Diagenesis of the Oligocene-Early Miocene Panjgur Formation, Paleocene Ispikan Formation and Wakai Exotic Blocks in the Makran Accretionary Belt, Southwest Pakistan*

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Abstract

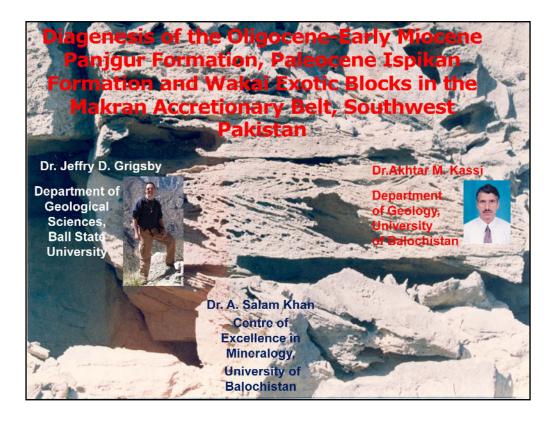
The Oligocene-Early Miocene Panjgur Formation is exposed in the Makran accretionary wedge in southwest Pakistan. Isolated, pre-Oligocene exotic blocks, the Ispikan Formation and Wakai exotic blocks, also crop out in the area. Petrographic and geochemical evidence indicate that the Panjgur and Ispikan detritus was transported from the Himalaya-Karakorum-Hindukush orogen southward and deposited as a deep marine turbidite complex in the northeast-southwest aligned Katawaz remnant ocean and east-west aligned Makran trench. The Bela-Muslimbagh ophiolites, associated mélanges to the east, and the Chaghai-Raskoh volcanic arc to the west contributed mafic to ultramafic detritus to the Panjgur and Ispikan sediments.

Sandstones of the Wakai exotic blocks, rich in mafic volcanic fragments and associated with sheared mafic volcanic rocks, purple colored shale, chert, marble, and pelagic limestone, suggest obduction of oceanic crust of the Gulf of Oman. The tectonics of these sandstones has led to a complex diagenetic history. Cements include ferroan>nonferroan calcite, ankerite, dolomite, and chlorite>illite with carbonate cements dominant in the Panjgur and Ispikan formations, and chlorite dominant in the Wakai exotic blocks. Physical compaction preceded cementation in the Panjgur and Ispikan Formations. Early nonferroan calcite inhibited physical compaction in the Wakai exotic blocks. Secondary porosity developed during dissolution of nonferroan calcite, feldspars, and rock fragments and was later destroyed by precipitation of chlorite, ferroan calcite, ankerite, and dolomite. Chlorite predates ferroan calcite, and both show signs of a second stage of dissolution with chlorite filling secondary pores in a nonchlorite-like I-beam morphology. Ankerite and dolomite formed late, replaced earlier stages of calcite, framework grains, and filled secondary pores.

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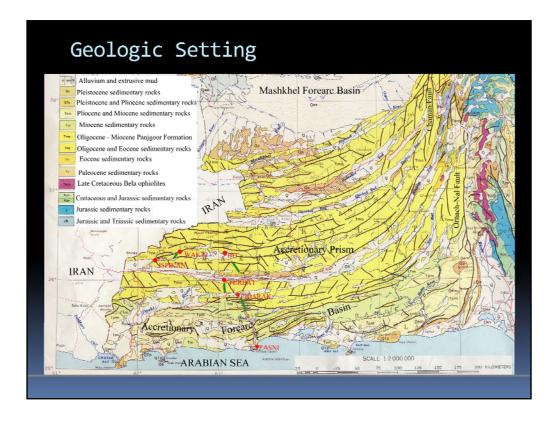
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Sample Locations





Geological map of the southern Balochistan (modified after the Geological Map of Pakistan, Bakr and Jackson 1964), showing positions of the studied sections (green colored lines) near Wakai, Bit, Turbat, and Pidarak.

Stratigraphy

Stratigraphy			
Age	Group	Formation	Lithology
Oligocene-Miocene		Panjgur Fm	SS interbedded with Shale
U. Eocene-L. Oligocene		Hoshab Shale	Shale w/thin bedded SS
	Turbat	Thrust	
Eocene		Wakai LS	Fossiliferous LS
		Thrust	
Paleocene		Ispikan Fm	Cong., SS, shale
		Thrust	
Cretaceous-Paleocene		Wakai Exotic Blocks	Agglom, purple sh, marl, chrt, pelagic LS, marble, mafic rxs

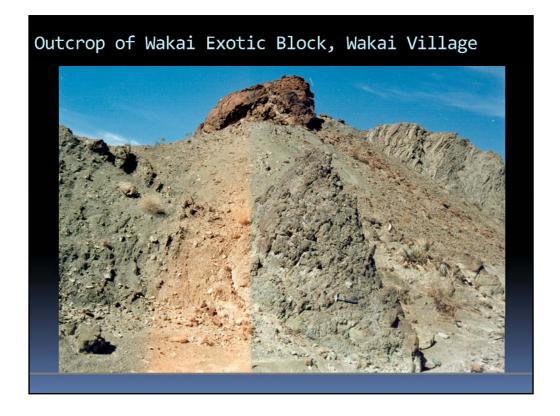
Travel



Field work

Wakai Exotic Blocks





Field photograph of the exposure of Wakai agglomerate near the Wakai village

Close up of Wakai Exotic Blocks - subangular boulders & pebbles of mafic volcanic rocks



lose up view of the agglomerate near Wakai village showing angular to subangluar boulders and pebbles of mafic volcanic rocks.

Paleocene? Ispikan Formation



Paleocene? Ispikan Formation



Ispikan Sandstone with Cross Bedding



Well rounded pebble-boulder conglomerate of the Ispikan Formation



Purple colored shale at the base of the Ispikan Formation

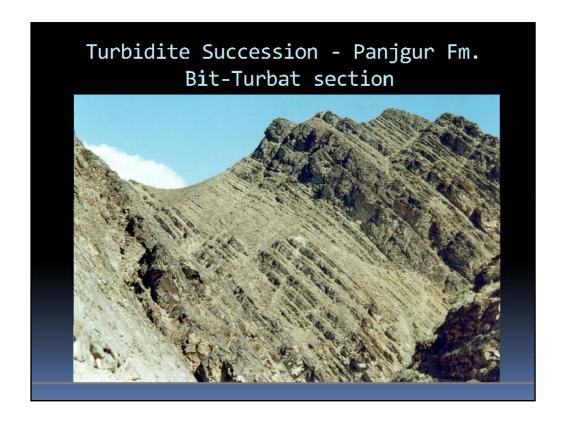


Bit-Turbat Section

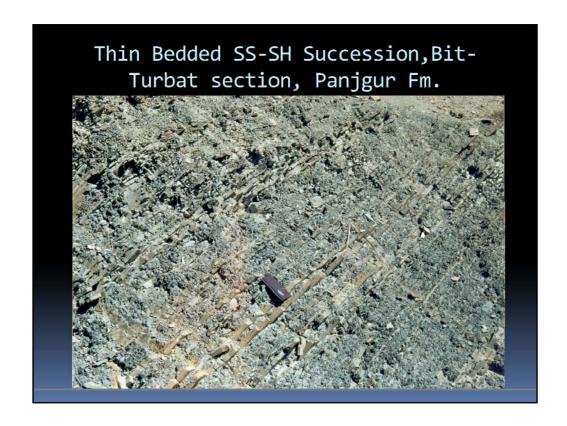


Oligocene-Miocene Panjgur Fm.





Field photograph of the thick-bedded sandstone-dominant succession at Bit-Turbat road section. Rhythmic alternations of sandstone and shale beds are characteristic of turbidite succession. Panjgoor Formation has turbidite successions which show deposition in submarine fan environments. Note second-order cyclicity within the succession.



Field photograph of the thin-bedded shale-dominant succession at Bit-Turbat road section northeast of the town of Turbat. Note lateral continuity of thin sandstone and shale beds.

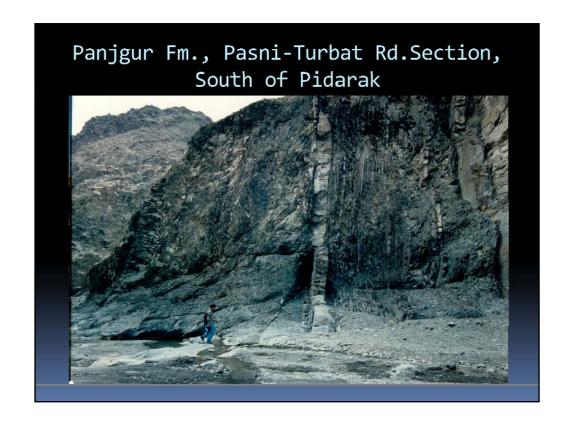


Field photograph of the flute marks at the base of a sandstone bed of the Panjgur Formation--commonly found in the turbidite successions.

Turbat-Pasni Rd. Section







Field photograph of the thin-bedded shale-dominant succession at Pasni-Turbat road section south of Pidarak. Note lateral continuity of thin sandstone and shale beds.



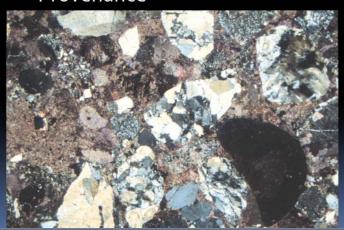
Field photograph of the thin-bedded shale-dominant succession at Pasni-Turbat road section south of Pidarak. Note lateral continuity of thin sandstone and shale beds.

Imbricated Sandstone Bed, Panjgur Fm.



PETROGRAPHY

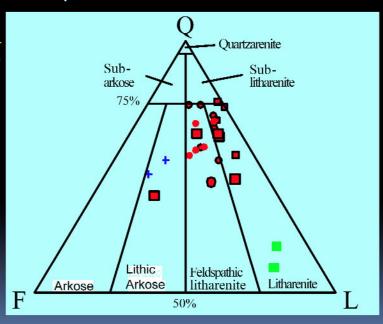
Provenance

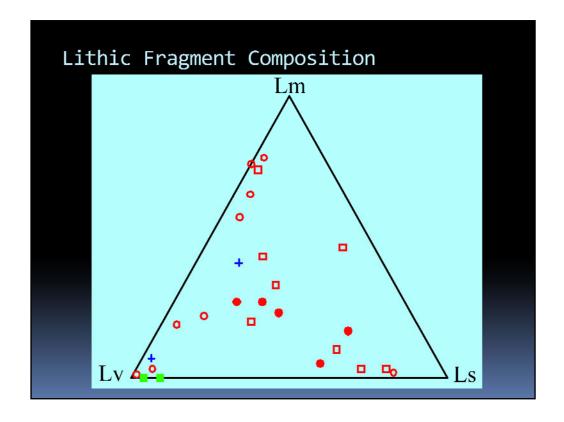


Sandstone Composition

Red = Panjgur Fm. Blue = Ispikan Fm. Green = Wakai

After Folk, 1980

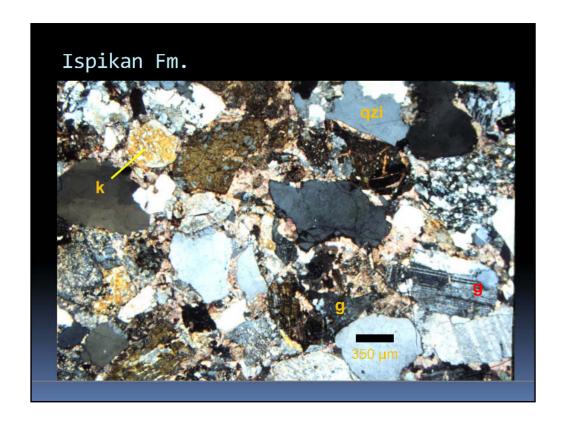




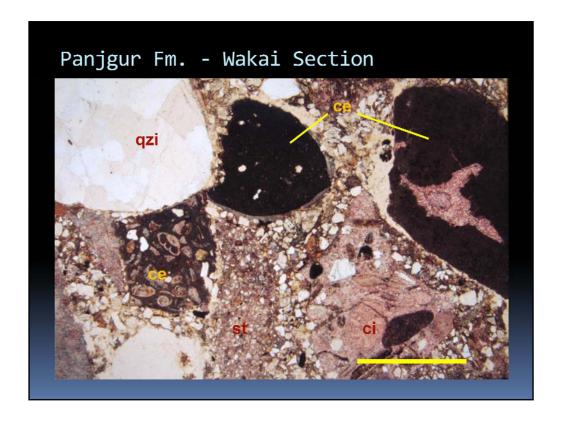
LmLvLs compositional diagram (after Dickinson and Suczek 1979) of sandstones of the Panjgur Formation, Ispikan Formation, and Wakai mélange showing a mixed igneous, sedimentary, and metamorphic provenance. Symbols are same as in previous figure.



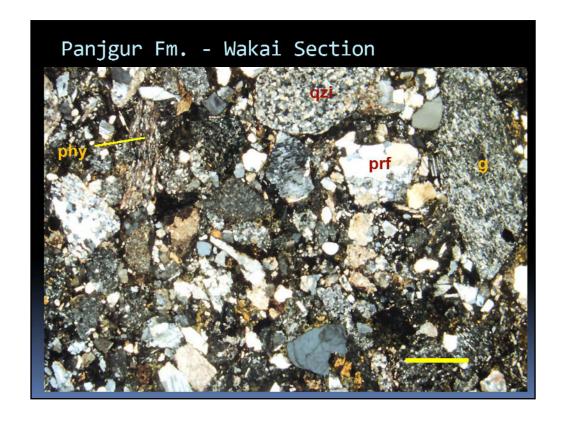
Photomicrograph of the volcaniclastic sandstone of the Wakai mélange showing abundant mafic volcanic fragments within micritic cement; crossed polars.



Photomicrograph of the sandstone of the Ispikan Formation showing granitic and metamorphic fragments. Note the presence of k-feldspar (stained yellow) within the granitic rock fragments; crossed polars.

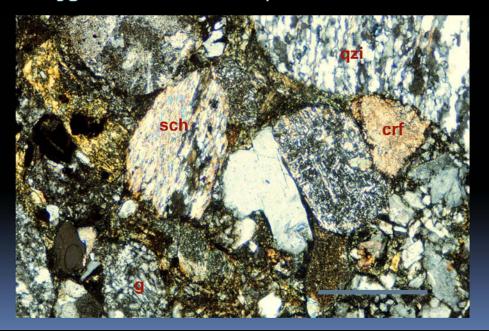


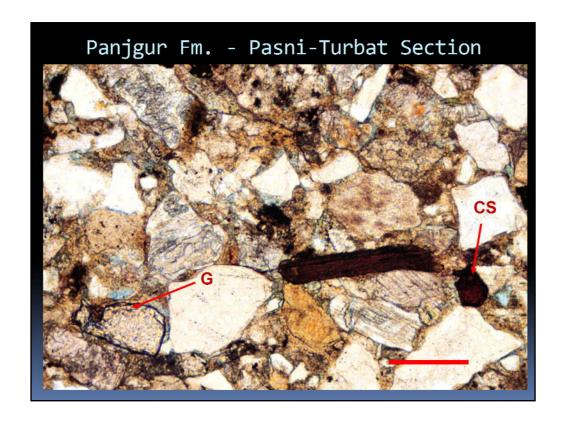
Panjgur Formation from Wakai section showing extrabasinal limestone (ce), some being fossiliferous containing pelagic foraminifers, siltstone (st), and quartzite fragments (qzi); crossed polars. Bar scale equals $250~\mu m$ in all photomicrographs unless otherwise stated



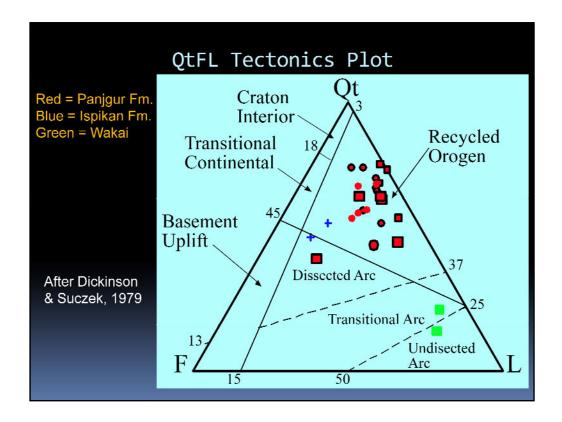
Photomicrograph of the sandstone of the Panjgur Formation from Wakai section showing various types of igneous and metamorphic rock fragments; crossed polars.

Panjgur Fm.- Wakai-Ispikan Section





Photomicrograph of the sandstone of the Panjgur Formation from Pasni-Turbat road section showing various types of mineral and fossil fragments. Note grains of garnet (lower left corner), chrome spinal (lower right corner), and partial replacement of feldspar grain by calcite (left center); plane polarized light.



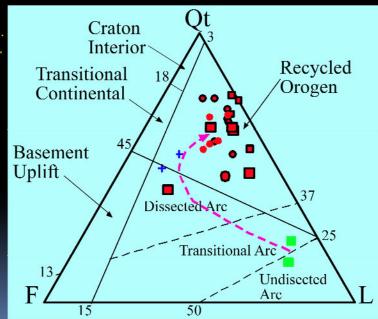
QFL compositional diagram (after Dickinson and Suczek 1979) of sandstones of the Panjgur Formation, Ispikan Formation, and Wakai mélange. Most of the samples of the Panjgur Formation and Ispikan Formation plot in the field of recycled orogen, whereas, the two sandstone samples

of the Wakai mélange plot in the fields of undissected and transitional arc. Symbols are same as in previous figures.

QtFL Tectonics Plot

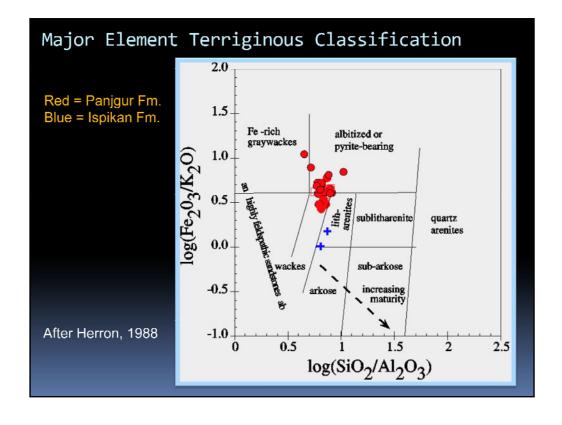
Red = Panjgur Fm. Blue = Ispikan Fm. Green = Wakai

After Dickinson & Suczek, 1979

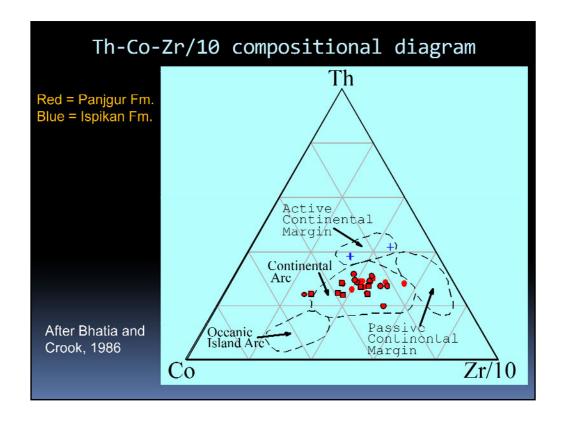


GEOCHEMISTRY

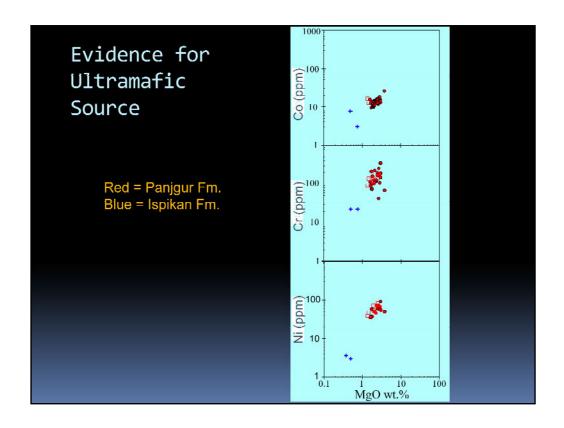
Provenance



Major-element classification of terrigenous sandstones reflecting sedimentary maturity (Herron 1988). Symbols the same as in previous figures.

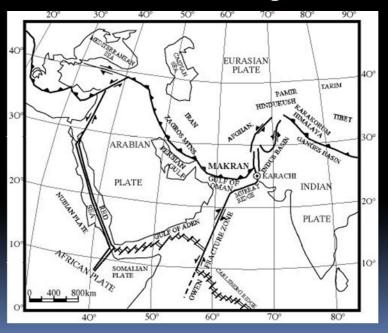


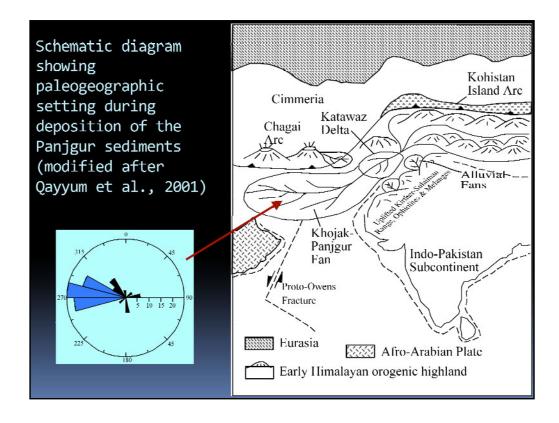
Th-Co-Zr/10 discrimination diagram for sandstones (Bhatia and Crook 1986). Symbols the same as in previous figures.



Log-log plots of Co, Cr, and Ni against MgO. High abundances of Co, Cr, and Ni are indicative of contributions from ultramafic sources. Symbols the same as in previous figures.

Tectonic Setting



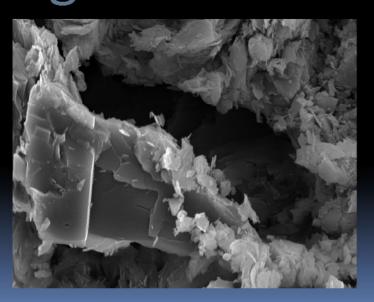


Schematic diagram (modified after Qayyum et al., 2001), showing paleogeographic setting of Indo-Pakistan subcontinent, Eurasia, and the Katawaz delta and Khojak-Panjgur submarine fan in the Katawaz Ocean.

We support the model that the Panjgur and Ispikan detritus was transported from the Himalayan collisional orogen southward across strike and deposited as a turbidite system in the Katawaz Ocean. The initial southward transport was controlled by the Chaman-Ornach Nal transform fault system. In the Makran region the submarine fan depositional system turned westward, roughly parallel to the present active margin of the Makran accretionary wedge.

The Bela-Muslimbagh ophiolites and associated mélanges in the east and Chaghai-Raskoh volcanic arc to the west concurrently shed mafic to ultramafic detritus that is also preserved in the study area.

Diagenesis

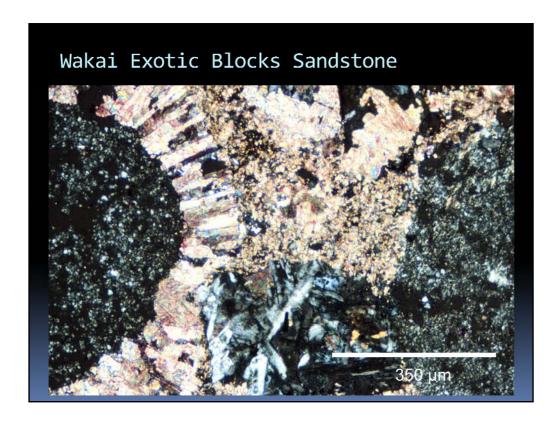




Photomicrograph of the volcaniclastic sandstone of the Wakai mélange showing abundant mafic volcanic fragments within micritic cement; crossed polars.



Chloritized mafic volcanic rock fragments. Micritic calcite. Fe-poor calcite replacing grains.

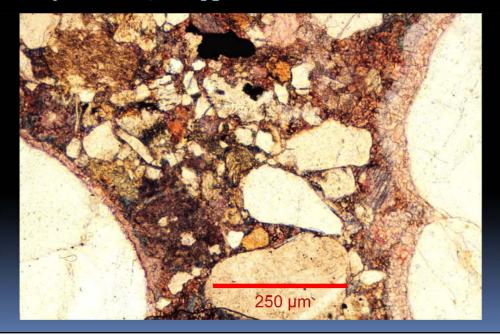


Photomicrograph of the volcaniclastic sandstone of the Wakai agglomerate showing growth of bladed rim of calcite around chert fragment and development of finely crystalline cement in the interstices; crossed polars.

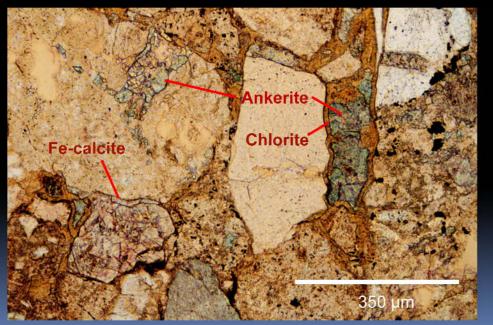


Fractured grains as a result of compaction. Fe-calcite fills fractures and displaces biotite along cleavage planes.

Early calcite, Panjgur Fm. - Wakai Section



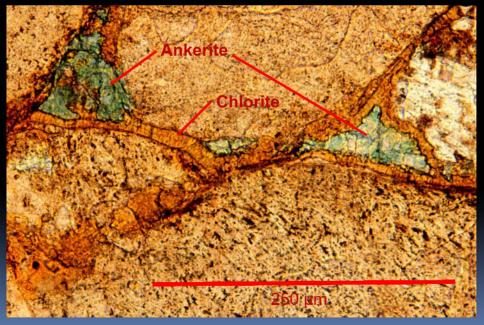
Panjgur Fm. - Wakai Section

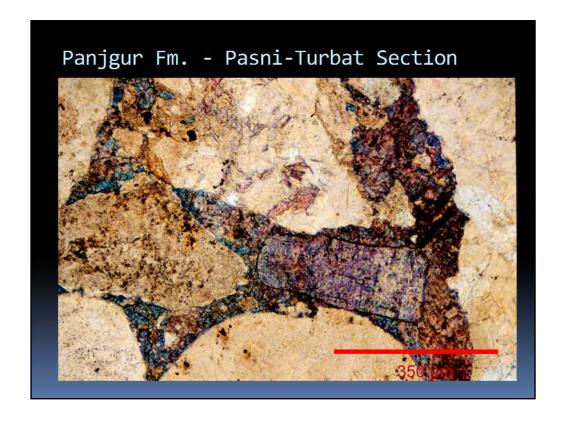


Fe-calcite & Chlorite - Panjgur Fm.

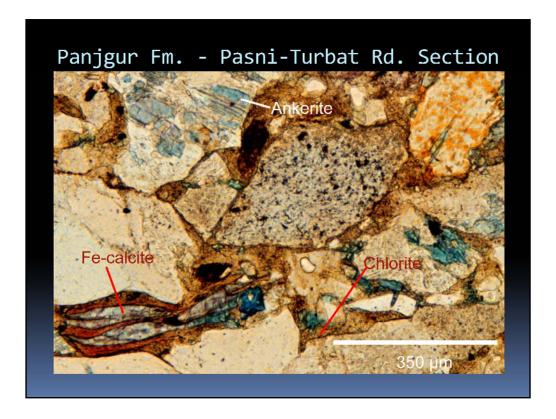


Panjgur Fm. - Wakai Section





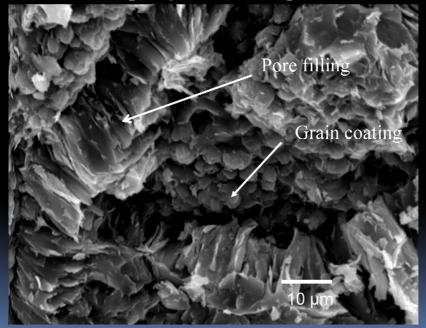
Calcitized and kaolinized feldspar, chert, and metamorphic r.f. in a calcite ferroan cement.



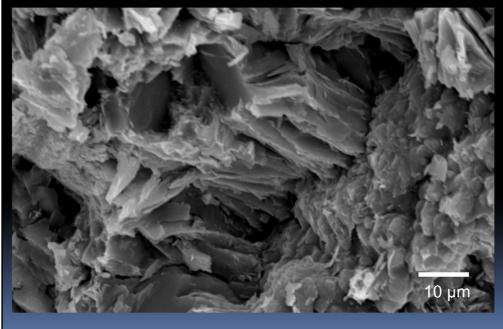
Displacive ferroan calcite (mauve stained) disrupting biotite grain and patchy calcite replacing plagioclase and orthoclase (yellow stained). Chlorite cement surrounds grains.

Grain coating chlorite - rounded by dissolution?

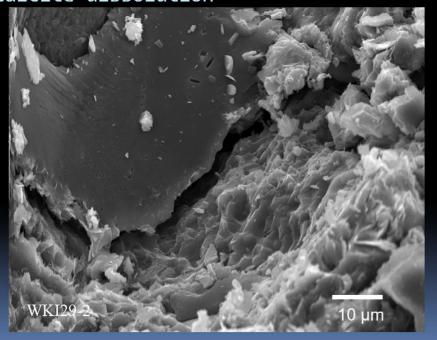
Grain coating & pore filling chlorite



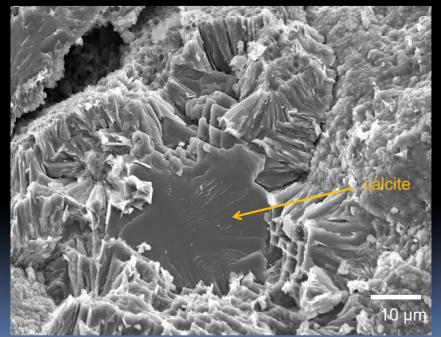
Chlorite growth

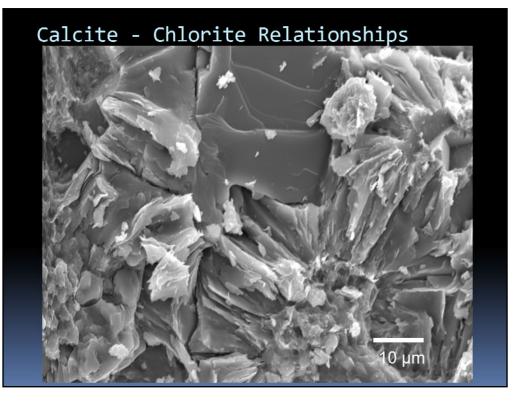


Calcite dissolution

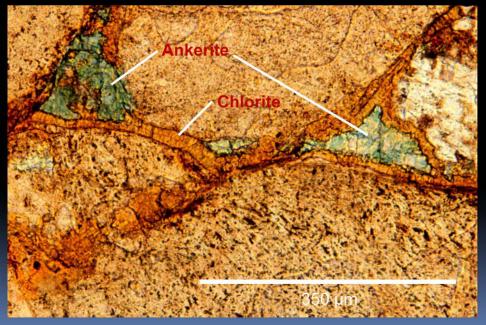


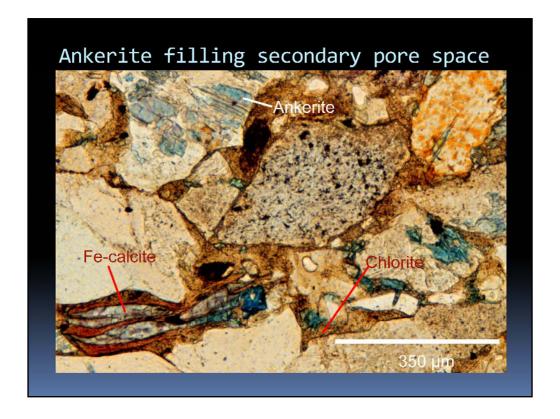
Grain coating & pore filling chlorite & calcite



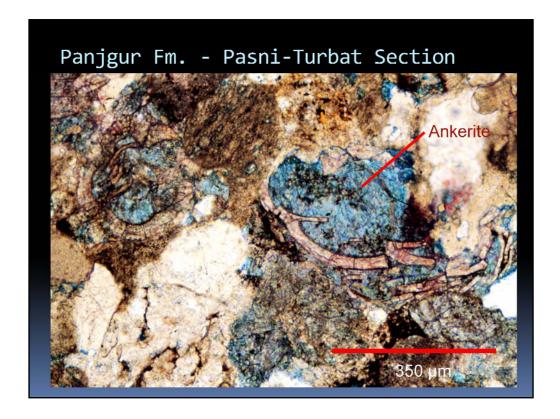


Panjgur Fm. - Wakai Section

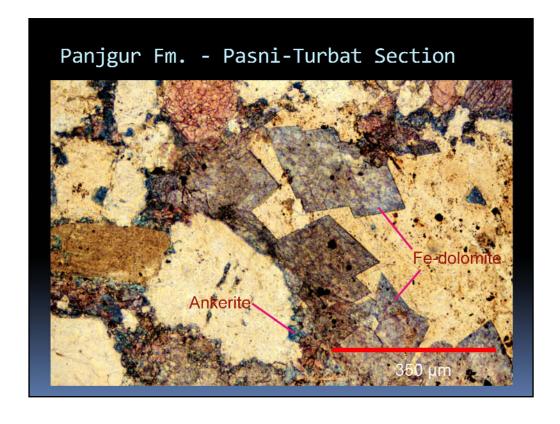




Displacive ferroan calcite (mauve stained) disrupting biotite grain and patchy calcite replacing plagioclase and orthoclase (yellow stained). Chlorite cement surrounds grains.



Secondary porosity filled by late stage ankerite.



Calcite ferroan (mauve stained) and ankerite (blue stained) replacing various r.f. (e.g, chert)

Conclusions (at the moment)

Sandstone composition of the Wakai exotic blocks, and the Ispikan and Panjgur Formations of the Makran belt indicates a transition from undissected arc to dissected arc to recycled orogen during Cretaceous to Miocene time.

Geochemical data support an active continental margin to continental arc tectonic setting.

Detritus was derived from the Himalayan orogenic belt and associated suture belts, ophiolite complexes, and volcanic arcs.

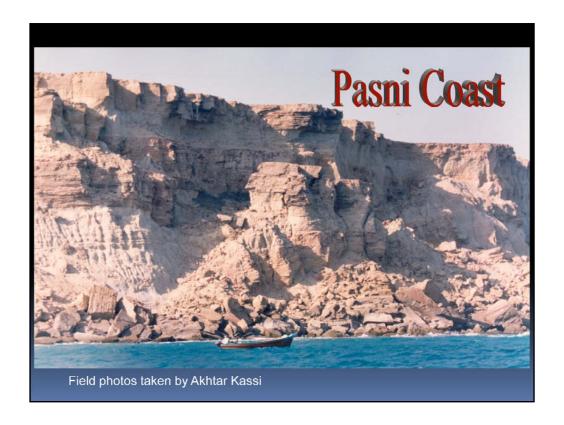
Conclusions (continued)

Early cementation (calcite &/or chlorite) retards compaction.

Wakai and Ispikan sandstones are dominated by early Fe-poor calcite that is replaced by a later ferroan calcite.

Diagenesis of the Oligocene-Miocene Panjgur Fm is complex.

- a) Minor compaction with formation of chlorite grain coats & early Fe-poor calcite.
- b) Fe-calcite precipitation replacing grains, Fe-poor calcite (& chlorite?)
- c) Carbonate dissolution & chlorite grain coating dissolution, precipitation of pore filling chlorite.
- d) Dissolution of feldspars
- e) Precipitation of ankerite filling secondary pore spaces
- f) Precipitation of Fe-dolomite



The Ispikan and Panjgur Formations of the Makran accretionary belt are composed of high proportions of quartz and metamorphic, igneous, and sedimentary rock fragments.

- 1) The QtFL plot suggests a dissected arc to recycled orogen for these sandstones.
- 2) The rock fragments and associated heavy minerals, like serpentinite and chrome spinel, suggest that the detritus was derived from the Himalayan orogenic belt and associated suture belts, ophiolite mélanges, and volcanic arcs.
- 3) Geochemical data support an active continental margin to continental arc tectonic setting for these sandstones with the Panjgur Fm. getting contributions from the ophiolite mélanges. This contribution is minor or absent in the Ispikan Fm.
- 4) Sandstone of the Wakai mélange is very rich in mafic volcanic fragments, which were derived from a nearby volcanic source. Its association with sheared mafic volcanic rocks, purple colored shales, cherts, marbles, and pelagic limestones suggest that they were possibly derived from the ophiolitic mélanges obducted due to the subduction of oceanic crust of the Gulf of Oman,
- 5) We support the model that the Panjgur and Ispikan detritus was transported from the Himalayan collisional orogen southward across strike and deposited as a delta-submarine-fan system in the Katawaz Ocean. The initial southward transport was controlled by the Chaman-Ornach Nal transform fault system. In the Makran region the submarine fan depositional system turned westward, roughly parallel to the present active margin of the Makran accretionary wedge.
- 6) The Bela-Muslimbagh ophiolites and associated mélanges in the east and Chaghai-Raskoh volcanic arc to the west concurrently shed mafic to ultramafic detritus that is also preserved in the study area.

Selected References

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