PSStratigraphic Correlation of the Cherry Canyon and Bell Canyon Formations (Guadalupian, Middle Permian), Southern Delaware Mountains, Texas Using Fusulinid, Conodont, and Bentonite Data*

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Abstract

Five stratigraphic sections of the Cherry Canyon / Bell Canyon (Guadalupian) transition have been studied in basinal strata of the southernmost part of the Delaware Mountains, West Texas. From west to east, the sections are named WP, WT, ERC, SC and WCC, which range from 70 m to 162 m thick, and are located 40 km northeast of Van Horn and 5 km north-northwest of Texas FM Road 2185. The strata exposed are siltstone, fine-grained sandstone, limestone, and sparse bentonite. Many of the carbonate intervals are thin fining upward debris flows. Three prominent debris flows can be mapped across the area, from oldest to youngest: C-debris, B-debris, and A-debris. The C-debris, approximately 12 m thick, is present in all sections except in the easternmost WCC section. The B-debris, about 60-65 meters above the C-debris, is a mega-conglomerate from 5-15 m thick, and caps the ridges in the area. It is exposed on Texas FM Road 2185 and, based on conodonts, correlates to the basal interval of the Pinery Member, Bell Canyon Formation. The fusulinid, *Polydiexodina* first appears in beds approximately 50 m below the B-debris and approximately 10 m above the last occurrence of *Parafusulina*. *Polydiexodina* first appears in the Hegler Member of the Bell Canyon Formation and is known from the Pinery, Rader and McCombs Members of the Bell Canyon Formation in the Guadalupe Mountains. The A-debris at 2-5 m thick is only present in the WCC section, but can be traced south, southeast to Texas FM 2185. Conodont faunas in the sections are dominated by *Jinogondolella aseratta*. The WT, SC, and WCC sections have three thin, closely spaced distinct bentonite beds approximately 26-29 m above the C-debris and below the B-debris. Preliminary U/Pb radiometric dating suggests age equivalency to bentonites in the Manzanita Limestone Member, Cherry Canyon Formation, Guadalupe Mountains.

References Cited

Dutton, S., M. Holtz, T. Tremblay, and H. Zirczy, 2000, Expansion of Gas Reservoir Data Base, Permian Basin, Texas: The Permian Basin: Proving Ground for Tomorrow's Technologies: West Texas Geological Society Publication No. 00-109.

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Lambert, L., B. Wardlaw, and C. Henderson, 2007, Mesogondolella and Jinogondolella (Conodonta): Multielement definition of the taxa that bracket the basal Guadalupian (Middle Permian Series) GSSP: Palaeoworld, v. 16, p. 208-221.

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WP section above showing the C-debris outlined in red, and B-debris capping the hill

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West Texas Middle Permian paleo-map showing major geographic

features and the location of the WP Section (after Dutton et al., 2000)

WCC ve stratigraphic sections of the Cherry Canyon / Bell Canyo A Debris (Guadalupian) transition have been studied in basinal strata of the southernmost part of the Delaware Mountains, West Texas \subseteq Silt to fine-grained sandstone 0 rom west to east the sections are named WP WT SC FRO Silty limestone cated 40 km northeast of Van Horn and 5 km north-northwe Limestone of Texas FM Road 2185. The strata exposed are siltstone, fine **ERC** Covered interval arbonate intervals are thin fining upward debris flows. Thre prominent debris flows can be mapped across the area, from oldest to youngest: C-debris, B-debris, and A-debris. The C-debris, approximately 12 m thick, is present in all sections except 120 n the easternmost WCC section. The B-debris, about 60-65 i soove the C-debris, is a mega-conglomerate from 5-15 m thich and caps the ridges in the area. It is exposed on Texas FM Roa 2185 and, based on conodonts, correlates to the basal interval of he Pinery Member, Bell Canyon Formation. The fusulinic lolydiexodina first appears in beds approximately 50 m below B-debris and approximately 10 m above the last occurrence CA-TIMS U-Pb isotopic ages of Parafusulina, Polydiexodina first appears in the Hegle ember of the Bell Canyon Formation and is also known from he Pinery, Rader and McCombs Members of the Bell Canyo ormation in the Guadalupe Mountains. The A-debris at 2-5 r 10 9797979797979797 acecececececece WP-60 hick is only present in the WCC section, but can be traced south touth-east to Texas FM 2185. Conodont faunas in the section are dominated by *Jinogondolella aseratta*. The WT, SC, and WCC sections have three thin, closely spaced distinct bentonite eds approximately 26-29 m above the C-debris and below the B-debris. Preliminary U/Pb radiometric dating suggests ag uivalency to bentonites in the Manzanita Lime 0 Cherry Canyon Formation, Guadalupe Mountains. $265.20 \pm 0.13 (0.18) [0.33] \pm 2\sigma int.$ ± 0.13 (0.18) [0.33] ± 95% c.i There are several correlative beds in the 5 sections (WP WT SC, EDC, WCC) measured across the southern part of the Delaware Mountains. The major debris flows, C and B are resent in all sections as well as the WP-60 carbonate debris $265.36 \pm 0.086 (0.16) [0.32] \pm 2\sigma int.$ ed. Unfortunately these stratigraphic markers are not known e Guadalupe Mountains region. $\pm 0.086 (0.16) [0.32] \pm 95\% c.i$ The correlation of the transitional strata of the Cherry Canyon and Bell Canyon Formations in the Guadalupe Mountains to uivalent age strata in the southernmost part of the Delaware Ranges of Mountains can be made biostratigraphically by identifying several key microfossils. Conodonts recovered from a number Important Fossils Standard stratigraphic intervals are currently being studied. The only 70 onodont discovered in these sections is Jinogondolella \mathbf{r} Capitaniar elong to the Wordian Stage. Rader Meters The fusulinacean Parafusulina is a key microfossil for the Cherry Canyon Formation and it is abundantly present in these Pinery neasured sections. Parafusulina is present in the strata in and elow the WP-60 carbonate debris bed. The fusulinacean Hegler Polydiexodina first occurs approximately 10 m above the WP-60 Wordia ed in sample W-63. This genus is a key microfossil for the r, Pinery, Rader and McCombs Members of the lower par of the Bell Canyon Formation. he bentonites exposed in the measured sections ha rocessed for zircons. Currently, only the zircons from the WCC ection have been radiometrically dated using CA-TIMS U-Pb nethods to yield dates between 265.36 to 265.17 MY which orrelate to the age of the bentonites exposed in the Manzanita Member of the Cherry Canyon Formation in the Guadalupe Meters Delaware Basin stratigraphic units of the Guadalupian Dutton, S., Holtz, M., Tremblay, T., Zirczy, H., 2000, Expansion o Gas Reservoir Data Base, Permian Basin, Texas. The Permia Series with international GSSP key ages and index conodont P elements for the Capitanian, Wordian and Basin: Proving Ground for Tomorrow's Technologies: West Roadian Stages (modified from Lambert et al., 2007). ambert I. Wardlaw B. Henderson C. 2007 Mesogondolella ambert, L., Wardiaw, B., Fienderson, C., 2007, Mesogonoleidand Jinogondolella (Conodonta): Multielement definition of the taxa that bracket the basal Guadalupian (Middle Permian Series) GSSP. Palaeoworld vol. 16, pp. 208-221. PERMIAN RASIN C Debris

Sample WP-63. This bed has the first occurrence of Polydiexodina in the WP-section.