Click to download movie, synthetic stratigraphy (60mb).

Sequence Stratigraphic Expression of an Evolving Experimental Fluvial-to-Marine Depositional Profile* By John Martin¹, Chris Paola¹, Vitor Abreu², Jack Neal², and Ben Sheets³

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Abstract

Sequence stratigraphy is founded on identifying stratal discordances, a geometrically scale-independent procedure amenable to experimental stratigraphy. Here we perform a sequence stratigraphic analysis on an experimental fluvio-marine deposit for which the boundary conditions and depositional history are known. Our goals are to (1) evaluate when sequence stratigraphic markers are created and how they are structured, and (2) quantify the extent to which the bounded strata honor the known basinal sedimentation patterns. The generation of sequence stratigraphic surfaces is through shifts in sediment mass balance, which is expressed at the surface as an evolving fluvio-marine surface profile. By direct comparison of measured surface topography with preserved stratal surfaces, we show that marine disconformities are especially robust indicators of relative base-level fall. These surfaces, along with maximum flooding surfaces, correlate closely with specific geomorphic surfaces and thus are nearly time synchronous. Erosional surfaces, however, are diachronous and not associated with any instantaneous topographic surface.

Although sequence stratigraphic surfaces are recurrent features in the experimental deposit, their areal correlation and properties are closely related to the shape of the associated stratigraphic cycle, erosion from subsequent cycles, and autogenic overprinting. These effects prohibit a basin-wide correlation of any one horizon. However, by mapping the stacking arrangement of the bounded strata and applying time constraints on the stratigraphic surfaces, we find remarkable agreement between the stratigraphic and known depositional history. This demonstrates sequence stratigraphic horizons accurately capture shifts in the basinal mass balance, and form largely independently of the speed and manner of the changing depositional profile.

Sequence stratigraphic expression of an evolving fluvial-to-marine depositional profile

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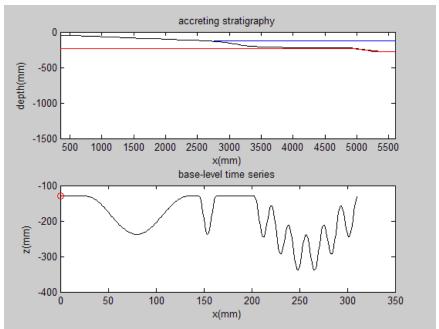


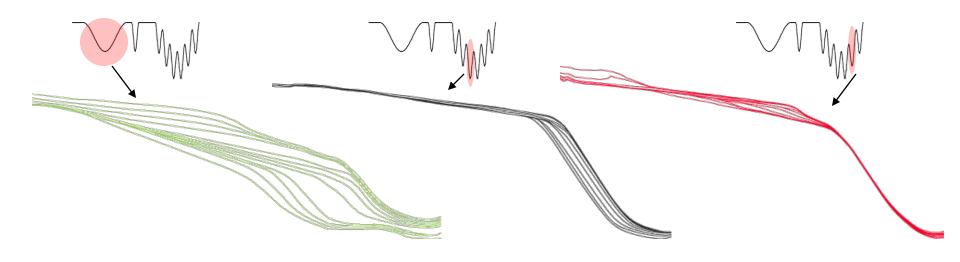
Sequence stratigraphic (time-integrated) imprint of an evolving depositional profile

Creation of stratigraphic disconformities (i.e. time domain of key stratal surfaces)

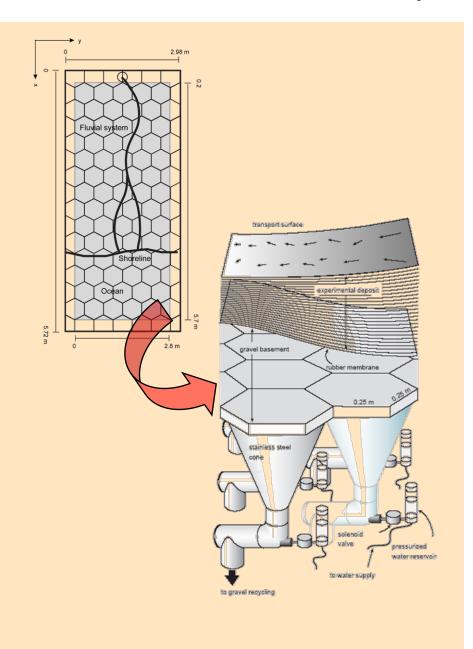
Spatial association between disconformities and paleo-topographic surfaces

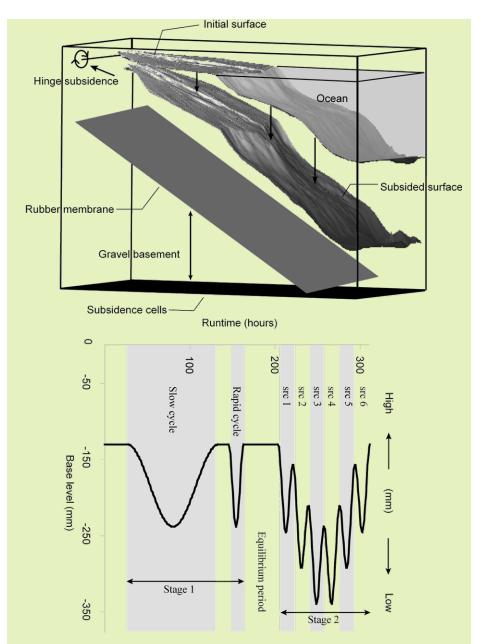
The degree to which the arrangement of sequence stratigraphic surfaces capture the depositional history



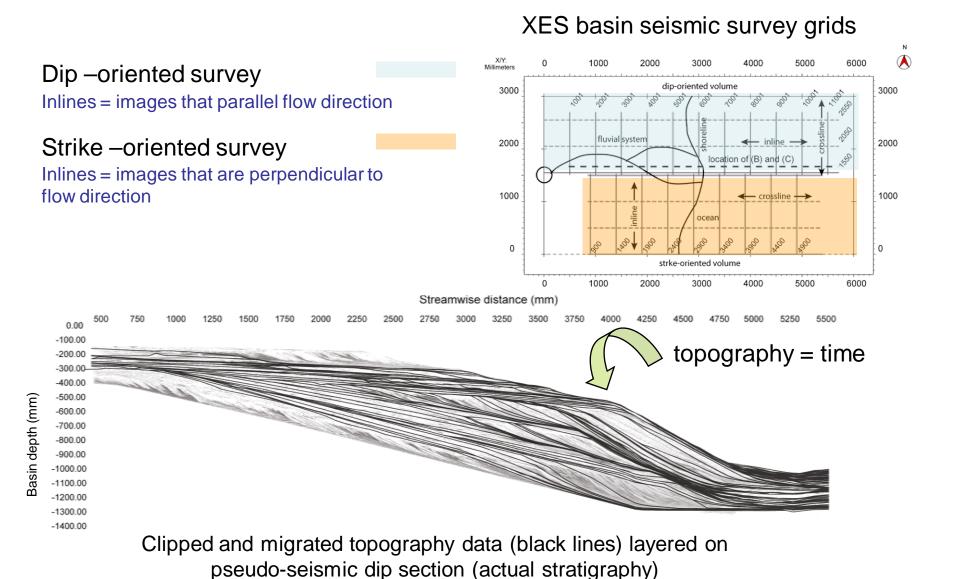


XES Basin and XES 02 experiment design

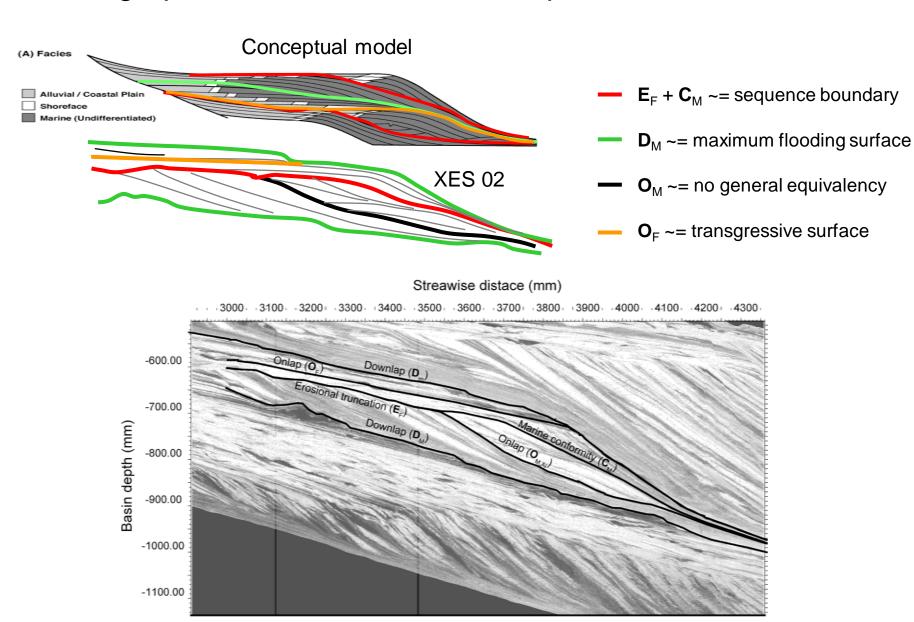




Stratigraphic methods: data integration



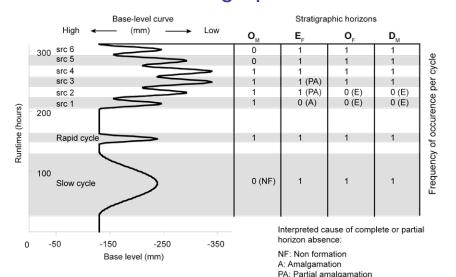
Stratigraphic methods: horizon interpretation



Stratigraphic methods: horizon interpretation

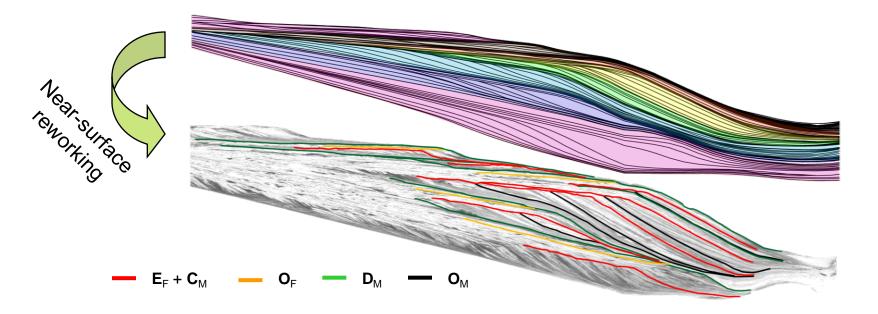
E: Erosion

Full suite of stratigraphic horizons is not always present for each cycle



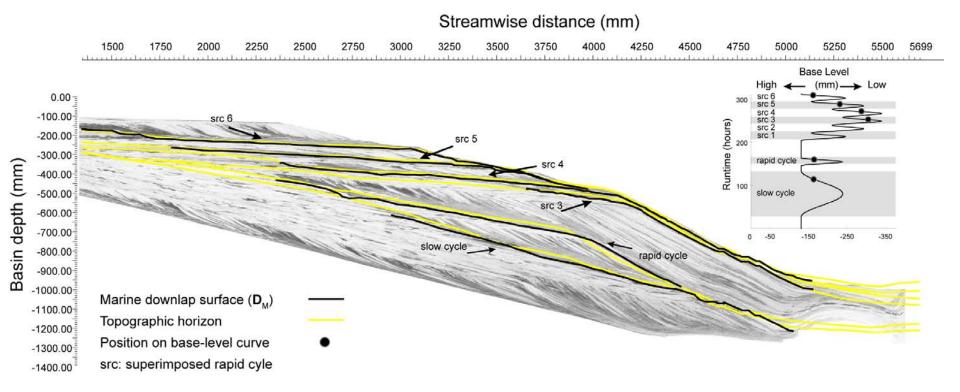
Destruction of original topography: Difficult to state unequivocally that 8 stratigraphic cycles are present

Scale independent form of complexity: Inevitably, the stratigraphy is an incomplete and composite record of deposition



Stratigraphic surfaces: time significance: **D**_M

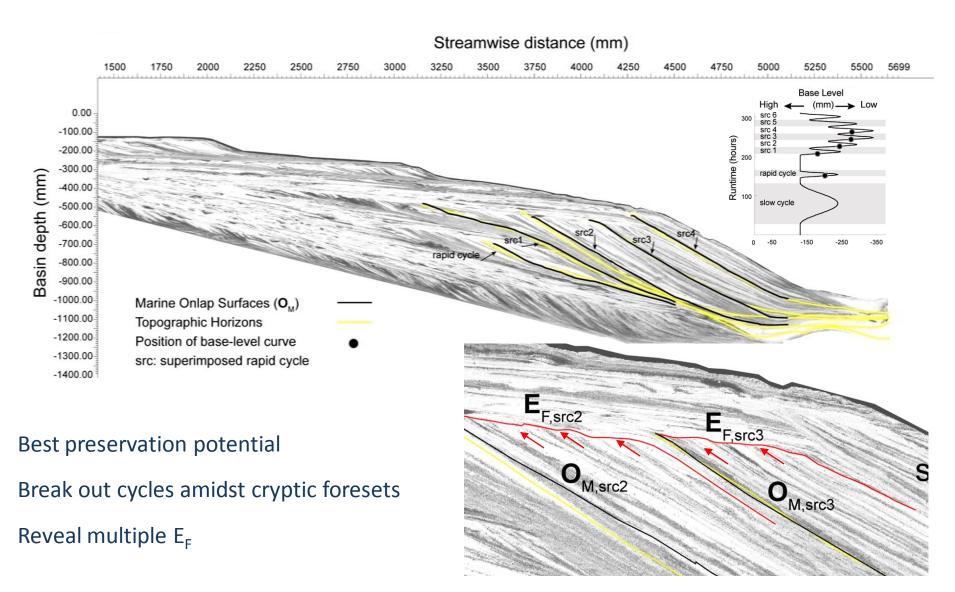
(Chronostratigraphic significance is proportional to geomorphic significance)



Closely correlate with topography around inflection point of base-level rise Consistent despite unique base-level forcing scenarios

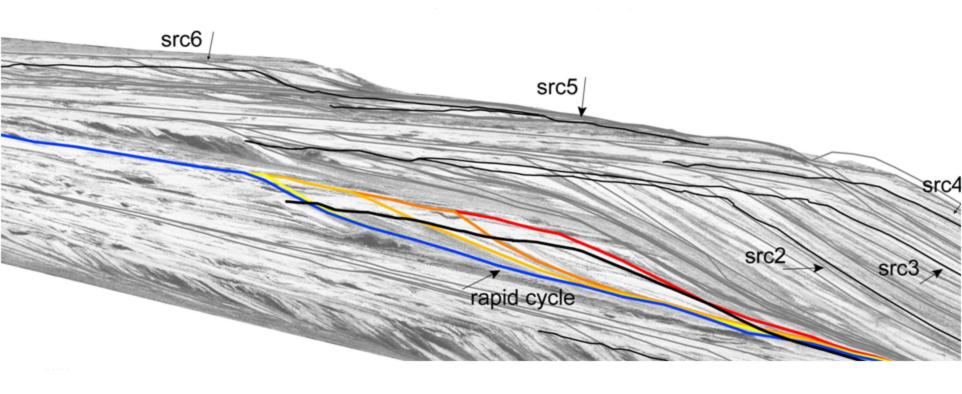
Stratigraphic surfaces: time significance: O_M

(Chronostratigraphic significance is proportional to geomorphic significance)



Stratigraphic surfaces: time significance: E_F

(Chronostratigraphic significance is proportional to geomorphic significance)



To topographic resolution, E_F spans the base-level fall period

At base-level minimum E_F below topography (by up to 3 channel depths!)

Deep erosion etches E_F locally

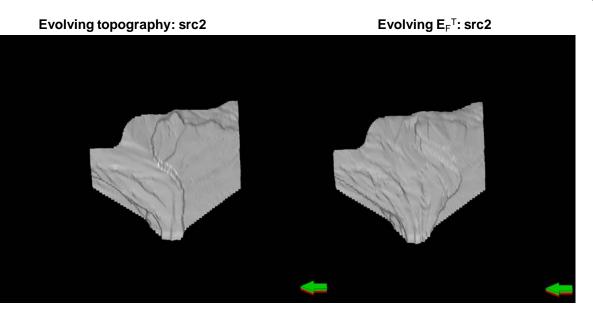
No clear geomorphic equivalency

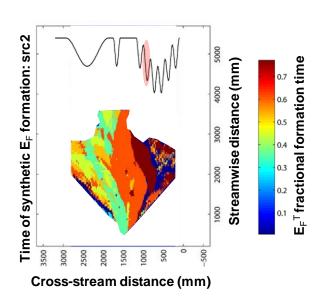
Stratigraphic surfaces: time significance: E_F

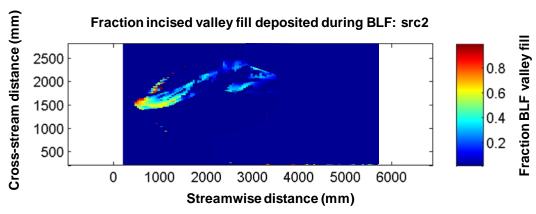
Quantify geomorphic significance of **E**_F using topography data

 $\mathbf{E}_{\mathsf{F}}^{\mathsf{T}}(t) \neq \text{topography } (t)$

E_F^T forms over entire base-level cycle

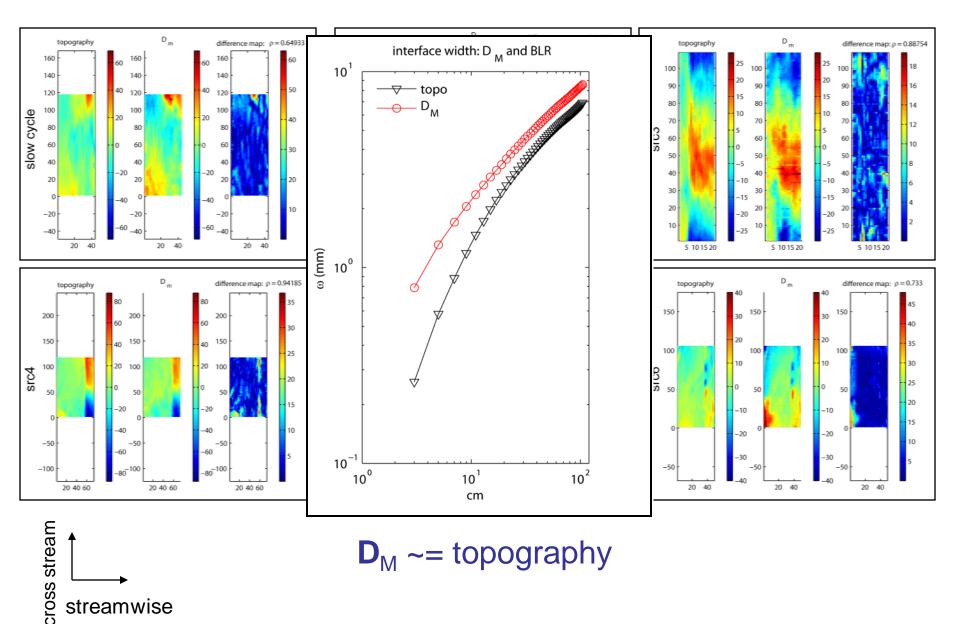






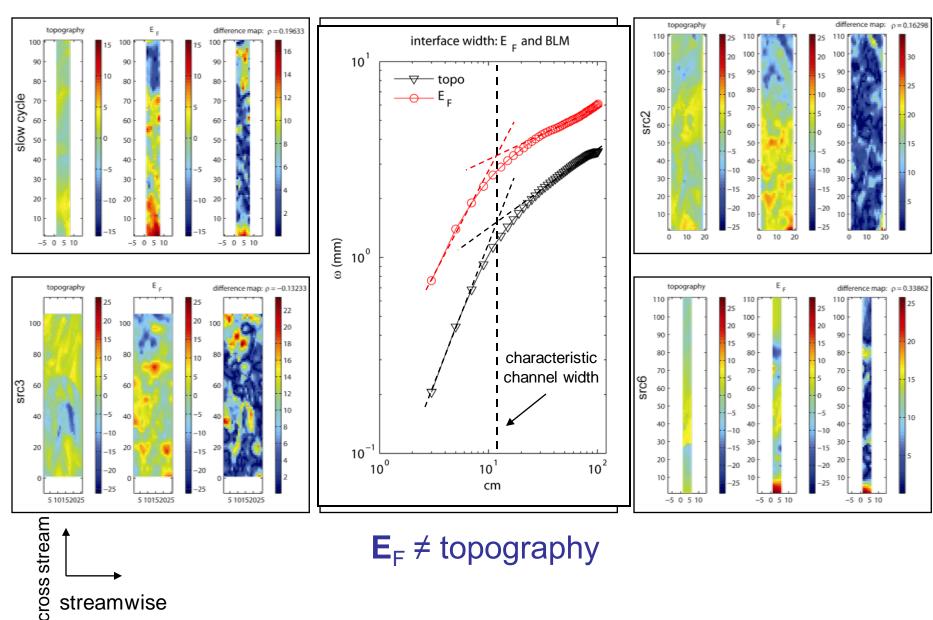
Stratigraphic surfaces: spatial structure: **D**_M

streamwise



Stratigraphic surfaces: spatial structure: **E**_F

streamwise



E_F ≠ topography

Sediment mass migration

XES02 stratigraphic surfaces result from and are primary indicators of shifts in the basinal mass balance

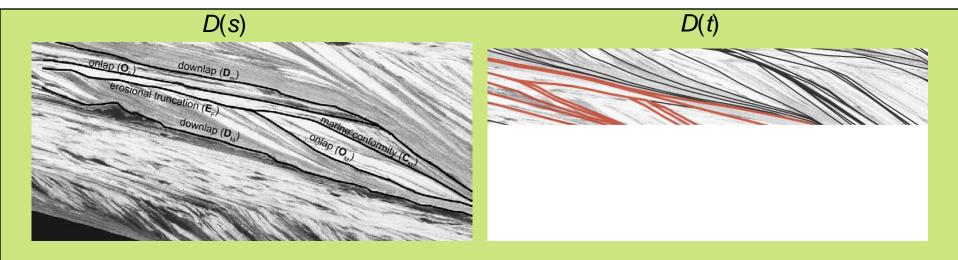
How well does the bounding strata stacking arrangement honor the known depositional history?

$$\overline{x} = \frac{1}{v} \int_{R} x \delta(x, y) dA$$

$$D(s), D(t) = \text{centroid} =$$

$$\overline{y} = \frac{1}{v} \int_{R} y \delta(x, y) dA$$

Average depositional position



C(src 6) ←B'(src 6) 300 - src 6 ← A'(src 6) Sediment mass migration - C(src 5) src 5 ← B'(src 5) A'(src 5) D (src 4) src 4 ← B(src 4) 250 src 3 B(src 3) src 2 B(src 2) A"(src 2)src 1 Runtime (hours) 200 equilibrium period rapid cycle (rc) <--- A(rc) 100 Shoreline slow cycle (sc) Continuous sediment mass migration Stratal package location 50 Time constraint Onlap (O_F) 1500 2000 2500 3000 3500 4000 4500 5000 5500 Erosional truncation (E, Streamwise distance (mm)

Key results

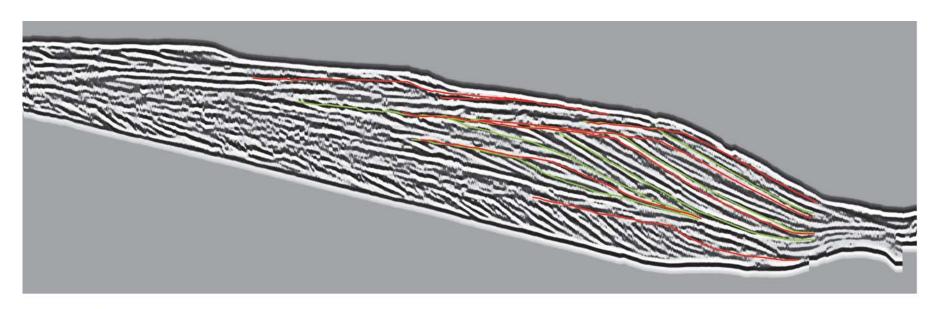
Mass balance effects from variable base level produce similar stratigraphic discordances at field and experimental scales that permit discretization of deposition

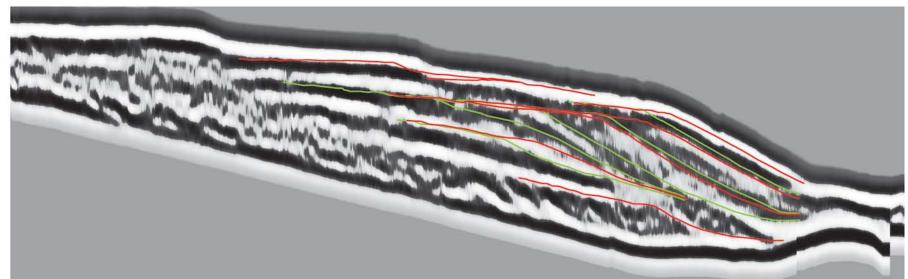
 \mathbf{O}_{M} and \mathbf{D}_{M} demonstrate closest time equivalence with the deposit surface

 O_M and D_M Are robust indicators of base-level fall and rise, respectively

E_F has no clear geomorphic or absolute chronostratigraphic significance

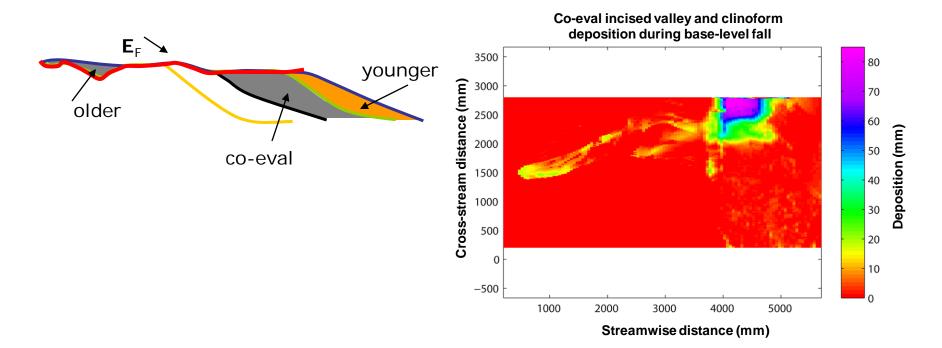
Stacking arrangement of strata honors the known preserved depositional history





Stratigraphic surfaces: time significance: E_F

(Chronostratigraphic significance is proportional to geomorphic significance)



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