

Effect of Froude Supercritical Flow on Fluvial Facies, Geometries and Architecture*

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Abstract

Froude subcritical flow is considered the geomorphically formative flow in rivers, as reflected in fluvial facies models, morphodynamic river discharge and sediment dispersal models, flood mitigation strategies, and in how we read the stratigraphic record. Froude supercritical flow is commonly assumed a transient occurrence, or limited to steep bedrock rivers with shallow flow. Yet, the sedimentary record of modern and ancient rivers with highly variable discharge displays an abundance or even dominance of Froude supercritical flow deposits.

This article shows that Froude supercritical flow may exert first-order control on river morphodynamics, such as the nature of the small- and large- scale bedforms and thus the resultant stratigraphy, as well as on sediment transport mode and rate. Such rivers characteristically lack well developed barforms, as supercritical flow is not advected by bars and bar migration thus not maintained. Transition to Froude supercritical flow significantly increases the proportion of sandy and gravelly sediment carried in suspension. Suspension transport critically increases downstream sediment transport rates, as bedload transport rates are linked to downstream bedform migration rates that are only a small fraction of the mean flow velocity. This article discusses the effects of supercritical flow on small- and large-scale bedform migration and the resultant depositional architecture in rivers where supercritical flow is the geomorphically formative flow.

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Effect of Froude Supercritical Flow on Fluvial Facies, Geometries and Architecture

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Problem:

- Froude supercritical flow ($Fr > 1$) sedimentary structures abundant on some
 - Deepwater slopes
 - Deltas
 - Rivers
- But $Fr > 1$ flow commonly considered a transient occurrence in a shallow flow and very high slopes
- How come we see an abundance or dominance of $Fr > 1$ features in some settings = geomorphically formative flow in some systems?
- What difference does it make?

Sub- vs supercritical flow



What difference does it make?

Subcritical flow $Fr < 1$



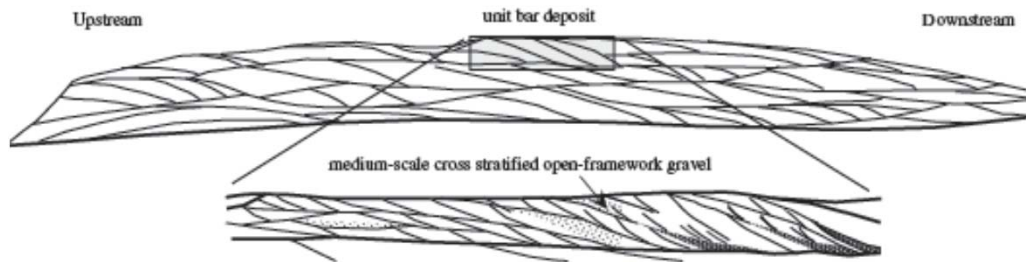
www.youtube.com/watch?v=cJoOfTpJypg

- Slow, stable flow
- Bedload transport
- Equilibrium conditions for bedform migration (ripples, dunes, bars)
- Downstream bedform migration with steep (angle of repose) foresets
- Sediment transport rates depend on bedform migration rates and lengths

What difference does it make?

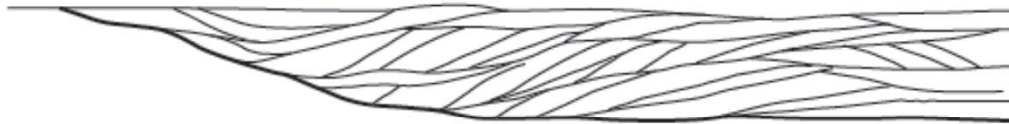
Subcritical flow $Fr < 1$

E Along-stream view through upstream end of large channel fill: lateral and downstream growth of compound bar

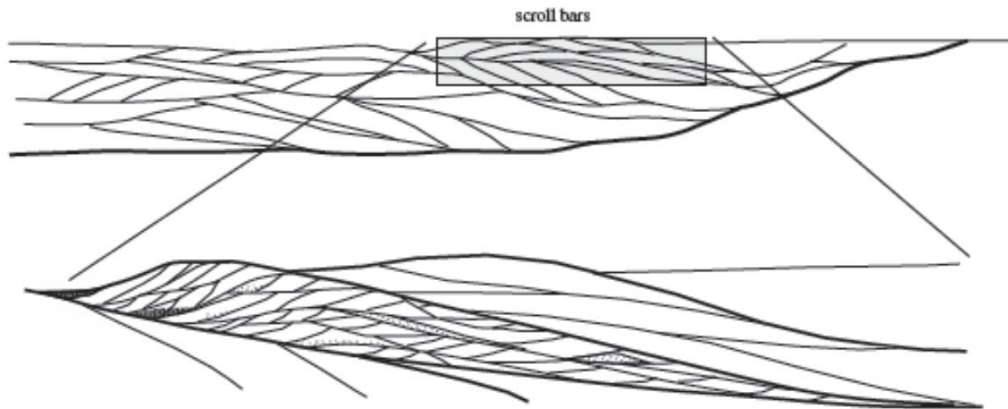


- Various scales of cross lamina and strata
- Indicate the downstream migration and lee side preservation of steep foresets

F Across-stream view of upstream end of large channel fill : lateral accretion and channel filling



G Across-stream view of downstream end of large channel fill : scroll bar accretion and channel filling



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What difference does it make?
Supercritical flow $Fr > 1$

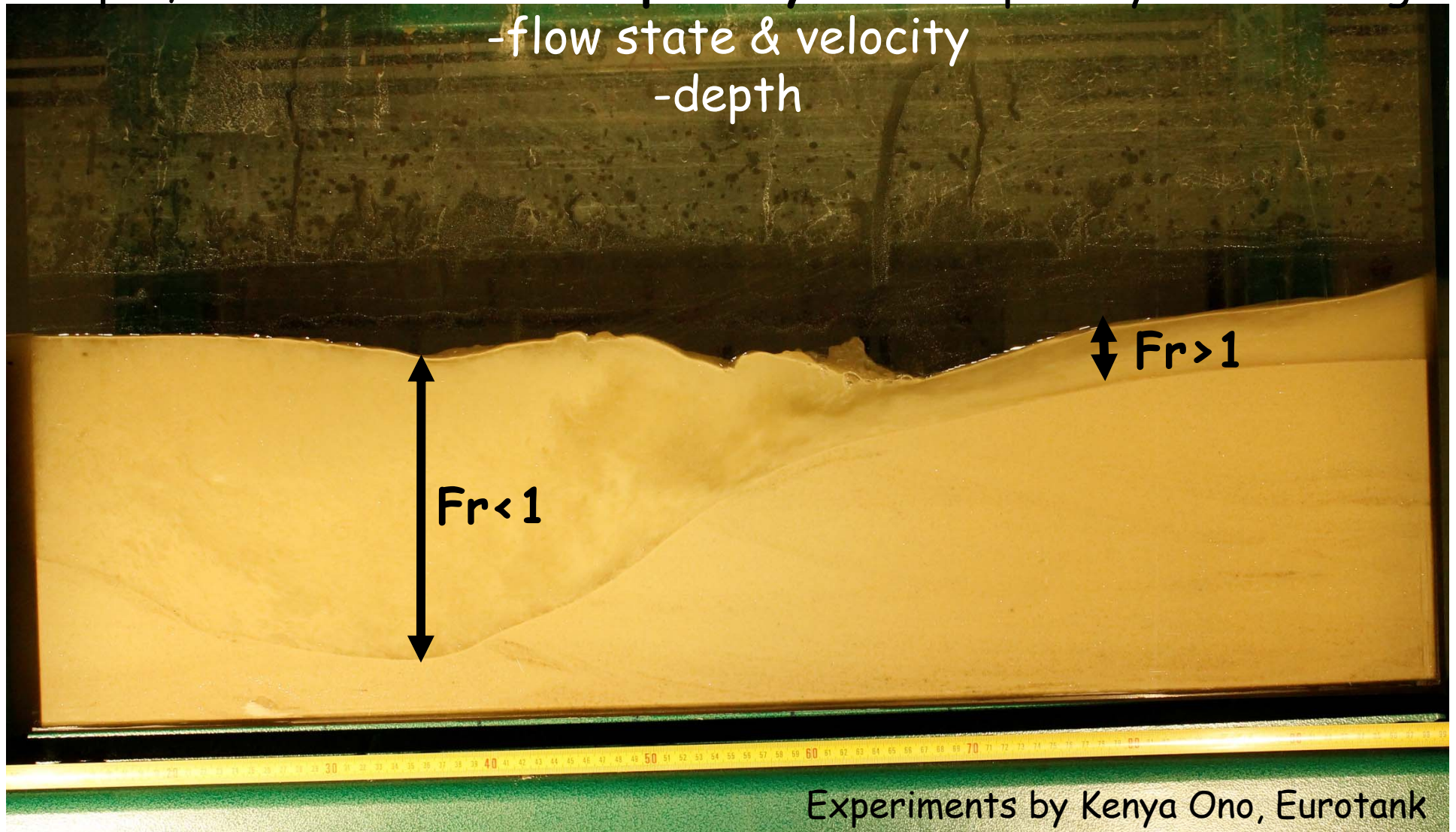
Experiments
by Kenya Ono,
Eurotank



- Rapid, unstable flow with highly fluctuating flow state, velocity and Fr nr
- Suspension transport of sand and gravel
- No equilibrium conditions or sustained bedform migration: deposition and erosion intermittent
- Up - or down-stream migration or vertical aggradation
- Sediment transport rates not linked to bedform migration rates or length

Unstable (large Fr variability) flow

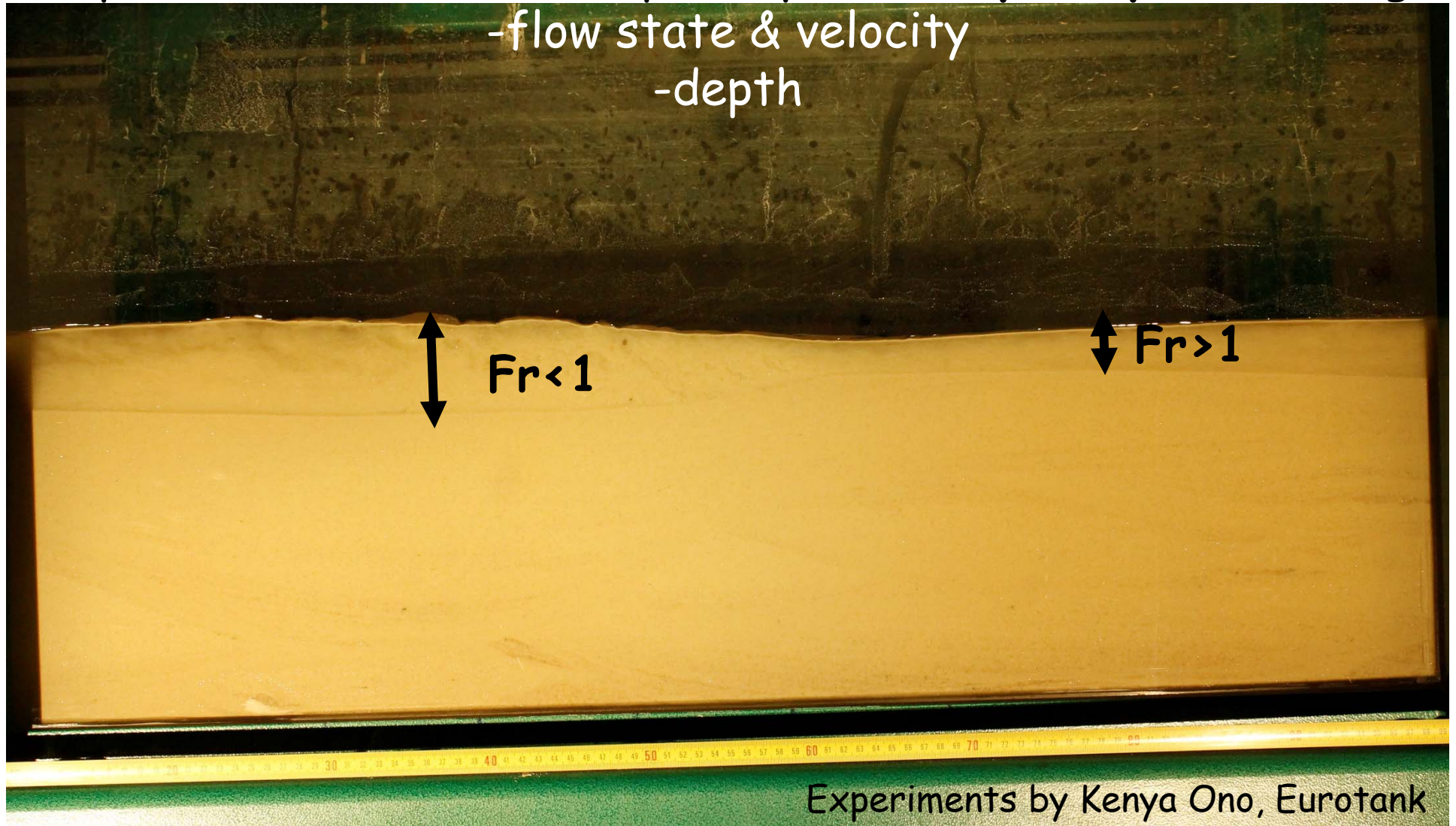
Rapid, unstable flow with **spatially** and temporally fluctuating
-flow state & velocity
-depth



Experiments by Kenya Ono, Eurotank

Unstable (large Fr variability) flow

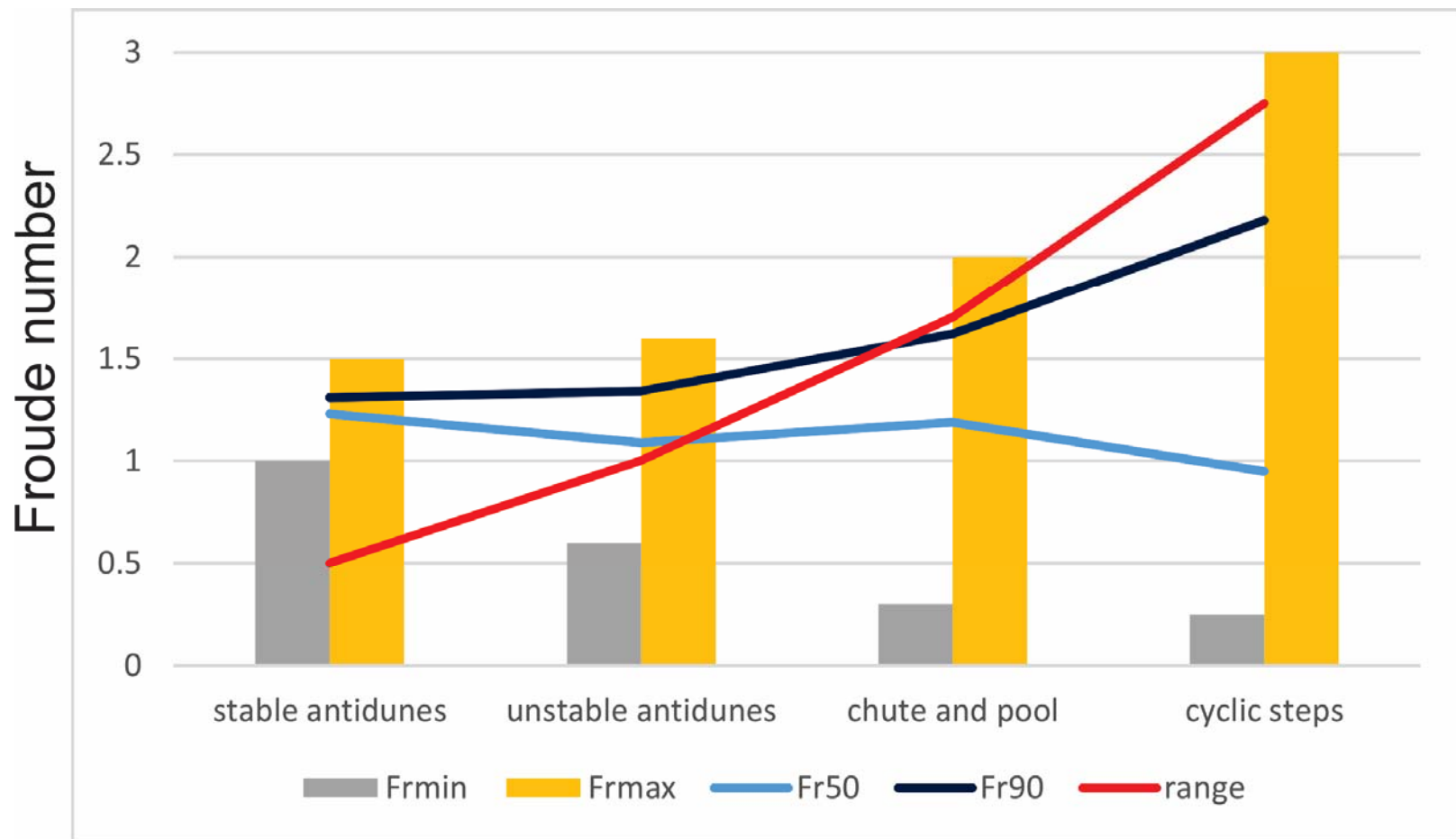
Rapid, unstable flow with spatially and **temporally** fluctuating
-flow state & velocity
-depth



Experiments by Kenya Ono, Eurotank

Unstable (large Fr variability) flow

Increasing instability with increasing Fr_{90}
= $Fr_{\max} - Fr_{\min}$ increases

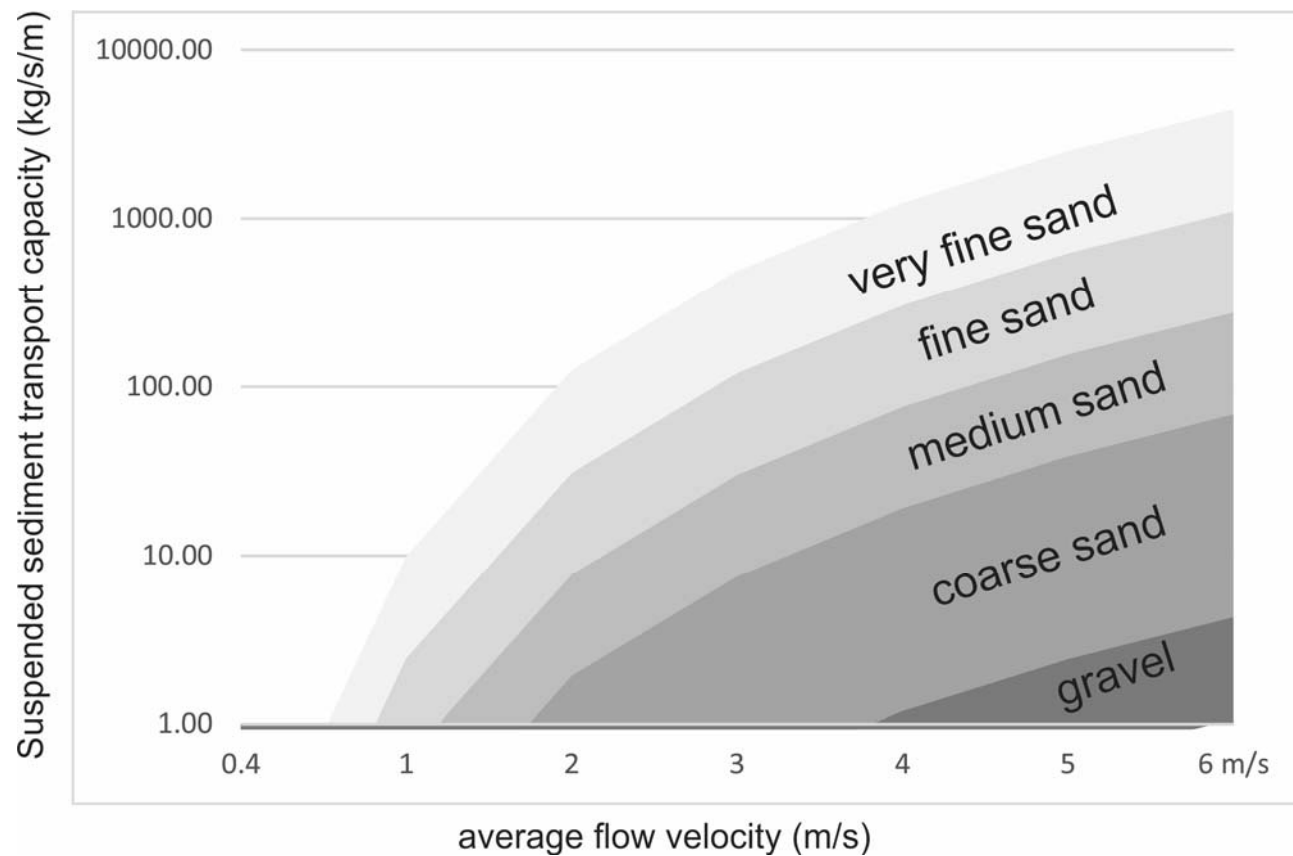


Data from Cartigny et al., 2014

So what?



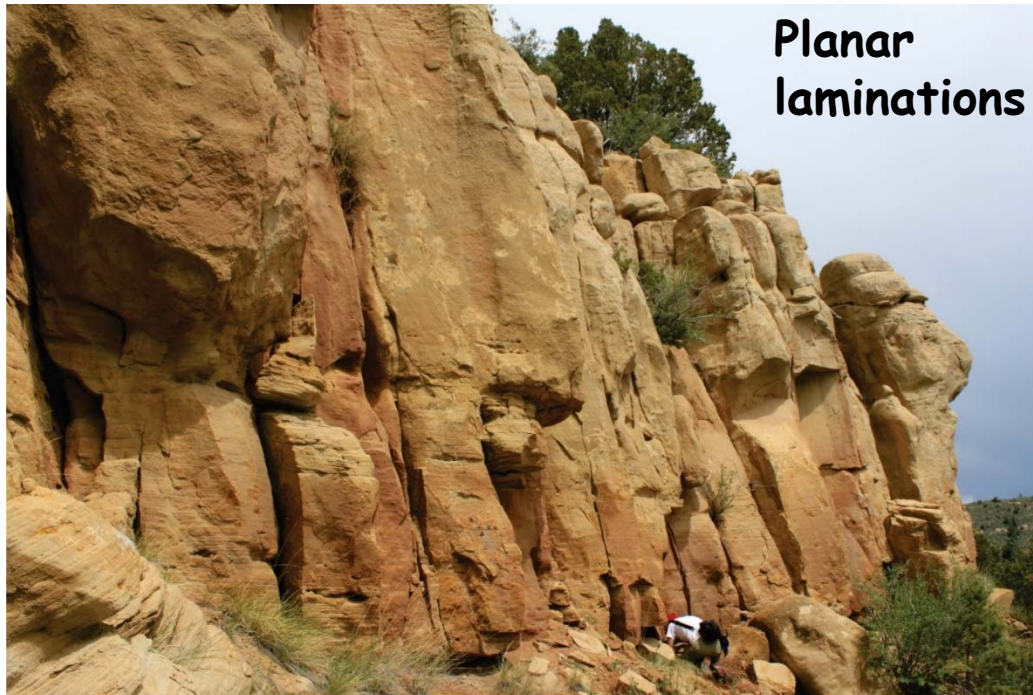
Effects: Suspension transport



Van Rijn, 2007

Effects: $Fr > 1$ sedimentary structures

Specific $Fr > 1$ sedimentary structures are difficult to interpret in some cases as they are not formed as sustained equilibrium bedforms but rather by intermittent deposition and erosion (except perhaps cyclic steps)



Planar laminations

**Convex-up
low-angle
strata**



**Scour and
fill**

Effects: "Poorly developed bars"

$Fr > 1$ flow not diverted around bars: low-angle downstream accreting sheets with lee-side erosion and backset bedding



Effects: "Poorly developed bars"

Erosionally based, low-angle sheets



Effects: Bars

Erosionally based, low-angle sheets



Experiments by Kenya Ono, Eurotank

Effects: "Poorly developed bars"

No sustained lateral accretion:
oblique downstream to vertical accretion and no upward fining



Effects: "Poorly developed bars"

$Fr > 1$ flow not advected around bars: upward coarsening



Effects: Architecture

Flow instability and high local deposition rates promote avulsions rather than channel migration



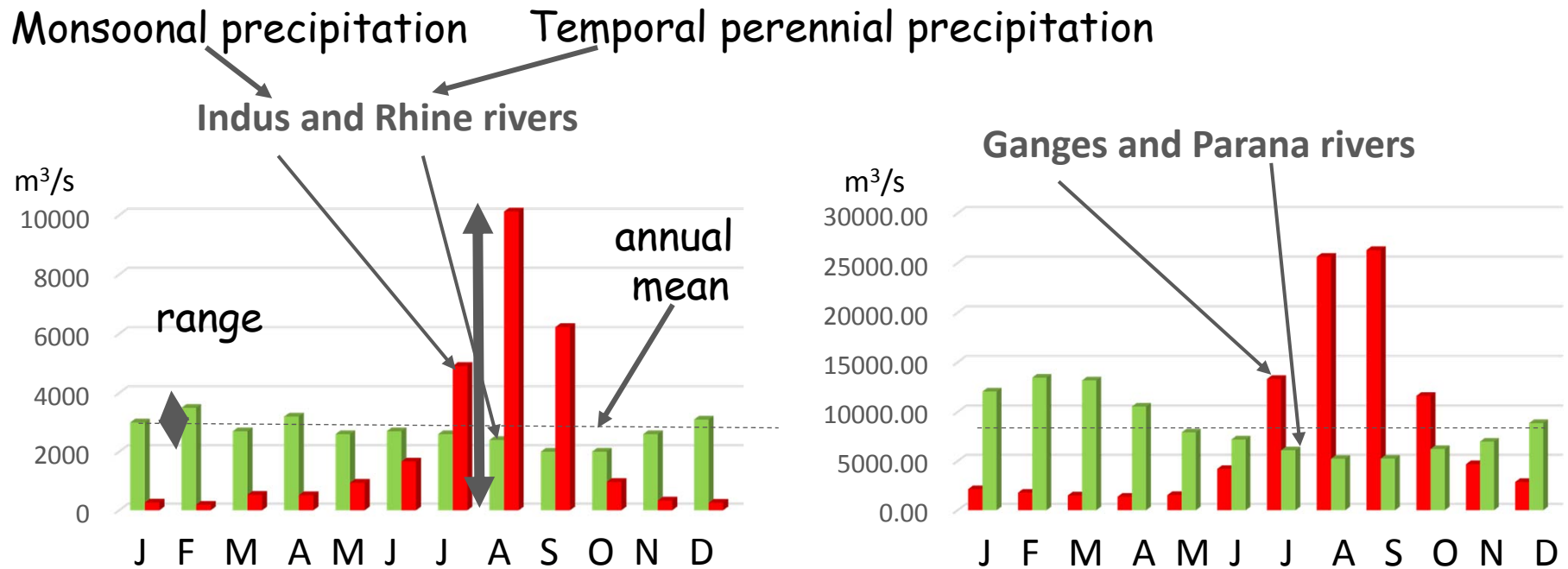
Effects: Architecture



At what conditions is $Fr > 1$ flow geomorphically formative?



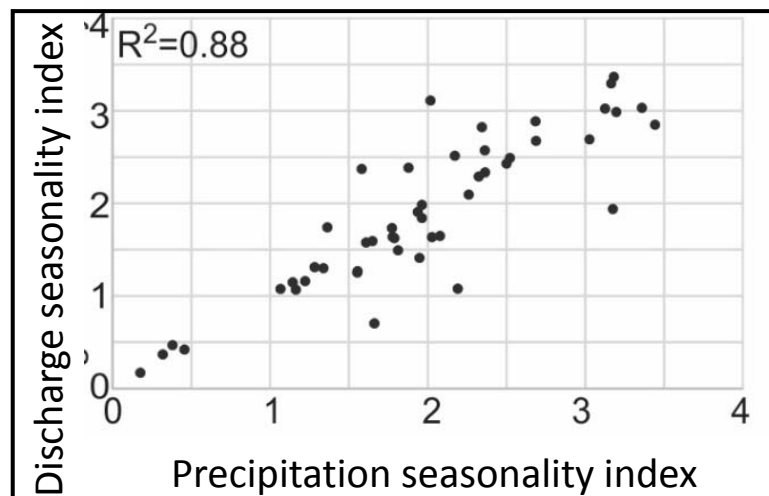
Unstable $FR > 1$ flow formative flow in some variable discharge rivers



Precipitation vs discharge seasonality index

Precipitation seasonality index:

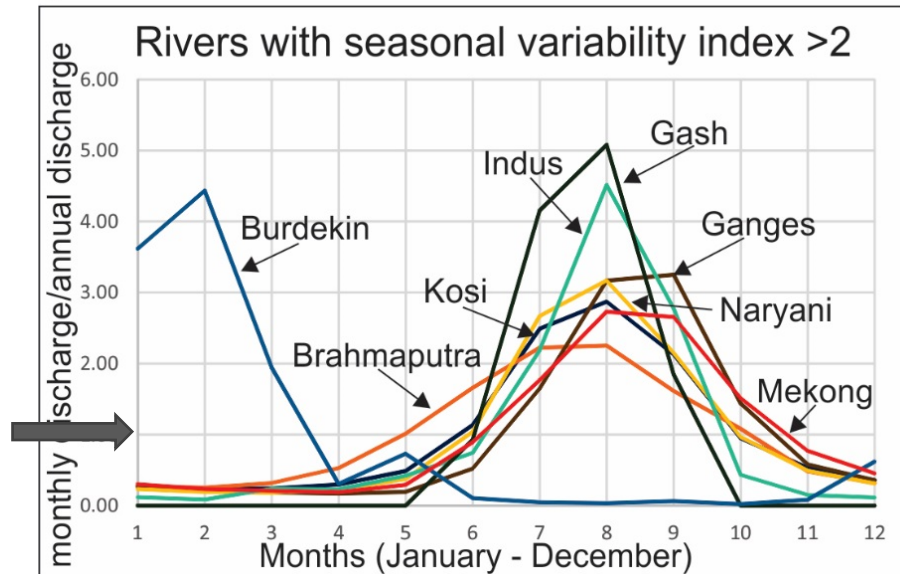
$$\frac{P_{\max} - P_{\min}}{P_{\text{mean}}}$$



Discharge seasonality index:

$$\frac{Q_{\max} - Q_{\min}}{Q_{\text{mean}}}$$

Variable discharge



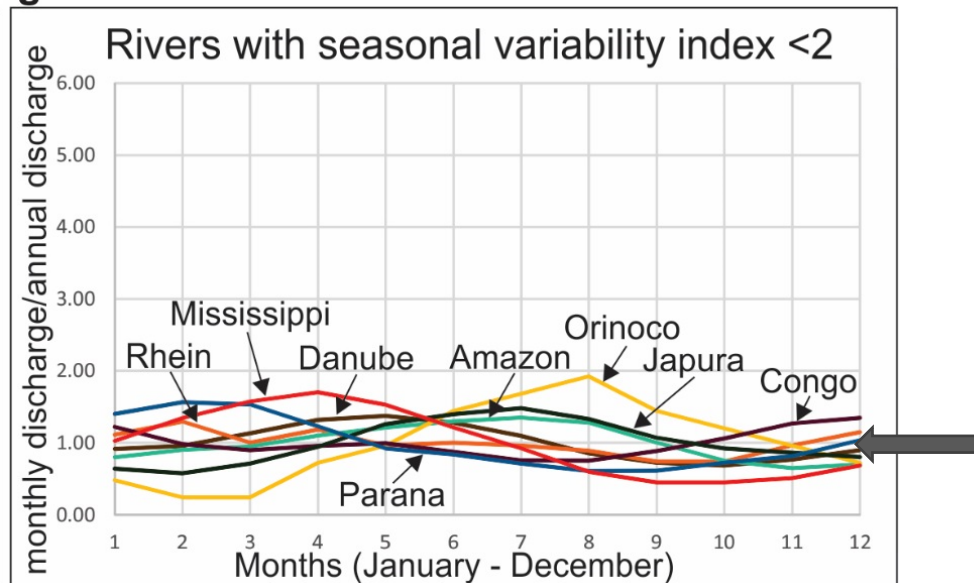
DVI >2

$Q_{\max} > 800\%$ of Q_{mean}

$Q_{90} > 500\%$ of Q_{mean}

Q_{50} ca 50% (0-10%) of Q_{mean}

Perennial precipitation zone



DVI <2

$Q_{\max} < 180\%$ of Q_{mean}

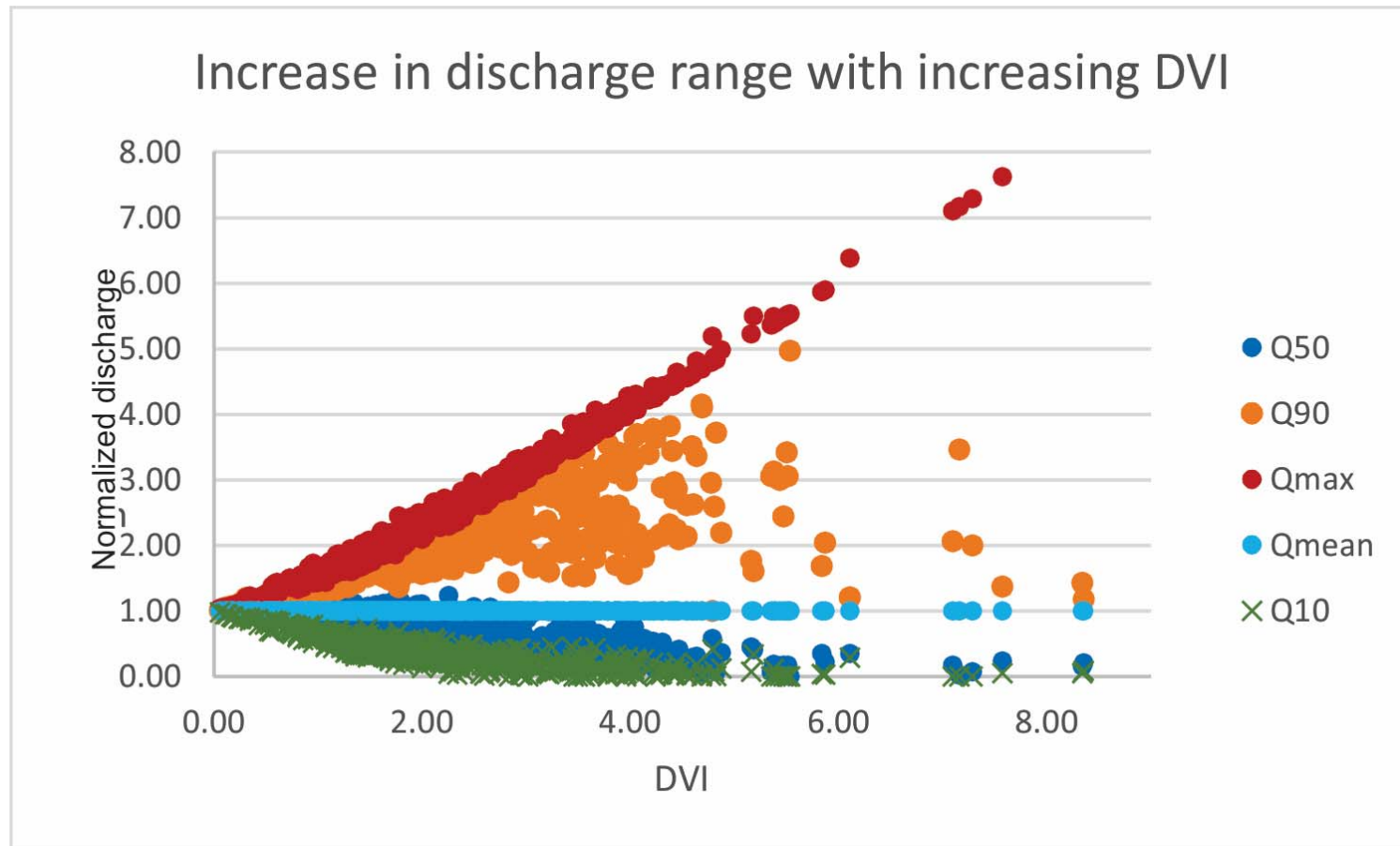
$Q_{90} < 160\%$ of Q_{mean}

Q_{50} within 92% of Q_{mean}

Q_{10} within 40-70% of Q_{mean}

Increased discharge range in variable discharge rivers

Monthly discharge data of 497 rivers <http://www.sage.wisc.edu/riverdata>



Plink-Bjorklund, 2017



Persistent discharge
ivers:

Base discharge =
ca Q_{mean}

Flood discharge =
ca Q_{mean}



Variable discharge rivers: Base discharge $\ll Q_{\text{mean}}$

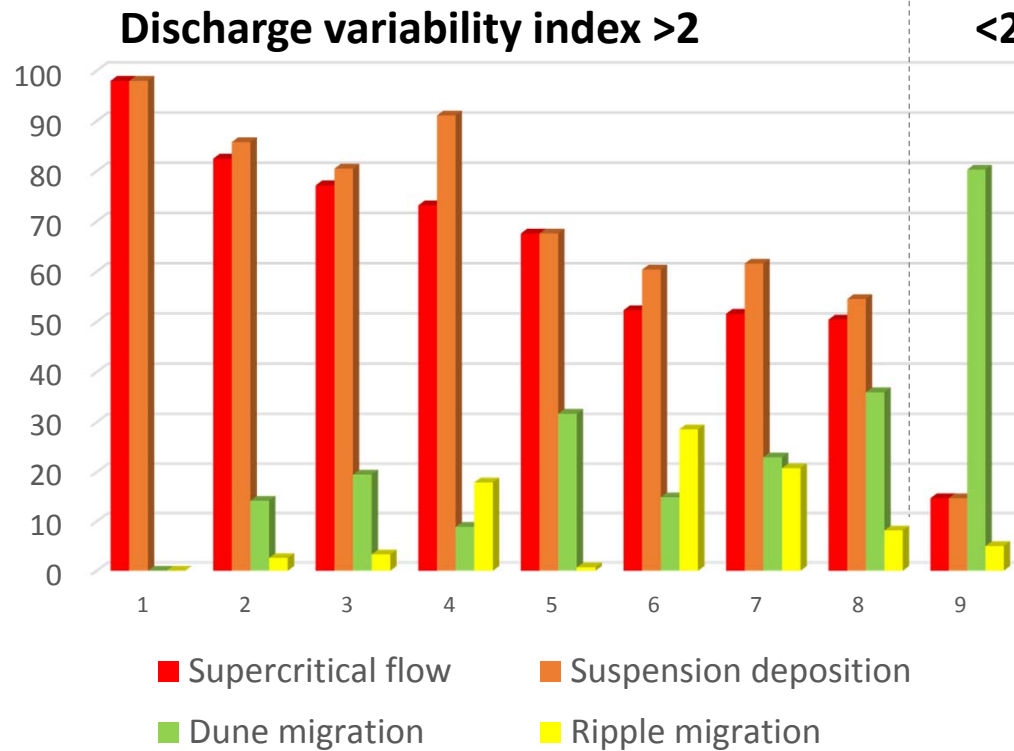


Variable discharge rivers: Flood discharge $\gg Q_{\text{mean}}$



Base flow = geomorphically inefficient
Flood discharge ($Fr > 1$) = short duration but the formative flow

Persistent perennial precipitation zone rivers

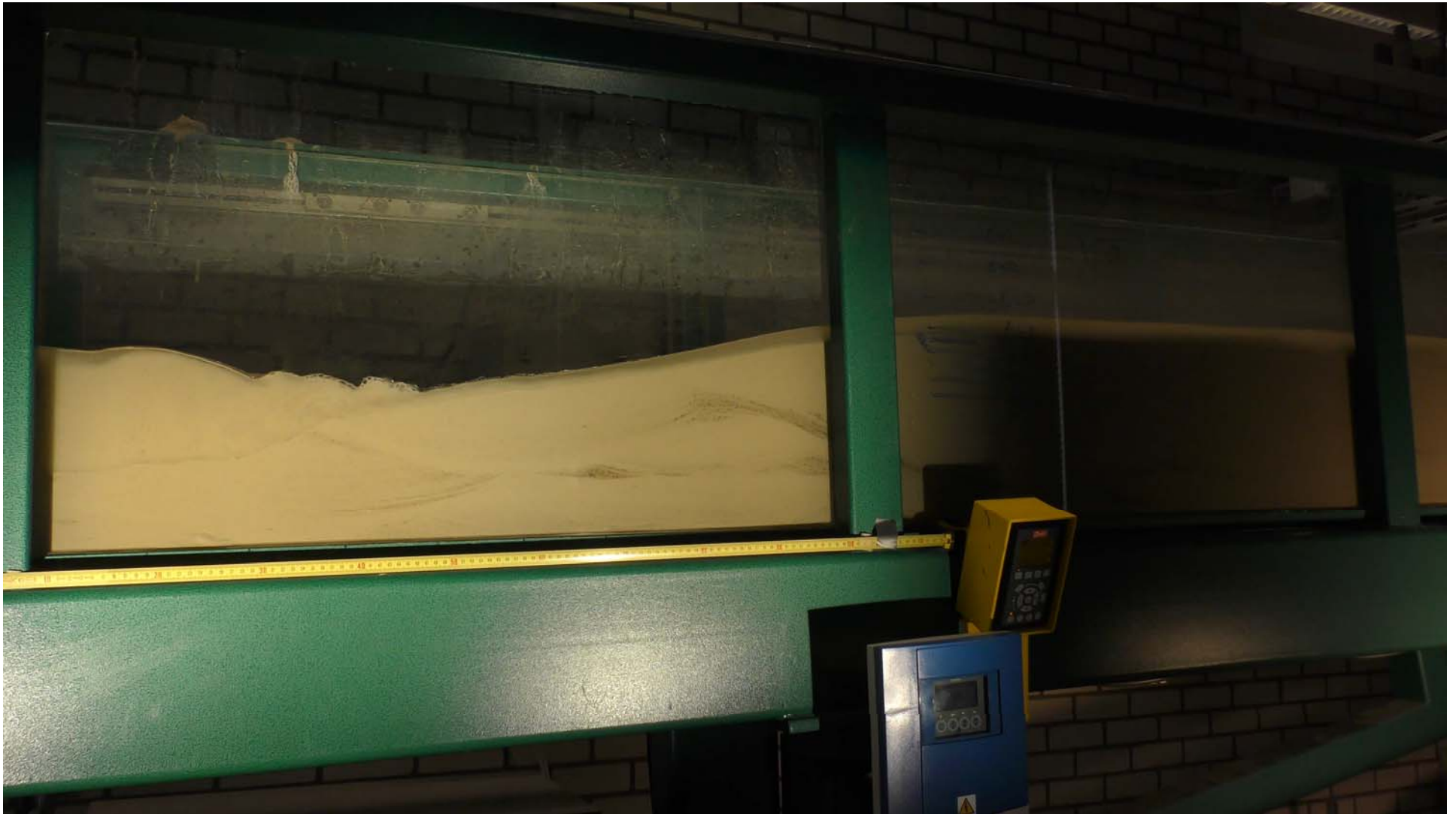


Plink-Bjorklund, 2015

Other things to consider?

Water depth? Advection length & sediment transport rates?

Environmental signal propagation? Autogenic processes?



Experiments by Kenya Ono, Eurotank

Conclusions

- Both the $Fr > 1$ and the large Fr range define the flow characteristics
- $Fr > 1$ in transient events, most deposition due to hydraulic jump and drop to $Fr < 1$ conditions: should be termed unstable or variable flow rather than $Fr > 1$ flow?
- Where it is the geomorphically efficient flow, it affects morphodynamics: not just sedimentary structures, but the transport mode, downstream sediment transport rate, mode of deposition, flow depth, advection lengths, autogenic processes, environmental signal transfer ...
- We need more than just new facies models ...

