Non-Actualistic Carbonate Deposystems: Revising The Carbonate Factory-Depth Paradigm* By Giovanna Della Porta¹, Jeroen A.M. Kenter², and Paul M. (Mitch) Harris³

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Abstract

Non-actualistic (e.g., deviating from depositional models developed for recent deposystems) carbonate platforms are not the exception in the geologic record because of the wide range of biological and environmental factors controlling facies character and architecture.

A fundamental concept is the exponential decrease of carbonate production with increasing water depth from a maximum at shallow depths to the base of the photic zone. However, when the dominant factory is microbially induced, high rates of carbonate production extend on the slope down to dys-/aphotic depths and new production-water depth profiles must be developed.

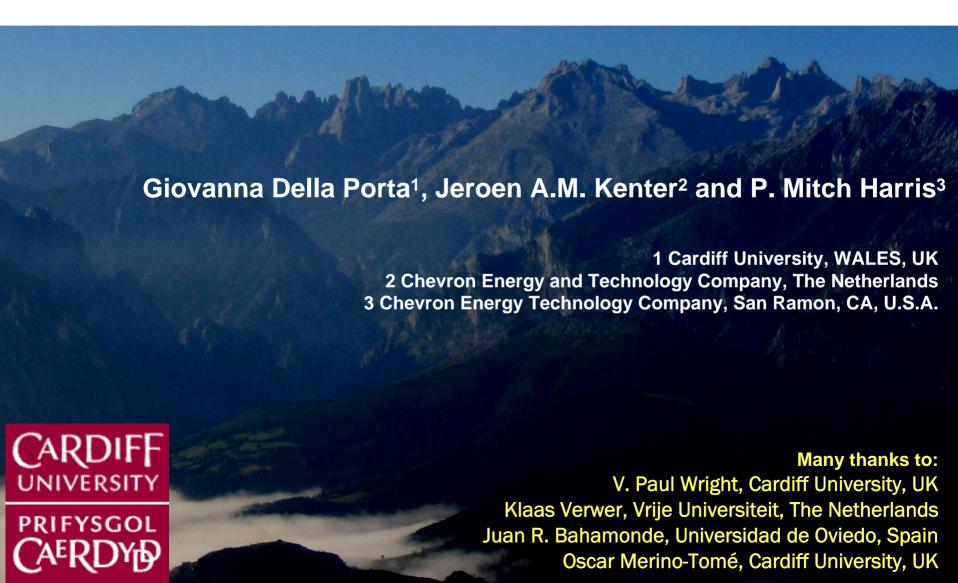
Studies of Carboniferous (Asturias, Pricaspian Basin) and Permian (Capitan Reef) high-relief carbonate platforms have shown that the microbial-boundstone production extends to 300m water depth: 1) the detrital lower slope consists mostly of matrix-free cemented rudstone sourced by the slope boundstone with subordinated platform-top-derived material; 2) carbonate production on the slope is controlled by environmental parameters (temperature, nutrients, oxygenation) that are water-depth dependent, but the microbial boundstone response to relative sea-level changes differs from modern reefs; 3) carbonate growth is not seriously reduced during sea-level falls because it can continue downslope, 4) progradation can take place at high rates despite the lack of platform-top shedding (slope vs. highstand shedding); 5) concepts of leeward progradational vs. windward aggradational margins have to be revised. Paleozoic high-relief platforms with microbial boundstone-dominated margins seem to have developed in mesotrophic, starved restricted basins with oxygen-depleted bottom waters that would not be suitable settings for the recent coral-reef rimmed platforms.

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NON-ACTUALISTIC CARBONATE DEPOSYSTEM: Revising The Carbonate Factory-Depth Paradigm



At the search for patterns

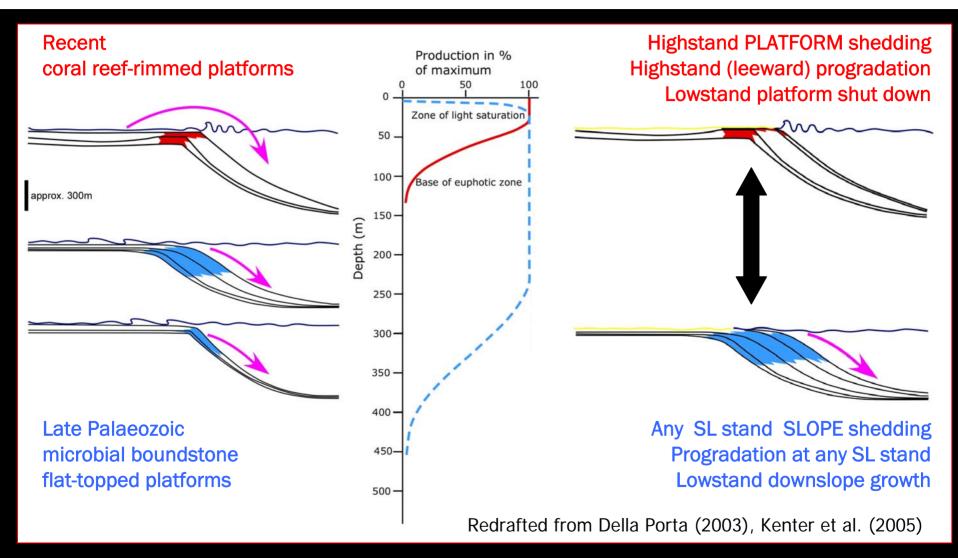
- Conceptual models are necessary for interpretation and prediction
- Model simplicity vs. stratigraphic complexity (space, time, incomplete record, diagenesis)
- Models need eventually to be revised

The carbonate factory represents the space where carbonate sediment is produced but also the processes that led to carbonate production

(from Schlager, 2003 after many sources)

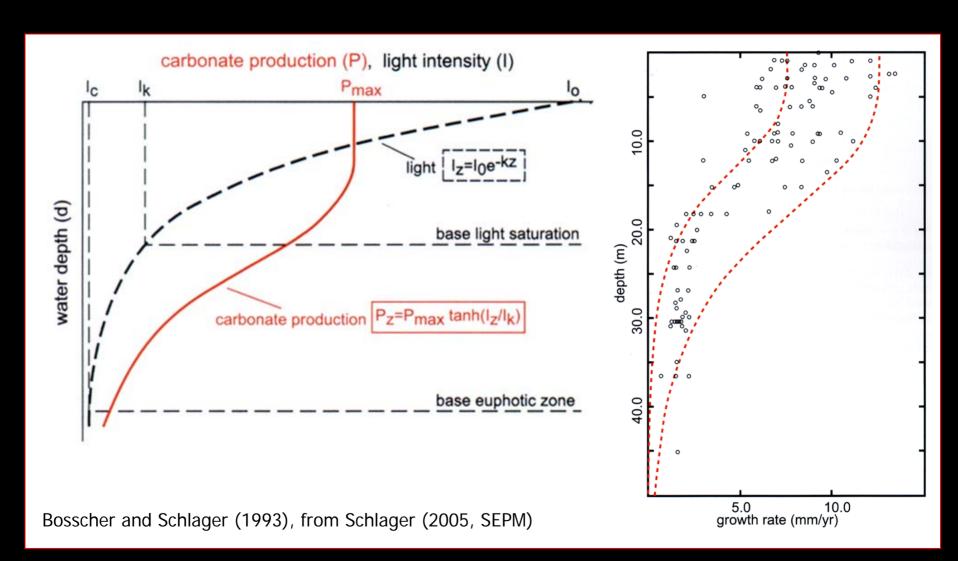
- Models developed for the Recent cannot always answer the complexity of the geologic record
- Little is known about many processes not in fashion in modern settings
- Several carbonate factories and lack of pure end-members (a continuum, spatially and temporally dynamic? cf. Wright and Burgess, 2005)
- Different <u>production rates</u>, <u>depth windows</u> and <u>DIFFERENT RESPONSES</u> to changes in accommodation space and environmental conditions

Non-actualistic carbonate platforms: the paradigm shift

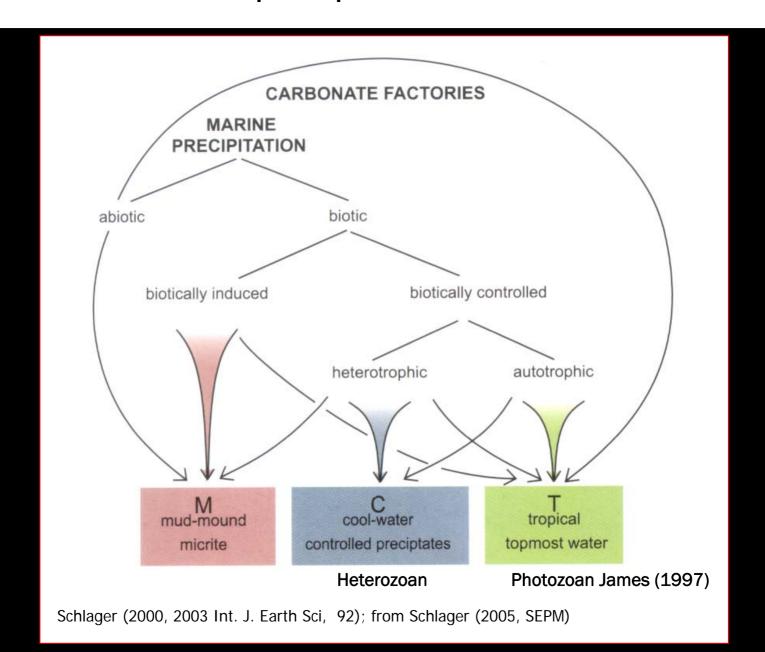


An exception? Carboniferous - Asturias, N Spain; Tengiz, Pricaspian Basin and Permian Capitan Reef, USA

(Coral reef) Carbonate Production vs. Depth



Modes of carbonate precipitation and carbonate factories



Microbial carbonates (biologically induced precipitation)

Organo-sedimentary deposits by benthic microbial communities

- trapping and binding detrital sediment
- forming the locus of mineral precipitation (Burne and Moore, 1987)

Processes

Biomineralization:

- Photosynthesis by cyanobacteria
- Heterotrophic bacteria (decay of organic matter via ammonification, nitrate and sulphate reduction).

Organo-mineralization: Non-living reactive organic substrates (bacteria, EPS, sponges)

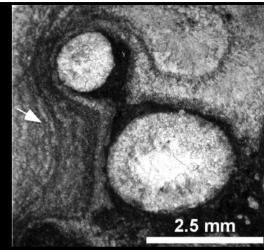
Products:

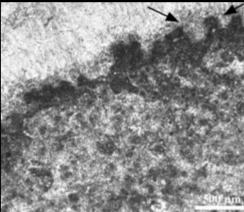
Microscale

- Aphanitic micrite (automicrite)
- peloidal
- laminae

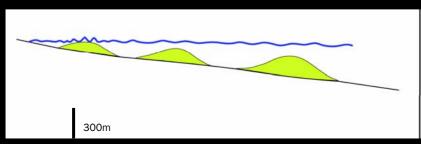
Macroscale

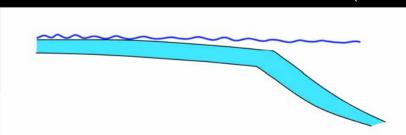
- Mounds
- Margin-slope



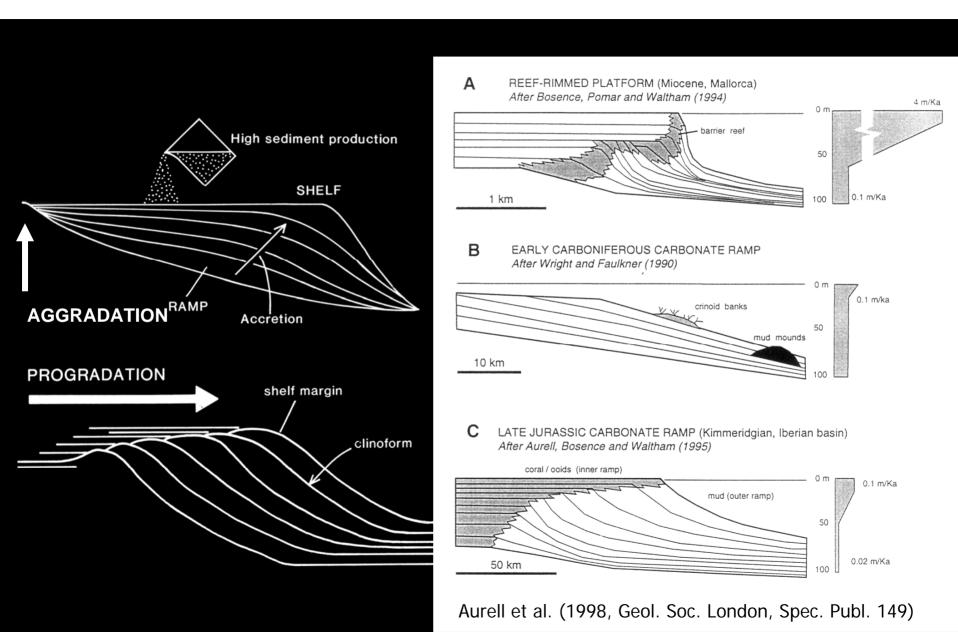


Carboniferous Asturias (Della Porta 2003)

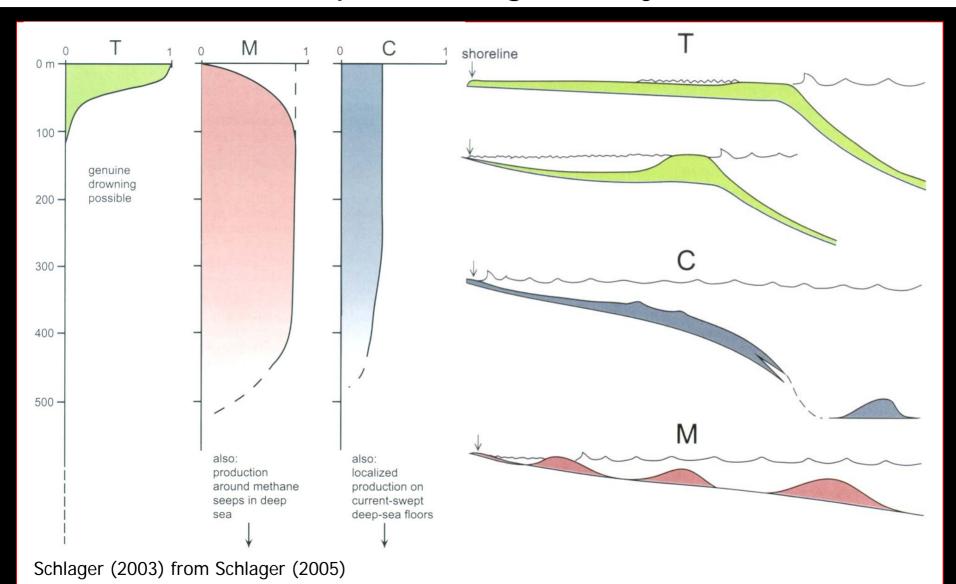




Modelling platform geometry



Production-depth profiles and depositional geometry



Types of carbonate platforms: a genetic approach

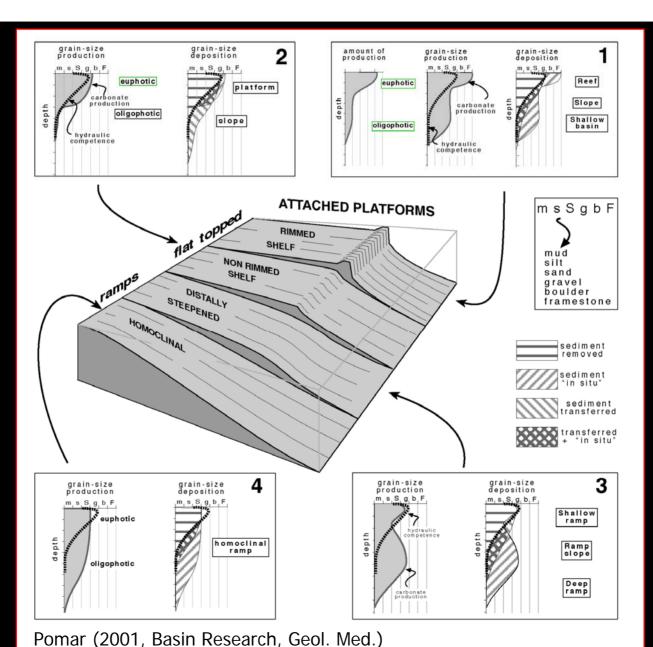
Euphotic gravel-size soft substrate dwellers (rudists)

Microbial Skeletal boundstone

Mud-producing biota

Euphotic removed by downshelf transport

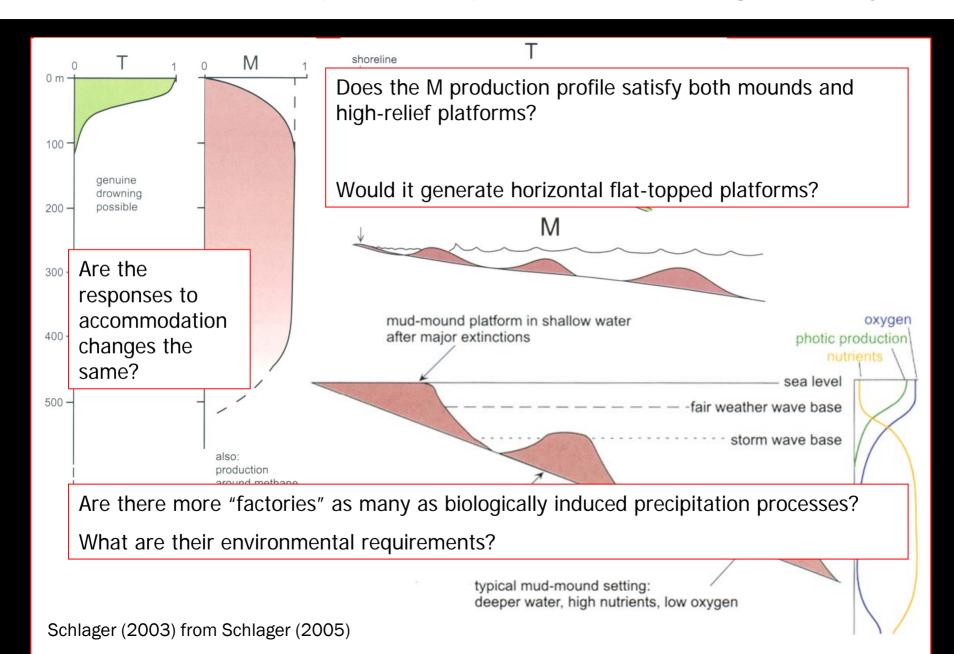
Photoindependent biota



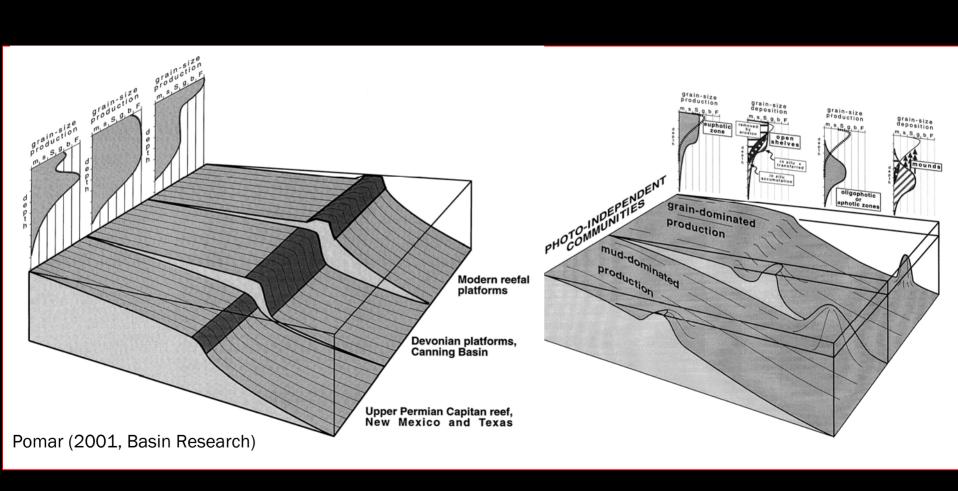
Euphotic frameworkproducing biota

Oligophotic coarse grained biota LBF, red algae

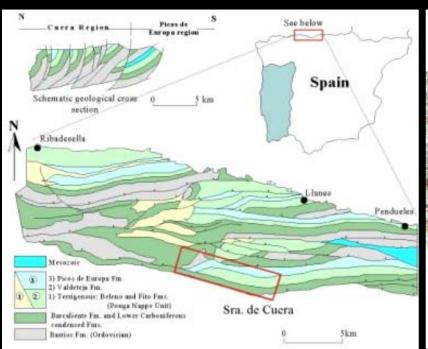
Non-actualistic platform production and geometry

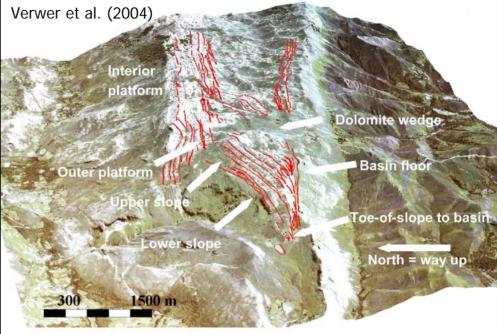


Non-actualistic platform production and geometry



Pennsylvanian high-relief platform (Asturias, N Spain)





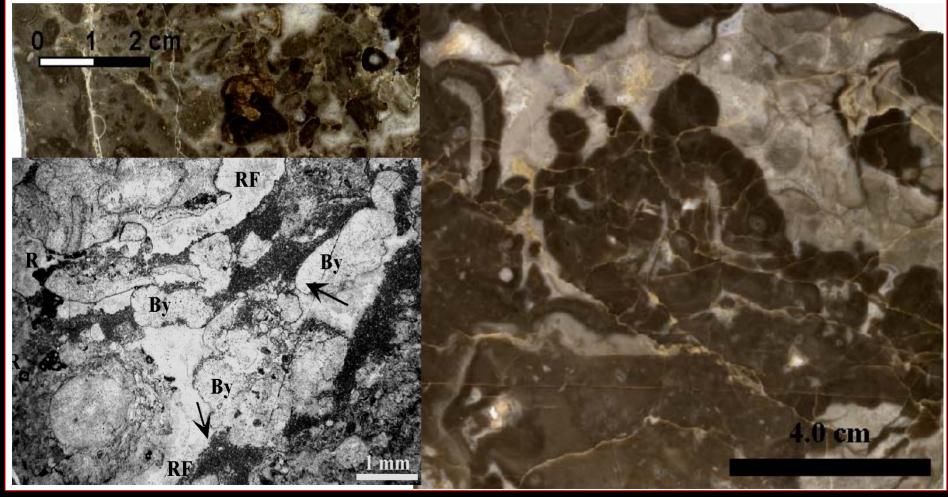


- high relief (600-800m)
- steep (30-40°)clinoforms
- >3 km of basinward progradation

Della Porta et al. (2003, 2004)

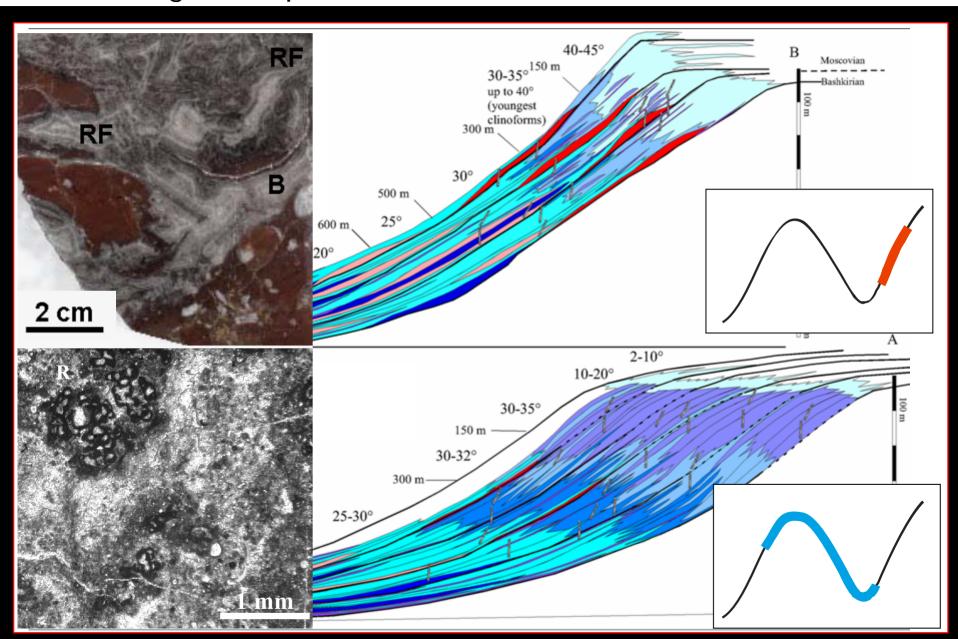
Self-nourished Slope

- Upper slope: Cement-rich microbial boundstone from 10-20m to 300-400 m depth (30-40% slope)
- Lower slope: Detrital matrix-free cemented boundstone breccias (50% slope)
- Minimum platform-top contribution (10-15% slope)

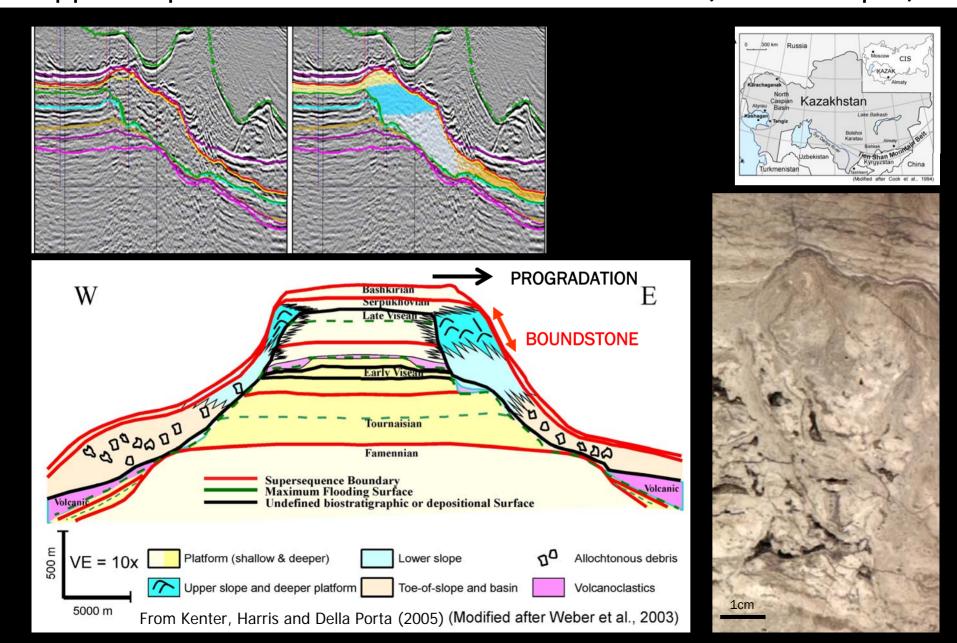


Della Porta et al. (2003, Facies); Kenter, Harris and Della Porta (2005, Sed. Geol.)

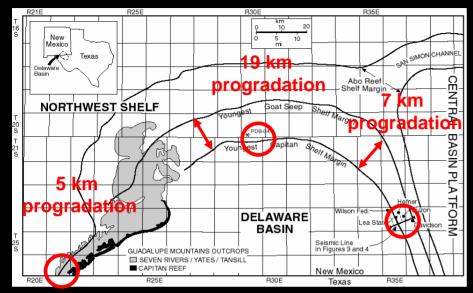
Progradational vs. aggradational slope built by light-independent microbialites: two factories?

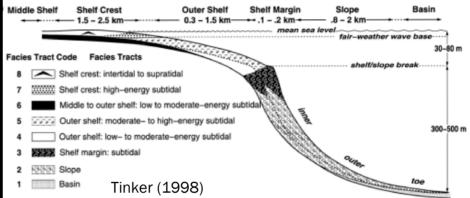


Tengiz (Carboniferous, Pricaspian Basin) upper slope cement-rich microbial boundstone (>400m depth)

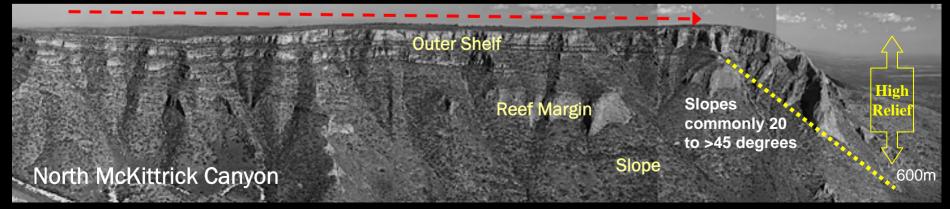


Capitan shelf margin (Permian, TX-NM) upper slope cemented microbial boundstone (30-150m depth)



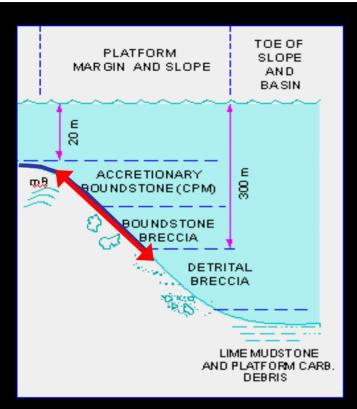


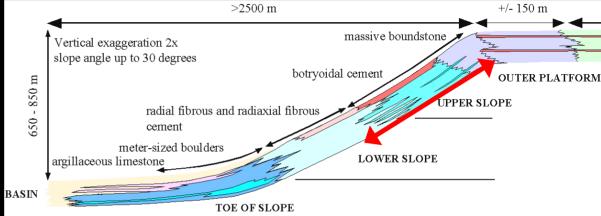
- Boundstone (calcisponges, bryozoans, Archaeolithoporella, botryoidal cement)
- Downslope in-situ boundstone breccias
- Toe of slope minimal platform top input
- Km's progradation



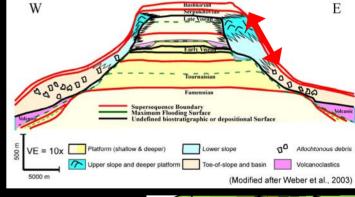
Harris and Saller (1999); Garber et al. (1989); Tinker (1998); Wood et al. (1994)

Depositional model late Palaeozoic steep high-relief microbial boundstone slopes

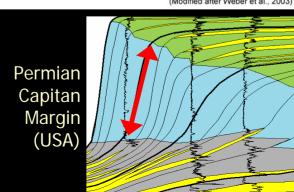




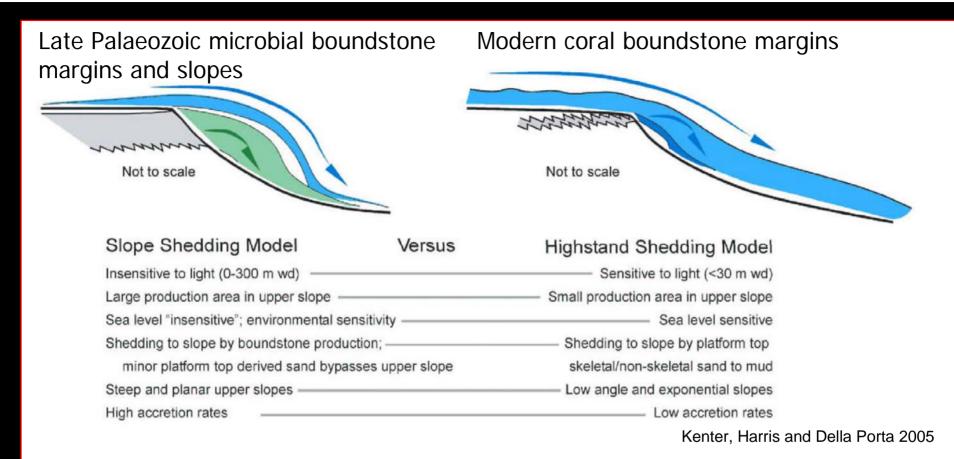
Carboniferous Northern Spain and Pricaspian basin



- Microbial dominated margin and slope factory
- Broad depth range (300-400m depth)
- Pervasive marine cementation (stable steep slopes)
- High production rate (400-1000m/My), km's progradation
- Detrital "talus" breccias cemented and matrix-free
- Minimal platform-top contribution



Slope shedding model

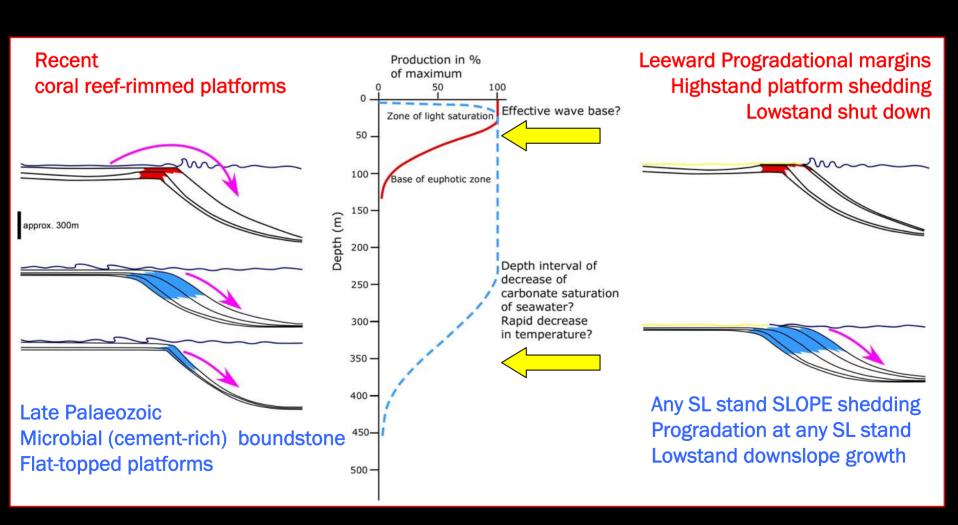


- Slope is "self-nourishing", little contribution from platform top
- High rates of steep slope progradation at any sea-level stand
- Progradation function of boundstone growth vs. off platform (Highstand) shedding

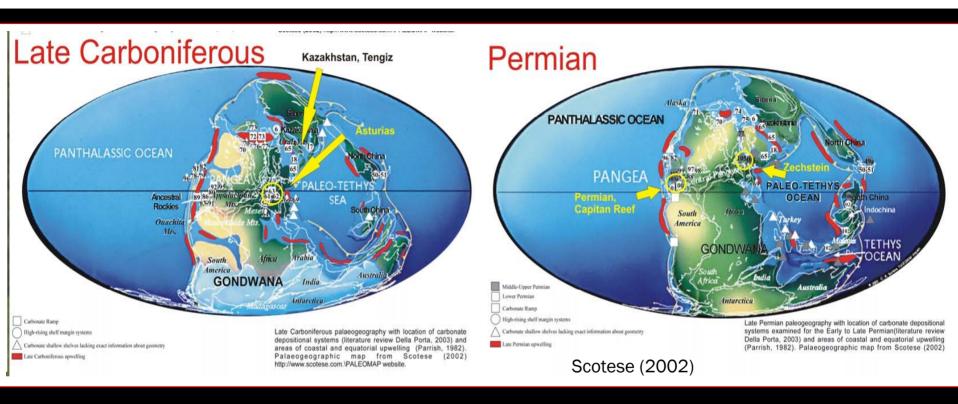
Model for microbial boundstone high-relief slope

Environmental controls on microbial boundstone growth rate and depth window?

- Tropical settings
- Highly supersaturated waters: high T, low PCO₂, high alkalinity, high pH
- Degradation reactive organic matter (bacteria, EPS, sponges), oxygen depletion.



What localizes microbial high-relief margins?



- high relief on antecedent topography
- Confined/restricted basins
- Oxygen-depleted sea floors (sulphate reduction, high alkalinity and increase saturation)
- Mesotrophic ... but phosphates inhibit aragonite and cyanobacteria-related precipitation
- Upwelling? Local overturning, no open ocean (deep cold waters reduce saturation)

Steep slopes vs. ramps: how many "microbial" factories?

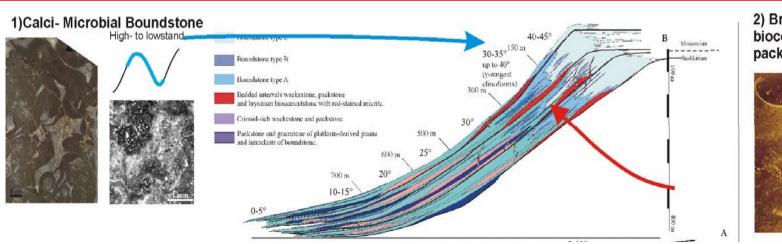
late Palaeozoic: many ramps with deep water mud-mounds, rare high-relief microbial boundstone slopes

Why many ramps with deep-water mud mounds did not evolve into high-relief?

- lower production rate confined in deeper water
- did not form in the ideal setting to promote high rates of growth

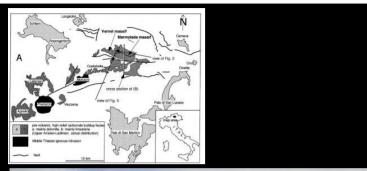
Did microbial high-relief platforms developed only after major extinction events and in specific oceanographic/environmental settings? How many microbial/ biologically induced carbonate factories?

- Calci-microbial vs. Heterotrophic bacteria vs. Organo-mineralization
- Or same "passive" precipitation processes have different rates according to physico-chemical conditions and organic substrates available?





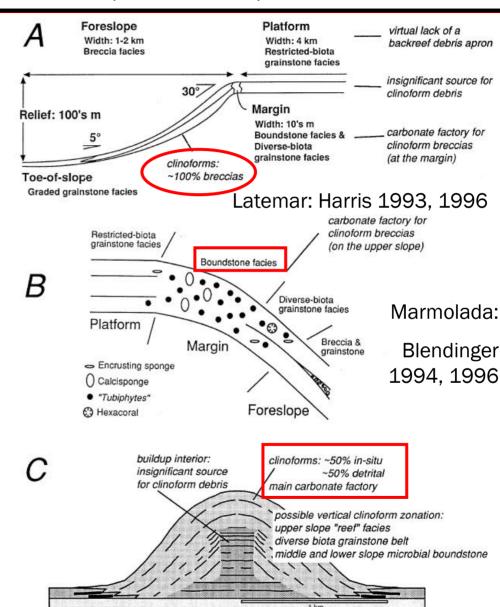
Where/What is the carbonate factory sourcing the progradational slope? Middle Triassic atolls (Dolomites)



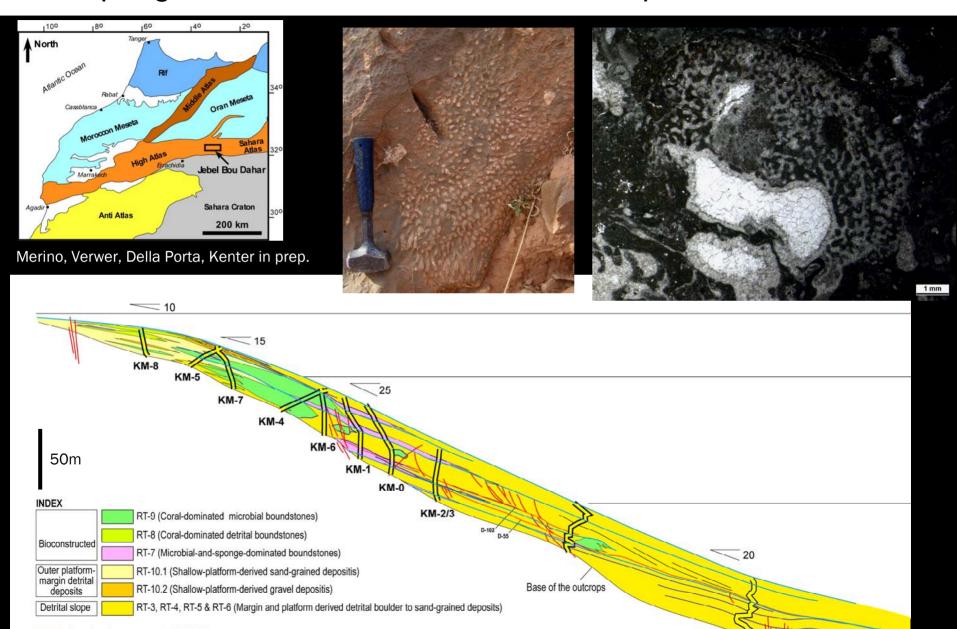


Sella Platform (Keim and Schlager 1999, 2001): Microbial micrite (automicrite) 20% upper slope, down to 200m depth

Cipit Boulders (Russo et al. 1997): Microbial micrite 60%



Lower Jurassic (High Atlas, Morocco) upper slope: sponge-microbial boundstones downslope of coral reefs



Summary

Depositional models based on the Recent need to be revised for margins with microbial cement boundstone or coral-microbial associations

- High production, extended depth window
- Carbonate growth not seriously reduced during sea-level falls
- Progradation of steep slope clinoforms independent from platform top shedding (slope vs. highstand shedding)
- Revision of leeward progradational vs. windward aggradational margins. Oceanographic control on margin growth?

The "model" provided for progradational, high relief microbial platform margins might be considered for other areas and times.

.....But many aspects of the microbial boundstone precipitation, cementation, and slope processes remain poorly understood.

For interpretation, modelling and prediction we need a better understanding of types, rates and the environmental controls of the "microbial" factories.