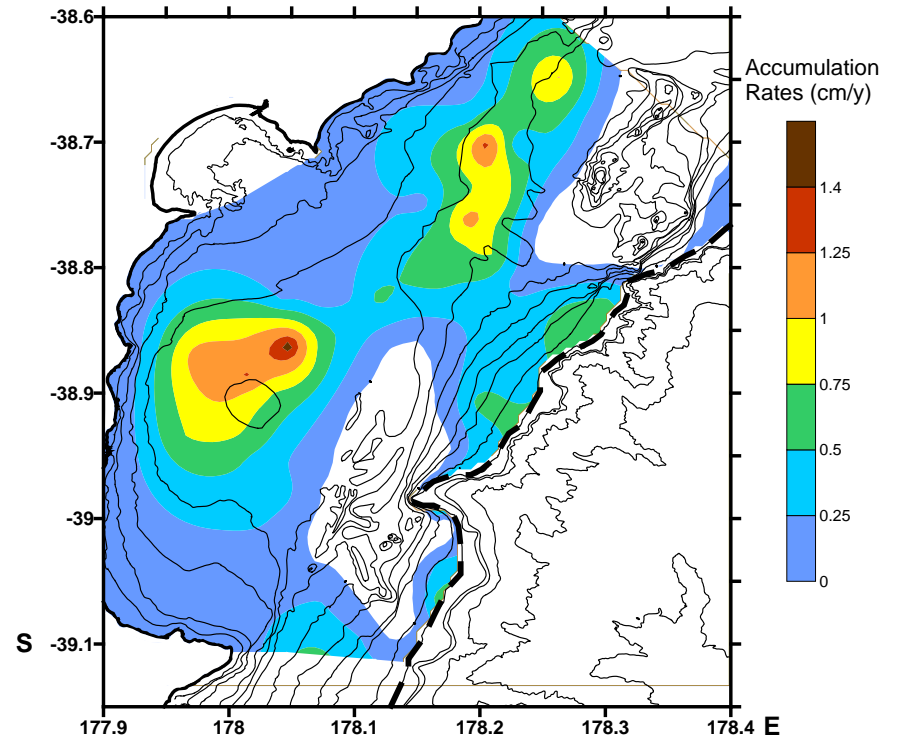
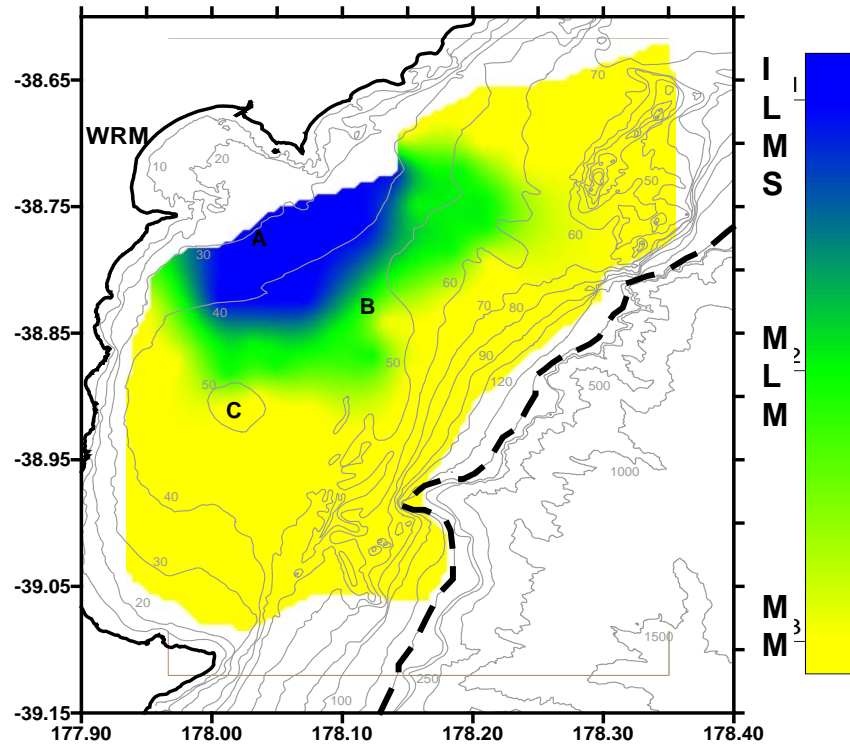
Foster and Carter, 1997



Distribution of Sedimentary Structures on Waipaoa Shelf

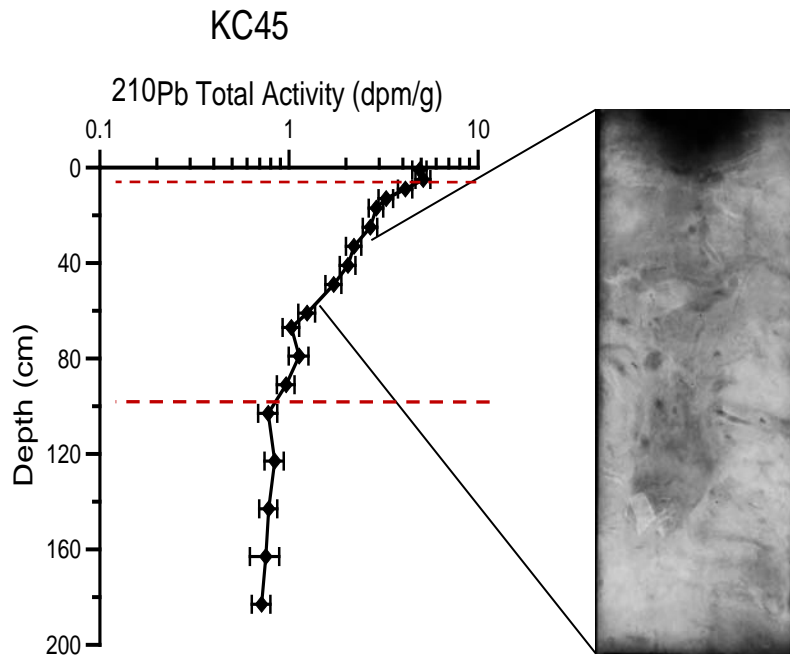


High Accumulation Rate \neq Preserved Structure!

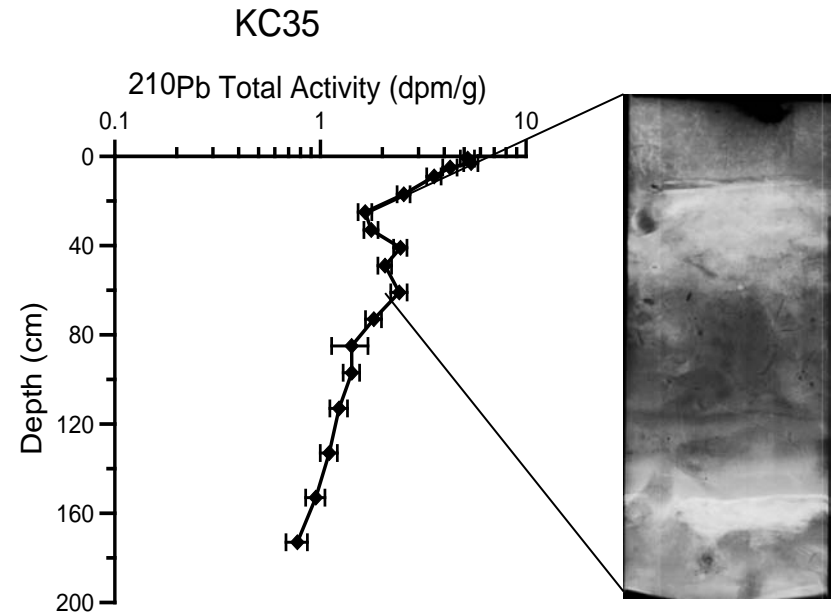


Preservation of Episodic Inputs

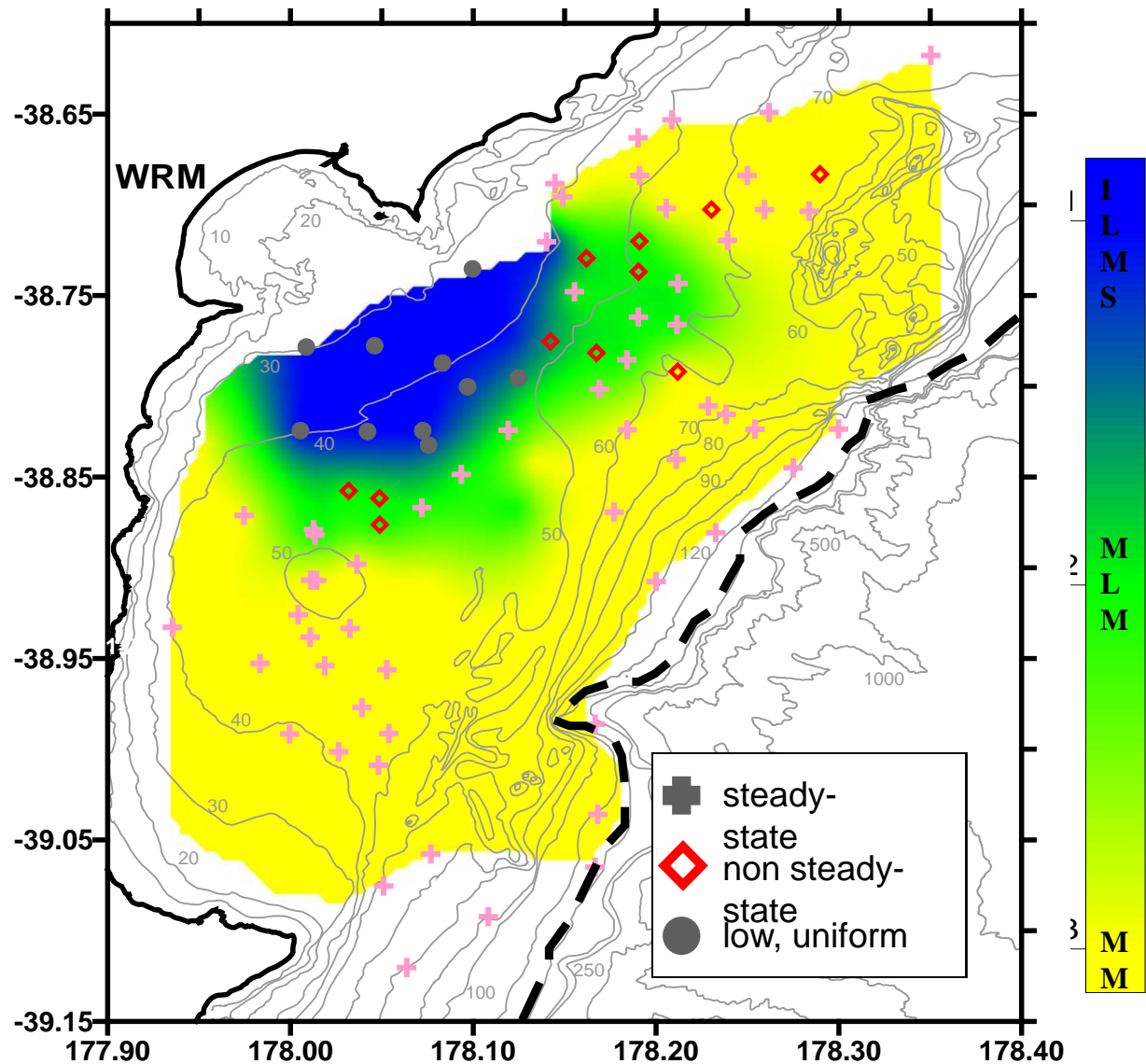
Steady State



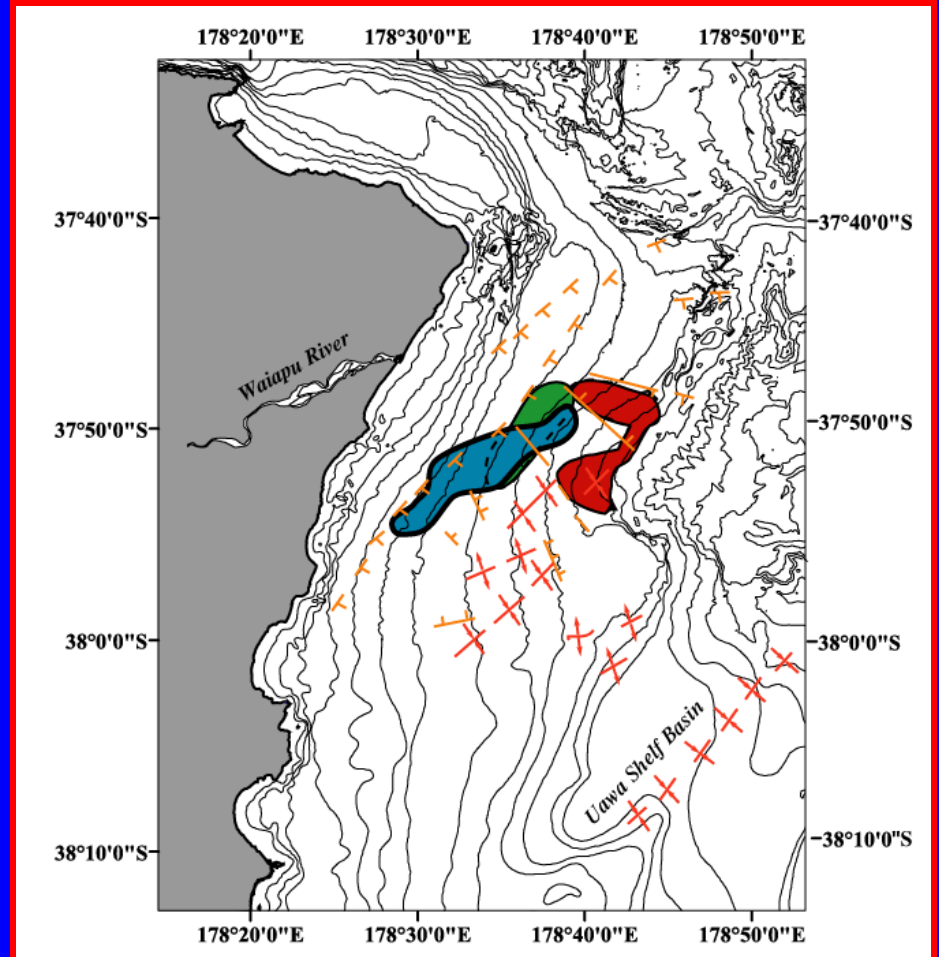
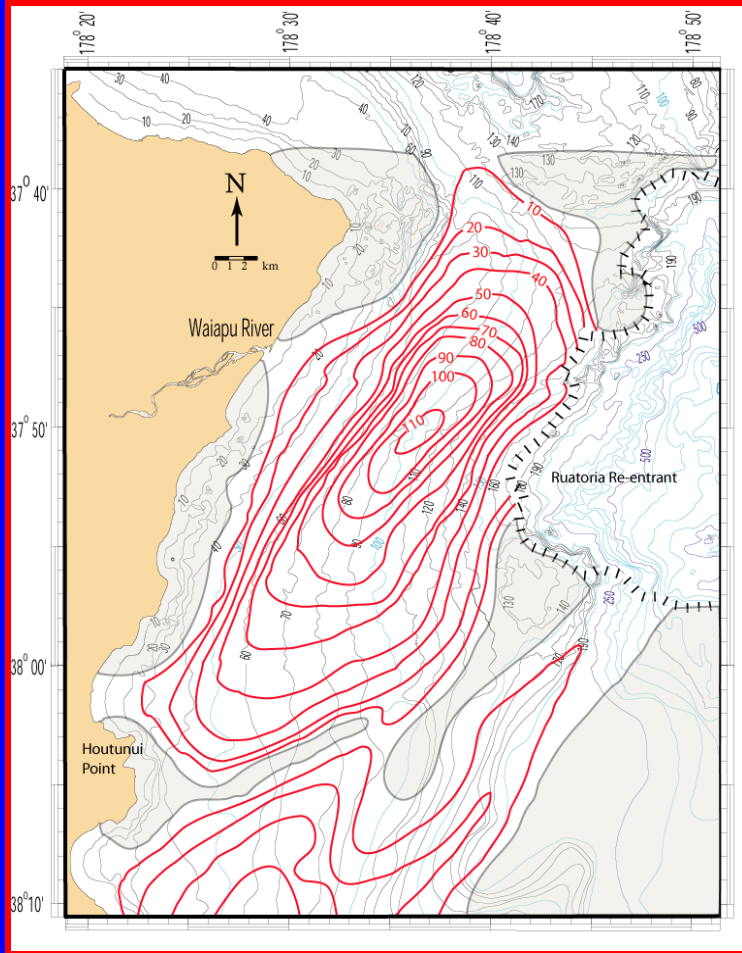
Non-Steady State



Episodic Inputs (Sediment Gravity Flows?) Preserved

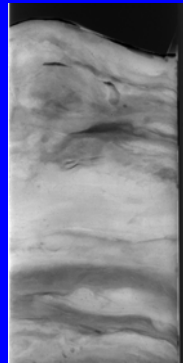
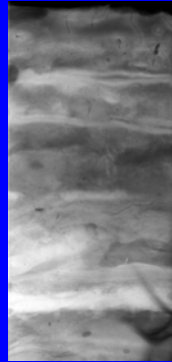


Waiapu Shelf Setting

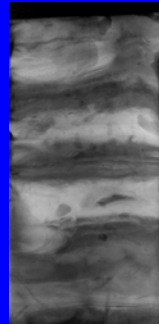
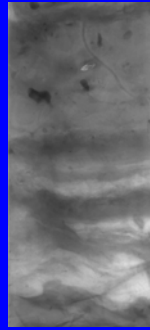


After Lewis et al., 2004

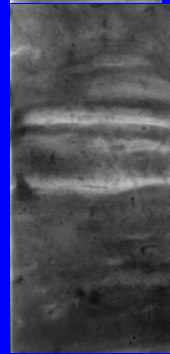
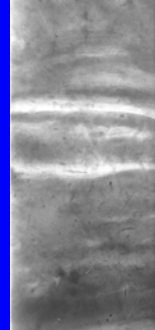
KC26



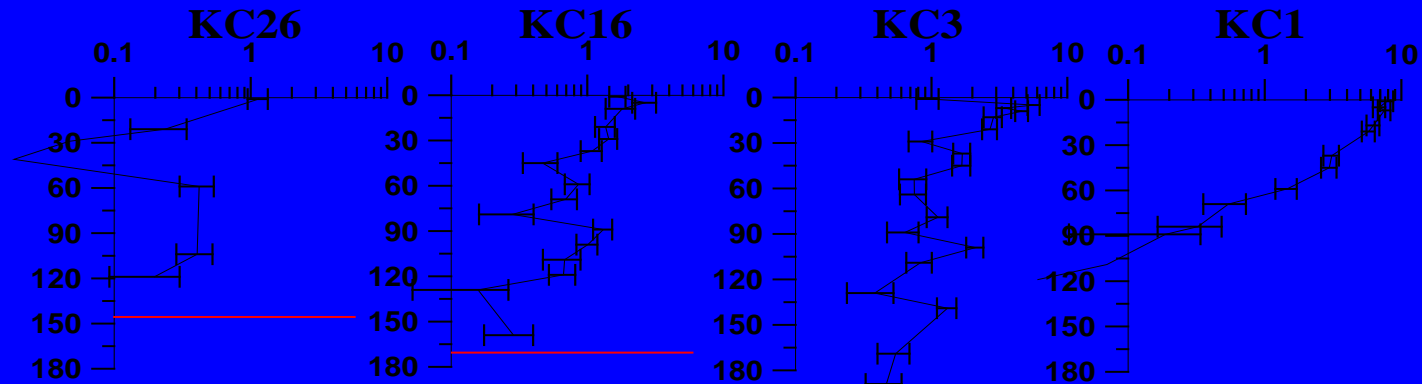
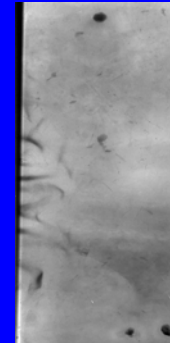
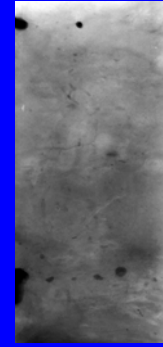
KC16



KC3

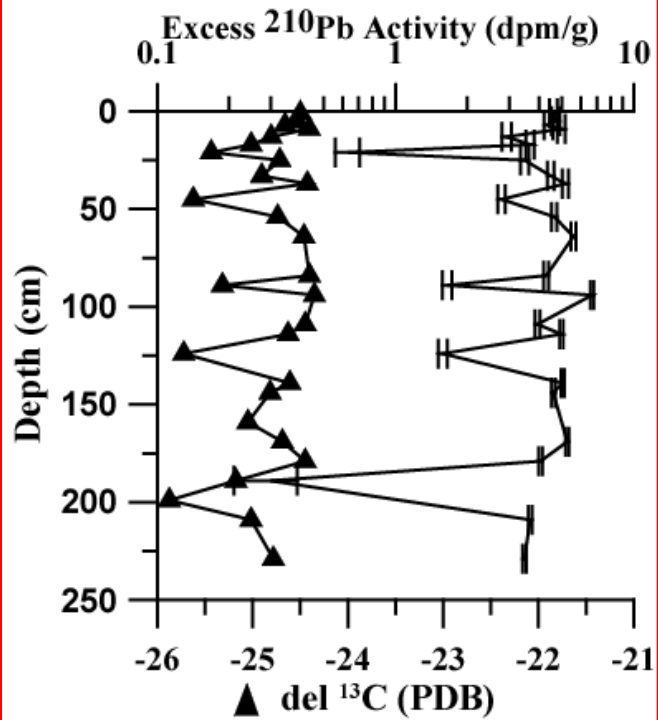


KC1

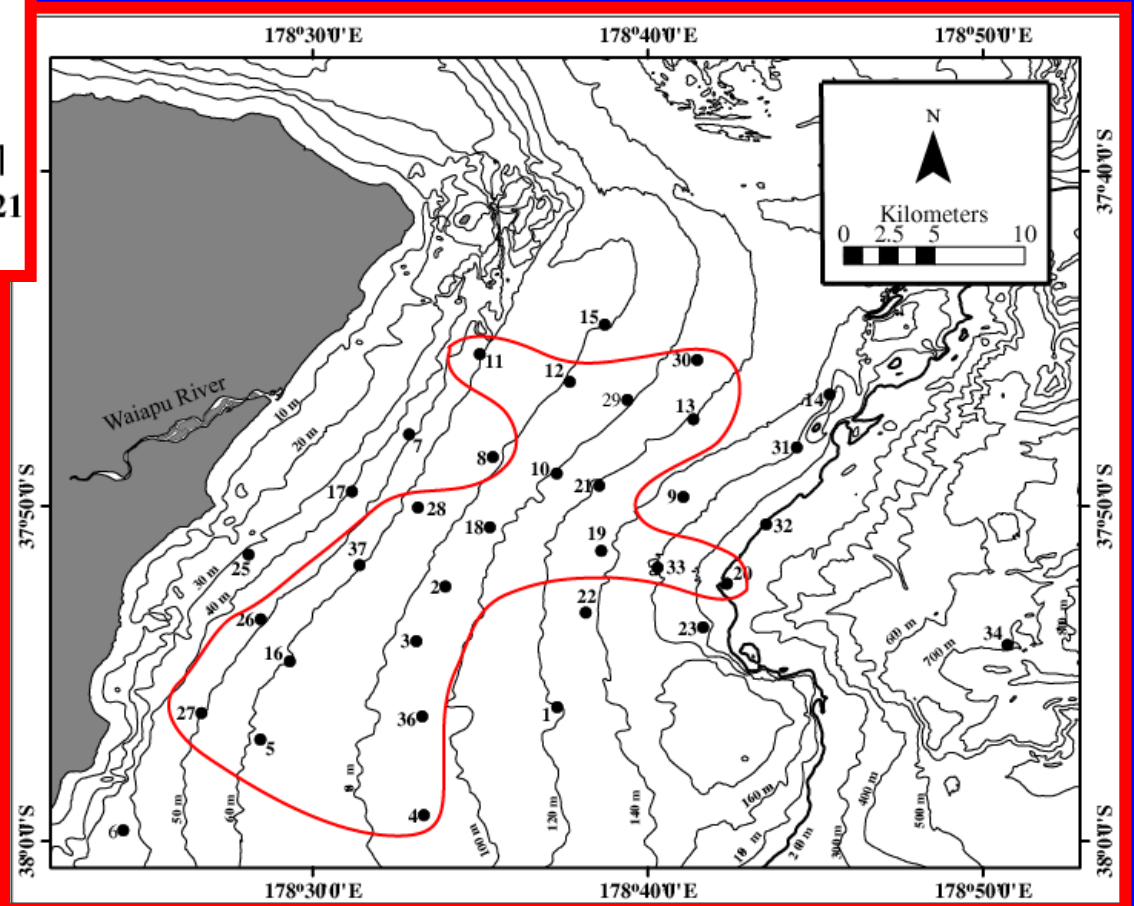
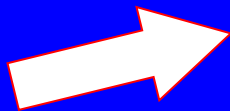


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Evidence of Multiple Sediment Transport Mechanisms: Wave/Current Supported Gravity Flows

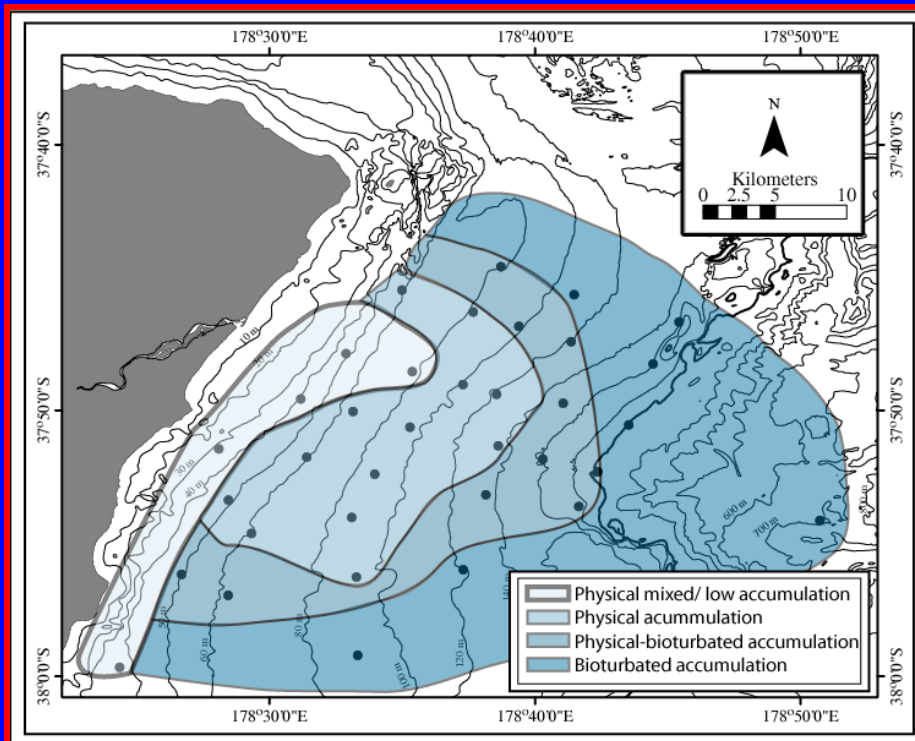


Non-steady state
excess ^{210}Pb
activity profiles

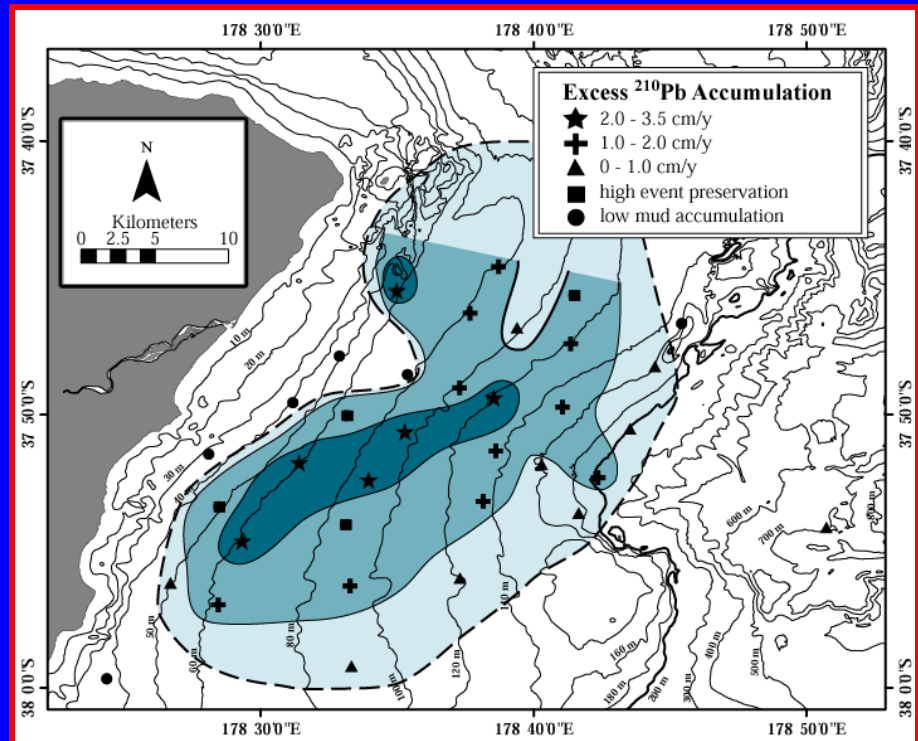


Physical Structures Well Correlated to Accumulation Rate and Episodic Input

Structure Distribution

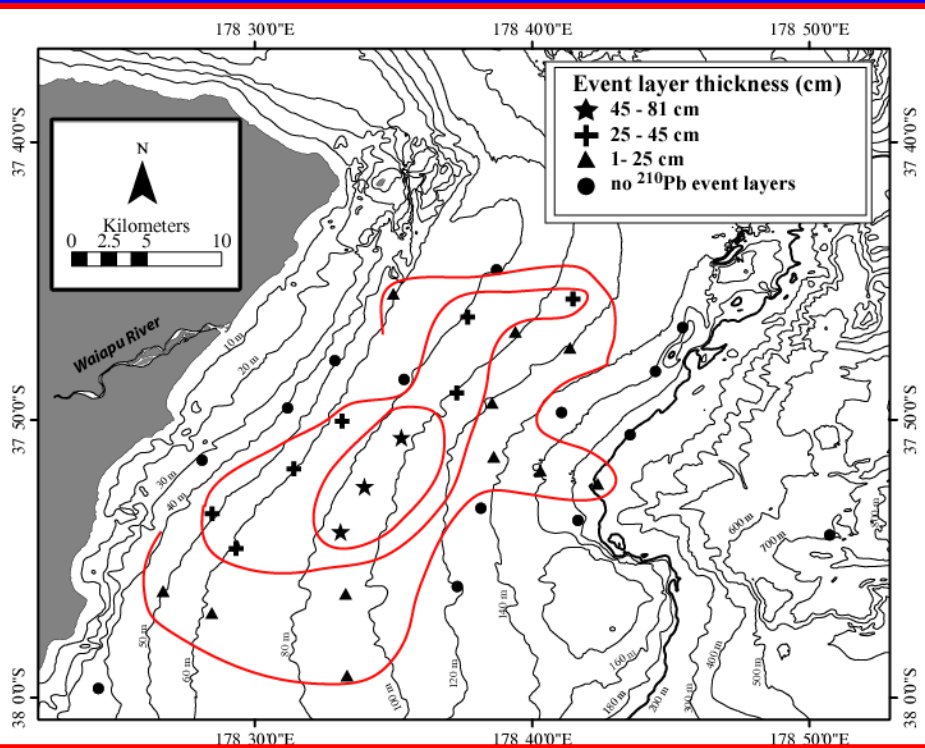


Pb-210 Accumulation Rates

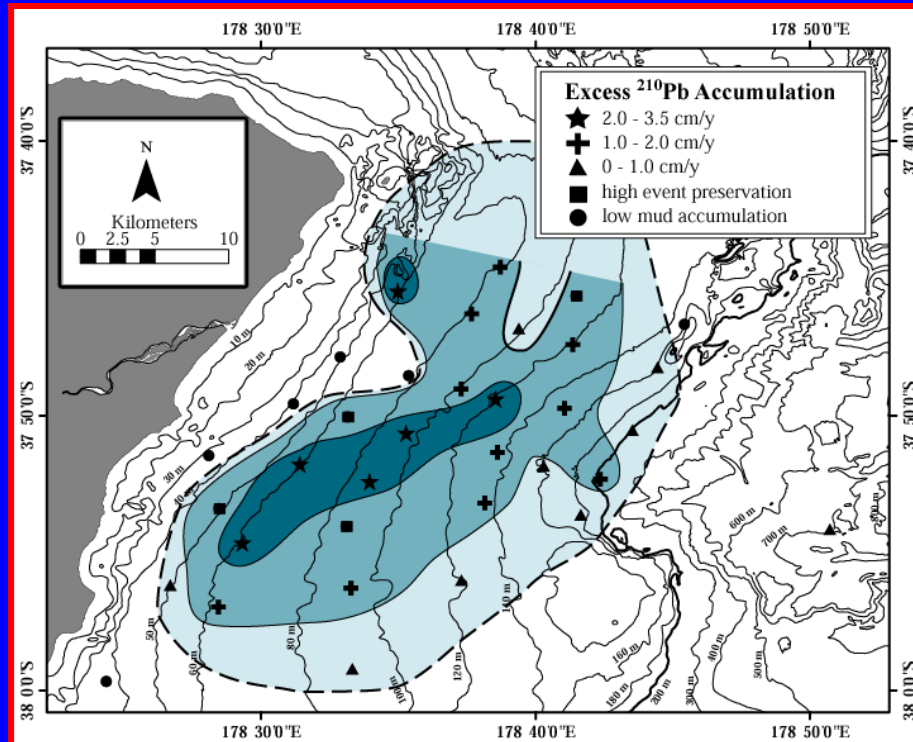


Underlying Tectonics: Wave and Current Driven Gravity Flow Direct Sediment Towards Shelf Basin

Event Layer Thickness



Pb-210 Accumulation Rate



Conclusions

- **Episodic rapid sediment inputs key to signal preservation**
- **Sediment gravity flows sensitive to subtle bathymetric gradients in slope**
- **Tectonic setting (i.e., accommodation) is dominant steering mechanism for sediment gravity flows**
- **Accumulation rate alone not sufficient to predict primary structure preservation**