Alternative Views for Common Assumptions: Reassessing the Origin and Significance of Sequence Boundaries Using Field and Flume*

John Holbrook¹ and Nikki Strong²

Search and Discovery Article #40446 (2009) Posted September 30, 2009

Abstract

Field data and recent experimental studies independently question long-held paradigms regarding the origin and time significance of fluvially carved sequence boundaries as well as genetic relationships between these surfaces and the strata they bind. These field data derive from an updip to downdip transect through the Cretaceous Dakota Group of the U.S. southern High Plains. The experimental data derive from repeated basin-scale runs of sequence development during relative sea-level change simulated in the Jurassic Tank at the University of Minnesota, St Anthony Falls. Both experimental and field data show that fluvial sand above sequence boundaries are deposited coexistent with the carving of the underlying sequence boundary. The field data do this by inference from mapped crosscutting relationships within observed stratigraphy and the experimental data through scaled reproduction of the processes inferred and products observed from the field. Both sources converge to reinforce assertions regarding sequence boundaries that require readjustment of some commonly held views. Namely, surfaces commonly mapped as sequence-bounding unconformities were not necessarily synchronously exposed, record no common age, and may not consistently separate older from younger strata. Also, fluvial strata above sequence boundaries do not necessarily reflect passive burial of these surfaces during subsequent transgression. Instead these strata may record co-generation of fluvial reservoir architecture and the underlying sequence-boundary over the full duration of the transgressive/regressive cycle because of close genetic links between the two. Furthermore, valley incision and sequence-boundary erosion need not reflect updip knickpoint migration from the shore; thus valleys and sequence-boundary continuity may commonly be lost down depositional dip.

^{*}Adapted from oral presentation at AAPG Annual Convention, Denver, Colorado, June 7-10, 2009

¹Earth and Environmental Sciences, University of Texas at Arlington, Arlington, TX (Holbrook@uta.edu)

²Geology and Geophysics, University of Minnesota, Minneapolis, MN

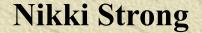
Alternative Views for Common Assumptions: Reassessing the Origin and Significance of Sequence Boundaries using Field and Flume





John Holbrook

University of Texas at Arlington

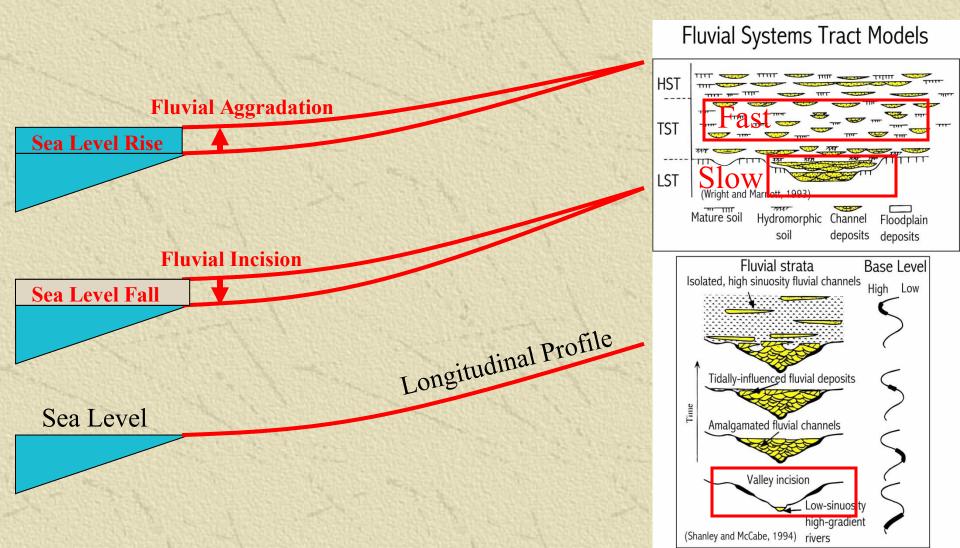


University of Minnesota

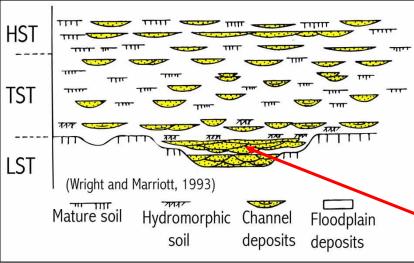


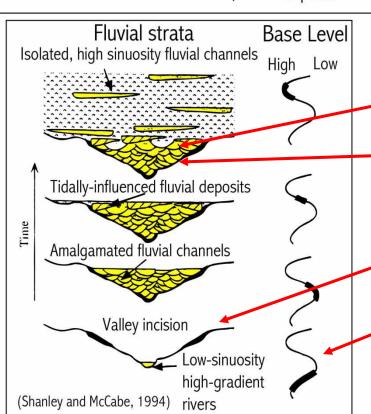


Fluvial Response to Base Level Change and Generation of Sequences



Fluvial Systems Tract Models





Upstream and Lateral?

Some Common Assumptions

Progressive Stacking in Sheets during Filling with little Valley Modification

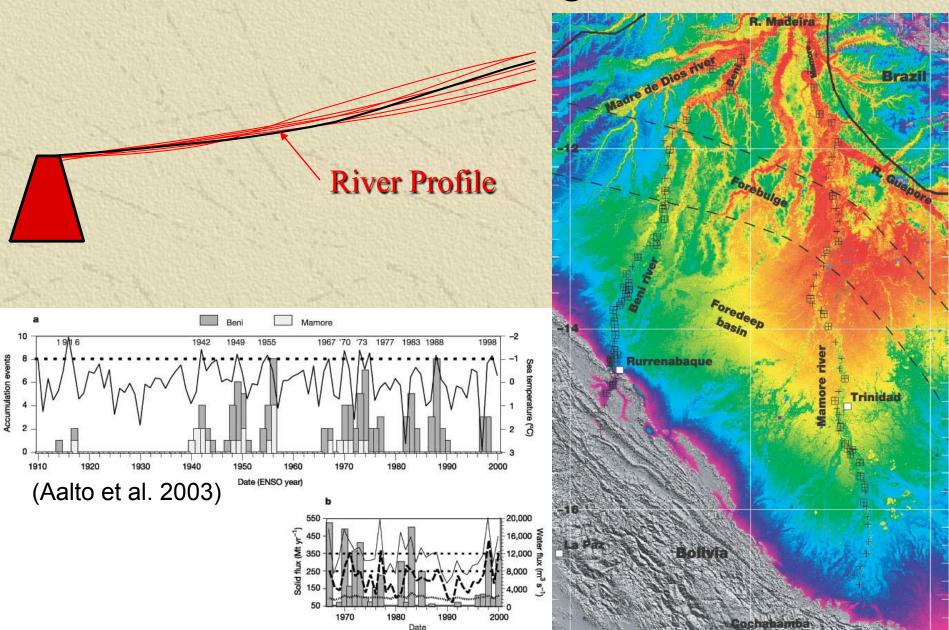
Stratigraphic Valley = Topographic Valley

Somewhat Synchronous Sequence Boundaries

Subaerial Lowstand Surface of Erosion

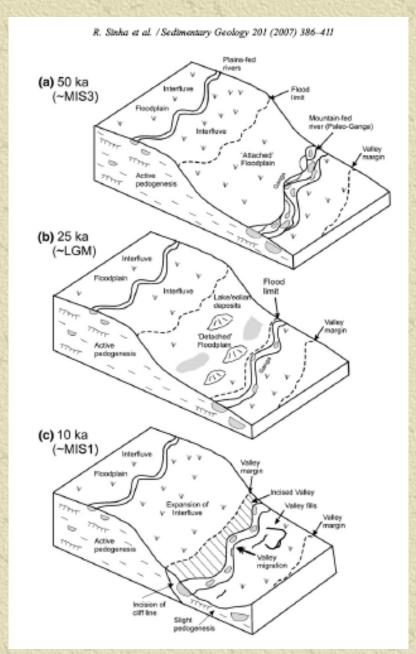
Incision during Falling Stage and Lowstand

Climate and Sediment Storage



Ganges Incision from Climate Change





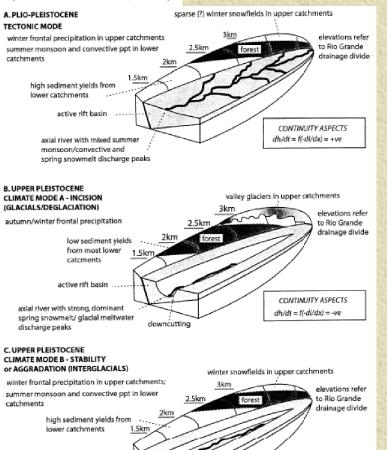
Rio Grande, NM

(Leeder and Mack, 2007)

active rift basin

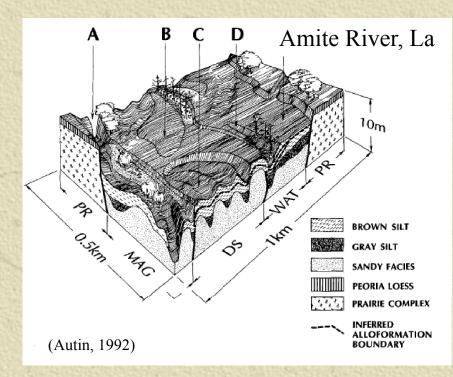
axial river with dominant

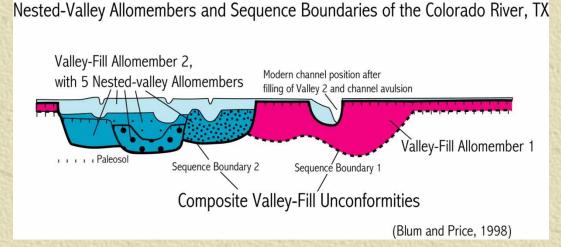
spring snowmelt discharge peak



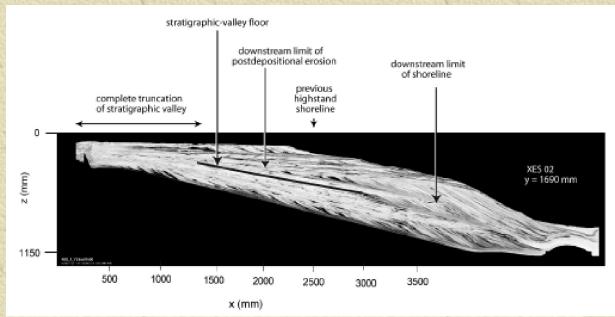
partial backfilling

CONTINUITY ASPECTS dh/dt = f(-di/dx) = +ve









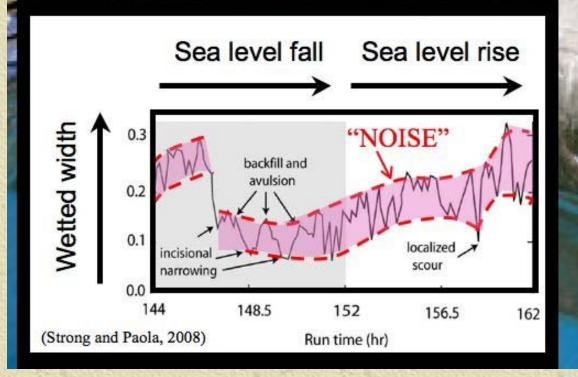


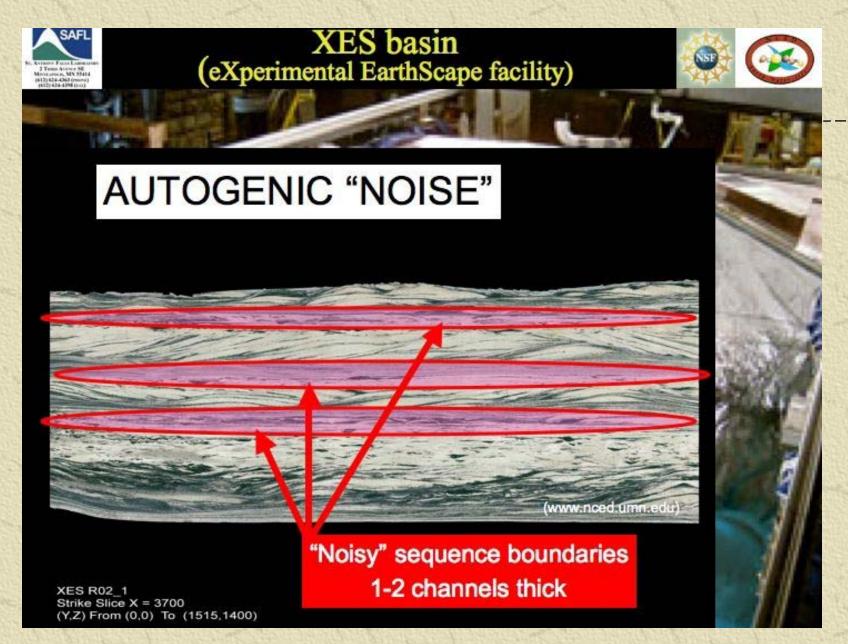
XES basin (eXperimental EarthScape facility)

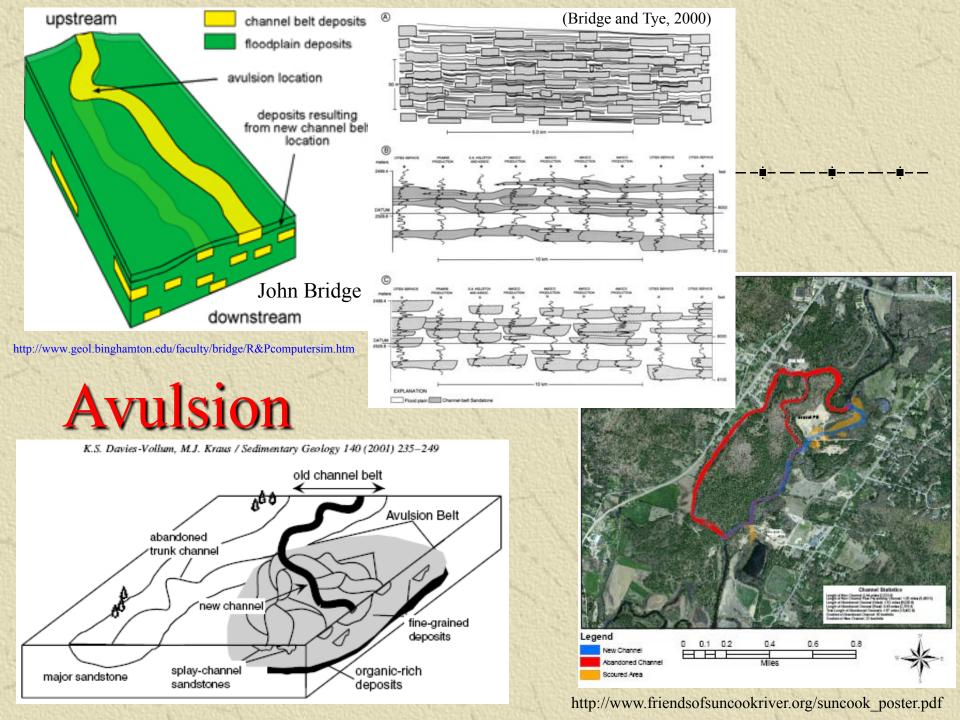




AUTOGENIC "NOISE"





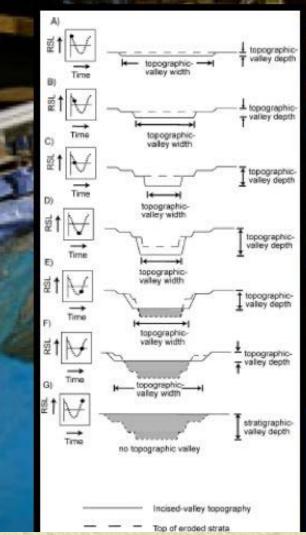




XES basin (eXperimental EarthScape facility)

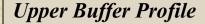






AUTOGENIC "NOISE" Autogenic-driven episodes of erosion and deposition both create and destroy fluvial terraces

Base Level Buffers and Buttresses



Transport Capacity = Min Sediment Influx = Max Uplift Rate = Max

Buffer Zone

Instantaneous

Profile

Lower Buffer Profile

Transport Capacity = Max Sediment Influx = Min Uplift Rate = Min

Preservation Space

Sea Level

Buttress

(Sea Level, Cataract, Lake Level, etc.)

Determiners of "Graded" Profile Elevation

Buffers

Sediment Influx/Transport Capacity = 1 (eq. 1) $dz/dt + dq_s/dx = 0 (eq. 2)$

Where:

 q_s = Sediment Discharge= $f(\omega)$, substrate erodability) Sediment Influx = q_s delivered at x_i = f (drainage basin) Transport Capacity = q_s that can be transported at x_i dz/dt = 0

Variables

Z Profile elevation

X Stream distance

t _{Time}

 $\omega = \gamma Q_w S$

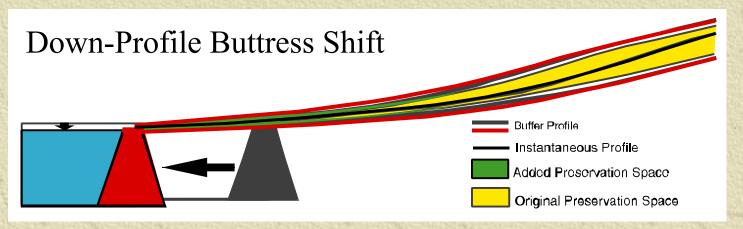
W Stream Power

γ Specific Weight

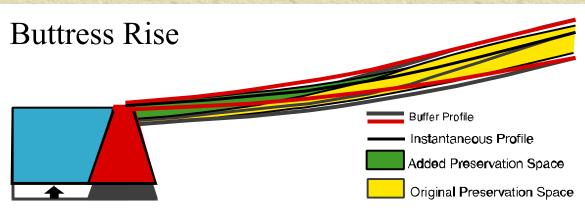
Qw Water Discharge

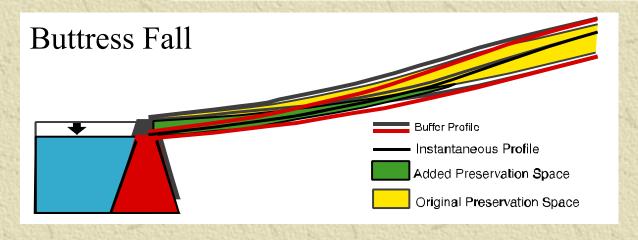
S Slope

(Holbrook et al., 2006)

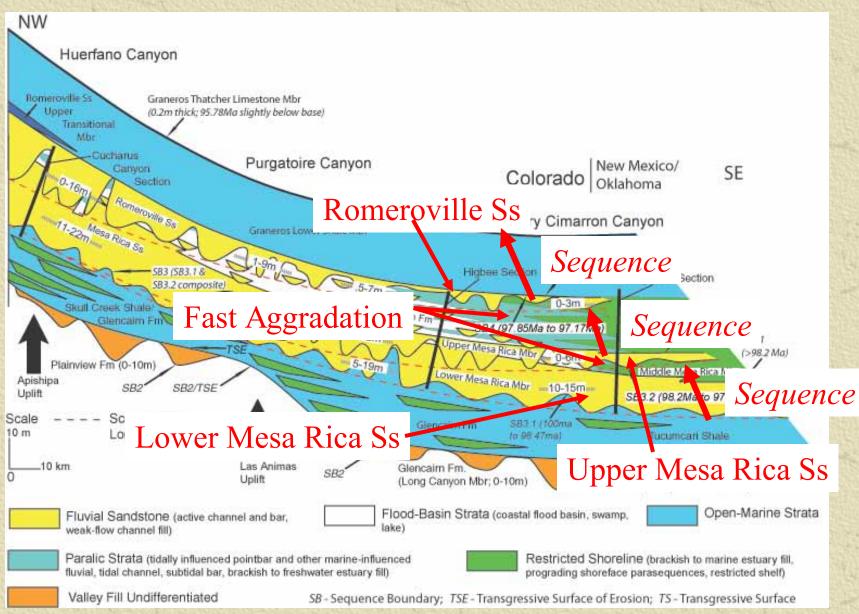


Some Effects of Buttress Shift



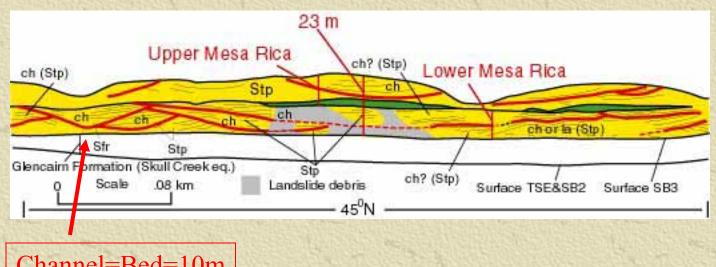


Cretaceous Dakota Group, US Western Interior

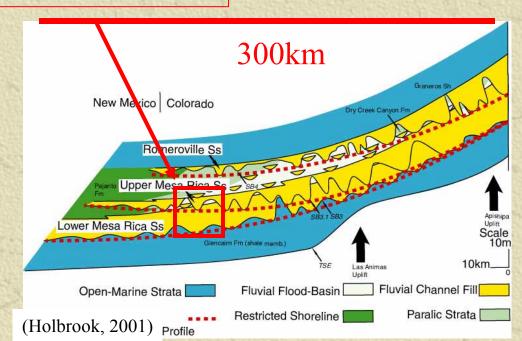


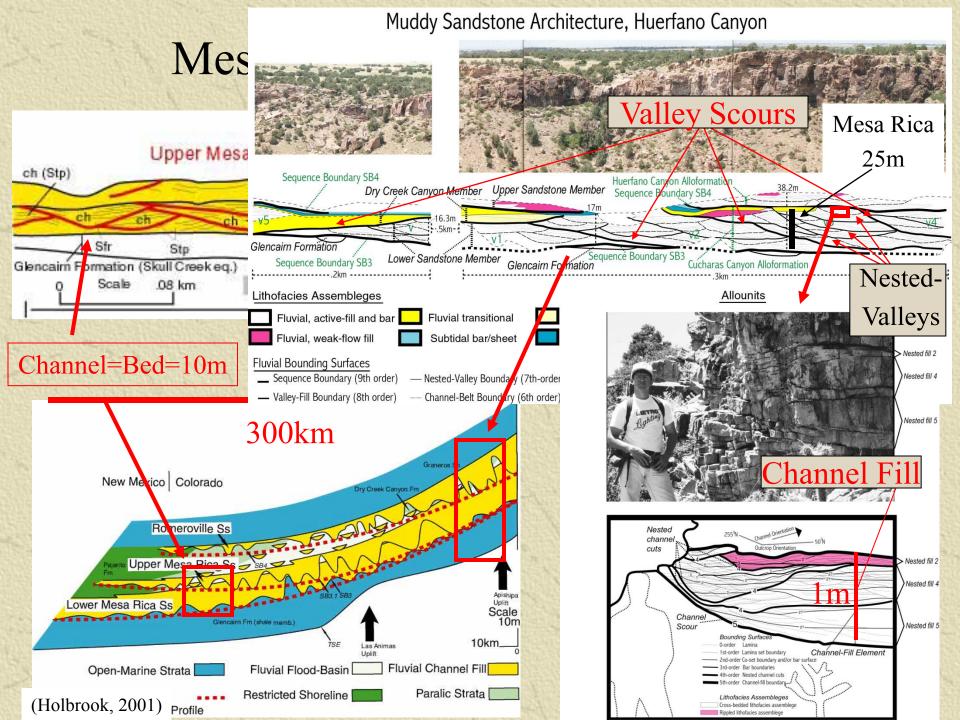
(Holbrook et al., 2006)

Mesa Rica Sandstone Architecture

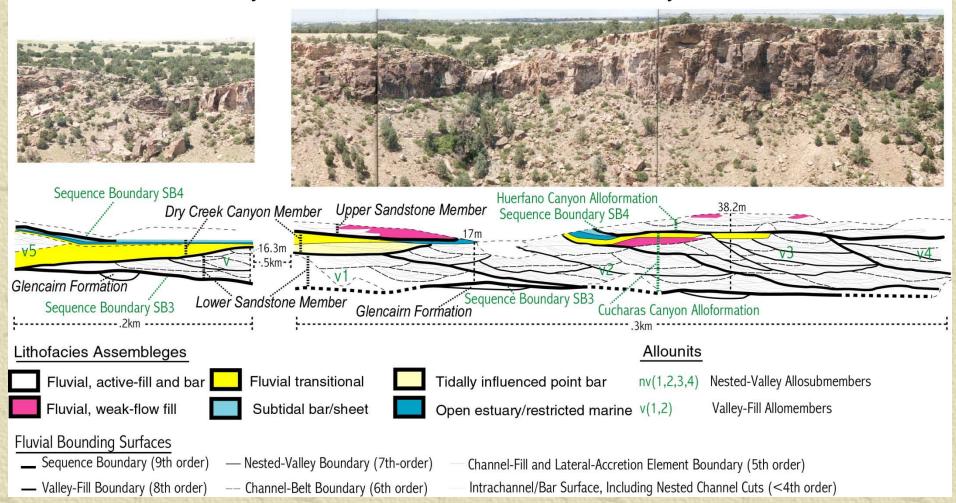




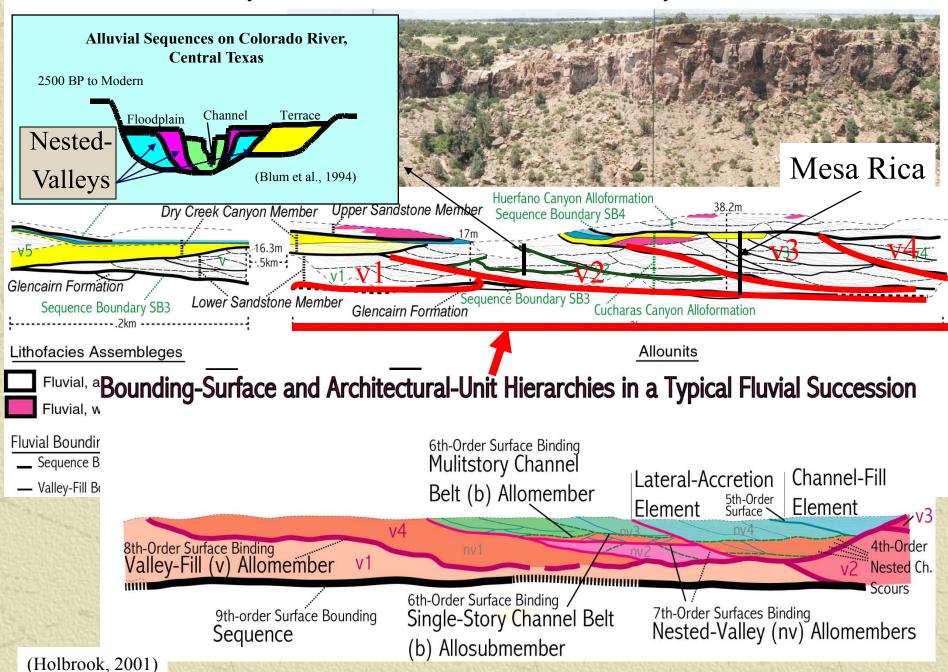




Muddy Sandstone Architecture, Huerfano Canyon

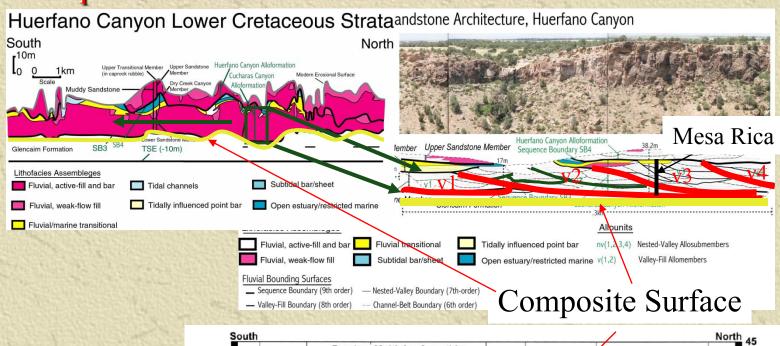


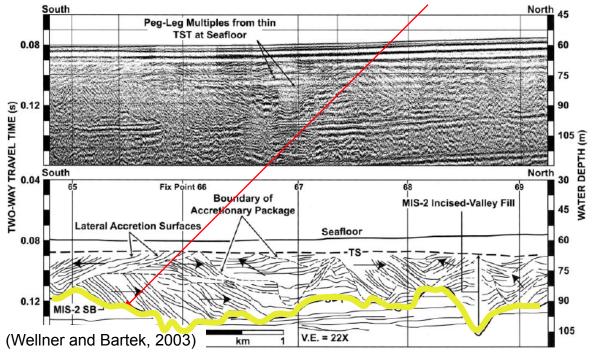
Muddy Sandstone Architecture, Huerfano Canyon



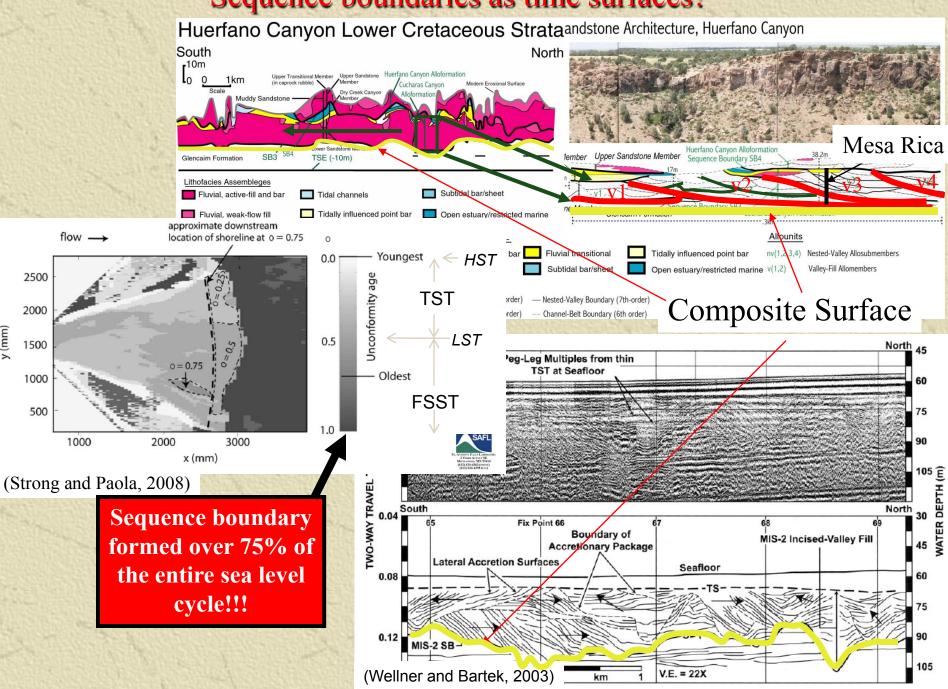
Implications for Sequence Boundaries, Sequence Architecture, and Sequence Construction

Sequence boundaries as time surfaces?

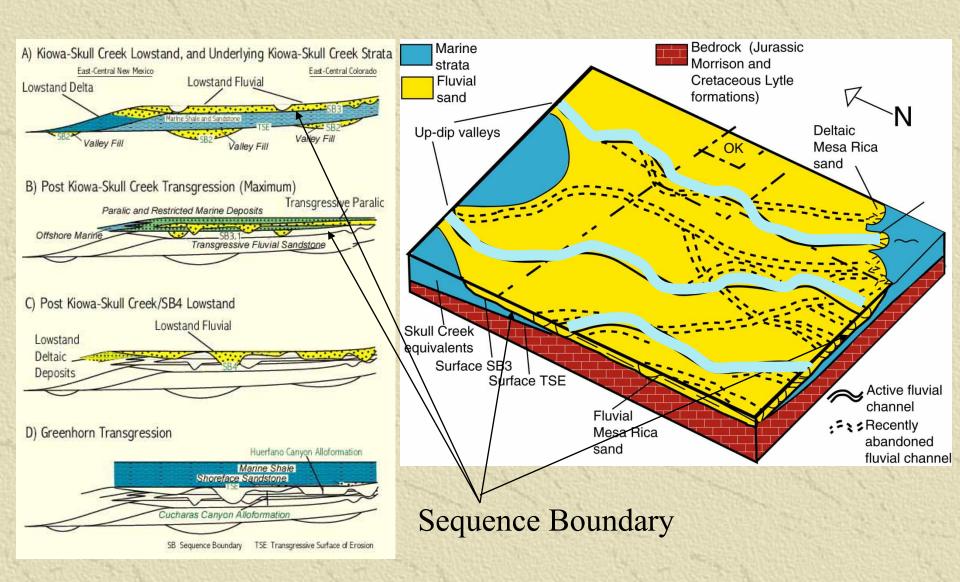




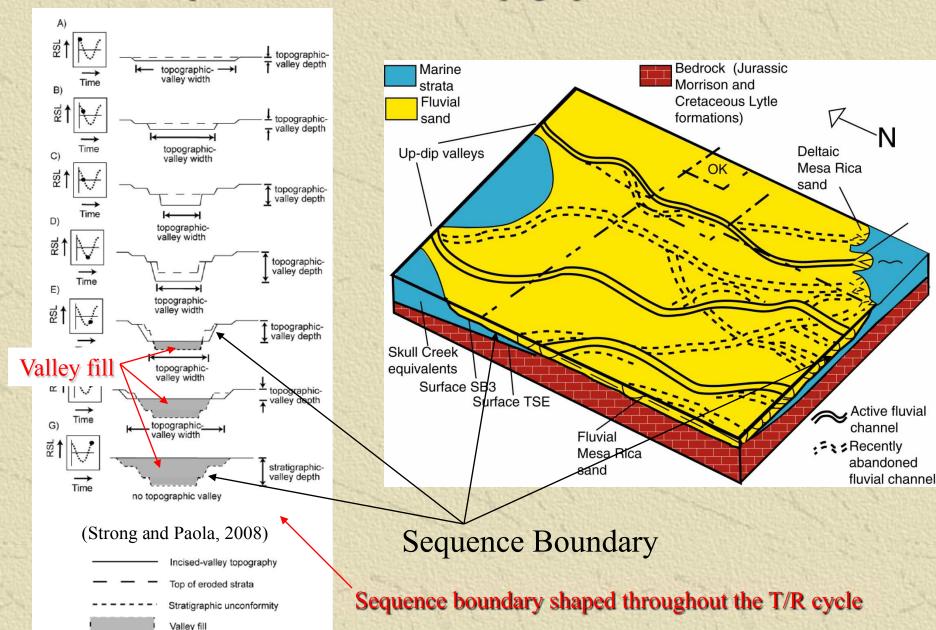
Sequence boundaries as time surfaces?



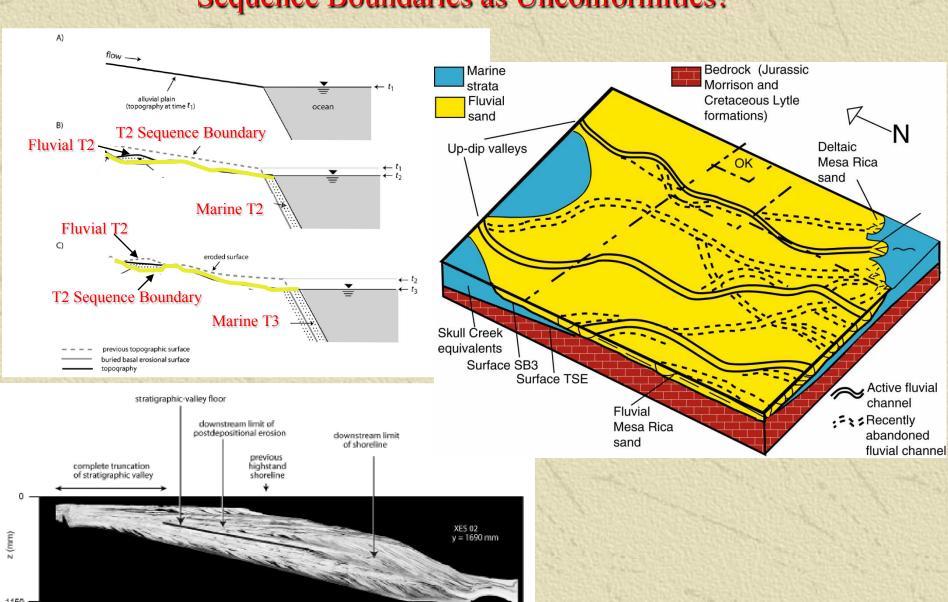
Sequence Boundaries as Topographic Surfaces?



Sequence Boundaries as Topographic Surfaces?



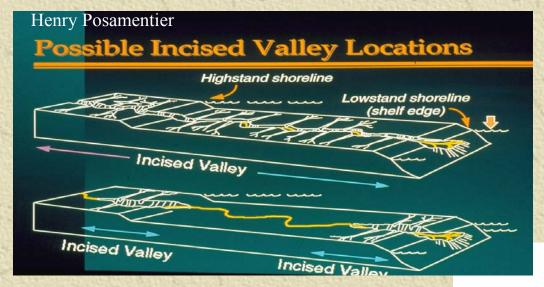
Sequence Boundaries as Unconformities?



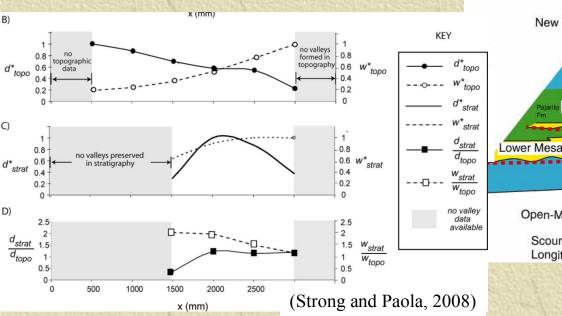
(Strong and Paola, 2008)

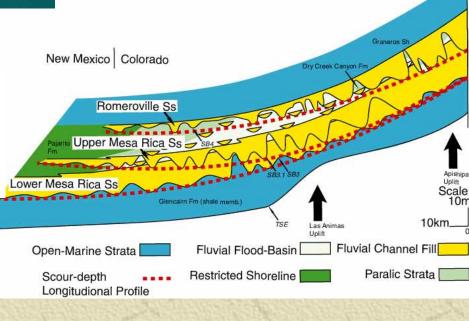
x (mm)

Valley Incision by Knickpoint or Buffer?



Knickpoint Valleys





Implications for Rapid Flooding

Deltaic Mesa Rica sand

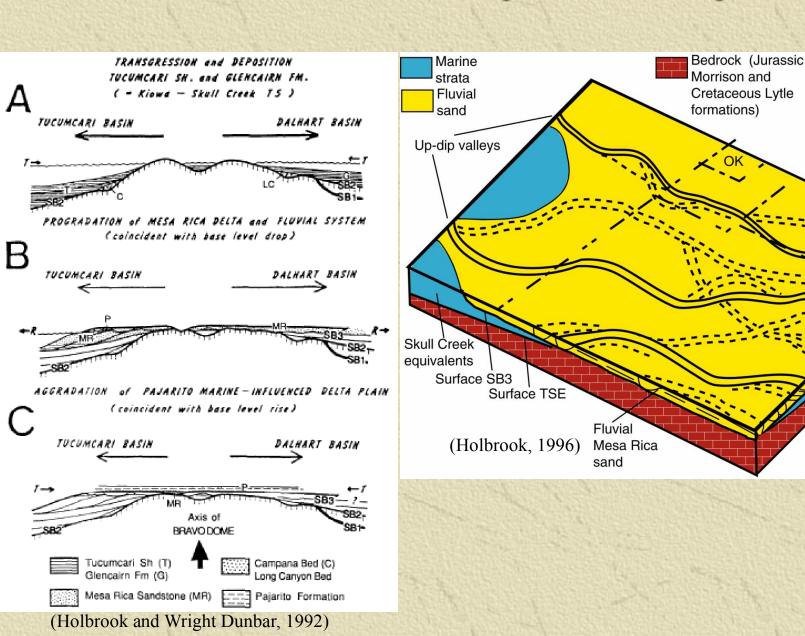
Active fluvial

abandoned

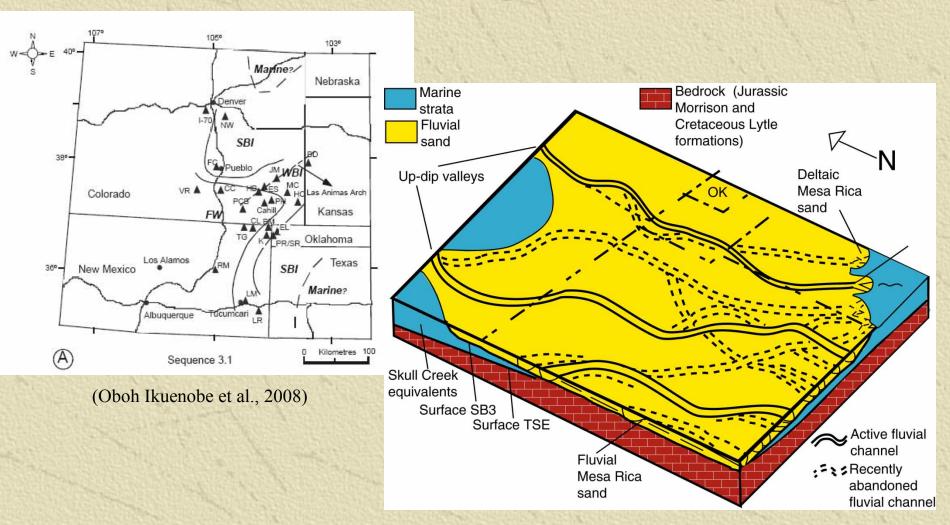
fluvial channel

channel

:=: Recently



Implications for Rapid Flooding



Conclusions

- 1) "Sequence boundaries" are time-transgressive composite surfaces formed over the duration of the T/R cycle...therefore...
- 2) "Sequence boundaries" rarely equate to topographic surfaces
- 3) "Sequence Boundaries" are not always unconformities
- 4) Valley erosion can initiate either in the proximal or distal region of the basin ...Buffer vs. Knickpoint valleys
- 5) Sand sheets above "sequence boundaries" are prone to rapid transgression

References

Aalto, R., L. Maurice-Bourgoin, T. Dunne, D.R. Montgomery, C.A. Nittrouer, and J.L. Guyot, 2003, Episodic sediment accumulation on Amazonian floodplains influenced by ENSO: Nature, v. 25, p. 493-497.

Autin, W. J., 1992, Use of alloformations for definition of Holocene meander belts in the middle Amite River, southeastern Louisiana: GSA Bulletin, v. 104, p. 233-41.

Blum, M.D. and D.M. Price, 1998, Quaternary alluvial plain construction in response to glacio-eustatic and climatic controls, Texas Gulf Coastal Plain, *in* Shanley, K.W., and McCabe, P.J., eds., Relative Role of Eustasy, Climate, and Tectonism in Continental Rocks: SEPM, Special Publication 59, p. 31–48.

Blum, M.D. R.S. Toomey, III, and S. Valastro, Jr., 1994, Fluvial response to late Quaternary climatic and environmental change, Edwards Plateau, Texas: Palaeogeography, Palaeoclimatology, Palaeoecology, v. 108, p. 1–21.

Bridge, J.S. and R.S. Tye, 2000, Interpreting the dimensions of ancient fluvial channel bars, channels, and channel belts from wireline logs and cores: AAPG Bulletin, v. 84, p. 1205 – 1228.

Davies-Vollum, K.S. and M.J. Kraus, 2001, A relationship between alluvial backswamps and avulsion cycles; an example from the Willwood Formation of the Bighorn Basin, Wyoming: Sedimentary Geology, v. 140/3-4, p. 235-249.

Holbrook, J.M., 1996, Complex fluvial response to low gradients at maximum regression: A genetic link between smooth sequence-boundary morphology and architecture of overlying sheet sandstone: Journal of Sedimentary Research, v. 66, p. 713–722.

Holbrook, J.M., 2001, Origin, genetic interrelationships, and stratigraphy over the continuum of fluvial channel-form bounding surfaces: An illustration from middle Cretaceous strata, southeastern Colorado: Sedimentary Geology, v. 124, p. 202–246.

Holbrook, J.M. and Wright Dunbar, R., 1992, Depositional history of Lower Cretaceous strata in northeastern New Mexico: Implications for regional tectonics and depositional sequences: GSA Bulletin, v. 104, p. 802–813.

Holbrook, J.M., R.W. Scott, and F.E. Oboh-Ikuenobe, 2006, Base-level buffers and buttresses: A model for upstream versus downstream control on fluvial geometry and architecture within sequences: Journal of Sedimentary Research, v. 76, p. 162-174.

Leeder, M.R. and Mack, G.H., 2007, Basin-fill incision, Rio Grande and Gulf of Corinth rifts: convergent response to climatic and tectonic drivers, *in* G. Nichols, E. Williams, and C. Paola, eds., Sedimentary Processes, Environments, and Basins: A Tribute to Peter Friend: International Association of Sedimentologists, Special Publication 38, p. 9–27.

Oboh-Ikuenobe, F.E., J.M. Holbrook, R.W. Scott, M.J. Evetts, D.G. Benson, S.L. Akins, and L.M. Pratt, 2008, Late Albian-Early Cenomanian flooding history: Southern U.S. Western Interior, *in* B.R. Pratt and C. Holmden (editors) Dynamics of Epeiric Seas: Geological Association of Canada Special Paper 48, St. John's, N.L.

Shanley, K.W. and P.J. McCabe, 1995, Sequence Stratigraphy of Turonian-Santonian strata, Kaiparowits Plateau, southern Utah, U.S.A.: AAPG Memoir 64 p. 103 – 136.

Strong, N. and C. Paola, 2008, Valleys that never were: time surfaces versus stratigraphic surfaces: Journal of Sedimentary Research, v. 78; no. 8; p. 579-593.

Wellner, R.W. and L.R. Bartek, 2003, The effect of sea level, climate, and shelf physiography on the development of incised-valley complexes: A modern example from the East China Sea: Journal of Sedimentary Research, v. 73, p. 926–940.

Wright, V.P. and S.B. Marriott, 1993, The sequence stratigraphy of fluvial depositional systems—the role of floodplain sediment storage: Sedimentary Geology, v. 86, p. 203–210.