Submarine mass transport complex evolution and control on overlying reservoir deposition, Permian Cutoff Formation, west Texas

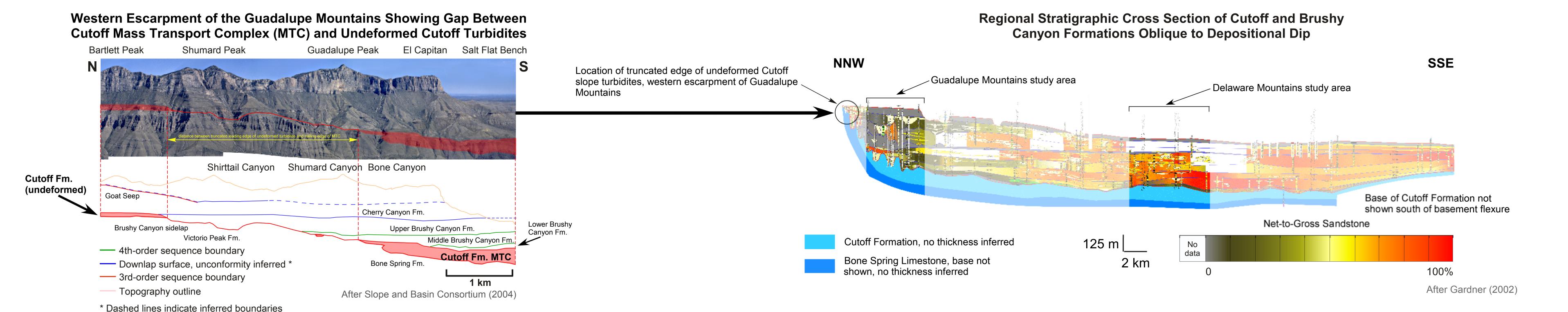
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Abstract

The Williams Ranch Member of the Cutoff Formation consists of six offlapping, basinward-stepping lithologic units of highstand carbonate turbidites deposited across a drowned Early Permian carbonate platform, then partially redistributed in slumps on the slope and basin floor. Slumps are intercalated with undeformed carbonate turbidites; the ratio of slumps to undeformed sediment increases basinward. Upslope evacuation scars correlate to downslope slump bodies. Gravity flow deposition and subsequent mass movement caused basinward thickening of the Williams Ranch Member and caused the toe of slope to shift basinward relative to the underlying Bone Spring Limestone. This shift controlled landward pinchouts of the overlying Permian Brushy Canyon Formation channel and sheet sandstone bodies.

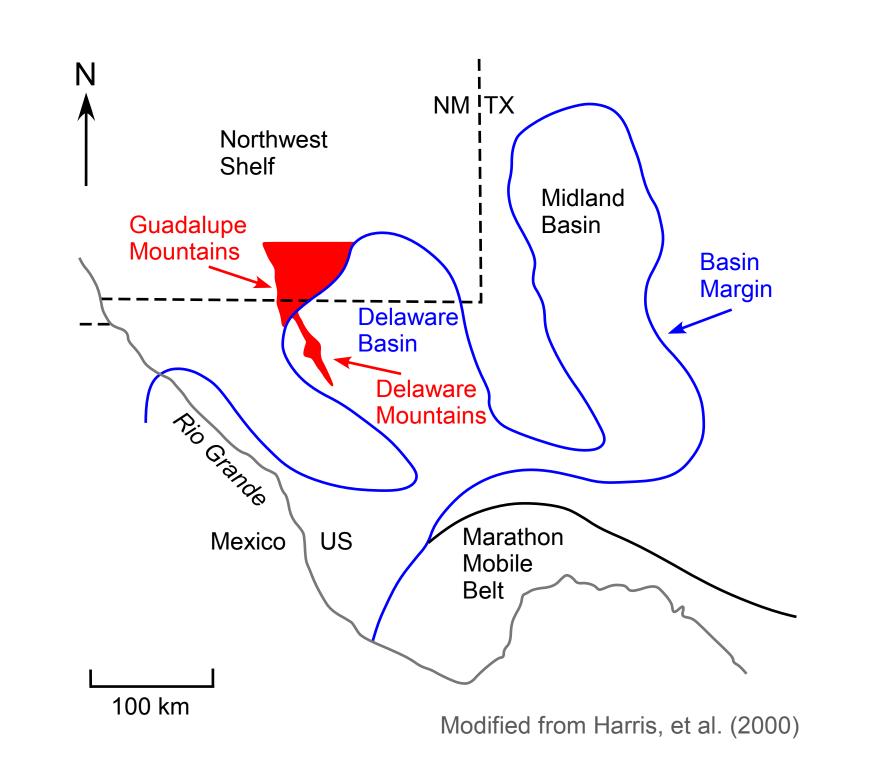
Williams Ranch Member deposition both responded to and modified inherited bathymetric relief. Williams Ranch isopach thicks correspond to larger underlying lows and smaller overlying highs and exhibit a higher ratio of undeformed to slumped sediment and a higher percentage of soft-sediment folds relative to softsediment truncation surfaces. These slump "pile-ups" appear to be concentrated in inherited lows. Slump bodies show a general southward transport vector, with significant local variation possibly reflecting underlying bathymetric influence. With repeated slump events, "pile-up" zones resulted in local positive bathymetry. Brushy Canyon sand fairways and ponded sheets are focused in bathymetric lows, and sands are sidelapped against highs atop the Williams Ranch Member. Increased understanding of mass transport complex evolution may lead to better prediction of overlying reservoir geometry, both within the Brushy Canyon Formation and in analogous reservoirs in other deepwater settings.



Research Questions

- 1) Can deformed Cutoff units in the southern Guadalupe and northern Delaware Mountains be correlated with the three formal members and five informal correlation units of Harris (1982b; 1987; 1988a; 1988b; 2000)?
- 2) Can a sequence stratigraphic framework be established for the Cutoff that correlates to work by Fitchen (1993; 1997), Kerans and others (1992; 1993), Kerans and Fitchen (1995), Sarg (1986; 1989), Rossen and others (1988), Sarg and Lehmann (1986a; 1986b), Sonnenfeld (1993) and Sonnenfeld and Cross (1993) on equivalent shelfal units?
- 3) How many mass transport events (MTEs) are represented in the Cutoff?
- 4) What were the MTE transport vectors?
- 5) How far did the MTE bodies (bodies of sediment transported to their present locations, each in a single MTE) travel?
- 6) Does the nature of the MTEs evolve over time?
- 7) Do contractional and/or extensional domains exist, and where are they located?
- 8) How do microscale (mms-cms) features relate to mesoscale (ms) and macroscale (100s ms-kms) features?
- 9) How does the thickness of the mass transport complex (MTC) vary within the basin?
- 10) How did the Cutoff Formation influence development of paleo-bathymetry below the Brushy Canyon Formation?
- 11) How can internal Cutoff Formation structure be used to predict paleo-bathymetry below the Brushy Canyon Formation?

Paleo-Geographic Map of the Permian Basin



Geologic Map of Cutoff and **Lower Brushy Canyon Formations** Blue line: Location of regional cross section Red line: Location of study area cross section

Orange box: Guadalupe Mountains study area Green box: Delaware Mountains study area

Guadalupe

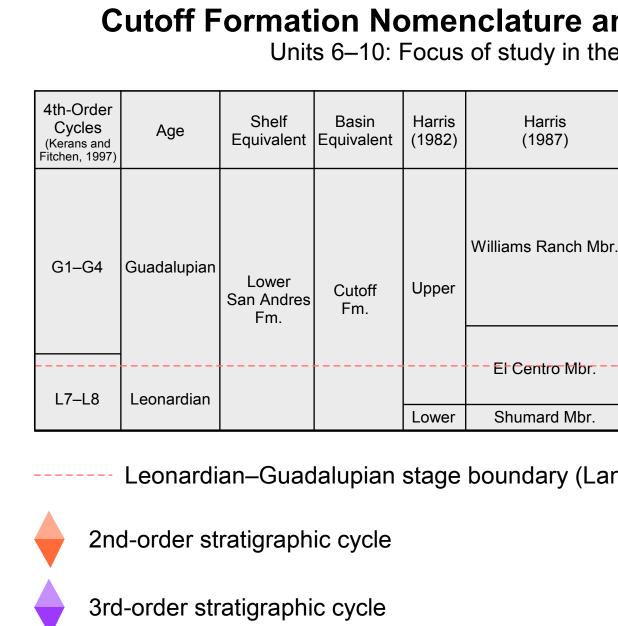
Mountains

Tertiary Normal Faults

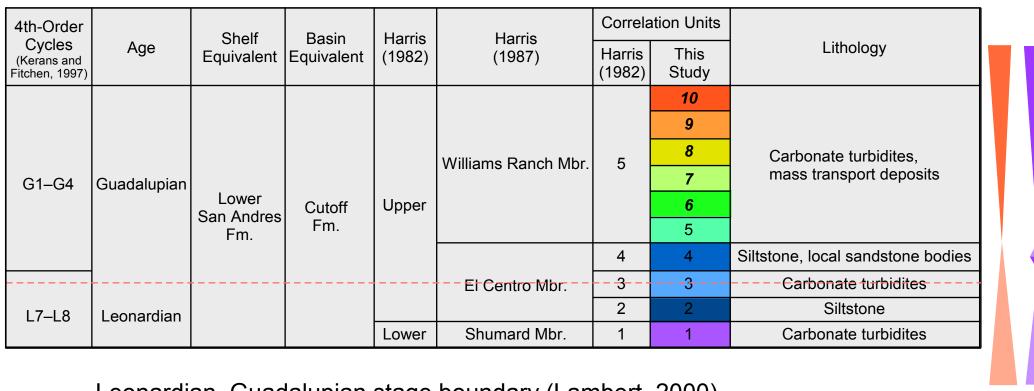
Brushy Canyon Fm.

from Slope and Basin Consortium (2003).

Fault locations from Kullman (1999). Lithologic contacts



Cutoff Formation Nomenclature and 2nd–4th Order Cyclicity Units 6-10: Focus of study in the Delaware Mountains



Leonardian–Guadalupian stage boundary (Lambert, 2000)

