So Different, Yet So Similar: Comparing and Contrasting Siliciclastic and Carbonate Slopes*

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Abstract

Carbonate slopes have a tendency to be steeper than their clastic counterparts. Commonly the stabilization potential by binding of slope sediment and early cementation of carbonates is evoked to explain this difference. However, differences and similarities between clastic and carbonate slope systems with respect to their gross development, curvature, and angle of dip are only expressed if one evaluates slope settings that are affected by comparable extrinsic and intrinsic processes. The likeness of clastic continental slopes and cool-water carbonate platforms is great where deep shelves, low-slope angles and usually sigmoidal-slope profiles are typical. Coarse-grained deltas compare with tropical carbonate platforms. Both have steep depositional slopes, exponential and linear slope profiles, and coarse sediments originating from shallow-water depths. Exponential profiles are common on rimmed platforms because reefs are resistant to erosion and the platform edge therefore relatively stationary vertically. This also accounts for ice-covered margins because the grounding level of the ice limits vertical fluctuations. A special case for carbonates is the in situ accretionary slope factory with abiotic and biotically induced precipitates stabilizing and building carbonate slopes. However, in situ slope accretion and stabilization in itself do not necessarily explain the largescale geometry of the platform flanks. It is more reasonable that the slope factory is insensitive to light and can therefore accrete during both low- and highstands. Thus, when a relative sea-level fall exposes the platform top and shallow-water carbonate production stops, in situ carbonate production continues in the slope realm. The combined effort of both types of sediment production and hence surplus allow the system to build-up to the angle of shear and constantly prograde. A direct comparison is coarse-grained fjord and lake deltas, where the inherent fast-prograding system, which is dominated by a mixture of coarse sand and rubble, obtains steep, planar slopes. Clearly, while sediment properties vary greatly, stark similarities in gross development, curvature, and angle are observed in comparable settings.

Selected References

Adams, E.W., 2001, Subaquatic slope curvature and its relation to sedimentary processes and sediment composition: Ph.D. Dissertation, Vrije University, Amsterdam, 135 p.

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Adams, E.W., W. Schlager, and F.S. Anselmetti, 2001, Morphology and curvature of delta slopes in Swiss lakes: lessons for the interpretation of clinoforms in seismic data: Sedimentology, v. 48, p. 661-679.

Carter, L., and R.M. Carter, 1988, Late Quaternary development of left-bank-dominant levees in the Bounty Trough, New Zealand: Marine Geology, v. 78/3-4, p. 185-197.

Eberli, G.P., F.S. Anselmetti, L.A. Melim, and J.A.M. Kenter, 1997, Facies, diagenesis and petrophysics of a prograding carbonate platform margin, Neogene, Great Bahama Bank, *in* B. Martindale, and J. Wood, (chairpersons), "CSPG-SEPM joint convention; Sedimentary events, hydrocarbon systems: Society of Petroleum Geologists, Canada, and Society for Sedimentary Geology, United States, p. 471-490.

James, N.P., D.A. Feary, F. Surlyk, J.A.T. Simo, C. Betzler, A.E. Holbourn, Q. Li, H. Matsuda, H. Machiyama, G.R. Brooks, M.S. Andres, A.C. Hine, M.J. Malone, and Ocean Drilling Program Leg 182 Scientific Party, 2000, Quaternary bryozoan reef mounds in cool-water, upper slope environments: Great Australian Bight: Geology, v. 28/7, p. 647-650.

James, N.P., and C.C. von der Borch, 1991, Carbonate shelf edge off southern Australia; a prograding open-platform margin: Geology, v. 19/10, p. 1005-1008.

Jansa, L.F., 1991, Lithostratigraphy 10, carbonate buildup morphology, *in* J.L. Bates, (ed.), East Coast Basin Atlas Series: Atlantic Geoscience Centre, Geological Survey, v. 69.

Kenter, J.A.M., 1990, Carbonate platform flanks; slope angle and sediment fabric: Sedimentology, v. 37/5, p. 777-794.

Kirkby, M.J., 1987, General models of long-term slope evolution through mass movement, in M.G. Anderson, and K.S. Richards, (eds.), Slope stability; geotechnical engineering and geomorphology: John Wiley & Sons Chichester, United Kingdom, p. 359-379.

Kuehl, S.A., B.M. Levy, W.S. Moore, W.S., and M.A. Allison, 1997, Subaqueous delta of the Ganges–Brahmaputra river system: Marine Geology, v. 144, p. 81-96.

Kuvaas, B., and G. Leitchenkov, 1992, Glaciomarine turbidite and current-controlled deposits in Prydz Bay, Antarctica: Marine Geology, v. 108, p. 365-381.

Stagg, H.M.J., 1985, The structure and origin of Prydz Bay and Mac-Robertson Shelf, East Antarctica: Tectonophysics, v. 114, p. 315-340.

Vail, P.R., R.M. Mitchum Jr., R.G. Todd, J.M. Widmier, S. Thompson, III., J.S. Sangree, J.N. Bubb, and W.G. Hatlelid, 1977, Seismic stratigraphy and global changes of sea level, *in* C.E. Payton, (Ed.), Seismic Stratigraphy-applications to hydrocarbon exploration: AAPG Memoir 29, p. 49-211.

Verwer, K., O. Merino-Tomé, J.A.M. Kenter, and G. Della Porta, 2009, Evolution of a high-relief carbonate platform slope using 3D digital outcrop models: lower Jurassic Djebel Bou Dahar, High Atlas, Morocco: JSR, v. 79, p. 416-439.

Verwer, K., J.A.M. Kenter, B. Maathuis, and G. Della Porta, 2004, Stratal patterns and lithofacies of an intact seismic-scale Carboniferous carbonate platform (Asturias, northwestern Spain): a virtual outcorp model: Geological Society of London, Speical Publications, v. 239, p. 29-41.



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AAPG 2012, Long Beach, California



In memory of Lorenz Keim

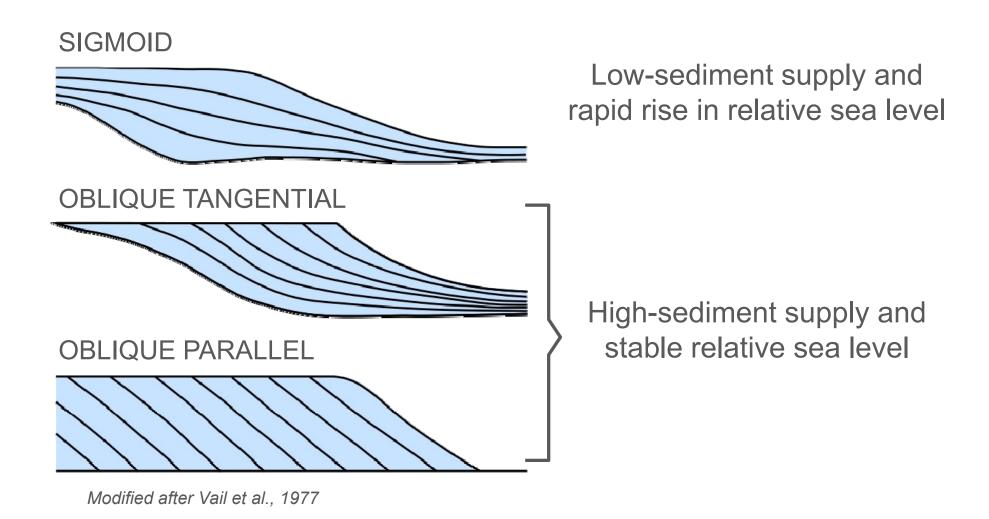
■ Lorenz Keim, outstanding Dolomites geologist, perished in a snow avalanche on February 4 and only 43 years old, leaving behind a wife and three children and a big hole in South Tyrolean geology.



Summary and take-away message

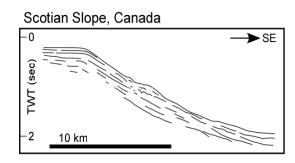
- Carbonates versus siliciclastic slopes:
 - Production near platform top (mostly during highstand shedding), or in situ on the upper slope (during all systems tracts) versus source-to-sink
 - Early lithification of carbonates
 - Petrophysical properties, i.e., shear-strength, pore systems, heterogeneity, acoustic properties
- The established view:
 - Carbonate slopes tend to be steeper than their clastic counterparts
 - The stabilization potential by binding of slope sediment, early cementation, and in situ growth of carbonates is evoked to explain this difference
- This talk:
 - Stark similarities in gross development, curvature, and angle are observed if similar settings, situations or processes prevail

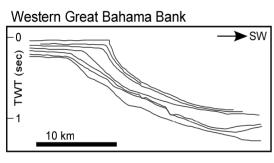
Classification of prograding clinoforms based on seismic

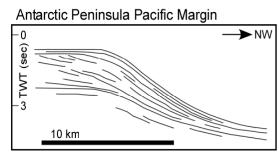


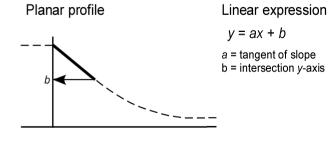
Quantifying slope curvature

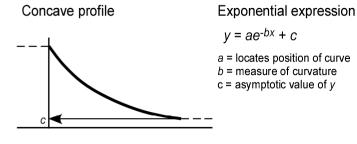
- Database comprising150 modern slopes
- Curve fitting on firstorder morphology
- Three equations quantify 90% of database
- Three basic types of slope curvature
 - Planar→ Linear
 - Concave → Exponential
 - Sigmoidal → Gaussian

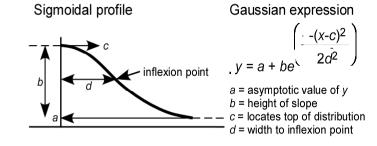








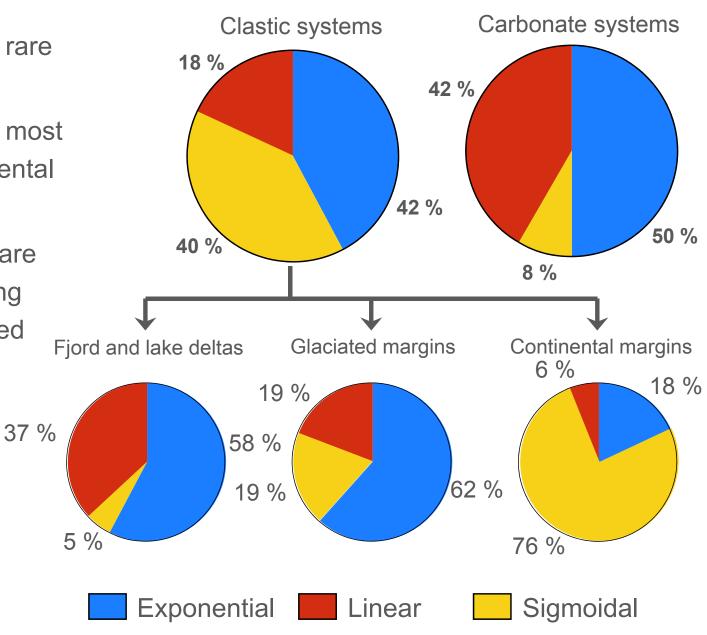




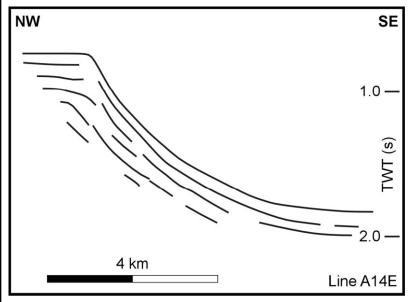
Modified from Adams and Schlager, 2001

Distribution of slope curvature

- Sigmoidal profiles rare in carbonates
- Sigmoidal profiles most common in continental margins
- Carbonates compare with fast-prograding deltas and glaciated margins



Exponential profile - Southeast South Island, New Zealand



shelfbreak

-0.3

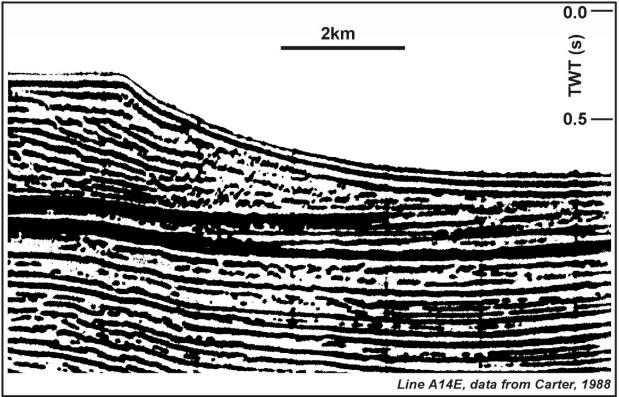
(Ex)
H Ld H Color of the shelfbreak

y=-0.37e^{-0.44x}+0.60

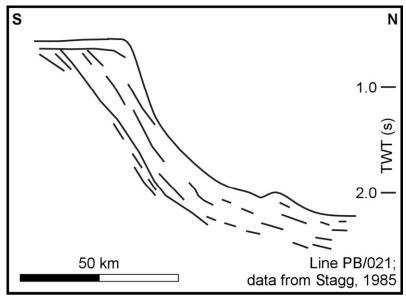
DISTANCE (km)

0 2 4

- Rapidly prograding, continental-shelf delta
- Terrigenous sand and silt
- Maximum inclination 10°



Exponential profile - Prydz Bay, Antarctica



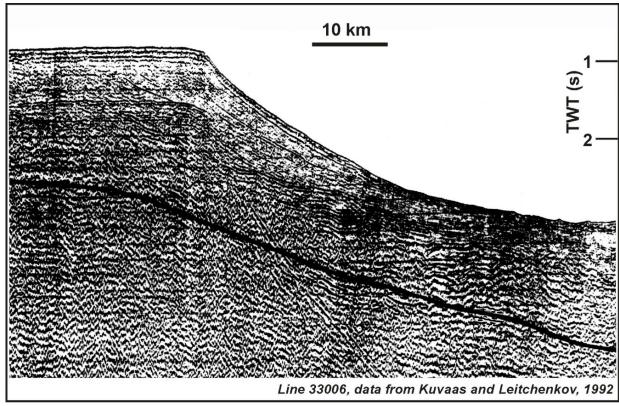
shelfbreak

-0.2 (w/y) HLd (w/y) = -2.06e -0.07x + 2.44

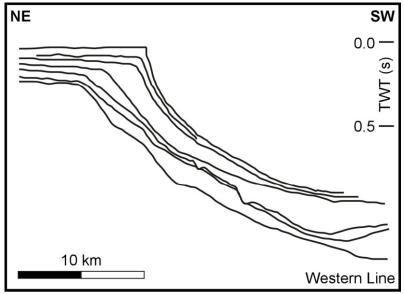
-0.6 DISTANCE (km)

0 20 40

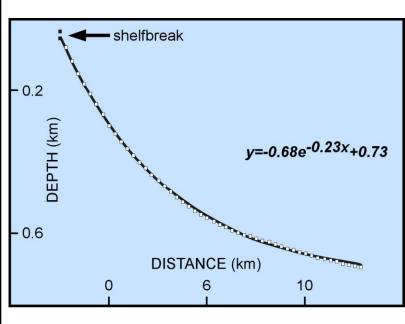
- Glaciated, continental margin
- Poorly sorted terrigenous diamictite
- Maximum inclination 8°

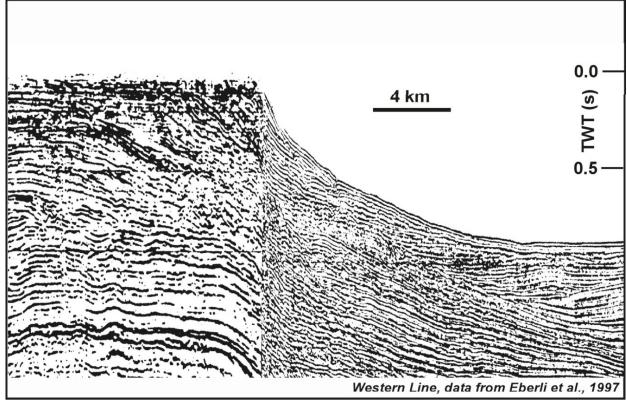


Exponential profile - Western Great Bahama Bank



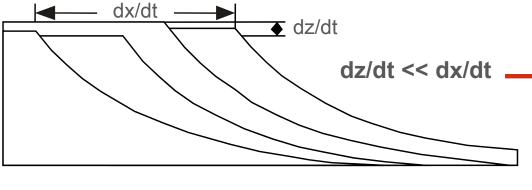
- Tropical, carbonate platform
- Carbonate mud and fine sand
- Maximum inclination 8°

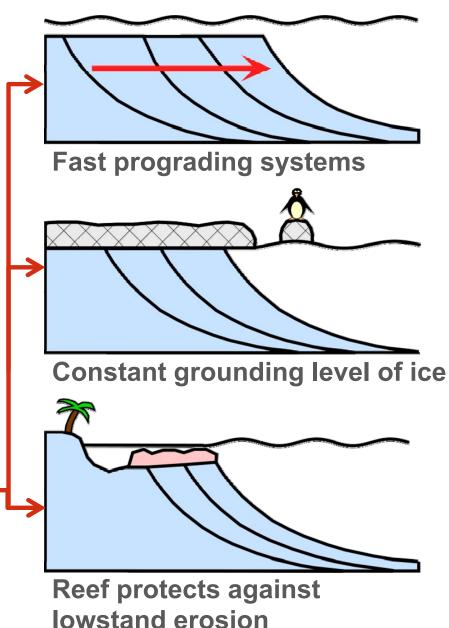




Exponential profiles

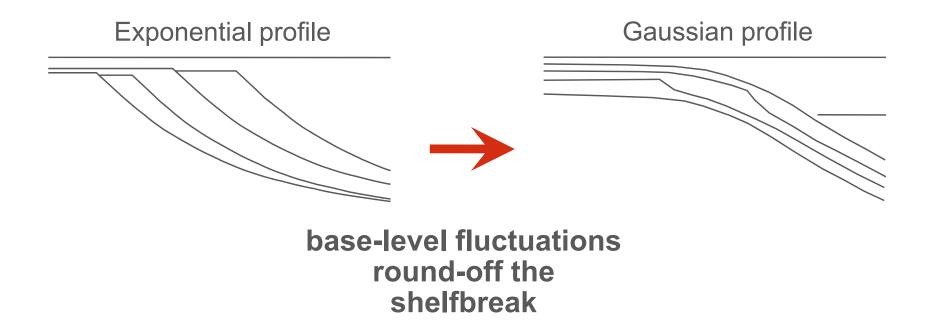
- Exponential curvatures and abrupt shelf breaks develop when the ratio between rate of vertical fluctuations of base level to rate of horizontal progradation is small
- Fast prograding systems with minor base-level fluctuations
- 2. Grounding ice sheets
- 3. Reef protection



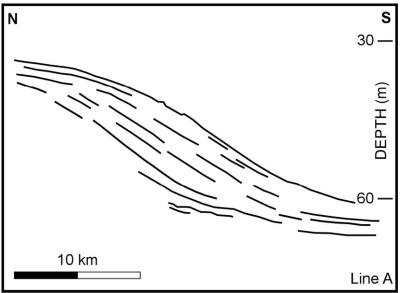


Exponential profiles and their modification to Gaussian

 Sigmoidal curvatures develop if base-level fluctuations round-off the shelfbreak



Gaussian profile - Ganges-Brahmaputra delta



10 km
Line A

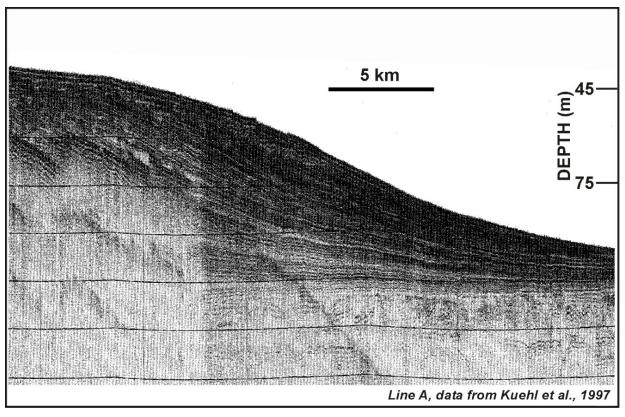
shelfbreak

-0.06 (w/y) H
-0.08 0

y=0.10-0.06e

-(x+0.93)²/_{504.99}

- Storm-dominated, continental-shelf delta
- Terrigenous silt and fine sand
- Maximum inclination 0.2°

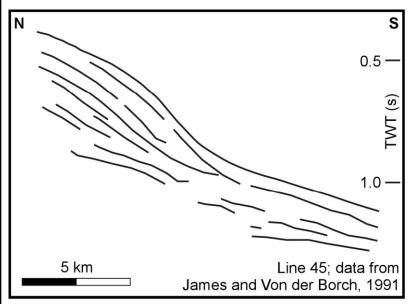


5

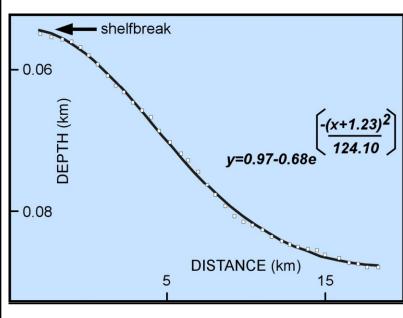
DISTANCE (km)

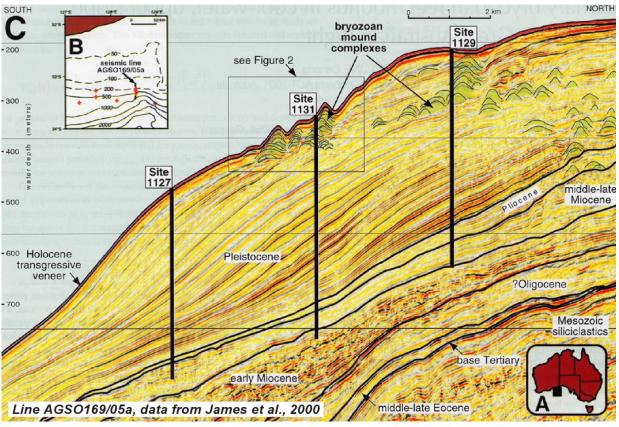
15

Gaussian profile - Great Australian Bight



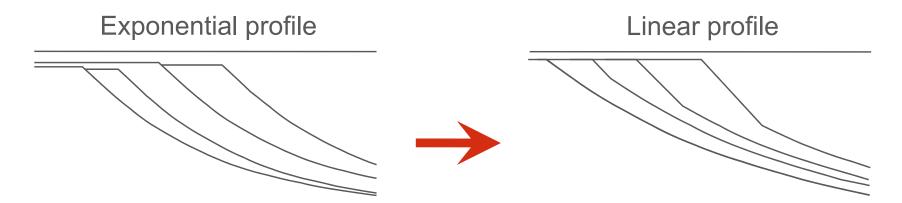
- Open shelf, cool-water carbonate
- Fine-grained skeletal carbonates
- Maximum inclination 3.0°





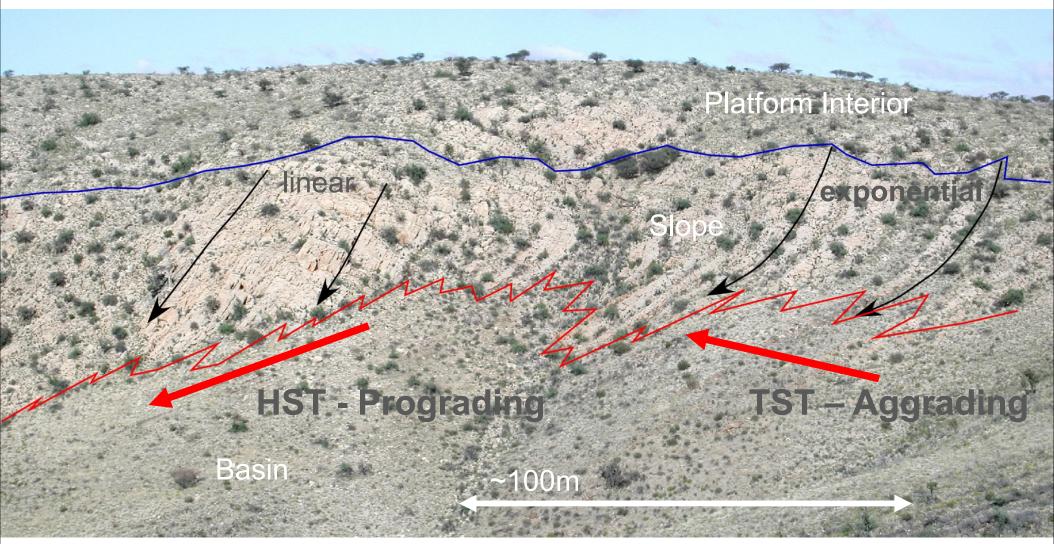
Exponential profiles and their modification to Linear

■ If a surplus of sediment is provided, the sediment fabric limits the slope inclination, and linear profiles develop.



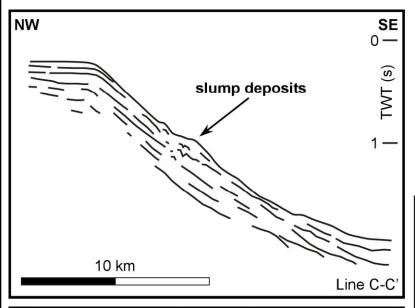
Increasing progradation rate steepens slope until sediment fabric limits the angle of dip and linear profile develops.

From TST exponential to HST planar profiles



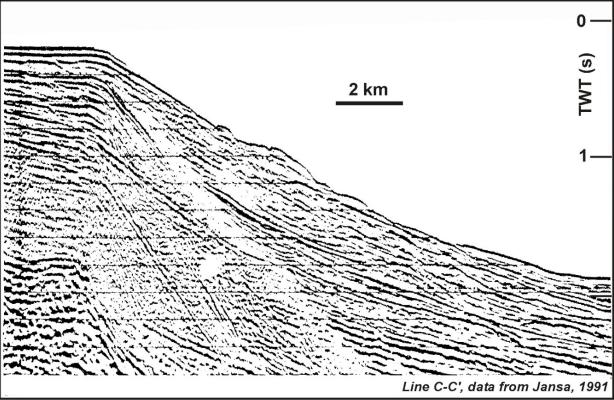
■ Terminal Proterozoic carbonate platform, interbedded grainstone and mudstone, maximum inclination 20-25 degrees

Linear profile - Scotian Slope, Canada

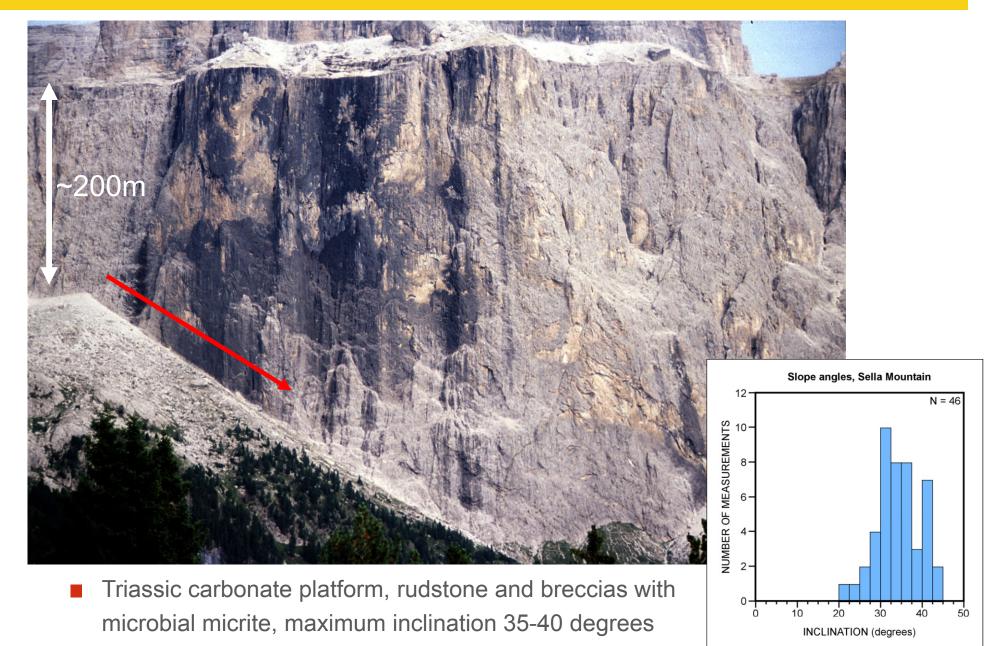


shelfbreak y=0.27x+1.92linear $-0.8 \stackrel{\text{H}}{=} 0$ exponential $y=-2.23e^{-0.13x}+4.11$ -1.2DISTANCE (km) $-4 \qquad 0 \qquad 4 \qquad 8$

- Continental margin
- Very fine-grained terrigenous sediments
- Maximum inclination 7°



Linear profile – Sella Mountain, Dolomites, Italy

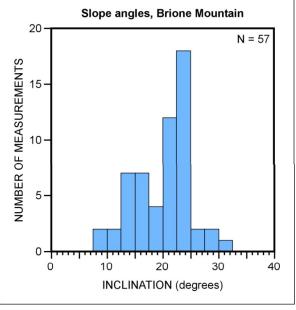


Linear profile – Brione Mountain, Italy

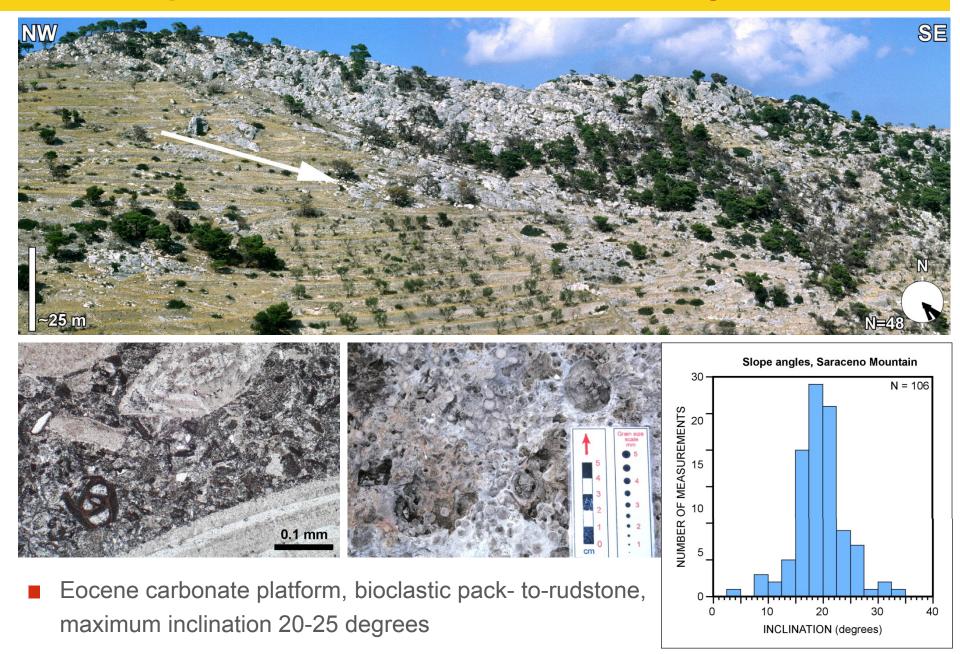




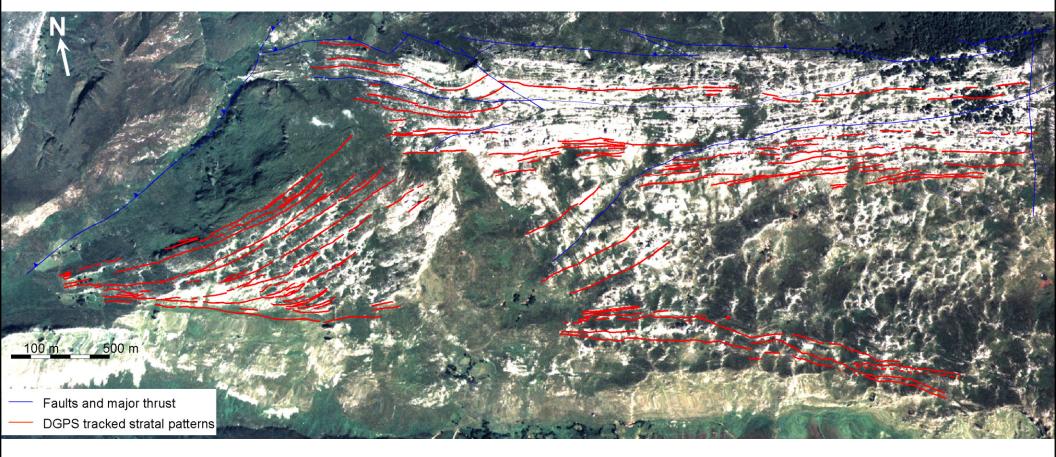
 Oligocene carbonate platform, bioclastic packstone, maximum inclination 20-25 degrees



Linear profile - Saraceno Mountain, Italy

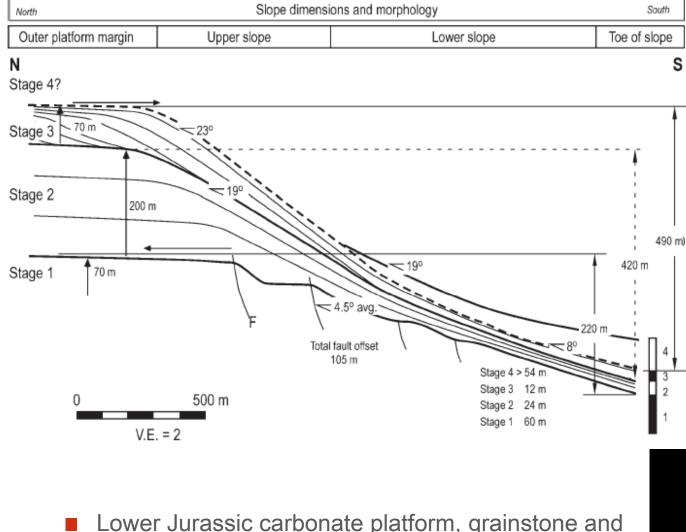


Linear profile – Sierra de Cuera, Spain

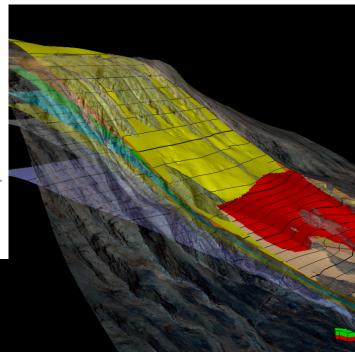


 Pennsylvanian carbonate platform, in situ microbial boundstone and breccia linear upper slope 30-35 degrees and exponential lower slope picking up mud From Verwer et al., 2004

Linear profile - Djebel Bou Dahar, Morocco



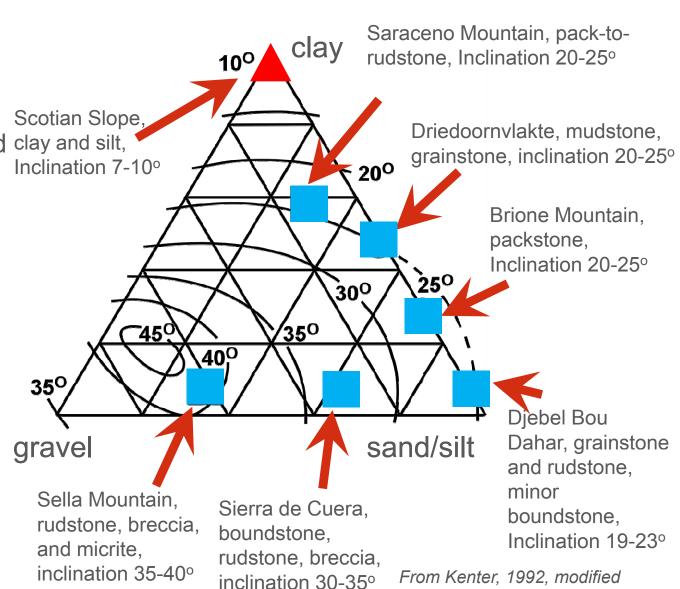
 Lower Jurassic carbonate platform, grainstone and rudstone stablilized by microbial micrite linear upper slope 19-23 degrees Slope model with surfaces fit through tracked bedding intersections in DEM (lower right) and interpreted cross section (left)



From Verwer et al., 2009 and Della Porta pers comm (2012)

Planar clinoforms rest at the angle-of-repose

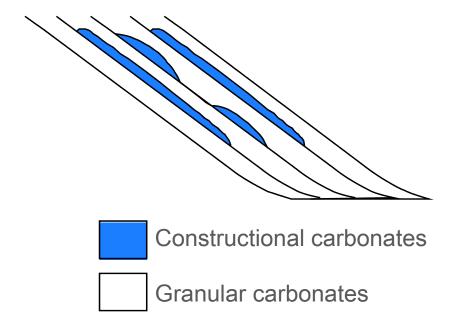
- In systems built to the angle of repose, slope inclination is directly correlated with grain size (Kenter, 1990) and the curvature is linear (Adams, 2001)
- Non-cohesive sediments have straight planar foresets
- Cohesive sediments
 have irregular
 surfaces due to creep
 and slumping



after Kirkby, 1987

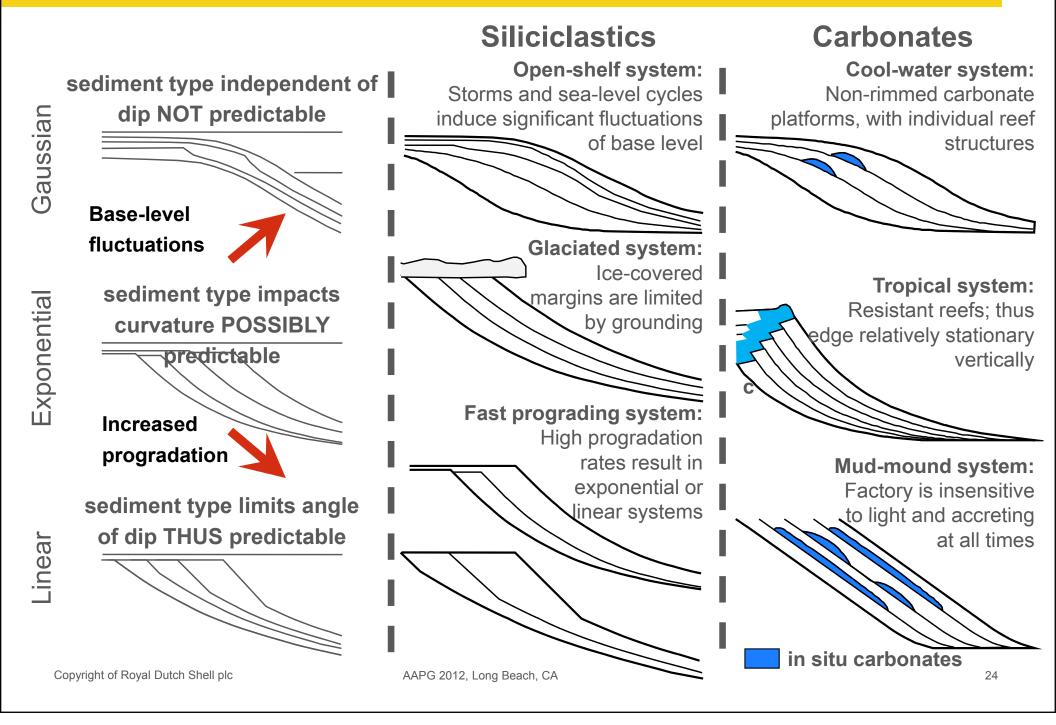
Steeply-inclined carbonates rest at the angle of repose

- Inclinations of linear carbonate slopes are in agreement with the angles of repose
- Supports assumption that linear slopes are normally inclined at the angle of repose, or the angle of shear approximates angle of repose for loose material
- Angle of repose is determined by noncohesive layers in alternating systems (i.e., Sella) or angle of shear in massive boundstone (i.e., Sierra de Cuera): carbonates fail, disintegrate and slide, resulting in similar declivities
- Critical is independence of light and "all time" shedding by large production area providing high accretion/production rates

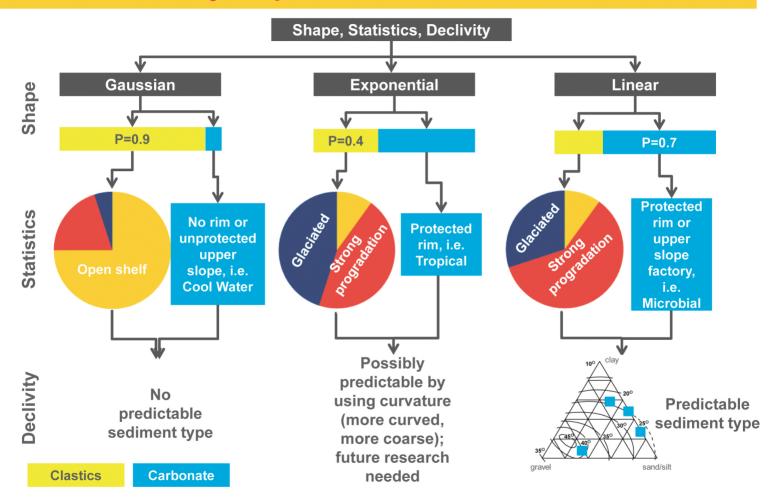




Unraveling Slope Origin and Composition



Or "Find My Slope Ahead of Drill"



Conclusions

- The established but biased view: carbonate slopes tend to be steeper than their clastic counterparts; in part because of stabilization through early cementation and microbial activity
- However, an unbiased dataset suggests similarities in gross development, curvature, and angle if similar settings, situations, or processes prevail
- The shelf edge is sharp and the profile exponential if sedimentary base level remains stationary during progradation
 - Sigmoidal: base-level fluctuations round-off the shelfbreak
 - Linear: excess sediment is piled up to the angle-of-repose
- Microbial slopes are insensitive to light and can therefore accrete during both lowand highstands, i.e., high accretion/production rates allowing systems to build-up to higher declivities
- Combining such rules may give a better handle on prediction of slope system,
 processes and type of sediment but needs additional refinement and validation