

Sparsity and Continuity Enhancing Seismic Imaging*

Peyman P. Moghaddam¹, Felix Herrmann², and Chris Stolk²

Search and Discovery Article #41532 (2015)

Posted March 2, 2015

*Adapted from extended abstract prepared in conjunction with a presentation given at CSPG/CSEG 2007 GeoConvention, Calgary, AB, Canada, May 14-17, 2007, CSPG/CSEG/Datapages © 2015

¹University of British Columbia, Vancouver, BC, Canada (ppoor@eos.ubc.ca)

²University of British Columbia, Vancouver, BC, Canada

Abstract

A non-linear singularity-preserving solution to the least-squares seismic imaging problem with sparseness and continuity constraints is proposed. The applied formalism explores curvelets as a directional frame that, by their sparsity on the image, and their invariance under the imaging operators, allows for a stable recovery of the amplitudes. Our method is based on the estimation of the normal operator in the form of an 'eigenvalue' decomposition with curvelets as the 'eigenvectors'. Subsequently, we propose an inversion method that derives from estimation of the normal operator and is formulated as a convex optimization problem. Sparsity in the curvelet domain as well as continuity along the reflectors in the image domain are promoted as part of this optimization. Our method is tested with a reverse-time 'wave-equation' migration code simulating the acoustic wave equation.

Introduction

Motivated by recent results on stable signal recovery for natural images from incomplete and noisy data (Candes et al., 2005), the seismic image recovery problem is formulated as a nonlinear optimization problem. After linearization and by ignoring the source and receiver signatures, the discretized forward model that generates seismic data can be written as

$$d = Km \quad (1)$$

In this single-scattering expression, $m(x)$ represents the (singular) fluctuations in the earth's acoustic properties with respect to an appropriately chosen smoothly varying background velocity model (the density of mass is assumed constant). These fluctuations are referred to as the model and seismic imaging aims to recover both the locations and the relative amplitudes of the velocity fluctuations from seismic data. Applying the adjoint of the linearized scattering operator to the data vector (d in Equation (1)) leads to the migrated image,

$$y = K^T d \quad (2)$$

An extensive literature has emerged on restoring the migration amplitude by inverting the normal matrix $\psi = K^T K$ (Nemeth et al., 1999; Kuhl and Sacchi, 2003), and involves the computation of the pseudo inverse of the scattering or demigration matrix K ,

$$m = (K^T K)^{-1} K^T d \quad (3)$$

Unfortunately, the normal operator is too big to be constructed explicitly and is too expensive to be evaluated as part of an iterative Krylov-subspace solver to invert the Hessian (see Nemeth et al., 1999).

Sparsity- and Continuity-Promoting Imaging

We address the above issues by exploiting recently developed curvelet frames. These frame expansions compress seismic images and consist of a collection of frame elements 'curvelets' that are invariant under pseudo-differential operators. These properties allow us to develop an approach similar to the so-called wavelet-vaguelette method (WVD), as proposed by Donoho 1995 and later by Candes and Donoho, 2000, where scale-invariant homogeneous operators are inverted using the eigenfunction-like behavior of multiscale transforms such as the wavelet and curvelet transform.

The solution is formulated in terms of a sparsity-promoting nonlinear optimization problem and can be seen as a formalization of earlier ideas on stable seismic image recovery. During the optimization, sparsity in the transformed domain as well as continuity along imaged reflectors, are jointly promoted. Both penalties are part of the following nonlinear optimization problem (see Herrmann et al., 2008)

$$P: \quad (4)$$

where the sparsity vector x is optimized with respect to the penalty functional $J(x)$ and the data misfit. We use the term sparsity vector for x to point out that this vector corresponds to the coefficients of a transform that is designed to be sparse on the model.

The penalty functional $J(x)$ is designed to promote sparsity and continuity. The above optimization problem solves for the model by finding a coefficient vector x that minimizes the penalty term subject to fitting the data to within a user-specified tolerance level ϵ . We reserved the 'tilde' symbol to denote estimates that are found through optimization. The recovered model m is calculated by computing the inverse of A^T , which represents the diagonally-weighted curvelet synthesis matrix. This synthesis matrix is designed such that

$$AA^T r \cong \psi(r) \quad (5)$$

with r an appropriately chosen discrete reference vector and ψ the discrete normal operator formed by compounding the discrete scattering and its transpose the migration operator. With $A = C^T \Gamma$ and C the curvelet transform, Equation (5) expresses the normal operator as a form

of 'eigenvalue' decomposition with curvelets as 'eigenvectors'. Γ is the square-root of the 'eigenvalues' and can be shown to be smooth in the curvelet domain.

Our algorithm for approximately inverting the normal operator involves the following sequence of steps:

- 1) Form the normal operator using one's favorite numerical implementation for the migration operator and its adjoint, i.e., $\psi = K^T K$ with the symbol. This discrete normal operator needs to be made zero order;
- 2) Select a relevant reference vector, r , that is close enough to the unknown image;
- 3) Estimate the diagonal approximation. This diagonal approximation defines the synthesis matrix A ;
- 4) Estimate x by solving the nonlinear optimization problem P . This program inverts the synthesis matrix. The discretized model vector m is calculated from the recovered coefficient vector x through the pseudo inverse of A^T .

Example

[Figure 1](#) and [Figure 2](#) show the evaluation of our method on a typical subsalt imaging problem. [Figure 1a](#) shows a sufficiently smooth background velocity for SEG-AA' salt model, with reflectivity shown in [Figure 1b](#). This reflectivity is de-migrated and migrated ($K^T K$) using the smooth SEG-AA' background velocity, [Figure 2a](#) shows the result of this process. [Figure 2b](#) shows the recovered image using our proposed recovery algorithm. [Figure 2c](#) shows the trace comparison between the migrated image (properly scaled) ([Figure 2a](#)), original reflectivity ([Figure 1b](#)) and recovered image ([Figure 2b](#)) along the horizontal line in the bottom of image (3500 m depth) for the offsets from 4.3 to 7.2 km. For this example, we used a two-way wave-equation reverse-time migration and modeling (Symes 2006). The dataset consist of 324 shots and 176 receivers for each shot.

Conclusion

The method presented in this paper combines the compression of images by curvelets with the invariance of this transform under the normal operator. This combination allows us to formulate a stable recovery method for seismic amplitudes. Compared to other approaches for migration preconditioning, our method (i) brings the amplitude correction problem within the context of stable signal recovery; and (ii) provides for a diagonal approximation of normal operator

References Cited

- Candes, E.J., and D.L. Donoho, 2000, Recovering Edges in Ill-posed Problems: Optimality of Curvelet Frames: *Ann. Statist.*, v. 30, p. 784-842.
- Candes, E.J., J. Romberg, and T. Tao, 2005, Stable Signal Recovery from Incomplete and Inaccurate Measurements: *Communications on Pure and Applied Mathematics*, v. 59, p. 1207-1223.
- Donoho, D.L., 1995, De-Noising by soft thresholding: *IEEEETIT*, v. 41, p. 613-627.
- Herrmann, F.J., P.P. Moghaddam, and C.C. Stolk, 2008, Sparsity- and continuity-promoting seismic imaging with curvelet frames: *Journal of Applied and Computational Harmonic Analysis*, v. 24/2, p.150–173. doi:10.1016/j.acha.2007.06.007
- Kuhl, H., and M. Sacchi, 2003, Least-squares wave-equation migration for {AVP/AVA} inversion: *Geophysics*, v. 68/1, p. 262-273.
- Nemeth, T., C. Wu, and G. T. Schuster, 1999, Least-squares migration of incomplete reflection data: *Geophysics*, v. 64/1, p. 208-221.
- Symes, W.W., 2006, Reverse time migration with optimal checkpointing: Technical Report TR06-18, Department of Computational and Applied Mathematics, Rice University, Houston, Texas, USA.

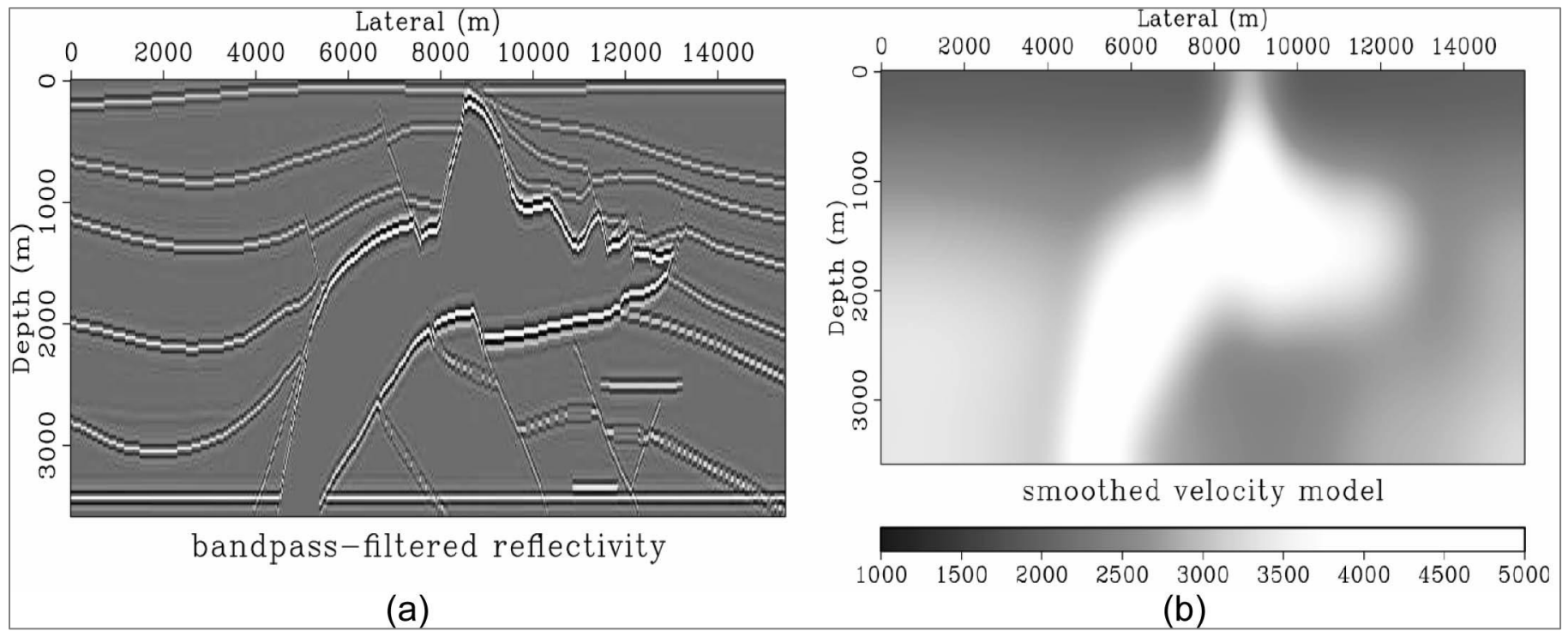


Figure 1. (a) Reflectivity model. (b) SEG-AA Sufficiently smooth background model.

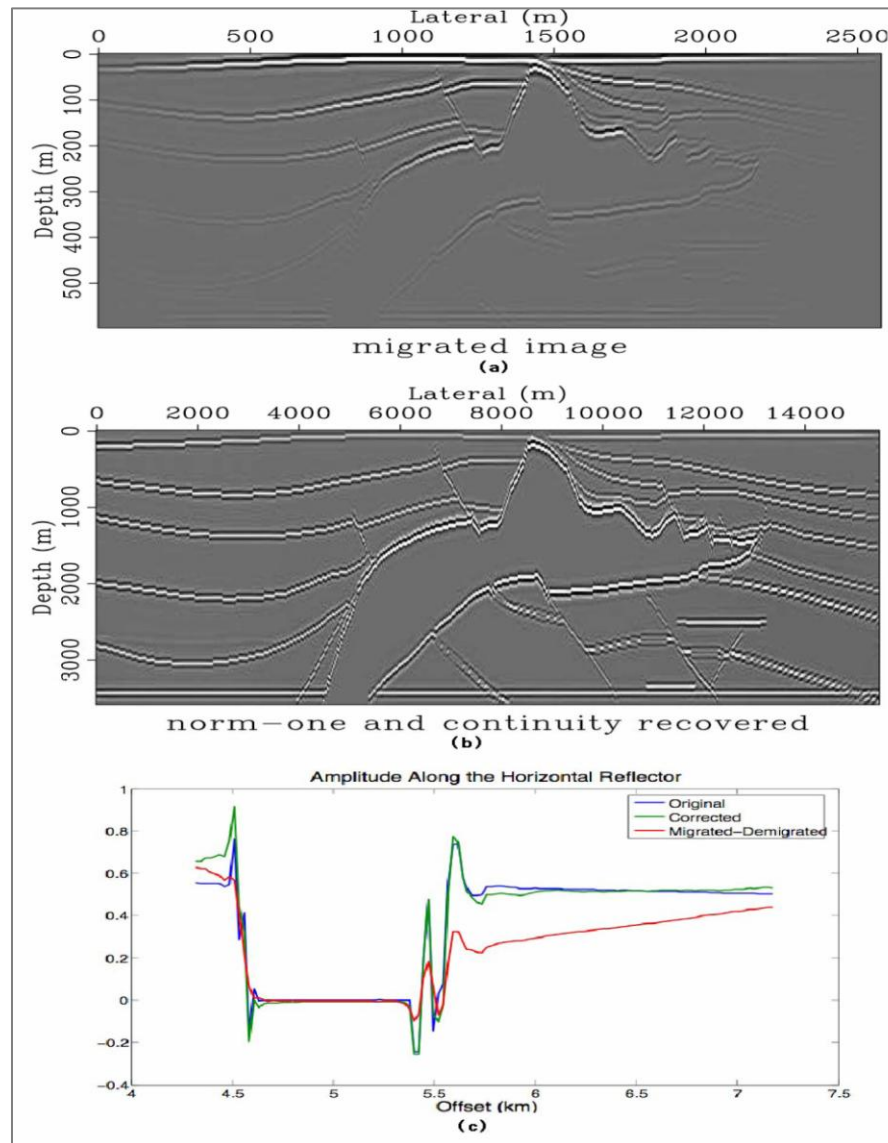


Figure 2. Example of seismic amplitude recovery. (a) Normal operator (modeling followed by migration) applied on the SEG-AA' reflectivity model. (b) Recovered image using our proposed method. (c) Amplitude comparison for the bottom reflector between original, migrated-demigrated and the recovered image.