

# **PS Seismic Signature of CO<sub>2</sub> Flooding in Tuscaloosa Sandstone\***

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## **Abstract**

Seismic technology offers promise to monitor the CO<sub>2</sub> flood front in reservoirs undergoing enhanced oil recovery and CO<sub>2</sub> sequestration. This laboratory study is focused on understanding the seismic signature of CO<sub>2</sub> flooding in rocks. Compressional (V<sub>p</sub>) and shear wave (V<sub>s</sub>) velocities were measured on several samples (fabricated glass bead, Berea Sandstone and Tuscaloosa Sandstone) as a function of effective pressure and pore fluids. In these experiments, phase change of CO<sub>2</sub> was achieved by varying pore pressure. Realistic subsurface conditions were simulated by having mixed phase (brine, hydrocarbon and CO<sub>2</sub>) in the rock samples. It is observed that both V<sub>p</sub> and V<sub>s</sub> decrease when CO<sub>2</sub> changes from gas to liquid, contrary to the general expectations. This behavior can be explained due to the smaller change in fluid compressibility (89%) as compared to fluid density change (800%) with phase change.

Shear velocity measurements as a function of effective pressure showed Biot effective pressure coefficient (n) is approximately equal to 1.0 and independent of the fluid phase. For compressional wave velocity, n is approximately 1.0 for gaseous CO<sub>2</sub> and much less than 1.0 for liquid CO<sub>2</sub>. The experimental work allows feasibility study of mapping the CO<sub>2</sub> front from a surface seismic survey.

## **Selected References**

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Christensen, N.I., and H.F. Wang, 1985, The influence of pore pressure and confining pressure on dynamic elastic properties of Berea Sandstone: *Geophysics*, v. 50/2, p. 207-213.

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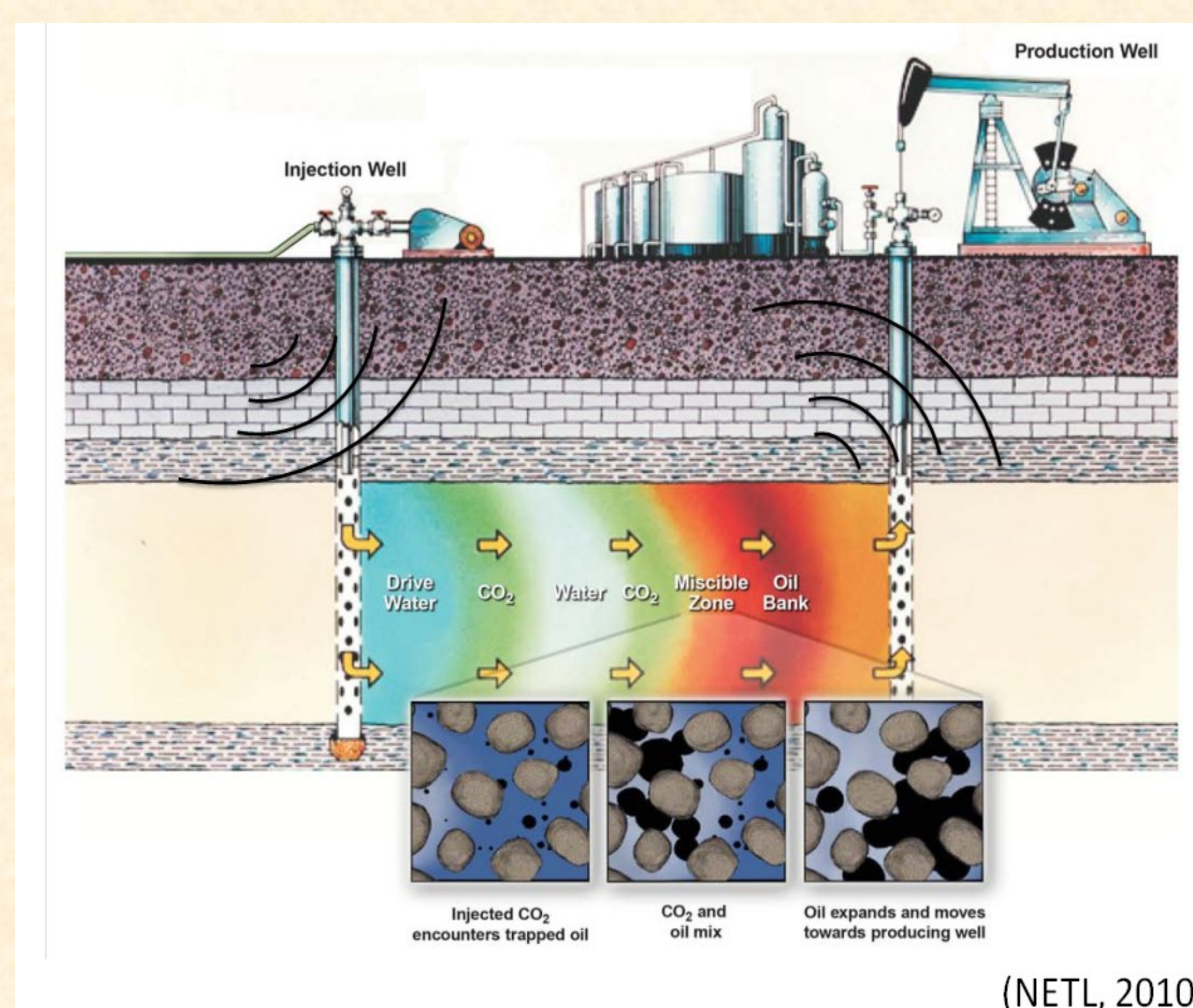
## 1. Abstract

Seismic technology offers promise to monitor the CO<sub>2</sub> flood front in the reservoirs undergoing enhanced oil recovery and CO<sub>2</sub> sequestration. This laboratory study is focused on understanding seismic signature of CO<sub>2</sub> flooding in rocks. Compressional (V<sub>p</sub>) and shear wave (V<sub>s</sub>) velocities were measured on several samples (fused glass beads, Berea sandstone and Tuscaloosa sandstone) as a function of effective pressure and pore fluids. In these experiments, phase change of CO<sub>2</sub> was achieved by varying pore pressure. Realistic subsurface conditions were simulated by having mixed phase (brine, hydrocarbon and CO<sub>2</sub>) in the rock samples. It is observed that both V<sub>p</sub> and V<sub>s</sub> decrease when CO<sub>2</sub> changes from gas to liquid. This behavior is due to smaller change in fluid compressibility (89%) as compared to fluid density change (800%) with phase change.

Biot-Gassmann modeling shows good agreement with experimental results for gas replaced by brine and oil replacing brine system but not for liquid CO<sub>2</sub> flooding. The experimental work allows feasibility study of mapping CO<sub>2</sub> front from surface seismic survey. Both pre-flooded and post-flooded well logs are compared and correlated based on the understanding of the laboratory scale behavior and Biot-Gassmann theory.

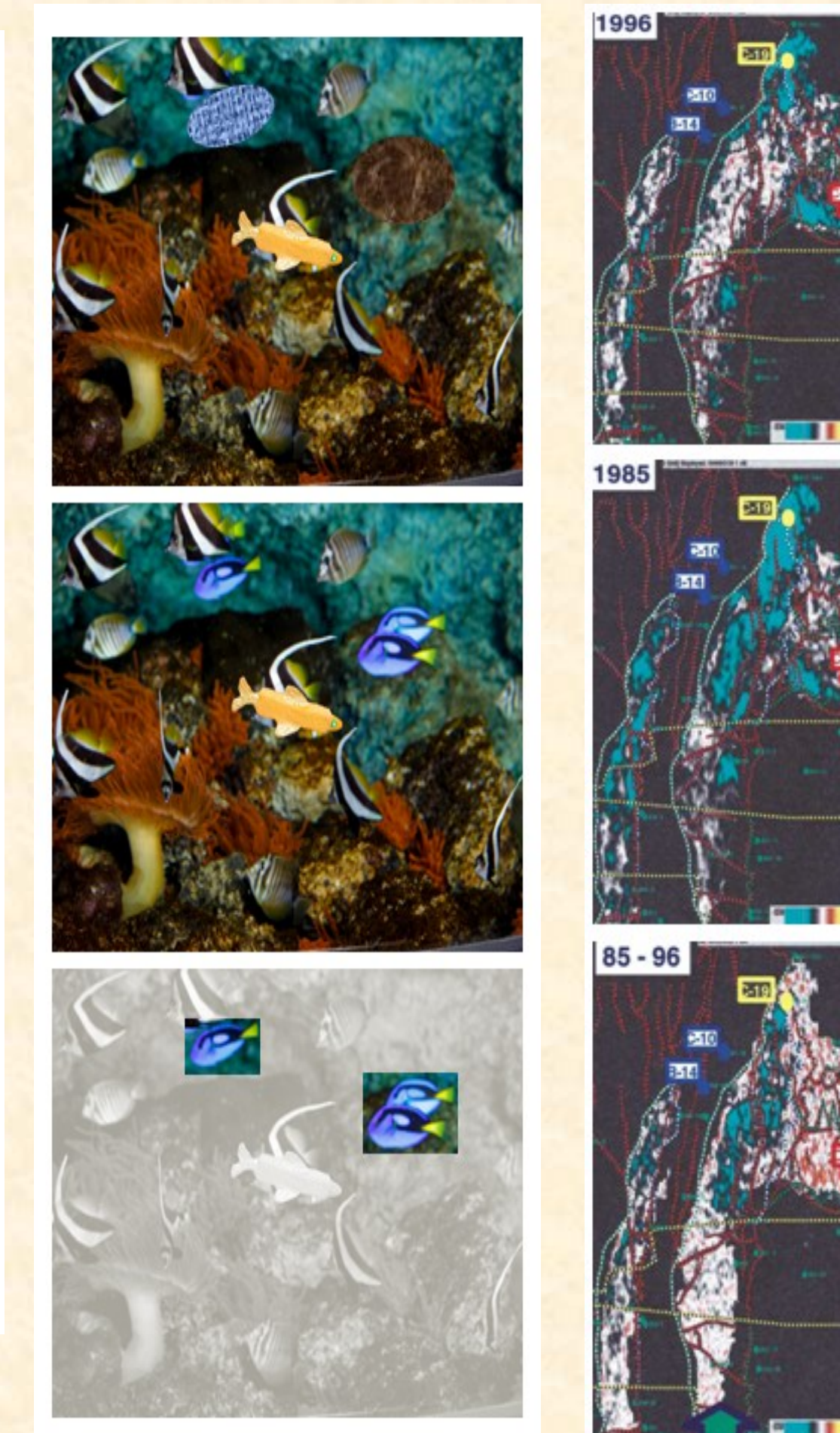
## 2. Introduction

### CO<sub>2</sub> injection



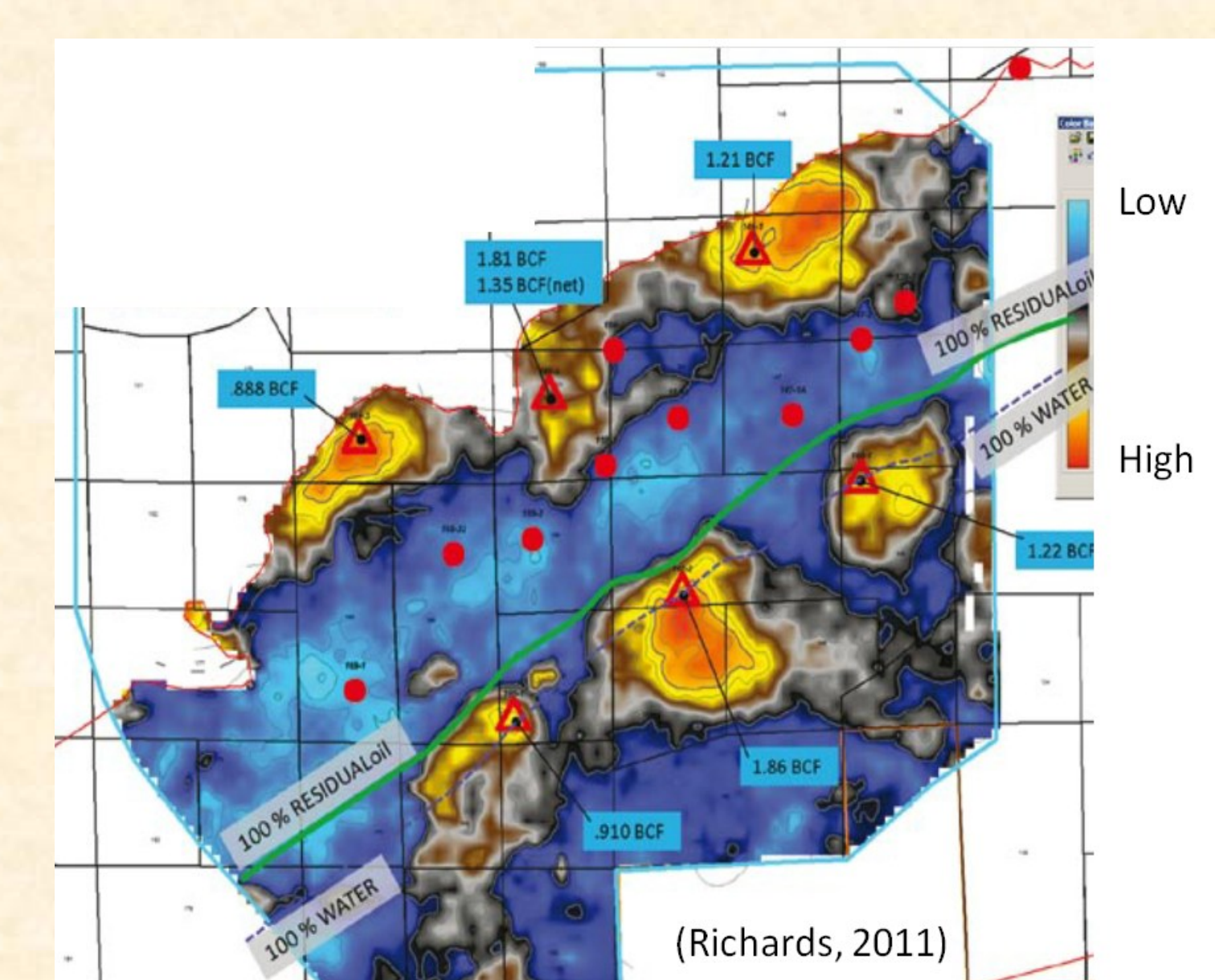
- Decreases the surface tension between rock pore surface and fluid.
- Decreases the viscosity
- Decreases the bulk modulus

### 4D Seismic



(Landro et al., 2001)

### Study field



- ▲ CO<sub>2</sub> injector wells
- Producing wells
- High amplitude difference indicating presence of CO<sub>2</sub>
- Low amplitude difference indicating presence water and oil

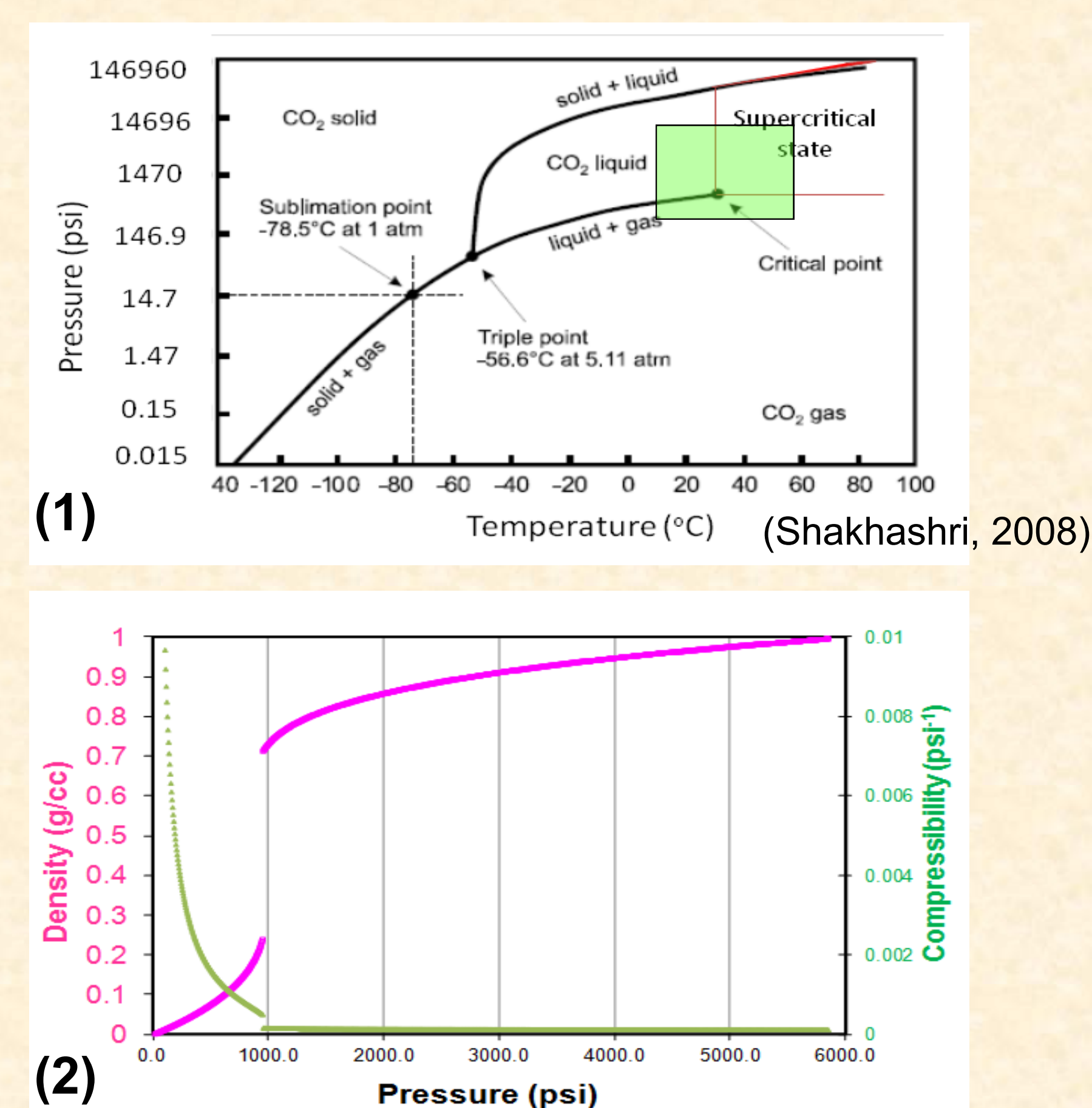
## 3. Objective

- Understand seismic response of CO<sub>2</sub> flooding in Tuscaloosa sandstone with the help of laboratory measurements

## 4. Lithology

- Holt Bryant reservoir of Delhi field, LA.
- Tuscaloosa (3204-3279 ft) sands underlie the Paluxy sands (3279-3283 ft) at well 159-2
- Tuscaloosa depositional environment is fluvial channel while Paluxy has distributary channel environment

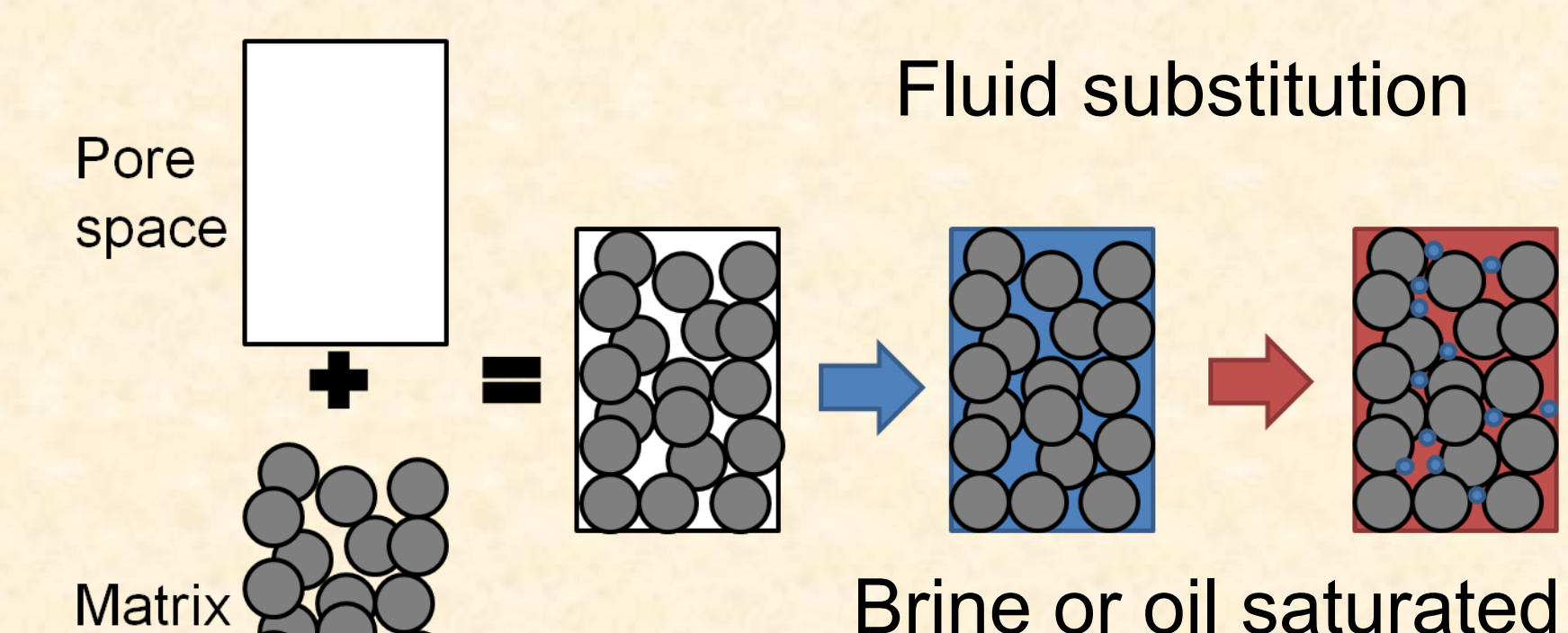
## CO<sub>2</sub> phase diagram:



At reservoir conditions (Figure 1), expected CO<sub>2</sub> phase is shown by green region. Figure 2 is plot of estimated density and compressibility using Peng Robinson EOS.

## 5. Literature review

### Biot Gassmann theory:



Effective bulk (K<sub>e</sub>) and Shear (G<sub>sat</sub>) modulus of the saturated sample given as

$$K_e = K_{dry} + \frac{n^2 K_f}{\phi + \frac{K_f}{K_g} (n - \phi)} \quad \text{.....(1)}$$

$$G_{sat} = G_{dry} \quad \text{.....(2)}$$

(Gassmann, 1951)

### Assumptions:

- Homogeneous
- Isotropic
- Connected pores
- Allow pore pressure equilibration
- No chemical alterations

### Variables:

**K<sub>dry</sub>**: Dry bulk modulus of the porous rock sample (drained condition). This is determined from the dry velocity (V<sub>p</sub> and V<sub>s</sub>) measurements.

**Φ**: Porosity of the rock at reservoir condition

**K<sub>f</sub>**: Fluid bulk modulus calculated from fluid Rues average

**K<sub>g</sub>**: Matrix bulk modulus calculated from mineral VRH average

### Variables (contd):

**n**: known as Biot effective stress coefficient. 'n' is crucial in determining the role of pore pressure on parameters governed by effective pressure. Previous published n values on brine and oil saturated sandstones. This coefficient is calculated using equations for static and dynamic basis respectively (Todd and Simons, 1972):

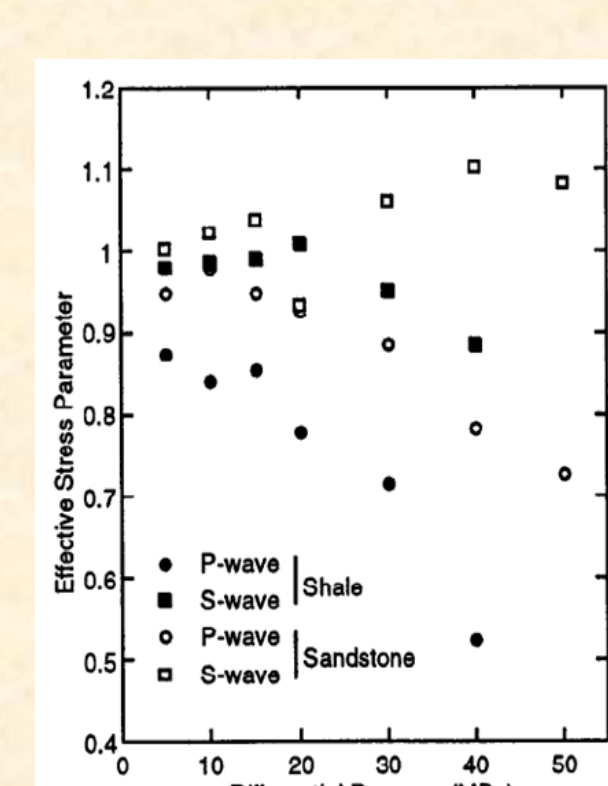
$$n = 1 - \frac{K_{dry}}{K_m} \quad \text{or} \quad n = 1 - \frac{\left| \frac{\partial V_p}{\partial P_p} \right|_{P_p = \text{const}}}{\left| \frac{\partial V_p}{\partial P_d} \right|_{P_p = \text{const}}} \quad \text{.....(3)}$$

Diff. press (psi)	Fluid	n <sub>p</sub>	n <sub>s</sub>
73	Brine	0.99	
725	Brine	0.93	1.02
2901	Brine	0.89	1.06
8702	Brine	0.84	1.07

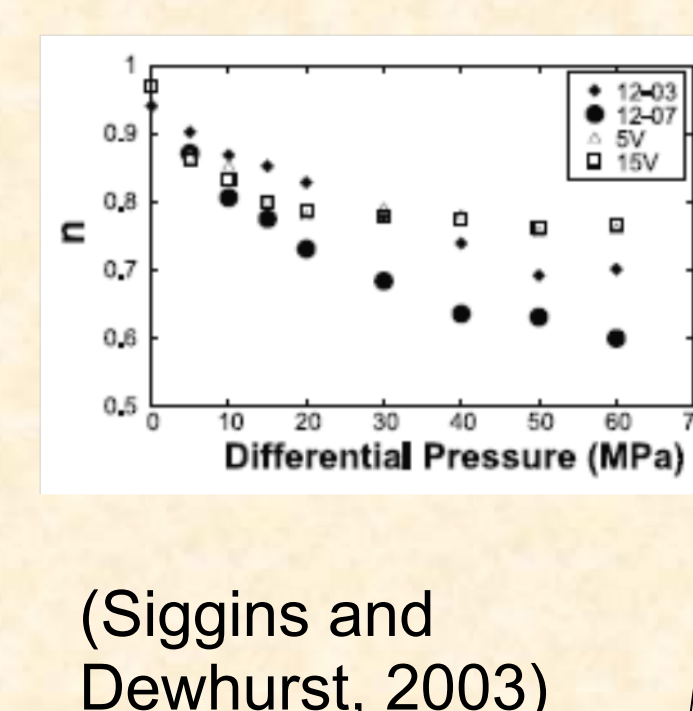
(Christensen and Wang, 1985)

Diff. press (psi)	Fluid	n <sub>p</sub>	n <sub>s</sub>
725	Brine	0.936	
1450	Brine	0.985	
2176	Brine	0.940	
2901	Brine	0.747	
3626	Brine	0.648	

(Prasad and Manghnani, 1985)



(Hornby, 1996)



(Siggins and Dewhurst, 2003)

## 6. Experimental algorithm

### Exp 1: Velocity equilibration study

Purpose:

- Study the time required for gaseous and liquid CO<sub>2</sub> to equilibrate in pores
- The velocity response of pure gaseous and liquid state CO<sub>2</sub> in room dry condition

### Exp 2: Velocity vs. effective pressure

Purpose:

- Velocity response with effective pressure
- n-values calculation using eq. (3)

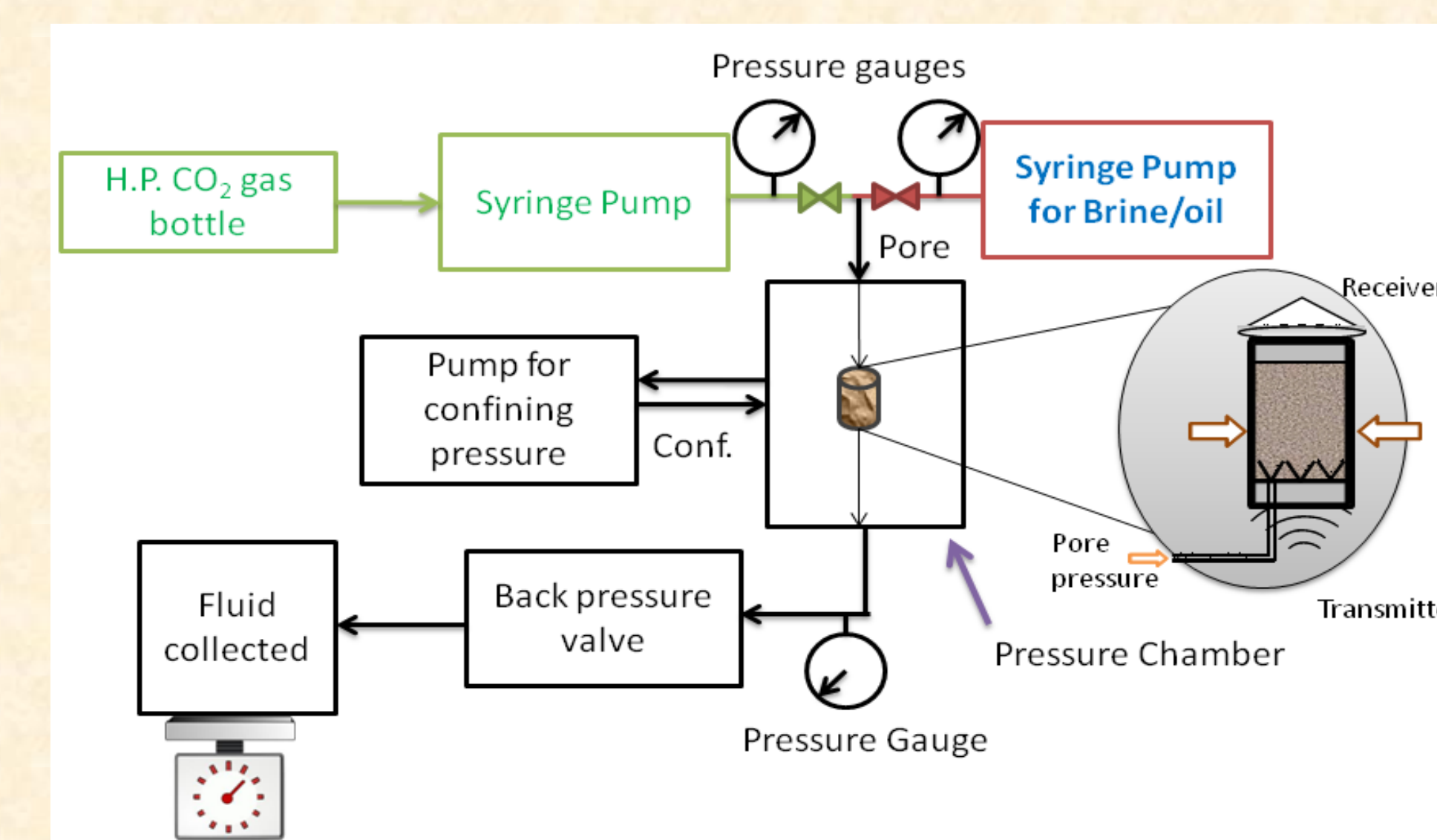
### Exp 3: Flooding Experiment with brine, oil and CO<sub>2</sub>

Purpose: To determine velocity as a function of fluid saturation at reservoir condition

Experimental elastic parameters compared with Biot Gassmann predicted parameters (V<sub>p</sub> & V<sub>s</sub>)

Empirical correlation proposed to be applied on well logs and verified with actual well logs

## 7. Procedure & Apparatus for Flooding experiment



### Sequence 1:

Confining and pore press. set to 1500 psi, 0 psi respectively → Brine pump attached to pore pressure line → Velocity measurements with brine saturation on dry sample

### Sequence 2:

Confining and pore press. increased to 3000 psi, 1500 psi res → Oil pump attached to pore pressure line, brine is pushed out → Vel. measurements with oil saturation on brine sat. sample

### Sequence 3:

Same pressure conditions (actual *insitu* conditions) → Liquid CO<sub>2</sub> is flooded in the pore pressure line to push oil out → Vel. measurements with CO<sub>2</sub> saturation on oil sat. sample

← Brine as saturant → ← Oil as saturant →



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## Sample details

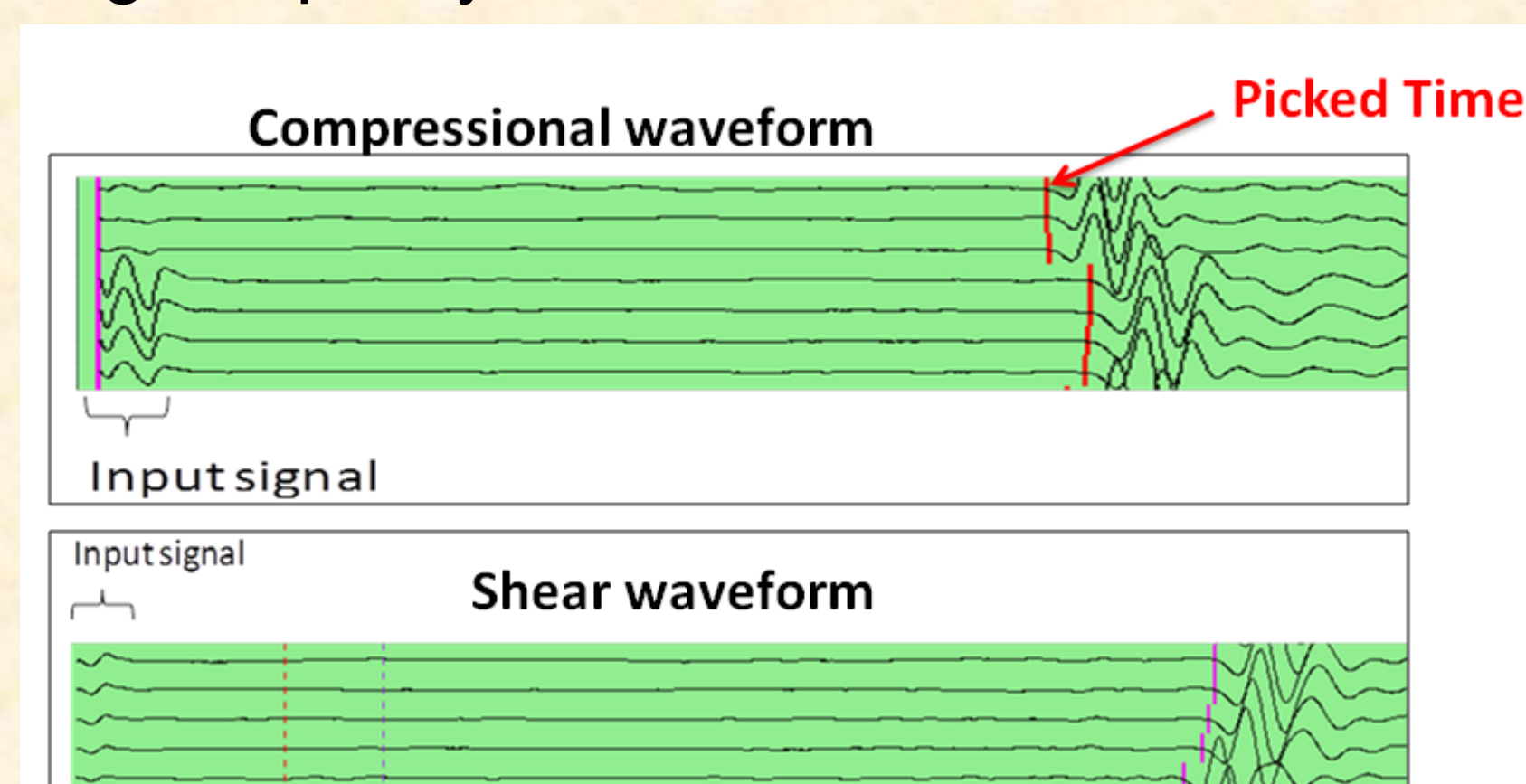
## 8. Sample and Fluid properties

Core interval (ft)	Unit	Characteristics	Environment	Mineral composition	Porosity (%)
3204-3207	Tusc 9	Lower fine grain sand, quartz rich	Lower Fluvial channel	Qtz: 84% Clay: 6%	28
3225-3228	Tusc 5	Middle fine grain sand, solely quartz	Middle Fluvial channel	Qtz: 93% Clay: 3%	23
3258-3261	Tusc 3	Lower very fine grain, quartz rich	Fluvial channel	Qtz: 82% Clay: 13%	25
3276-3279	Paluxy	Lower very fine grain, quartz rich	Distributary Mouth Bar	Qtz: 81% Clay: 12%	24
3282-3285	Paluxy	Middle very fine grain sand, quartz rich	Estuary / Distributary Channel	Qtz: 86% Clay: 7%	26

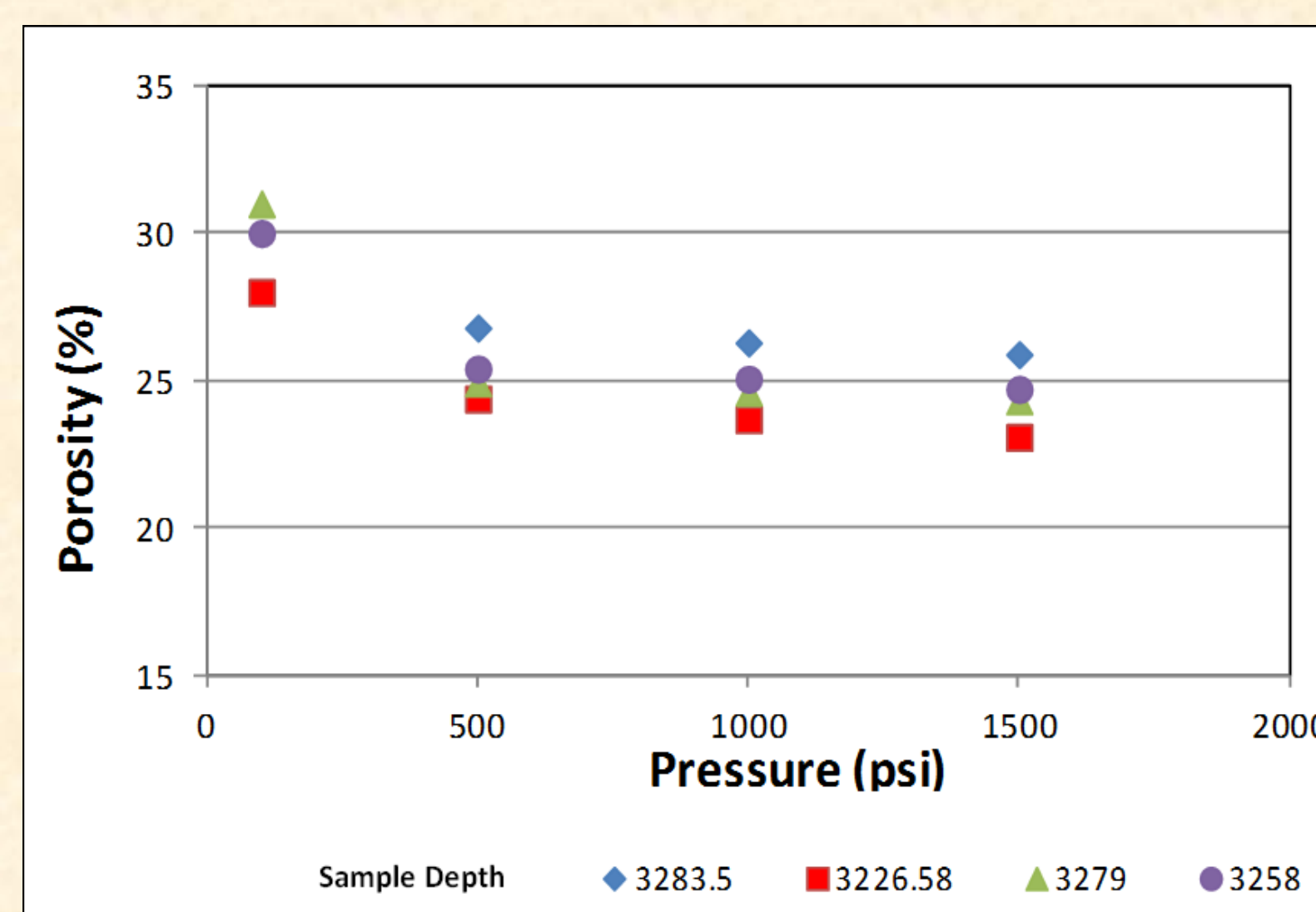
## Fluid properties

Fluid	Phase	Density (g/cc)	Bulk modulus (GPa <sup>-1</sup> )
CO <sub>2</sub>	gas	0.002	1.3 x 10 <sup>-4</sup>
CO <sub>2</sub>	Liquid (1500 psi)	0.733	0.01
Brine	Liquid	1.05	2.583
Oil	Liquid	0.81	1.375

## Signal quality

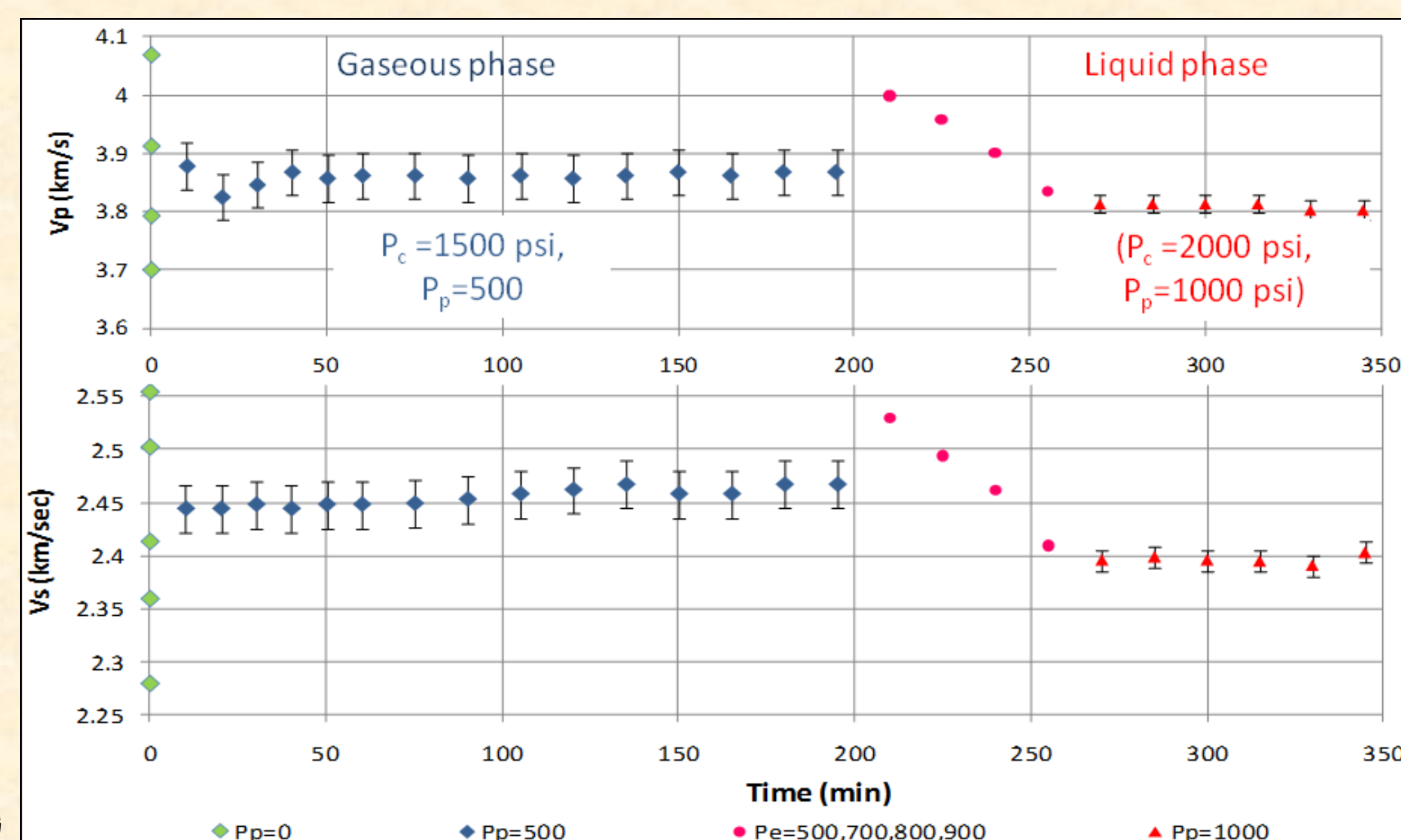


## Porosity as function of effective pressure



Drop of 18% (average) in porosity was observed with effective pressure when brought to reservoir conditions. This is one of the reasons why we perform the experiments at reservoir conditions.

## 9a. Experiment 1: Velocity equilibration study



Compressional and Shear velocities decreased after CO<sub>2</sub> phase change from gaseous and liquid. The increase in fluid modulus (80%) raised the Bulk modulus (K) while shear modulus (G) remains same (shows no dispersion). However, fluid density dominates due to higher change (800%) than its compressibility which decreased the elastic velocities (both P and S wave).

$$K_e = V_p^2 \cdot \rho_{bulk} - \left( \frac{4}{3} \right) V_s^2 \cdot \rho_{bulk}$$

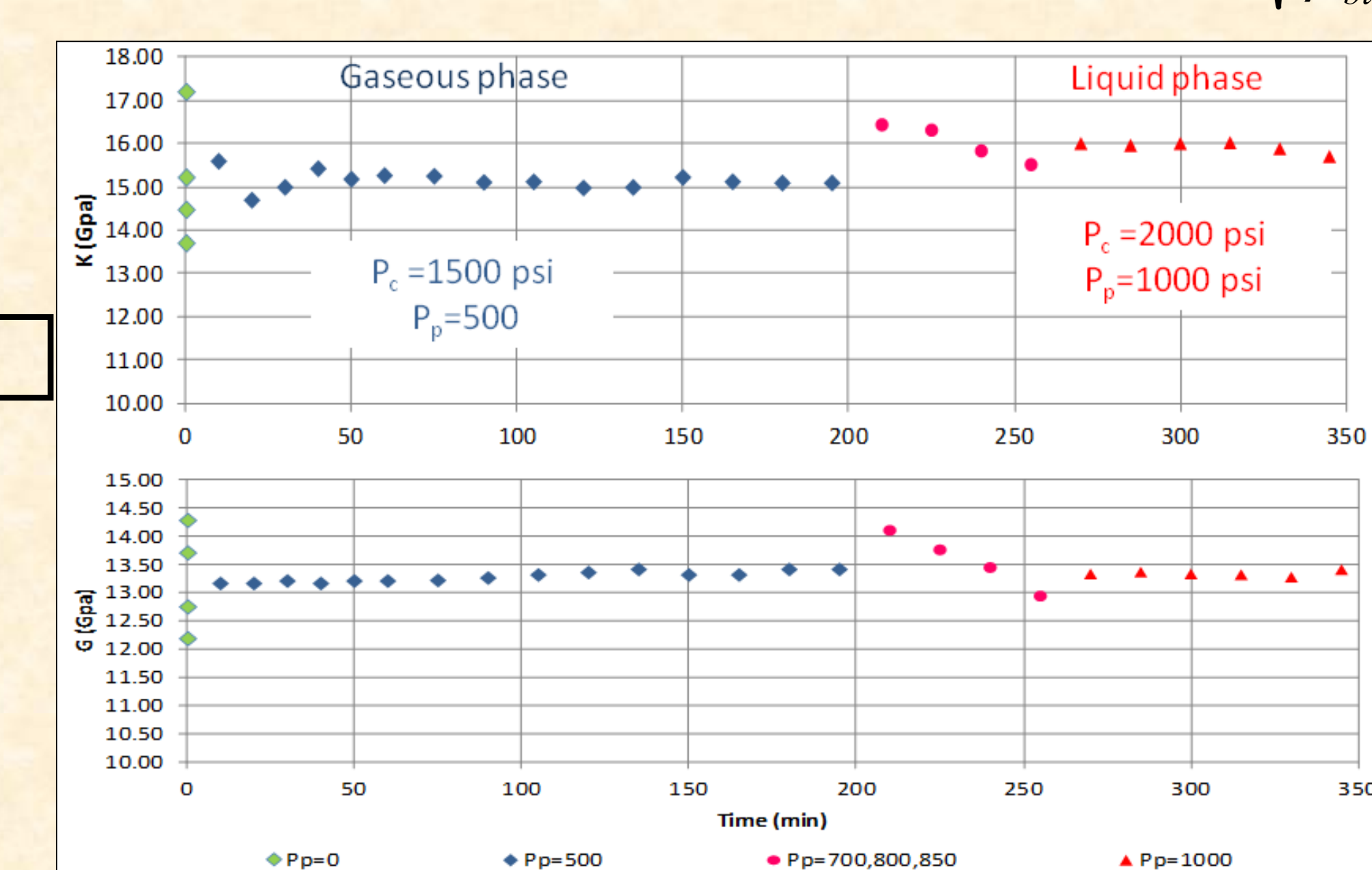
This equation rewrite as,

$$V_p = \sqrt{\frac{K + \frac{4}{3} G_{sat}}{\rho_{bulk}}}$$

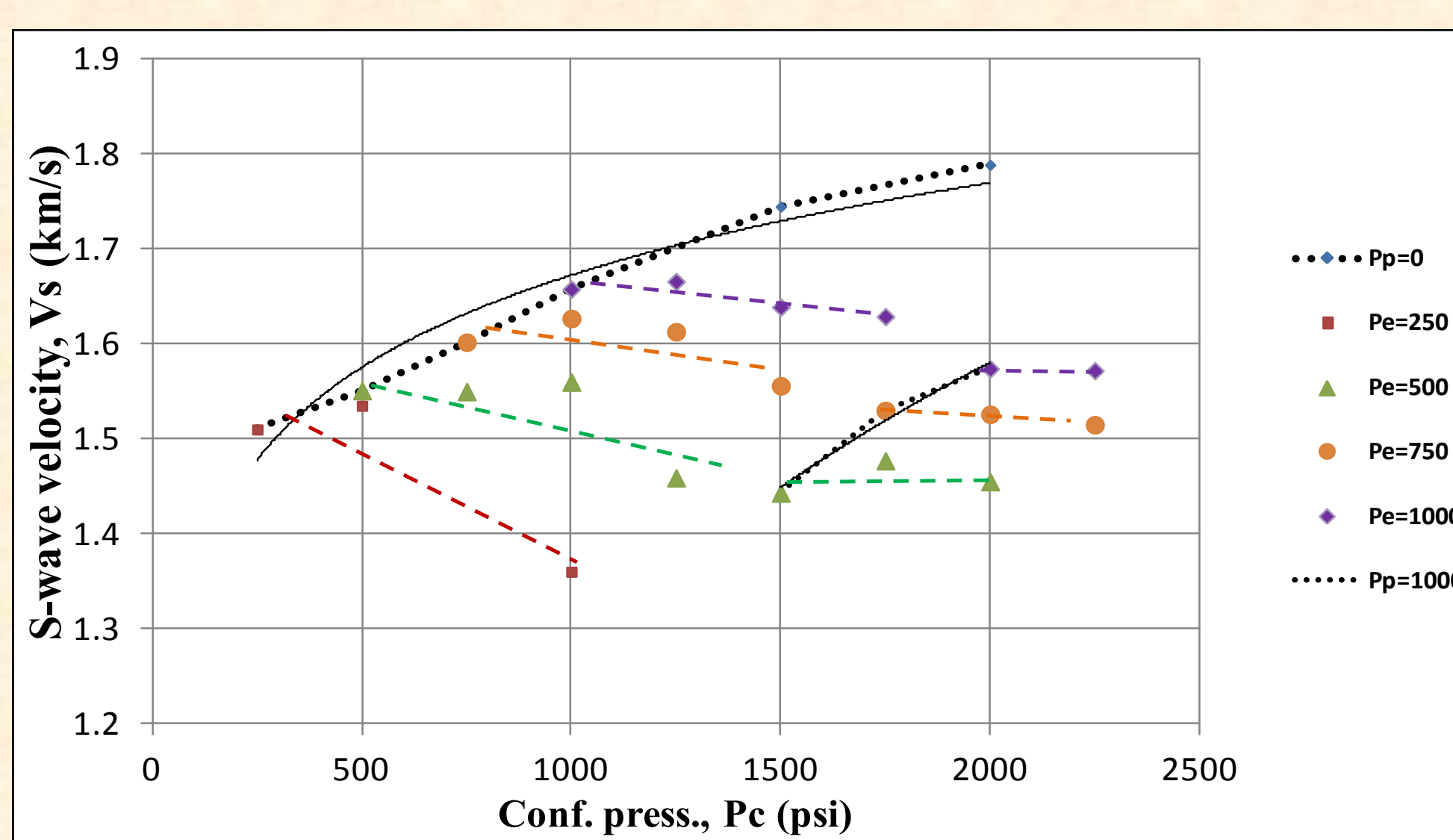
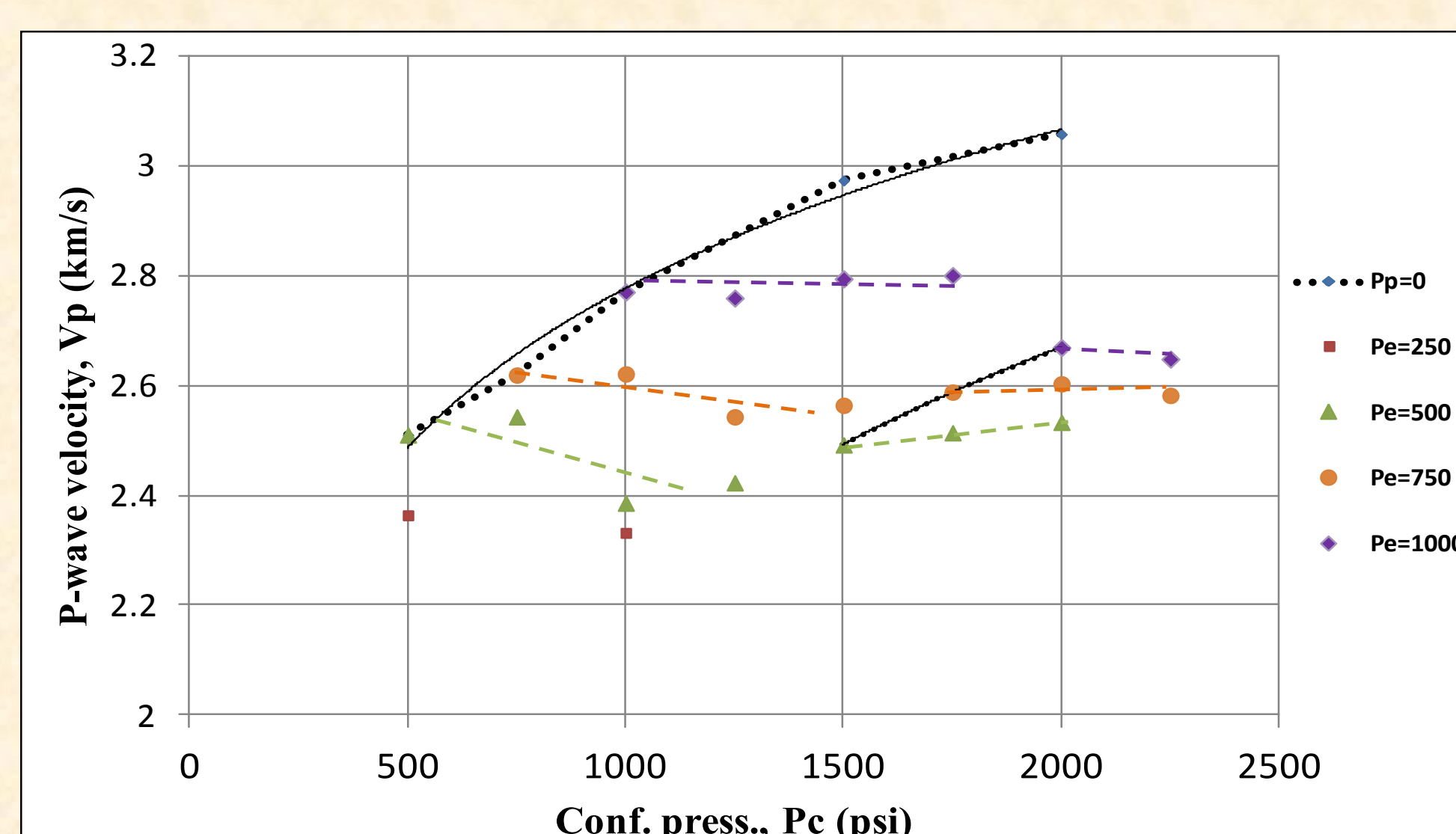
$$G_{sat} = V_s^2 \cdot \rho_{bulk}$$

This equation rewrite as,

$$V_s = \sqrt{\frac{G_{sat}}{\rho_{bulk}}}$$



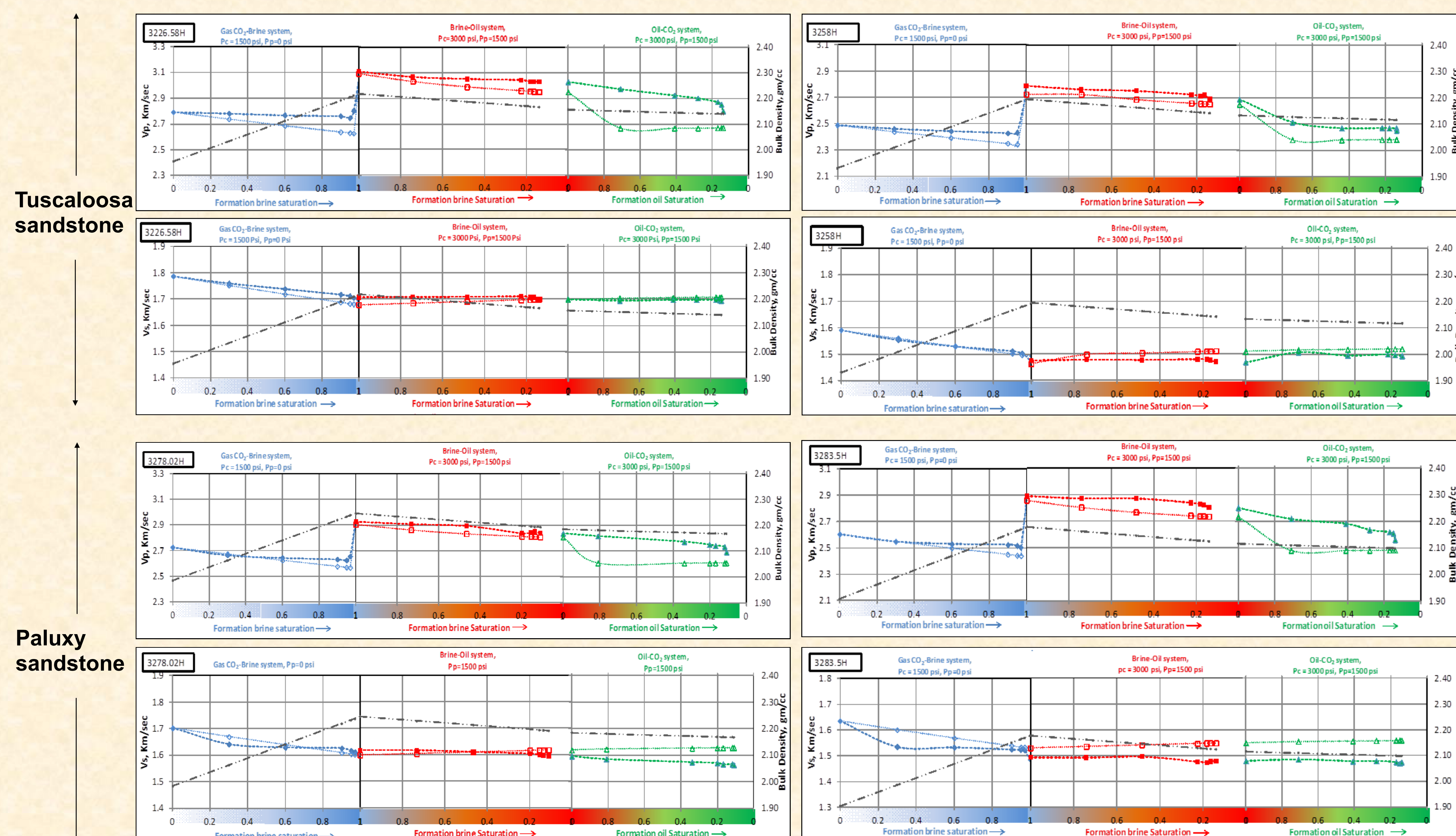
## 9b. Experiment 2: Velocity Vs. Effective pressure and 'n'



Elastic Velocities (P and S wave) are measured as a function of confining pressure and zero pore pressure. The pore pressure was incremented by 250 psi and variation of velocities with confining pressure were recorded, till pore and confining pressure of 1500 psi and 2250 psi respectively.

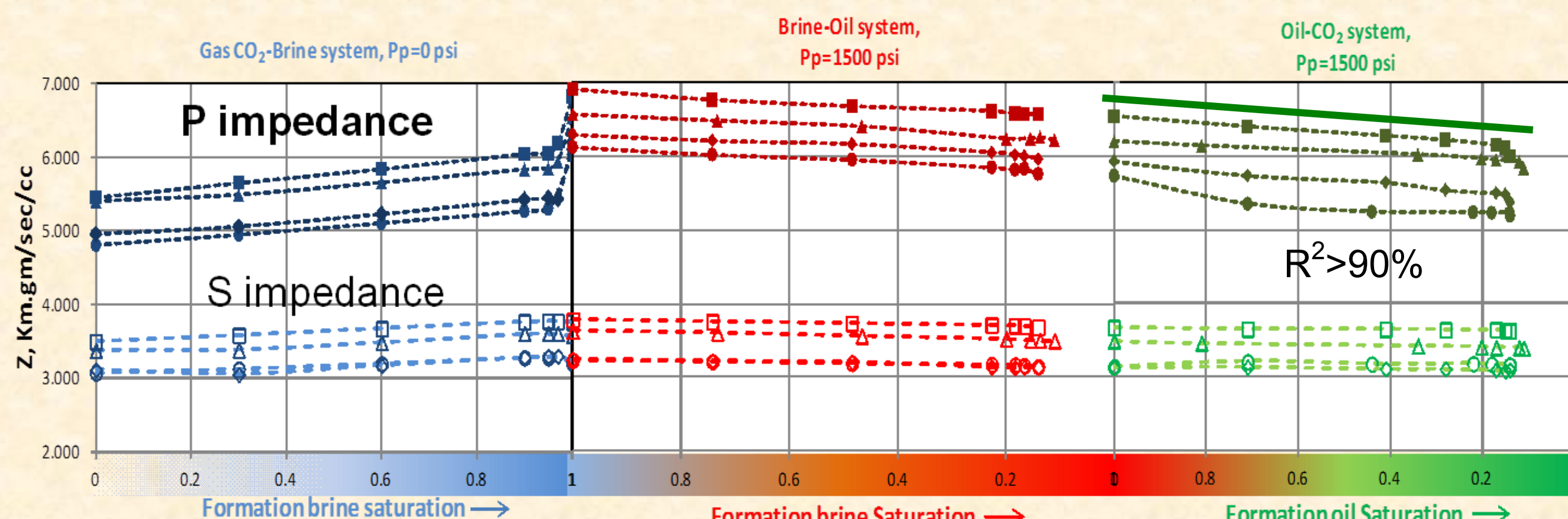
Diff. pressure (psi)	Fluid	$\eta_p$	$\eta_s$
500	CO <sub>2</sub> gas	1.20±0.13	1.38±0.26
750	CO <sub>2</sub> gas	1.18±0.09	1.33±0.28
1000	CO <sub>2</sub> gas	0.90±0.06	1.33±0.13
500	Liq. CO <sub>2</sub>	0.80±0.01	0.92±0.21
750	Liq. CO <sub>2</sub>	1.03±0.12	1.12±0.03
1000	Liq. CO <sub>2</sub>	1.27±0.00	1.04±0.00

## 9c. Experiment 3: Results of Velocity variation with fluid flooding



Biot Gassmann theory not in good agreement with CO<sub>2</sub> flooding (3<sup>rd</sup>) scenario

## 10. Empirical relationship and correlation with well logs

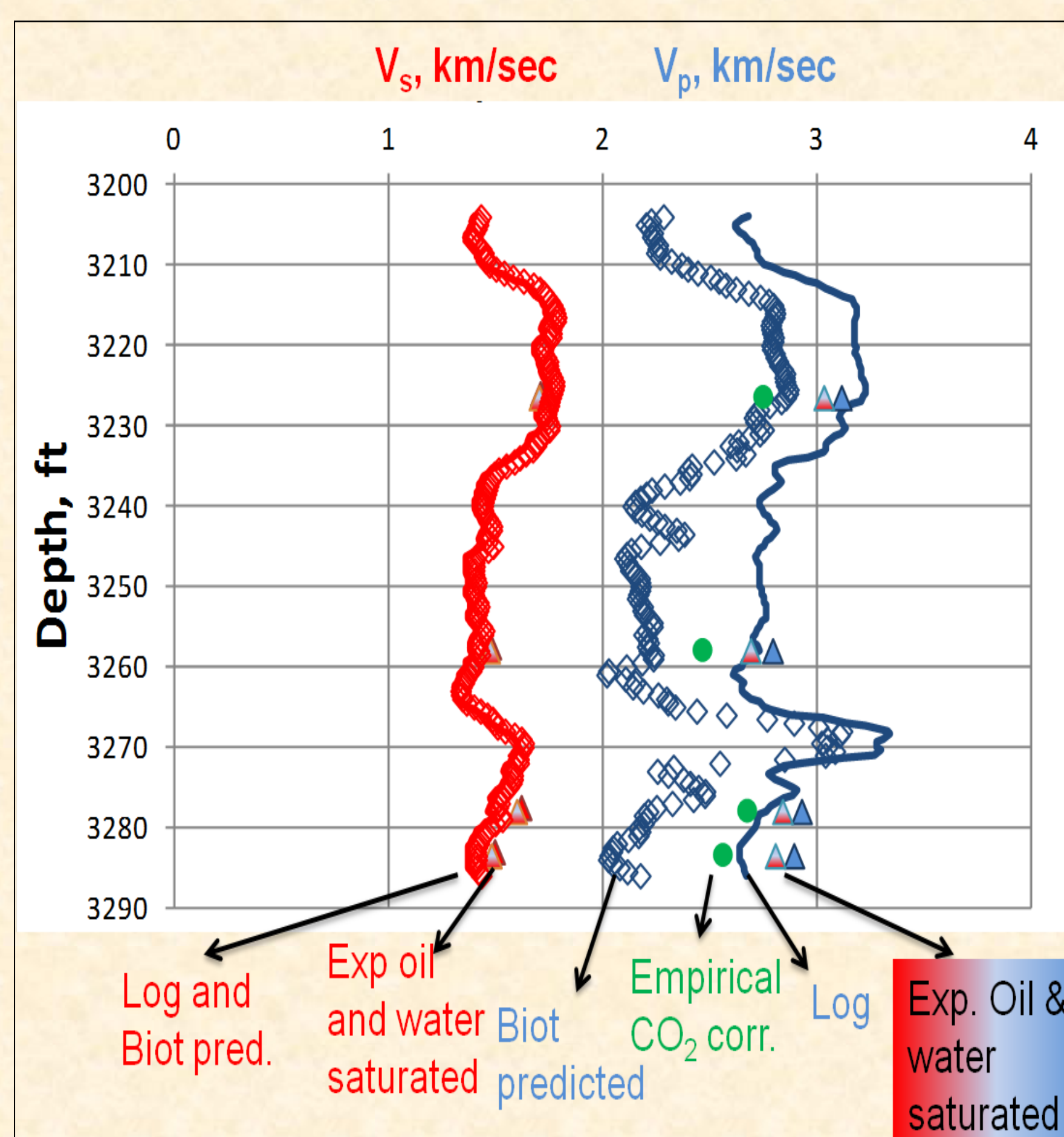


Linear regression analysis is performed on the experimental impedance data points in CO<sub>2</sub> flooding scenario to generate an empirical relationship.

$$Z_p = r S_o - 1.1 Z_{dry}$$

Where, r is found to be 0.5 km.gm/sec/cc, S<sub>o</sub> is the saturation of the oil and Z<sub>dry</sub> is the P-impedance calculated in the dry state. r is the slope of the velocity variation as CO<sub>2</sub> saturation changes. It depends on type of rock and pore fluid distribution while undergoing saturation. Some of the velocity vales during low CO<sub>2</sub> saturation (high oil saturation) may represent CO<sub>2</sub> in miscible state.

## 11. Pre-flooded actual and synthetic wells logs with 20% CO<sub>2</sub> fluid substitution



Pre-flooded log from a well 159-2 on the study field is plotted with empirical and Biot predicted elastic velocity values.

Shear parameters shows good match with Biot predicted values. However, the theory under predicts the compressional velocity values with saturation.

## 12. Conclusions

- P-impedance shows high sensitivity to the saturation.
- The Biot Gassmann predicted P-velocity is found to be good agreement with brine replacing dry state and oil replacing brine saturated conditions.
- The P-velocity generated using Biot Gassmann deviate by 10% from the experimental values in CO<sub>2</sub> flooding scenario.
- The proposed impedance model can be used for Tuscaloosa trend samples for better mapping of the CO<sub>2</sub> saturation.

## Acknowledgement:

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