

Intra-Reservoir Generation of Organic Acids and Late Stage Enhanced Porosity in Sandstones*

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Abstract

The source of organic acids is commonly attributed to the early maturation of kerogen in source rocks. Organic acids are frequently interpreted as the cause of burial-related secondary porosity. Major deficiencies of this concept are insufficient oxygen in the associated kerogen and likely deactivation of the acids already within the source rocks or during migration. Experimental studies by Eglinton et al. (1987) and Seewald (2001, 2003) proved the generation of significant amounts of water-soluble organic acids (carboxylic acids) by oxidation of n-alkanes. This implies that oxidation of oil compounds and generation of organic acids is likely to occur not only in source rocks but even more effectively along migration paths and in reservoirs. We aim at supplementary experimental work, based on a wealth of field observations and lab data from Permian red-bed reservoir sandstones in the Central European Basin. We want to evaluate the hypothesis that liquid hydrocarbons in hematitic reservoirs can generate reactive organic acids and/or CO₂ during post-emplacement thermal evolution. The expected outcome could allow a better understanding of late stage (syn- and post-oil) porosity enhancement in deep basinal settings with methane source/reservoir potential (tight gas plays).

Session: Sandstone Diagenesis and Reservoir Quality (SEPM)

Intra-Reservoir Generation of Organic Acids and Late Stage Enhanced Porosity in Sandstones

Reinhard Gaupp & Robert Schöner

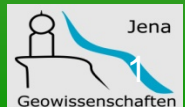
Earth Sciences Friedrich Schiller University Jena / Germany

with contributions by A. Meier, R. Lippmann, E. Kiefer

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A microscopic image of plant tissue, likely a cross-section of a stem or root, stained with a blue dye. The image shows various cellular structures, including large, irregularly shaped cells with thick walls and smaller, more densely packed cells. The blue staining highlights the cell walls and some internal structures, creating a complex, interconnected pattern. The background is a light, off-white color.

Dedicated to

Mingchou Lee

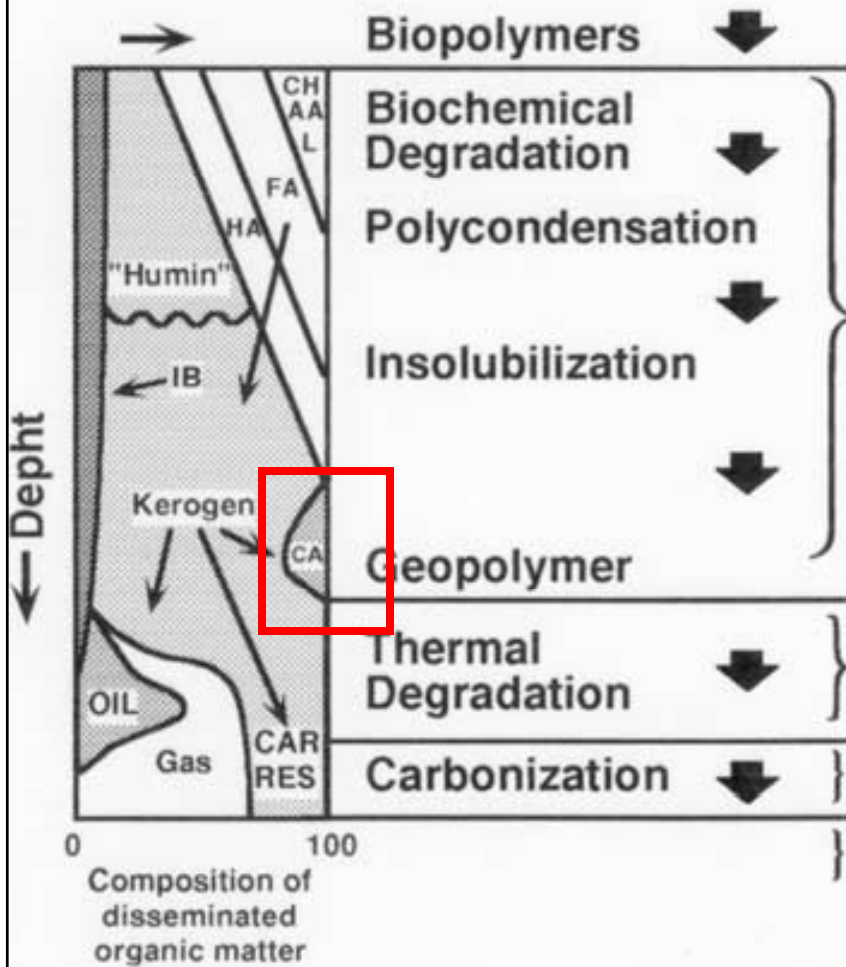
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An old topic?



Diagenesis

$R_0 = 0.5$

Catagenesis

$R_0 = 2$

Metagenesis

$R_0 = 4$

Metamorphism

CH = Carbohydrates

AA = Amino Acids

L = Lipids

FA = Fulvic Acids

HA = Humic Acids

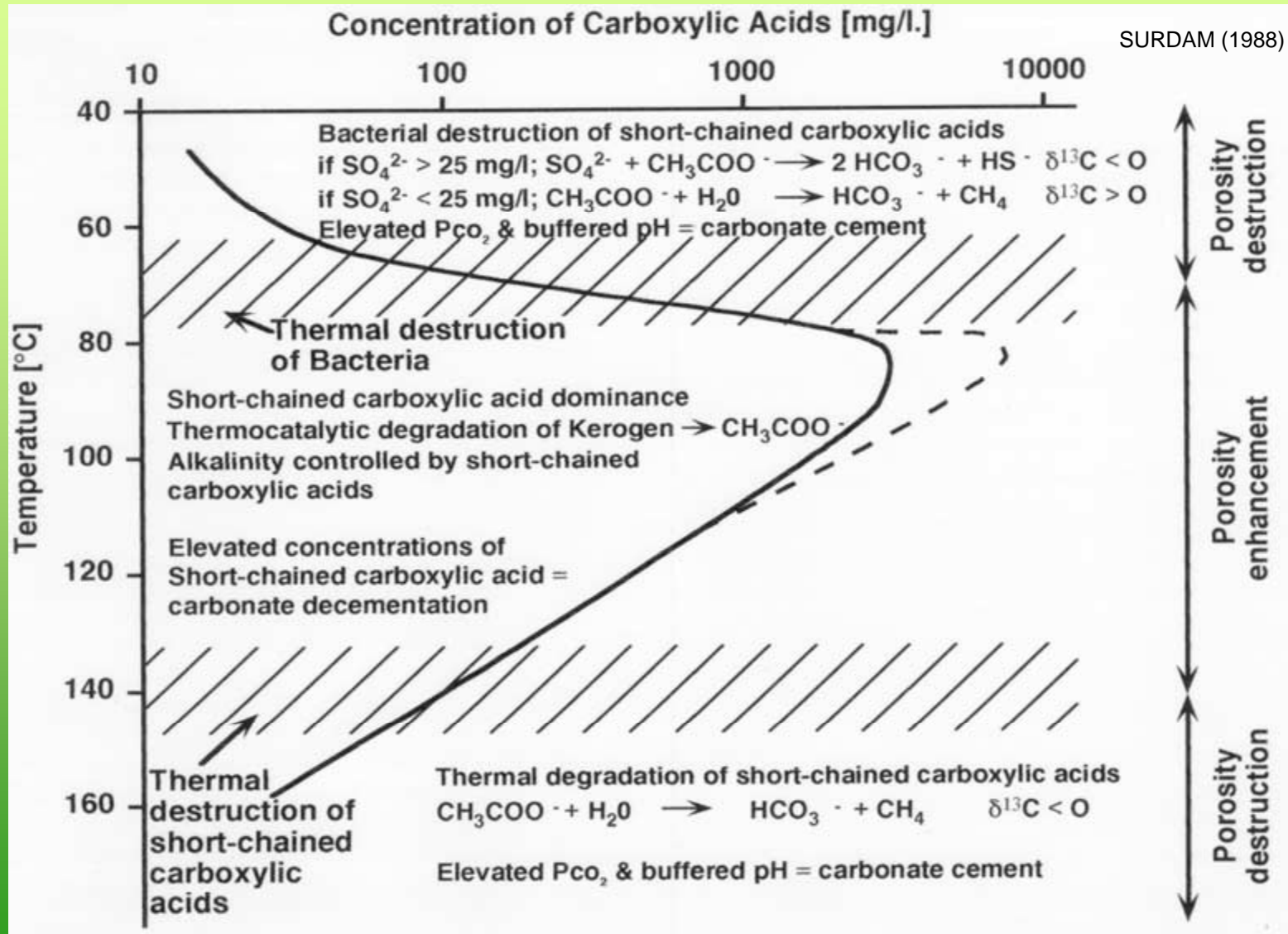
IB = Inherited Bitumen
HC + (N, S, O)

CA = Carboxylic Acids

CAR = Carbon Residue
RES

TISSOT & WELTE (1978)

An old topic?



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Carboxylic acids were considered as cause of secondary porosity (e.g. Surdam et al.1988)

Less attention paid to carboxylic acids since late 80s

Major **deficiency of concept:**

Organic acids form within source rocks and become deactivated by metall-cations during migration

✎ **We predict generation of organic acids within reservoir as reaction product between oil (n-alkanes) and oxidants**

Outline

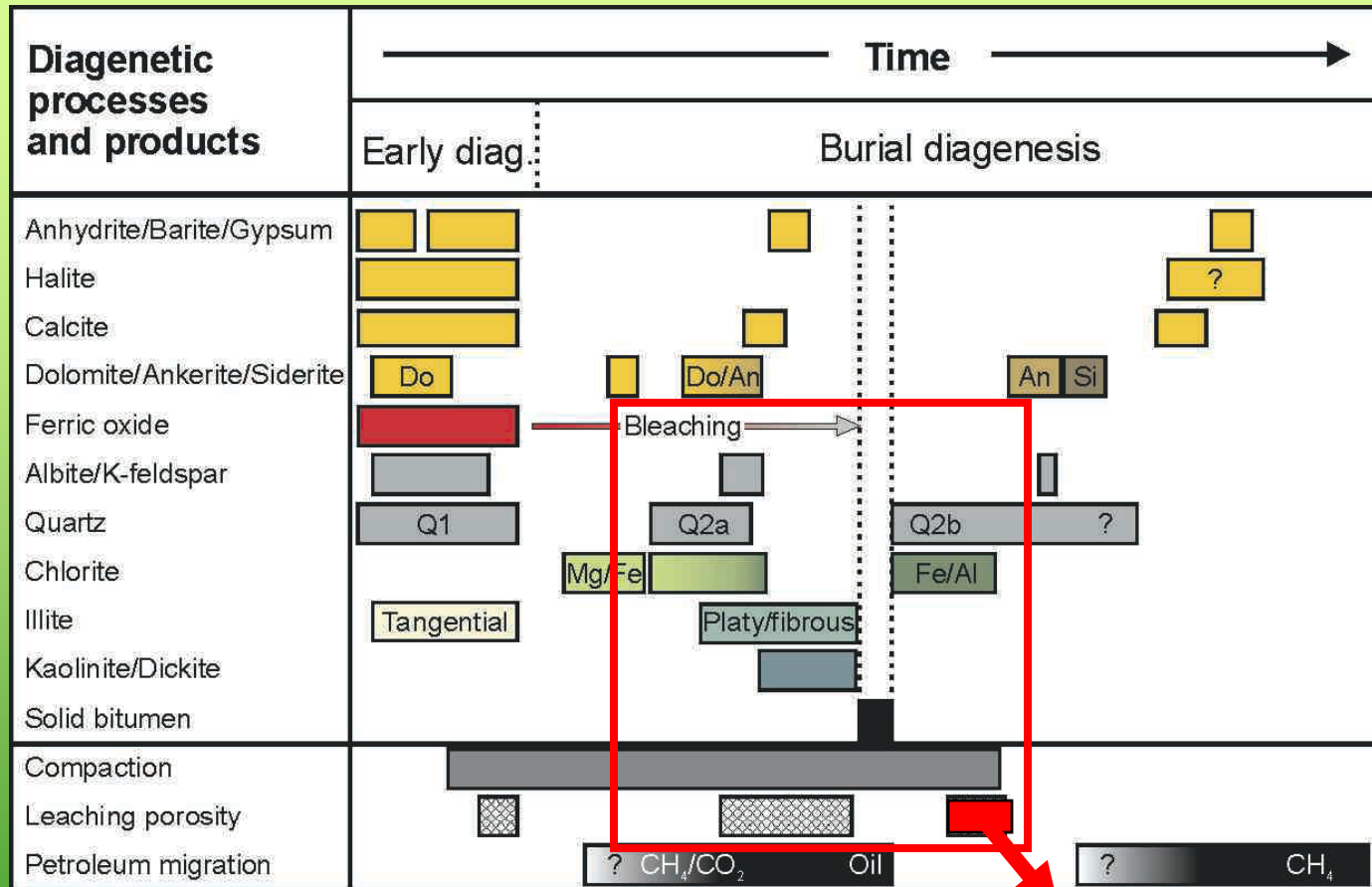
Cases where present gas reservoirs were former oil reservoirs or HC migration pathways

Examples from Permian and Triassic red beds
(North Sea Central Graben, NL on- and offshore, Germany, Poland)

Observations

1. Partial or total bleaching of red beds, clay growth
2. Oil impregnations \neq bitumen staining
3. Feldspar and lithoclast dissolution **after** the hydrocarbon immigration
4. Thermal and maturation modelling indicates $T > 120^{\circ}\text{C}$, K/Ar dating of authigenic illite points to **late burial stage**

Paragenetic sequences



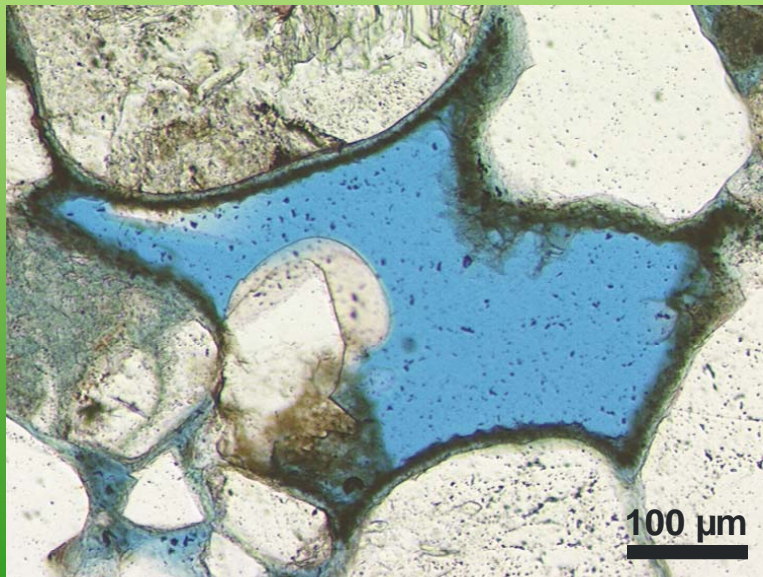
**Late secondary porosity
(enhanced porosity)**

Mineral reactions related to “maturing source rocks”

within the reservoir:

... microscopic scale

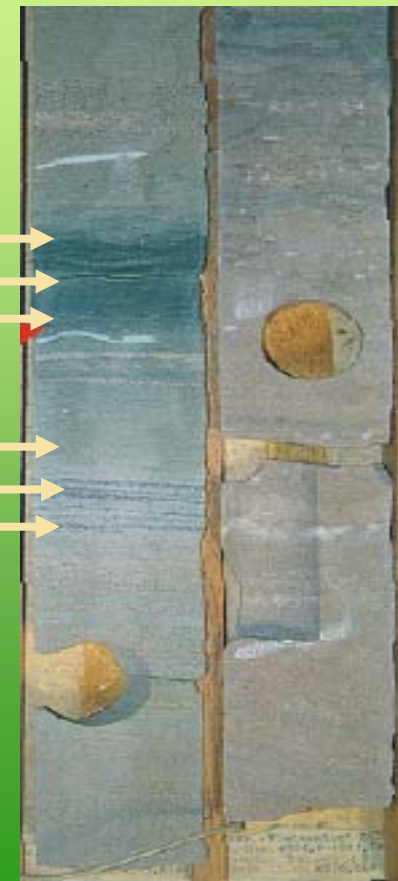
- ⇒ bleaching
- ⇒ kaolinite, illite, chlorite
- ⇒ bitumen
- ⇒ late secondary porosity



... core scale

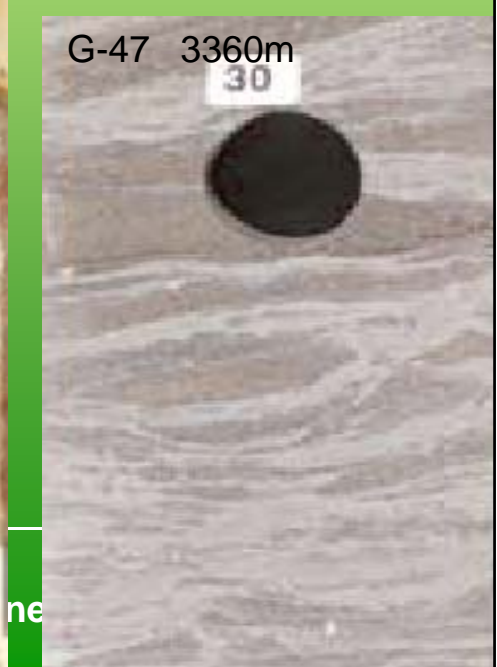
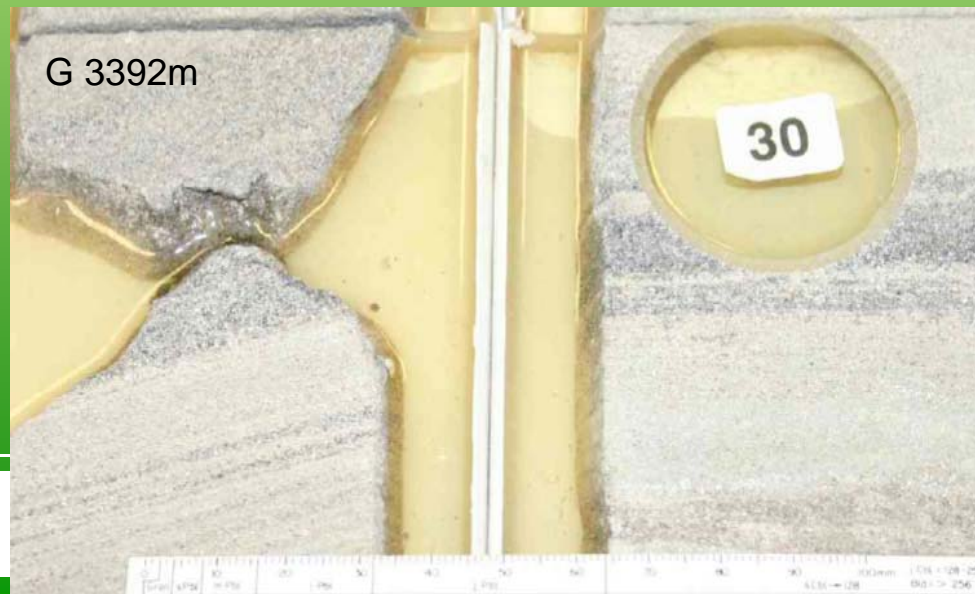
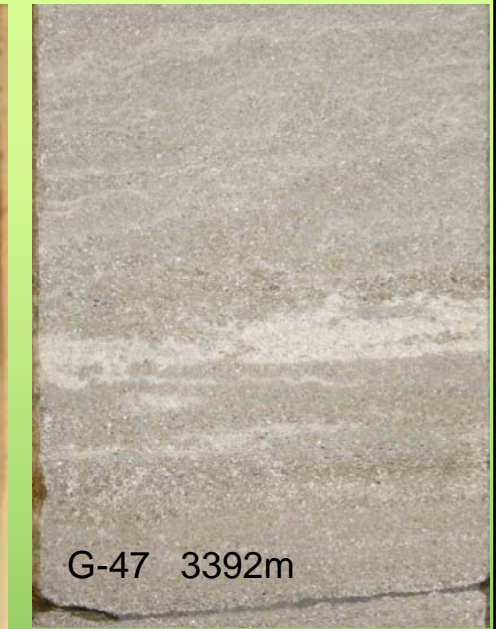
bitumen

bitumen



from Gaupp et al. (2005)

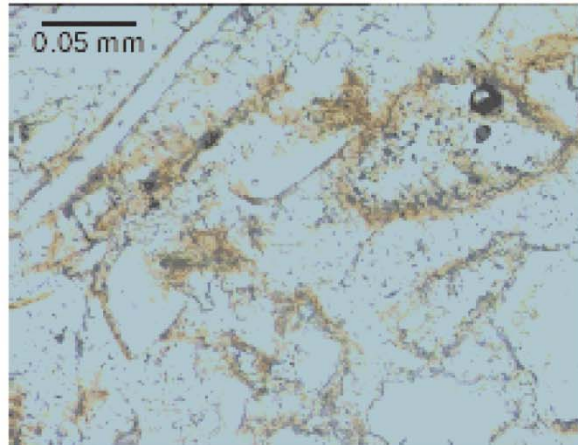
Bleaching in red beds:



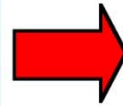
ne

Bleaching and Fe transfer

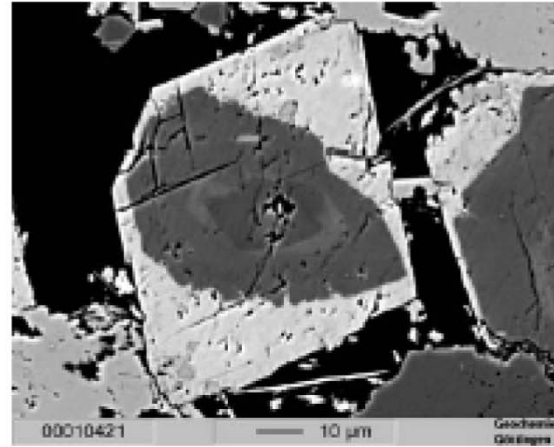
Reduction of Fe-oxides



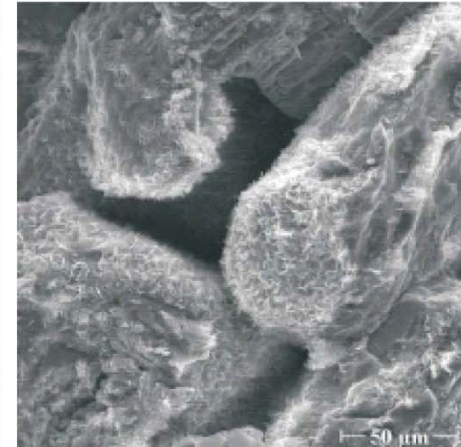
Hematite coatings preserved in a carbonate nodule



Fe(II) bearing mesodiagenetic phases



Fe-carbonate (white) postdates
Nonferrous carbonate (gray)



Euhedral grain-rimming
chlorite

Reduction/bleaching of iron (hematite coats)

✗ authigenic clays ✗ oil immigration

✗ porosity enhancement

Lippmann (2006)

Bleaching

Consequences of Bleaching:

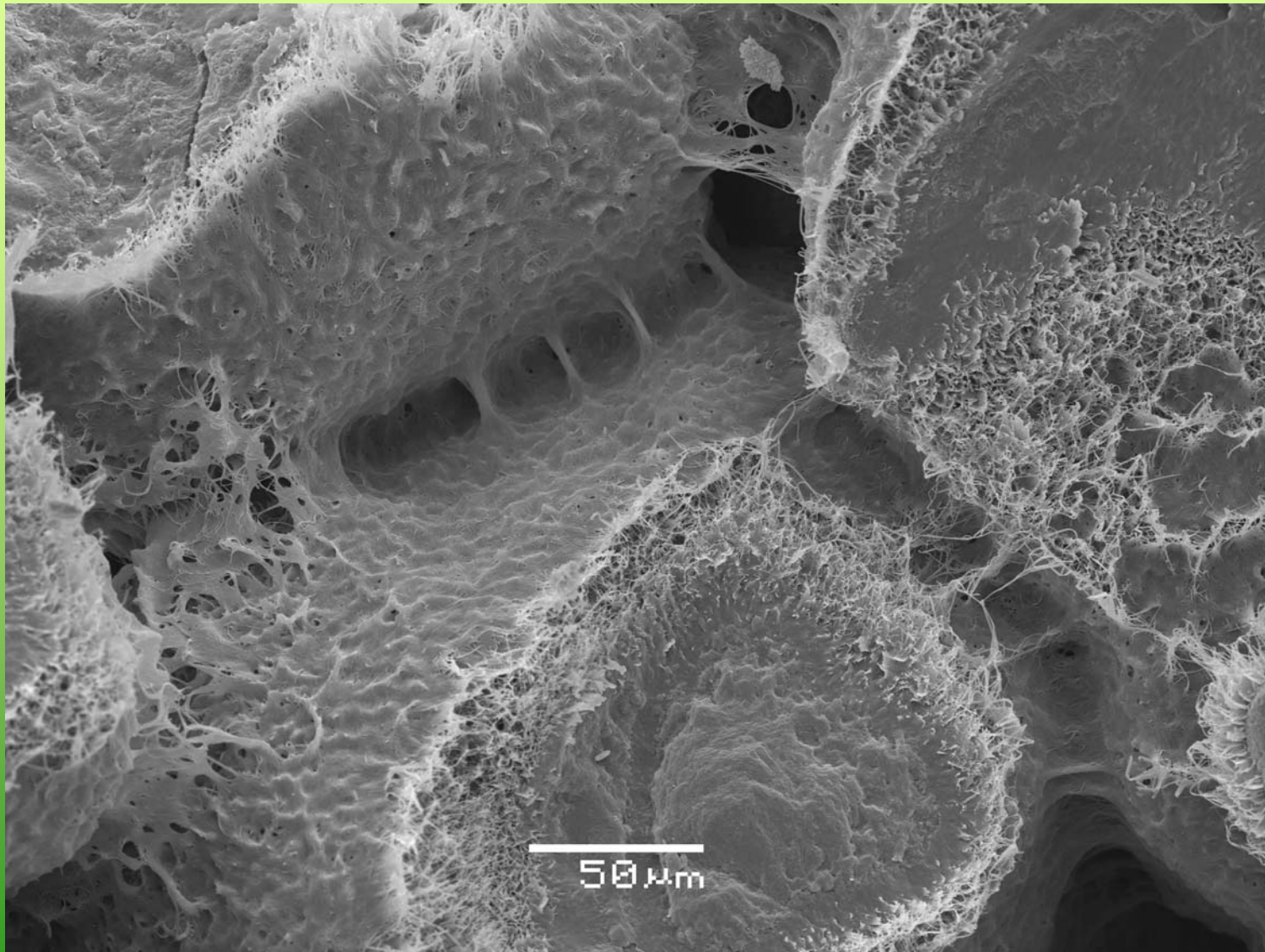
Cleaning of migration pathways and parts of the reservoirs

Feox reduction by hydrolysis occurs with a marked increase in alkalinity and solution iron activity

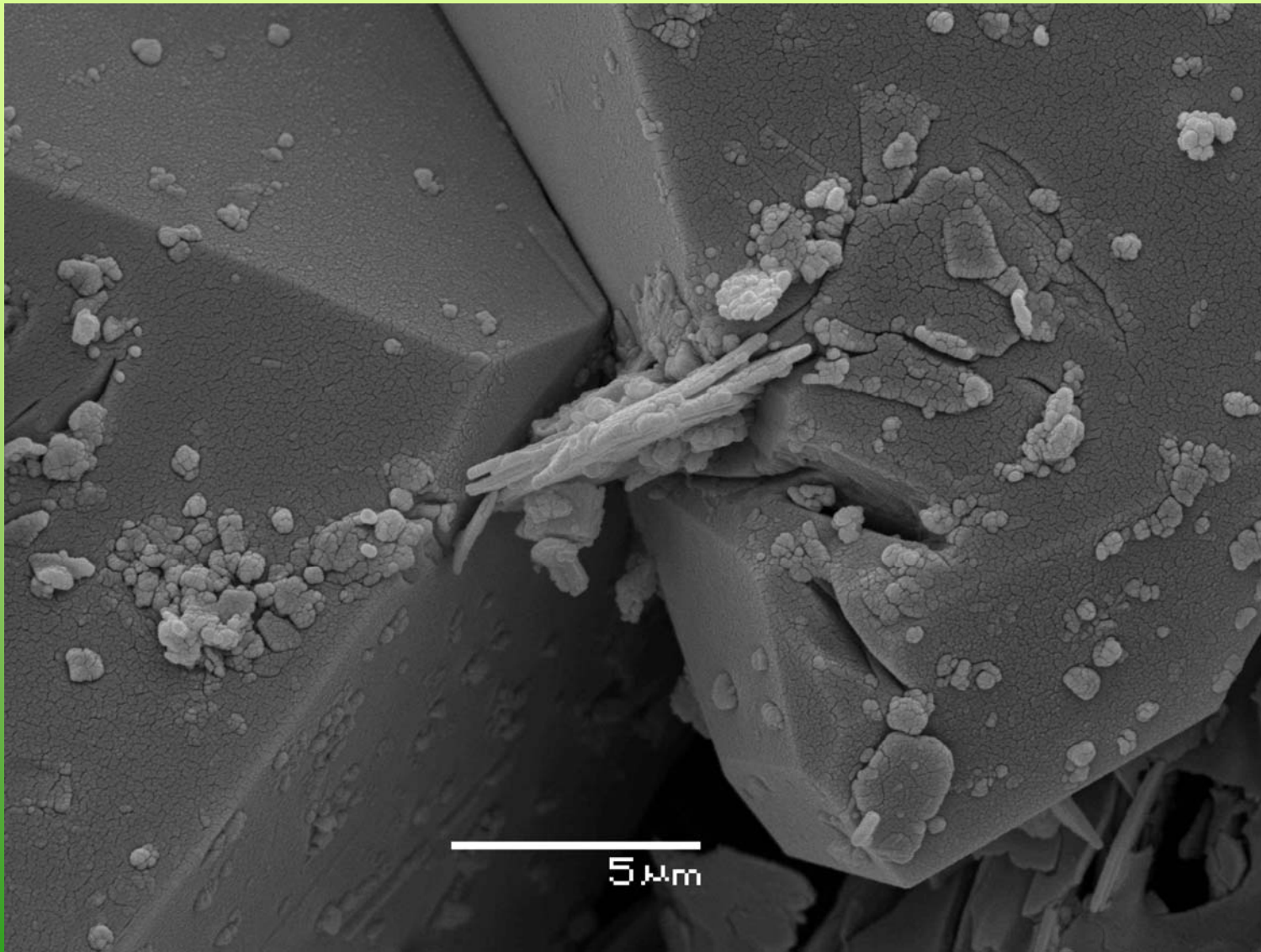
(Curtis 1985)

Fe transferred from oxides/hydroxides to carbonates, sheet silicates (chlorite, Fe-illite)

Fibrous illite with thin bitumen coating



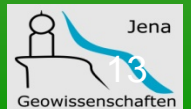
Bitumen coats on authigenic quartz and chlorite



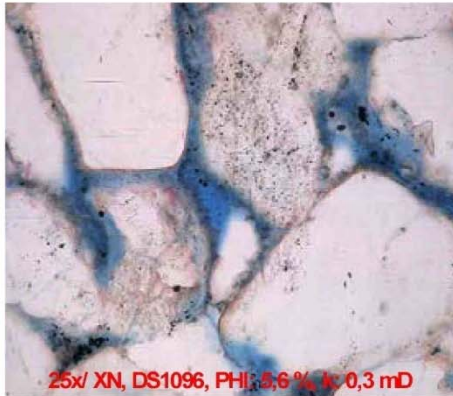
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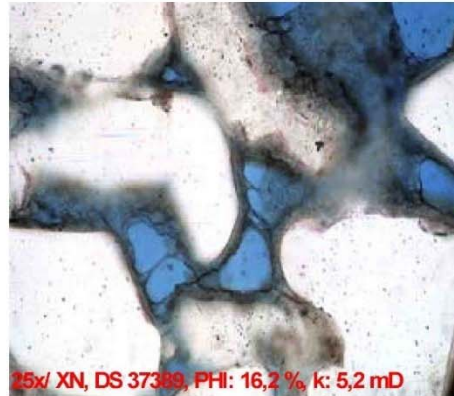
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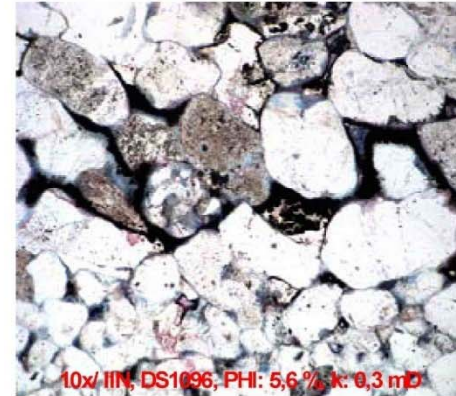
Bitumen types



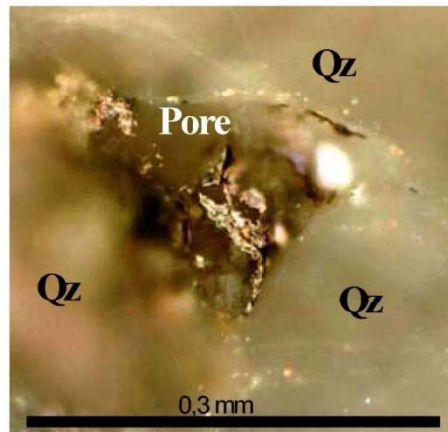
Bitumen type B1: Thin veils



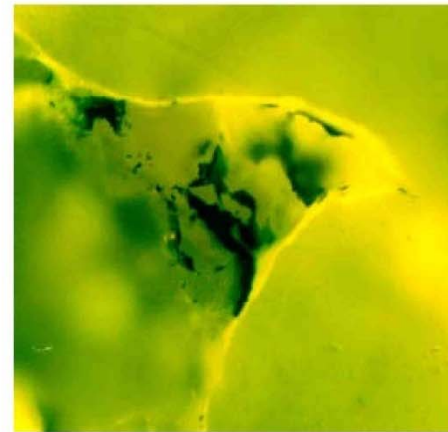
Bitumen type B2: Impregnation by clay minerals



Bitumen type B3: Solid bitumen in pores



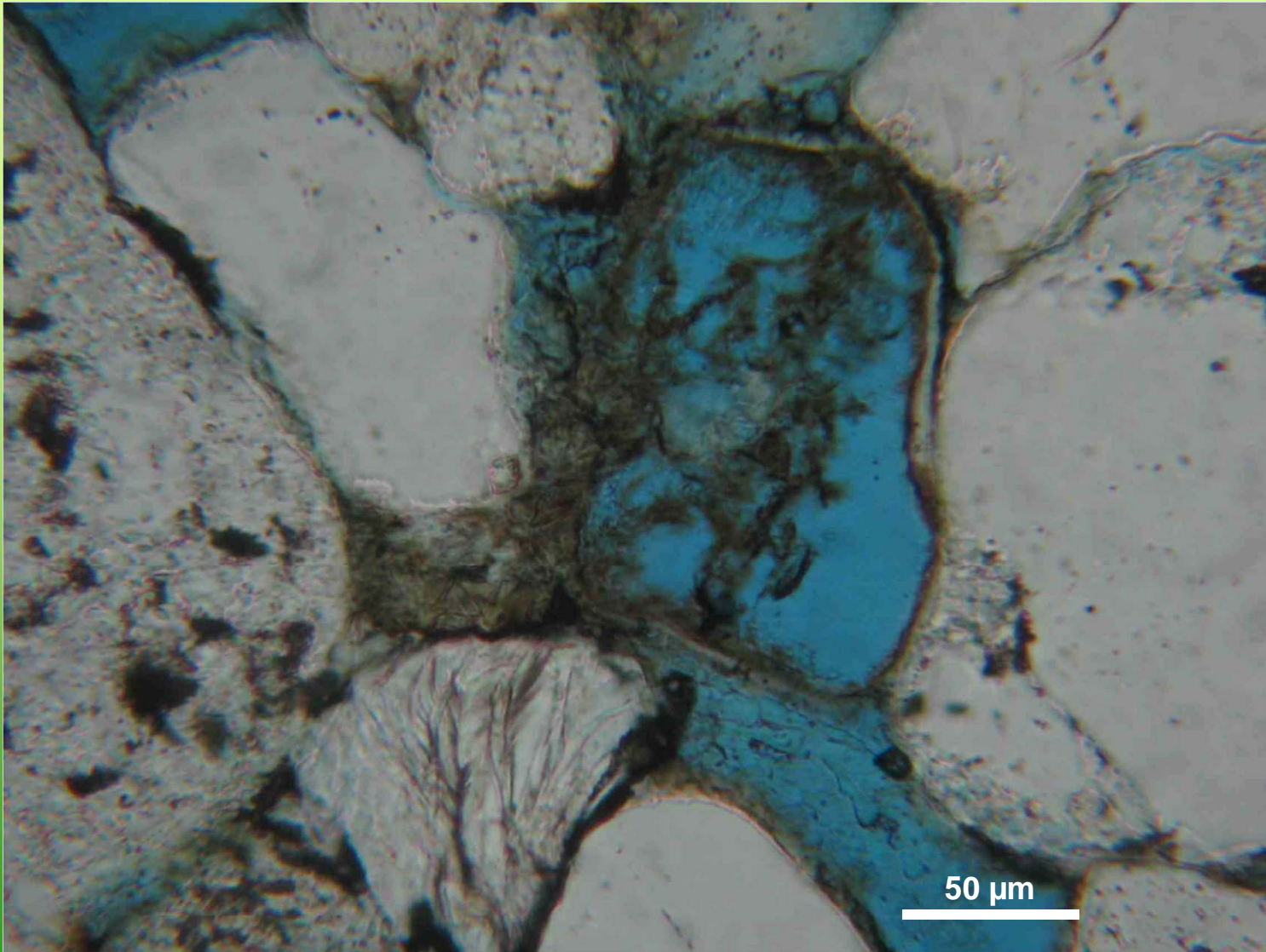
Incident light microscopy



Fluorescence

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Bleaching succeeded by clay and hydrocarbons



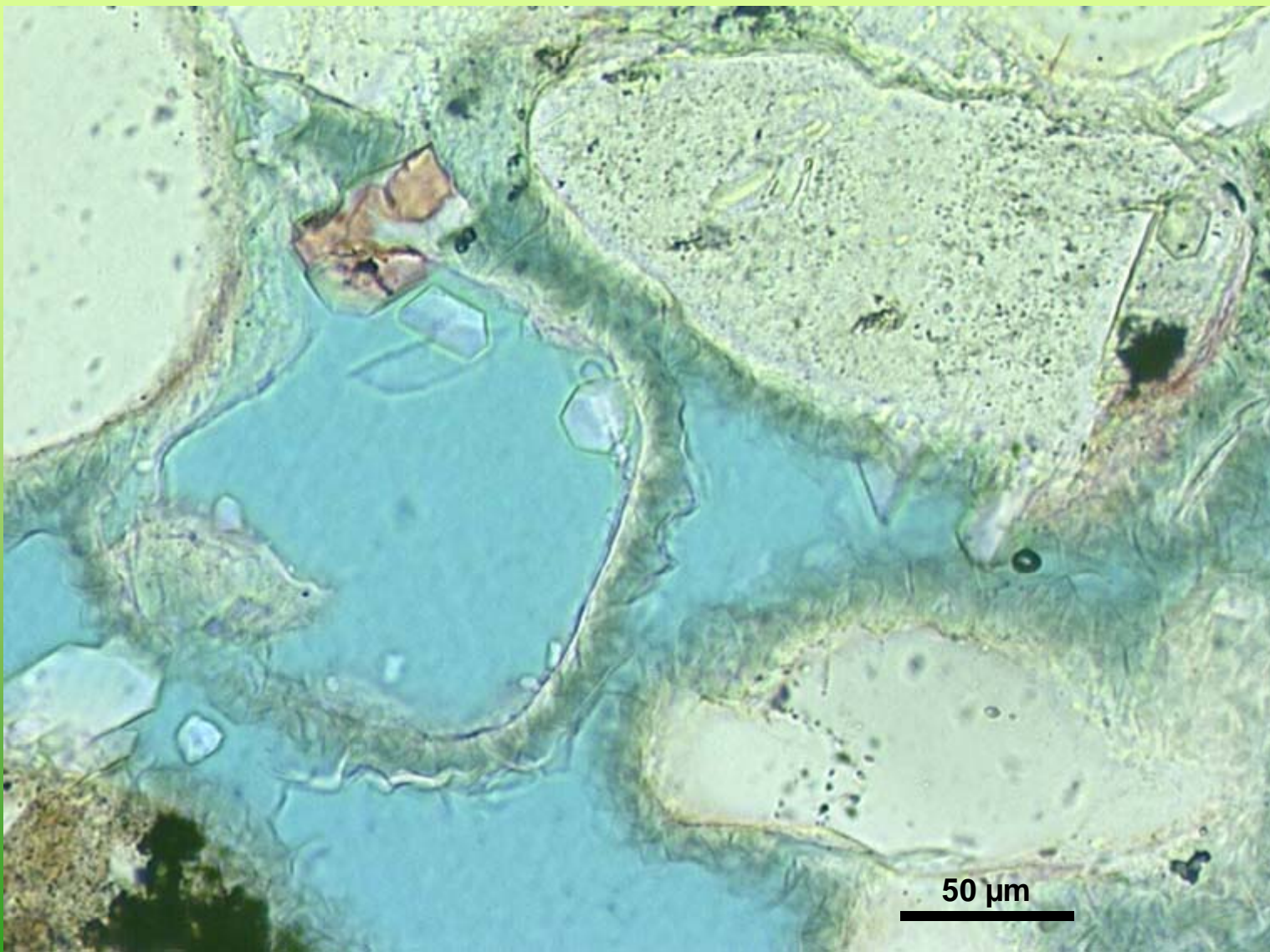
50 μm

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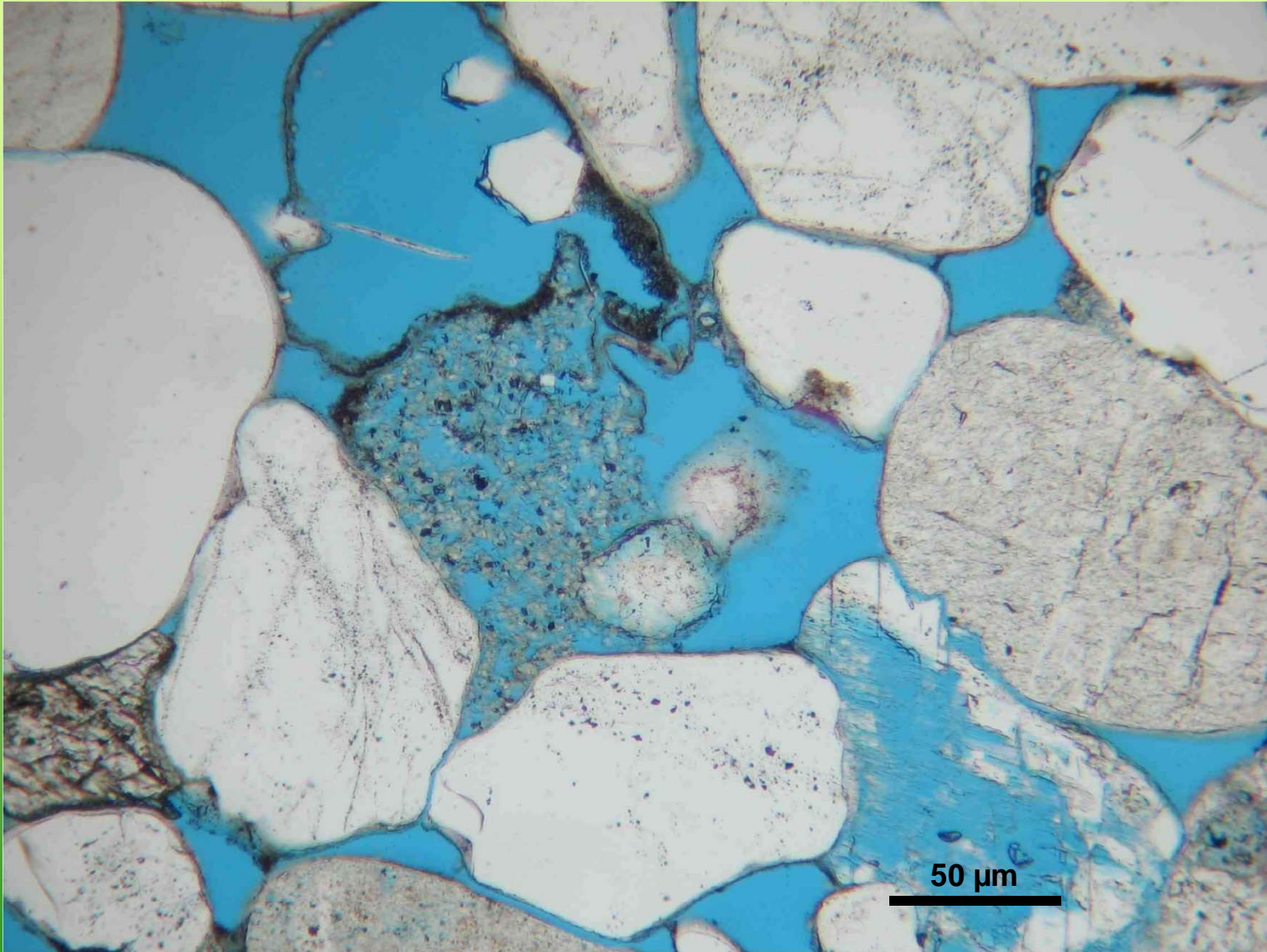
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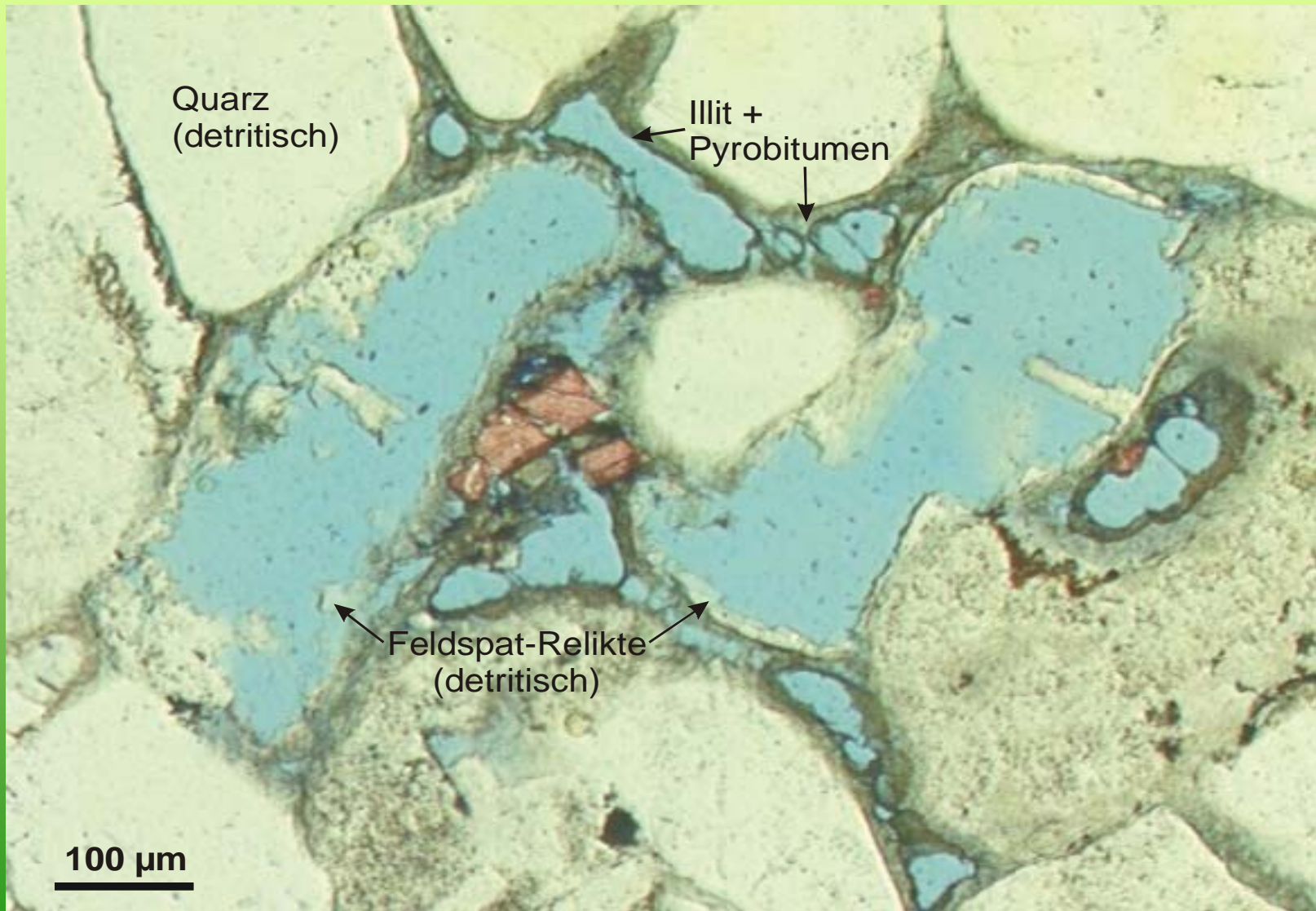
Late dissolution of clasts



Late dissolution of clasts



Late dissolution of clasts



Late dissolution of clasts



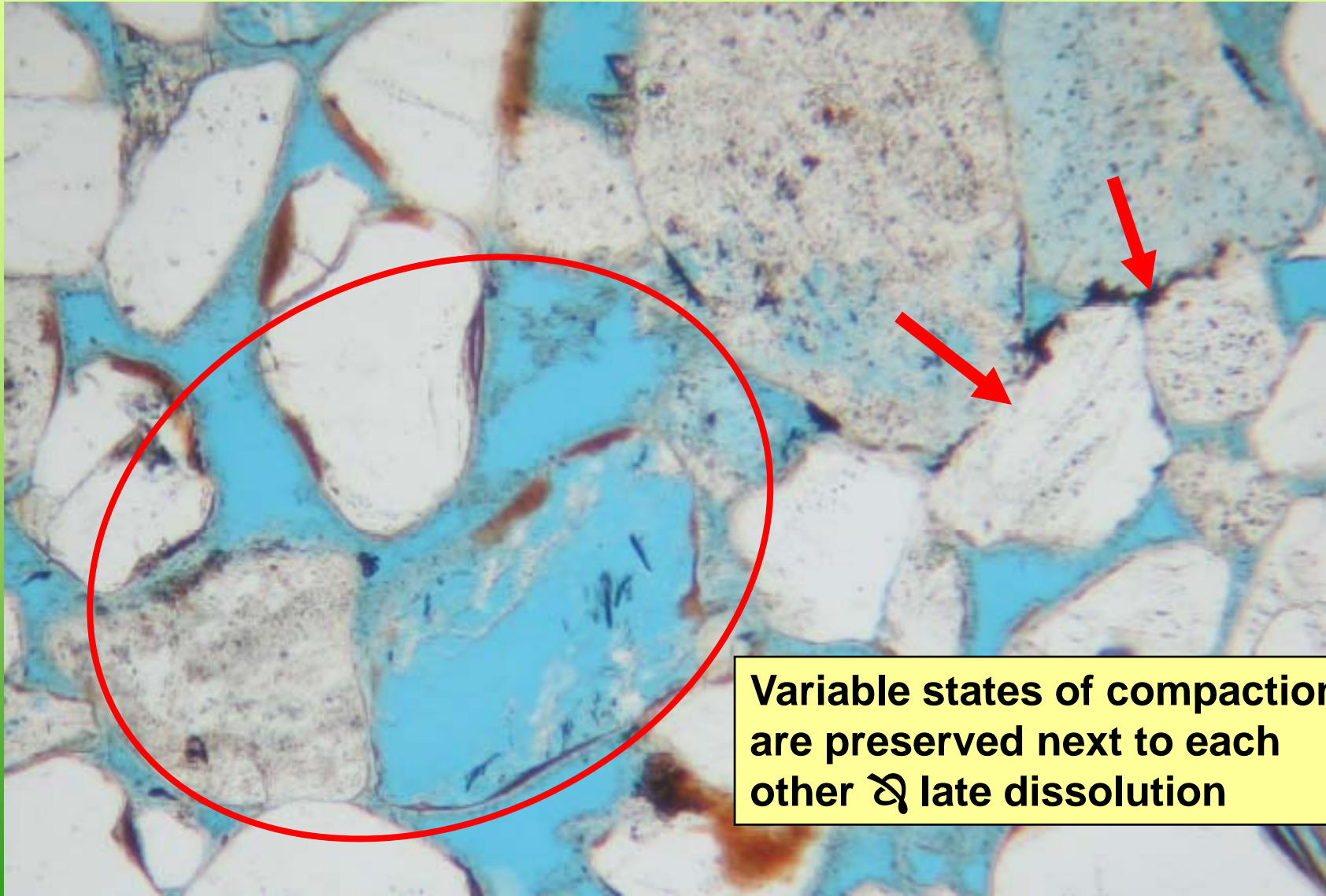
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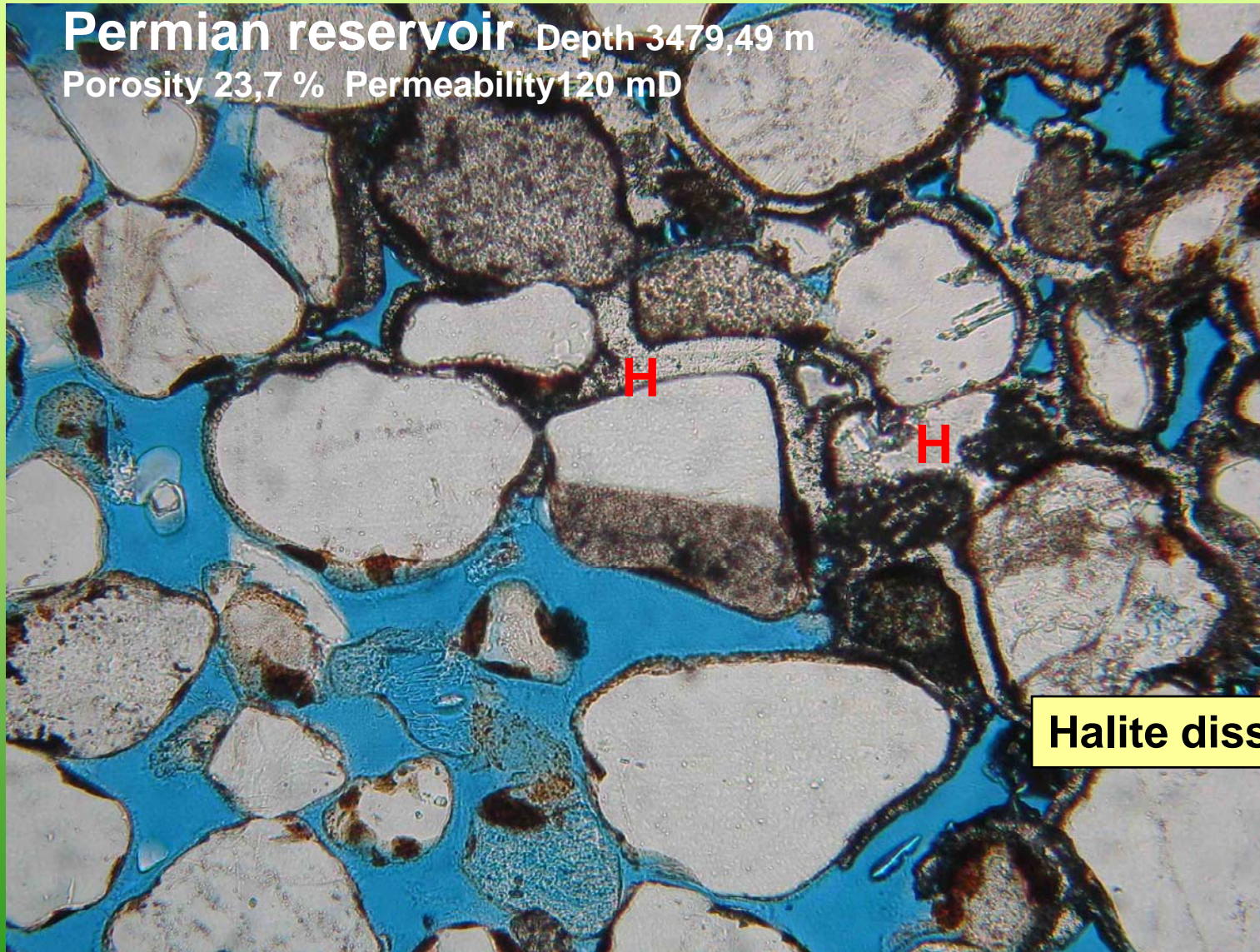
Late dissolution of blocky cements



**Variable states of compaction
are preserved next to each
other \Rightarrow late dissolution**

Late dissolution of blocky cements

Permian reservoir Depth 3479,49 m
Porosity 23,7 % Permeability 120 mD



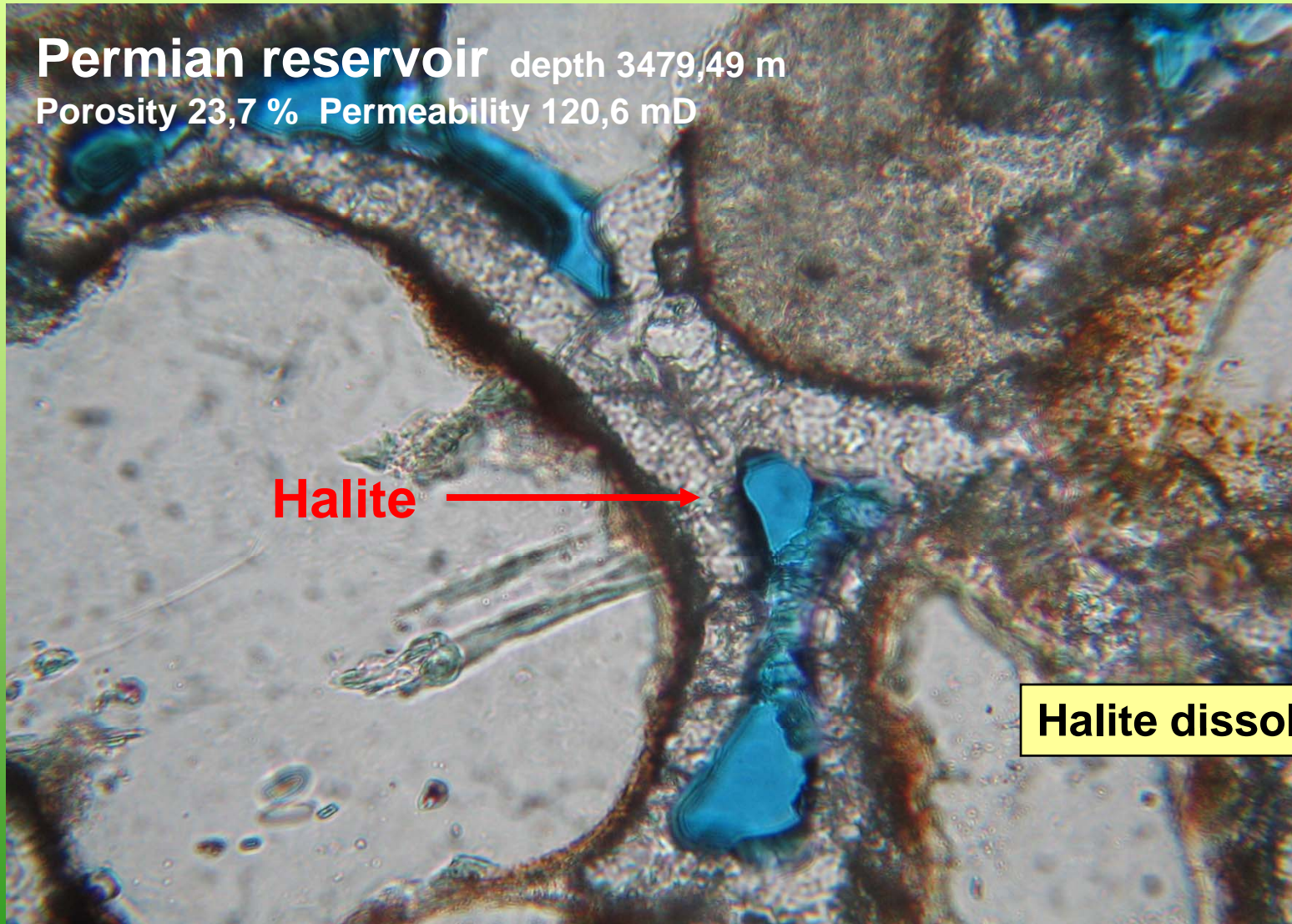
Halite dissolution

250 μm

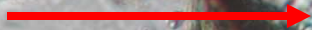
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Late dissolution of blocky cements

Permian reservoir depth 3479,49 m
Porosity 23,7 % Permeability 120,6 mD



Halite



Halite dissolution

50 μ m

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Late stage dissolution of clasts and cements

Dissolution is post-bitumen

In depths of > ca. 2,5km

Clasts: feldspar, volcaniclasts, cherts

Blocky cements: carbonates, anhydrite, halite

Porosity enhancement: + 2 to 15%, averages 3 to 10%

What caused dissolution??

Elements of a Concept

Recurrent patterns:

1. Partial or total bleaching in red beds
2. (First) mesodiagenetic phase of clay growth
3. Oil impregnation (today as bitumen)
4. Post-oil (late stage) **porosity enhancement**



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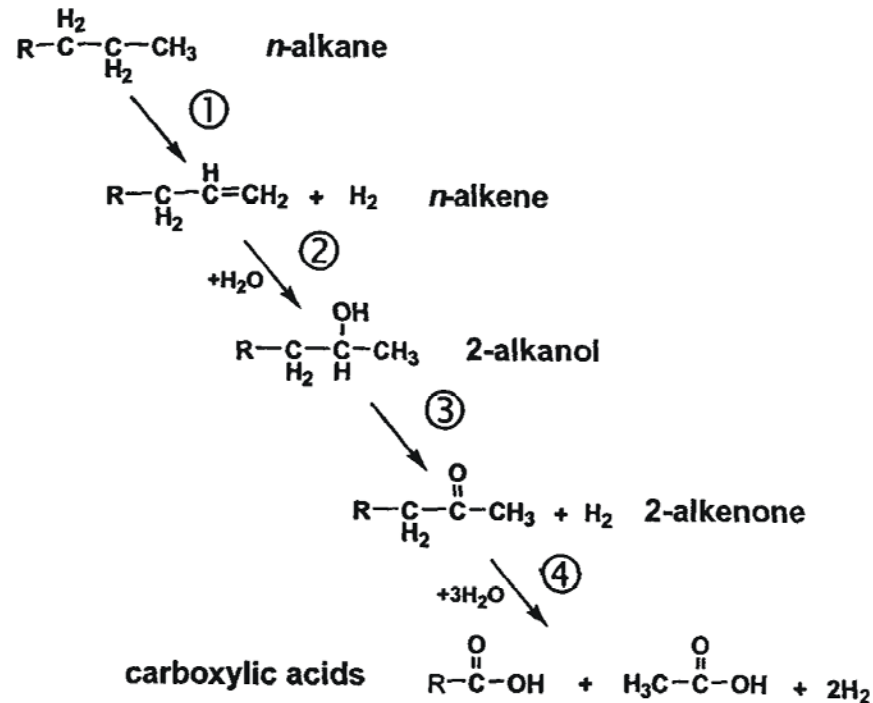
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All these phenomena are missing in reservoirs that had no hydraulic contact to maturing source rocks.

Elements of a Concept

Seewald (2003): *n*-alkanes oxidize stepwise to form carboxylic acids



Seewald (2003) suggests that water and minerals from the host rock are a source of hydrogen and oxygen during organic-inorganic reactions in the pore space of clastic reservoir rocks and that the generation of reductive organic fluids may occur before, during, and after the main stage of petroleum generation.

Elements of a Concept

Thermal alteration of oil in reservoir releases n-alkanes to the porewater

Oxidation of n-alkanes is facilitated by close vicinity of bleached and unbleached facies

Generation of carboxylic acids within reservoirs at $T > 100^{\circ}\text{C}$ allows dissolution of cements and clasts



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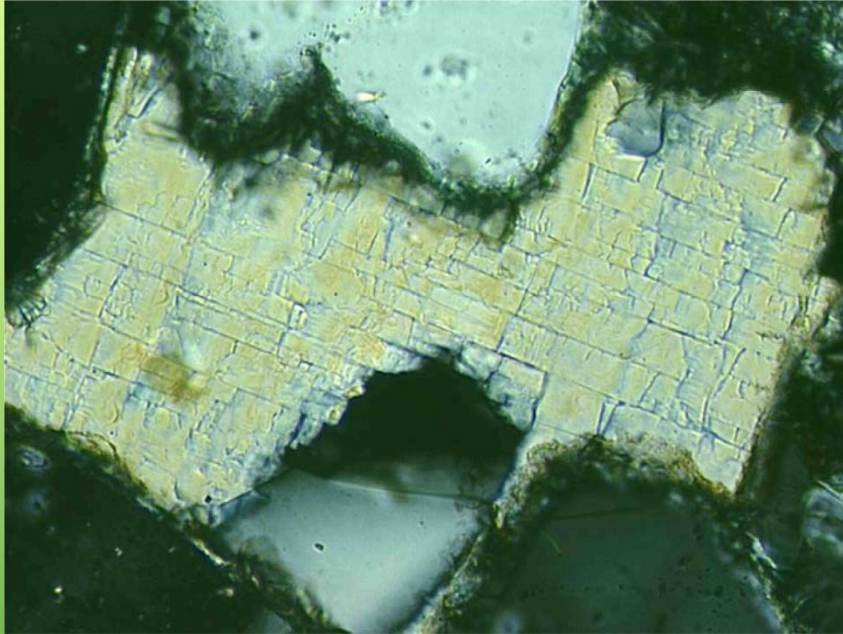
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All these phenomena are missing in reservoirs that had no hydraulic contact to maturing source rocks.

Elements of a Concept

Where is the dissolved material?



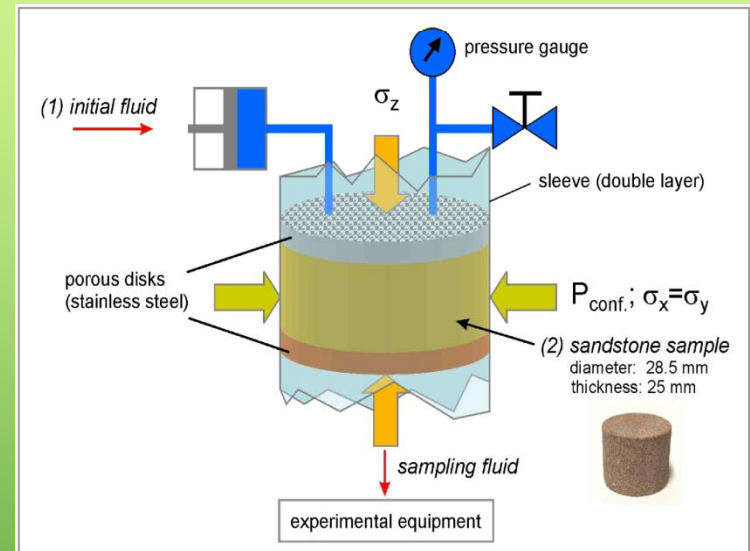
Mass balance calculation show only minor reprecipitation of dissolved material

net plus of stabilized porosity

Experiments

Flow Cell Experiments of organic compounds in clastic reservoir rocks

Meier et al.
(2008)



Summary

1. Recurrent diagenetic phenomena in gas reservoirs, that experienced **former oil** immigration / fill
2. Bleaching–Clay growth–Bitumen and **late Enhanced Porosity**
3. Oxidation of n-alkanes from former oil leads to generation of **carboxylic acids within reservoirs at $T > 100^{\circ}\text{C}$** .
Acids allow dissolution of cements and clasts



Thank you for your attention

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