Connecting Hyperpycnal Flow Deposits to River Flood Dynamics*

Michael Lamb¹, Brandon McElroy², David Mohrig², John Shaw², Bryant Kopriva², Steff Lazo-Herencia² and Jim Buttles²

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Abstract

Hyperpycnal flows are turbid river plumes that can plunge to form turbidity currents where they enter a water body of lesser density. Because these flows provide one of the most direct connections between terrestrial sediment sources and marine depositional sinks, their deposits preserve an important record across a variety of climatic and tectonic settings. A leading hypothesis assumes that hyperpycnal-flow velocity scales directly with river discharge, such that individual turbidites record the rising and falling discharge of a flooding river. Using a 1D numerical model and flume experiments, we test this hypothesis and find that turbid river flow must move through a backwater zone, depth-limited plume, and plunging zone before becoming a turbidity current. These zones can extend tens of kilometers offshore and significantly affect the transfer of momentum from river to turbidity current. Counter to the proposed hypothesis, our results indicate that local flow velocities within hyperpycnal flows can be uncorrelated or even anti-correlated with inlet river discharge because of translation of the plunge point resulting from temporal variations in discharge and sediment concentration through the course of a river flood. Furthermore, hyperpycnal flow deposits can be influenced by both local plunge-point dynamics and inlet river conditions, and the relative degree of influence depends on the advection length scale of settling sediment. Results also suggest that the criteria used to identify plunging hyperpycnal flows (a flow density in excess of the ambient fluid) is a necessary, but not sufficient condition. The basin also must be deep enough, in cases greater than tens of meters, in order for the plume to collapse and form a turbidity current.

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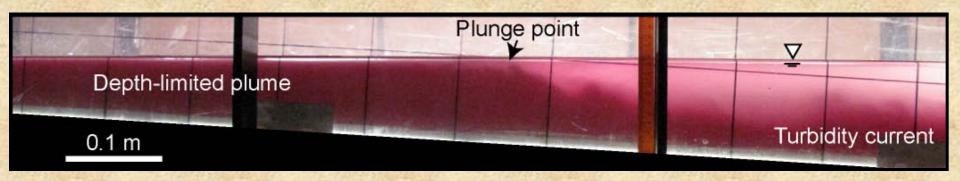
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Michael Lamb

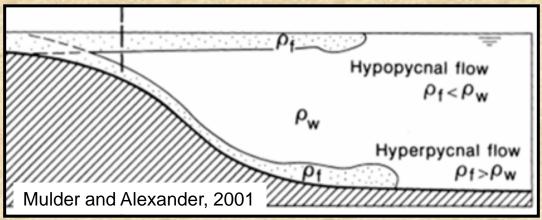
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Plunging hyperpycnal river plumes

Current Ideas:



- 1. River plumes plunge when denser than seawater
- 2. Depositional record of river floods (river response to climate change)

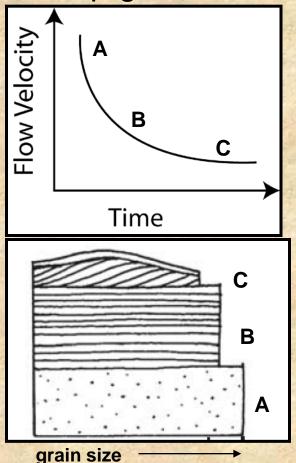


3. Deposits track the hydrograph of a flooding river

Do hyperpycnal-flow deposits record river flood dynamics?

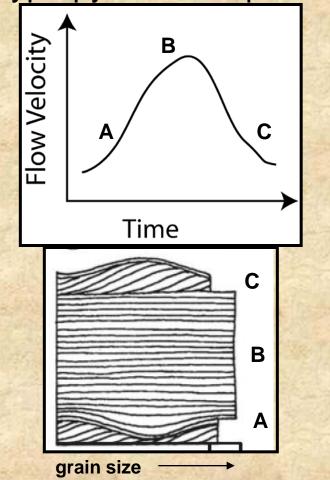
VS.

Slump-generated



Deposition under decelerating flow

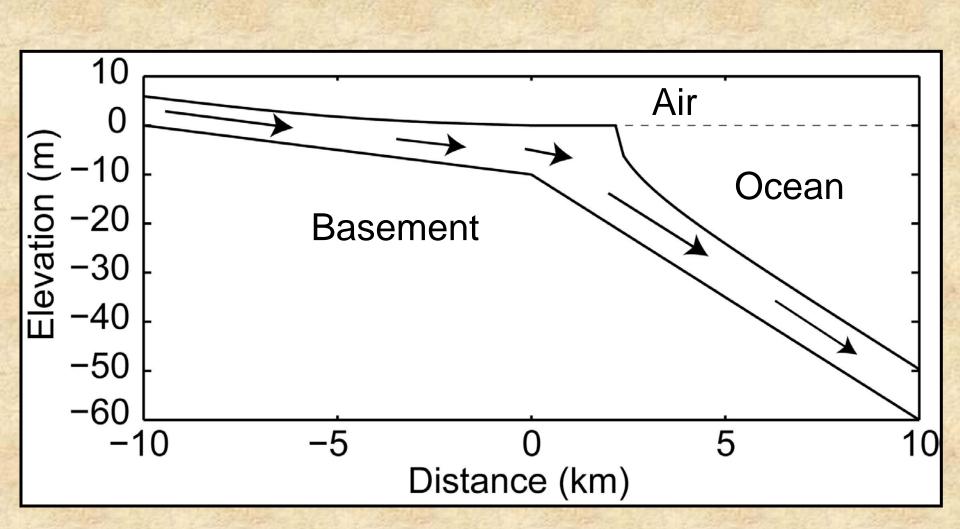
Hyperpycnal river plume



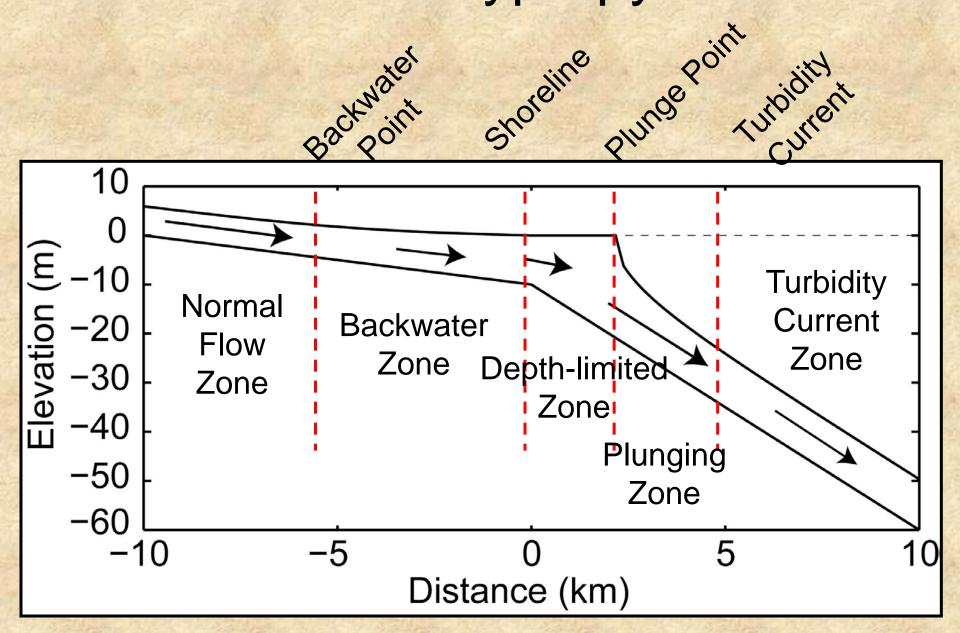
- Deposition following river hydrograph
- Reverse to normal grading

(after Mulder and Alexander, 2001)

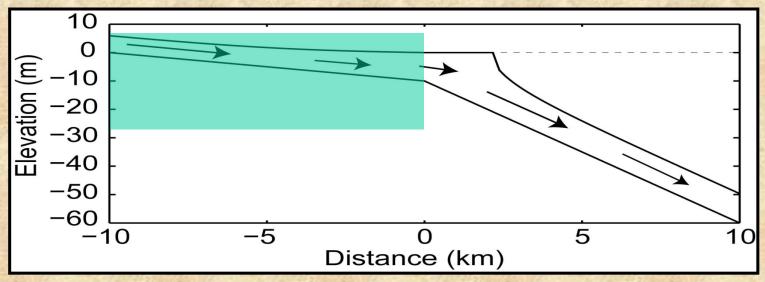
Zones of flow: Hyperpycnal Plume



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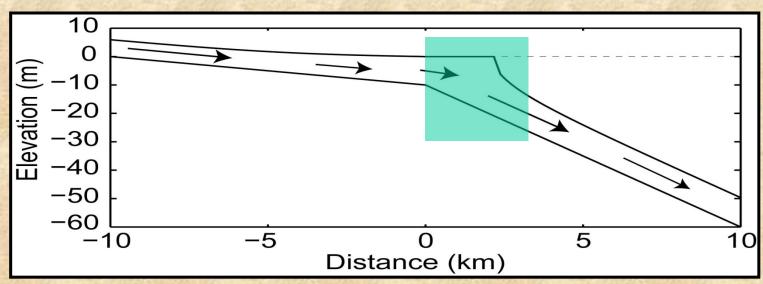
Numerical Formulation



- 1. 1D Layer averaged equations of motion (St. Venant)
- 2. Conservation of fluid mass and momentum
- 3. No erosion or deposition
- 4. Upstream of shoreline: backwater equation

$$\frac{dh}{dx} = \frac{S_b - C_f F^2}{1 - F^2}$$

Numerical Formulation



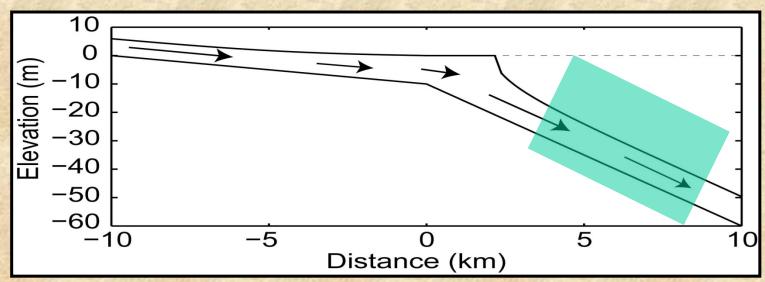
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5. Plunge zone: $F_{dp} = 0.5$ at plunge point (Akiyama and Stefan, 1984; Toniolo and Parker, 2007)

$$h_p = \left(\frac{q_p^2}{\left(\Delta\rho/\rho_a\right)gF_{dp}^2}\right)^{1/3}$$

Numerical Formulation

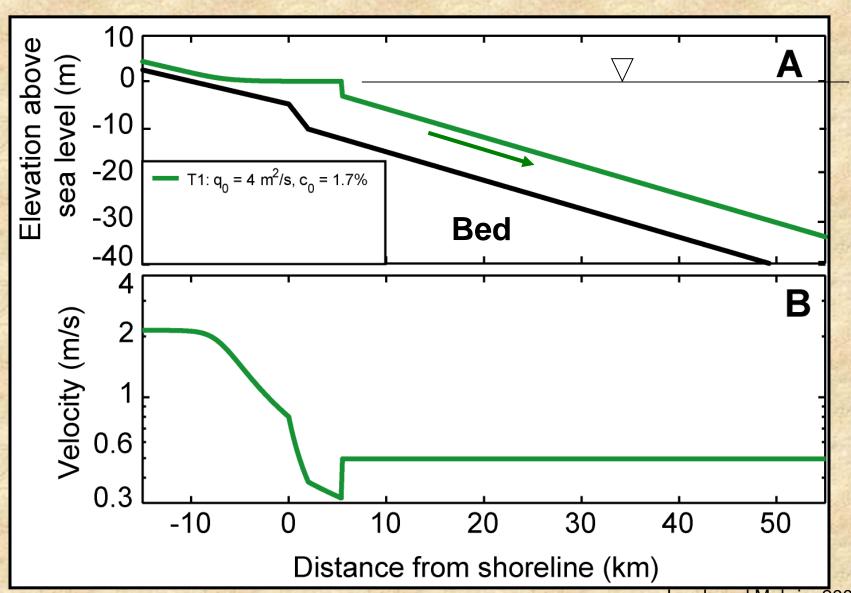


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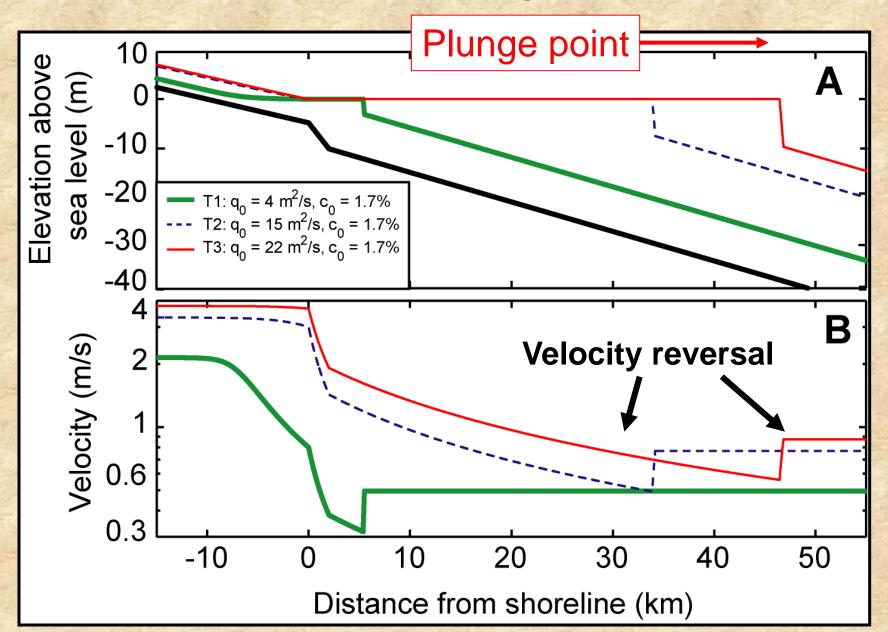
- 5. Plunge zone: $F_{dp} = 0.5$ at plunge point (Akiyama and Stefan, 1984; Toniolo and Parker, 2007)
- 6. Turbidity current zone: steady and uniform

Example case: Huanghe River, China

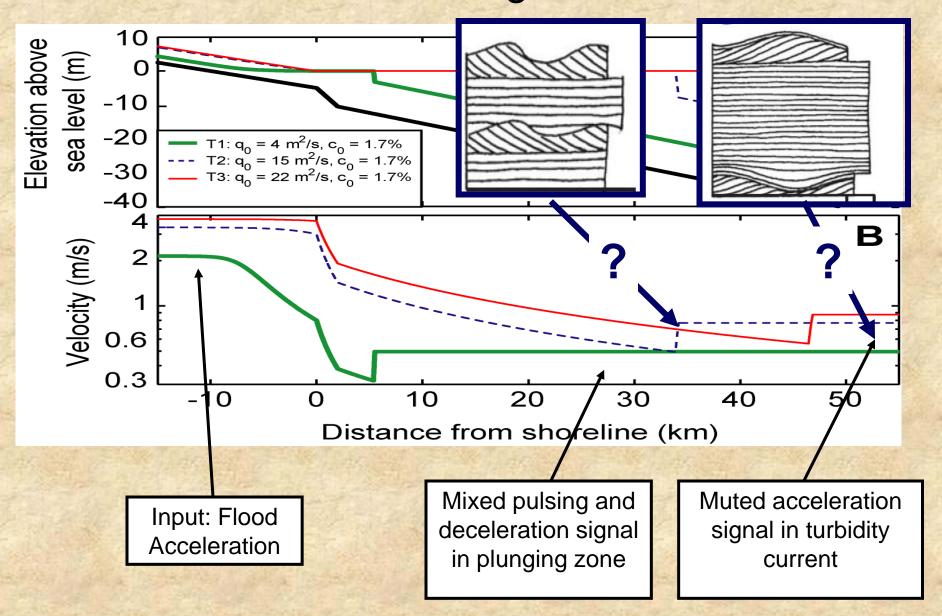


Lamb and Mohrig, 2009, Geology

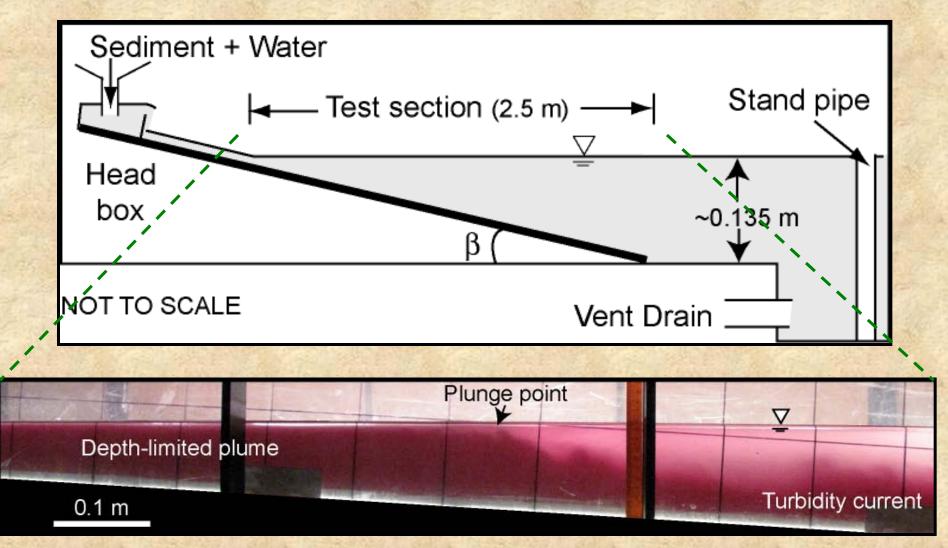
Source to sink signal transfer



Source to sink signal transfer



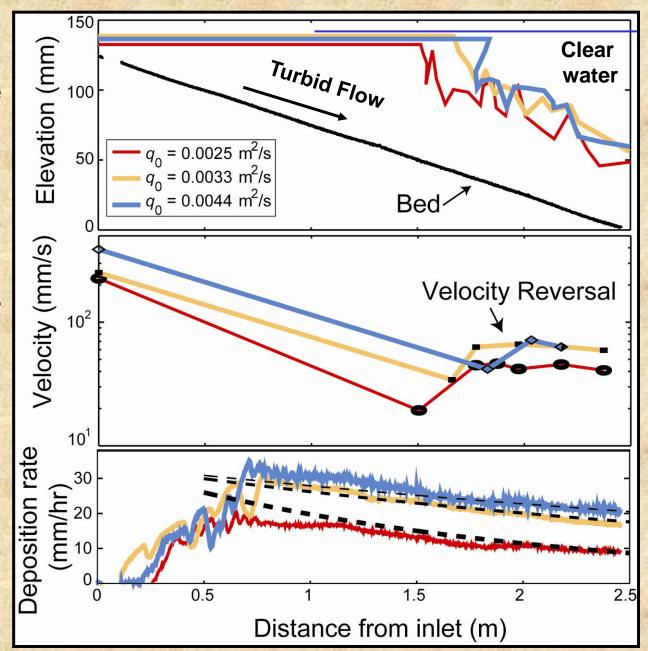
2-D experimental facility



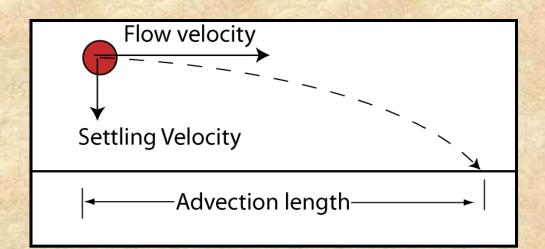
Flume Results

 Increasing discharge results in seaward movement of the plunge point.

- Local velocity is anticorrelated with discharge in plunge zone.
- Despite complex local flow dynamics, deposition rate is most sensitive to inlet discharge.



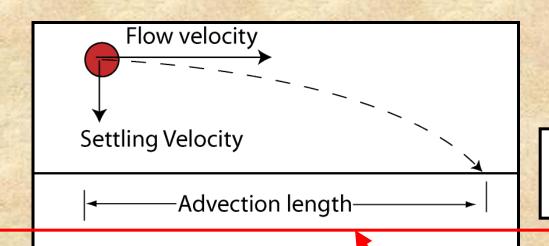
Advection length scale



$$l_a = \frac{UH}{r_0 w_s}$$

= Flow Velocity x settling time

Advection length scale



$$l_a = \frac{UH}{r_0 w_s}$$

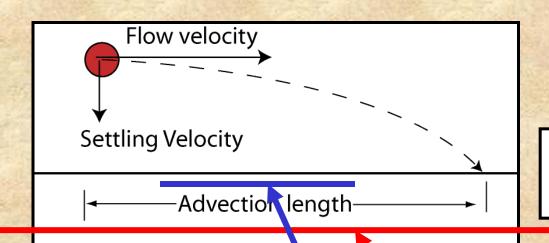
= Flow Velocity x settling time

If plunge point translation >
$$l_{_{\sigma}}$$

$$q_s = q_{sc}$$

Deposition governed by divergence in local transport capacity. Deposits record local plunge point dynamics.

Advection length scale



$$l_a = \frac{UH}{r_0 w_s}$$

= Flow Velocity x settling time

If plunge point translation >
$$l_{
m c}$$

$$q_s = q_{sc}$$

Deposition governed by divergence in local transport capacity. Deposits record local plunge point dynamics.

If plunge point translation <
$$l_a$$

$$q_s = q_{s0}$$

Settling sediment cannot respond to local flow. Deposition governed by inlet boundary conditions. Deposits record river flood dynamics.

Example calculations of advection length scale

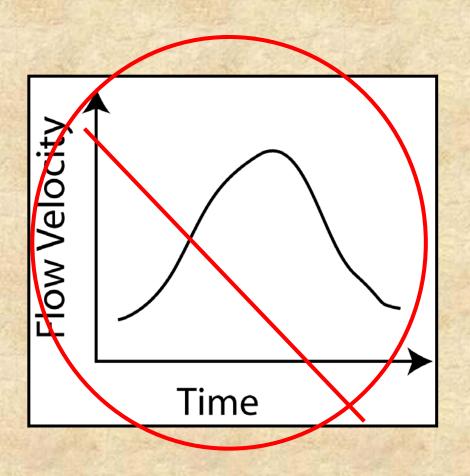
$$l_a = \frac{UH}{r_0 w_s}$$

		$q (m^2/s)$	w_s (m/s)	$\mathbf{l_a}$	Plunge poin translation	t
South Control of the	Laboratory	1 to 4 x 10 ⁻³	4.4 x 10 ⁻⁴	1.5 to 6 m	0.5 m	Inlet conditions
Control of the last	Field (silt)	5 to 20	4.4 x 10 ⁻⁴	7.5 to 30 km	10s km	Inlet conditions or local dynamics
では、これでは、これでは、これでは、これでは、これでは、これでは、これでは、これ	Field (sand)	5 to 20	9 x 10 ⁻³	0.3 to 1.2 km	10s km	Local dynamics

Sand is likely to record plunge point dynamics (not river flood discharge) at field scale

Conclusion 1: Flow Dynamics

Hyperpycnal plume *velocities* do not linearly follow the flood discharge curve.

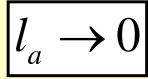


Depth-limited and plunging zones: anticorrelations, uncorrelations, and multiple accelerations and decelerations are possible due to the movement of the plunge point.

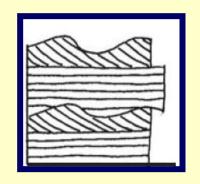
Conclusion 2: Depositional signature

Despite complex spatial changes in flow, plume deposits might still record inlet discharge and sediment concentration depending on the advection length scale.

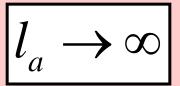
Low discharges and coarse sediment:



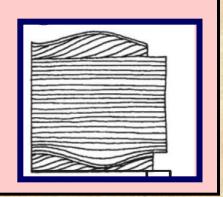
Deposits record local dynamics (plunge point translation – multiple pulses)



High discharges and fine sediment:



Deposits record inlet conditions (fluvial discharge and sediment concentration)



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