

Stratigraphic Framework for Basin-Margin, Sub-Unconformity Diagenesis Below the Acadian Unconformity in the Southern Williston Basin*

David M. Petty¹

Search and Discovery Article #10981 (2017)**

Posted August 14, 2017

*Adapted from oral presentation given at AAPG Rocky Mountain Section Annual Meeting, Billings, Montana, June 25-28, 2017

**Datapages © 2017 Serial rights given by author. For all other rights contact author directly.

¹Geological Consultant, Katy, TX (dpetty53@aol.com)

Abstract

The Acadian discontinuity is represented by a late Devonian (Famennian) unconformity that lies on top of the Famennian Three Forks Formation in the central portion of the Williston Basin and on truncated Devonian through Cambrian formations in basin-margin areas. Bakken-Pronghorn sediments were deposited in the basin-center and onlap the Acadian unconformity. A maximum flooding surface overlies the Bakken-Pronghorn system in the basin-center but this surface overlies the Acadian unconformity in basin-margin areas. The maximum flooding surface is overlain regionally by open-marine carbonate strata of the basal Lodgepole.

Sub-unconformity diagenesis was an important process in basin-margin areas around the southern Williston Basin; however, the diagenetic analysis in this study focuses on a portion of the southwest Williston Basin where good core data and numerous GR-neutron-density logs are present. Basin-margin, shallow-burial diagenesis beneath the Acadian unconformity formed a paleokarst characterized by evaporite dissolution and massive dolomitization. This diagenesis is illustrated in the Duperow and Birdbear formations, which consist of stratiform limestone, dolostone, and anhydrite beds in the basin-center. These layered lithologies transition to massive dolostone in subcrop areas beneath the Acadian unconformity. A sub-unconformity paleo-dissolution front extends as much as 50 m below the unconformity and former anhydrite beds are now represented by solution-breccia beds within the dissolution area. Massive dolostone cuts across stratigraphic units and extends as much as 70 m below the Acadian unconformity.

While they overlapped in basin-margin areas, massive dolomitization was not related to the fresh-water processes that produced evaporite dissolution during paleokarst formation. Since the massive dolomitization front is sub-parallel to the unconformity over large areas and there is no massive dolostone in overlying basal Lodgepole beds, it is likely that dolomitization occurred after unconformity development and before Lodgepole deposition. This indicates that basin-margin dolomitization took place while Bakken-Pronghorn sediments were being deposited, 10-20 million years after deposition of the Birdbear and Duperow. Basin-margin massive dolomitization probably occurred during one or more evaporative events that caused regional dolomitization within carbonates of the Middle Bakken. The “massive” characteristics, defined by the absence of a stratiform geometry, relate to dolomitization that occurred after mineral stabilization and initial lithification within the Duperow and Birdbear.

References Cited

Alexandre, C.S., S. Sonnenberg, and J.F. Sarg, 2012, Diagenesis of the Bakken Formation, Elm Coulee Field, Richland County, Montana: AAPG Annual Convention and Exhibition, Long Beach California, April 22-25, 2012, [Search and Discovery Article #20160 \(2012\)](#). Website accessed August 2017.

Boyd, H.A., 1997, “Devono-Mississippian” Discontinuities in the Northern Rocky Mountain Region and Their Regional Significance: Wyoming Geological Association 48th Annual Field Conference Guidebook, p. 3-28.

Johnson, K.S., 1997, Evaporite Karsts in the United States: Carbonates and Evaporites, v. 12, p. 2-14.

Loeffler, P.T., 1982, Depositional Environment and Diagenesis, Birdbear Formation (Upper Devonian), Williston Basin, North Dakota: Master’s Thesis, University of North Dakota, Grand Forks, North Dakota, 268 p.

Petty, D.M., 2010, Sequence Stratigraphy and Sequence Boundary Characteristics for Upper Tournaisian (Mississippian) Strata in the Greater Williston Basin Area: An Analysis of a Third-Order Cratonic Carbonate-Evaporite Depositional Cycle: Bulletin of Canadian Petroleum Geology, v. 58/4, p. 375-402.

Petty, D.M., 2006, Stratigraphic Development of the Madison Paleokarst in the Southern Williston Basin Area: The Mountain Geologist, v. 43/4, p. 263-281.

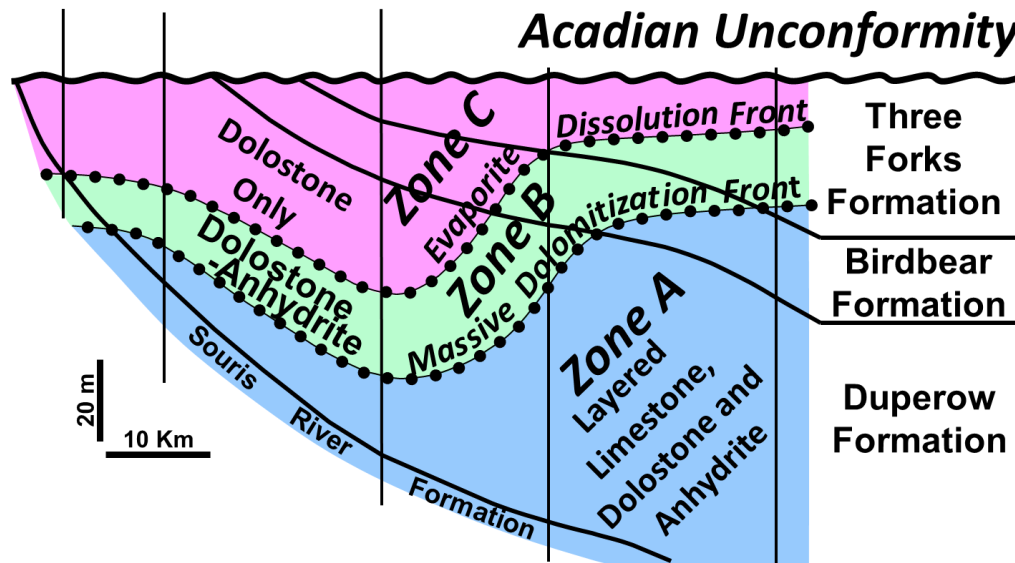
Skinner, O., L. Canter, M.D. Sonnenfeld, and M. Williams, 2015, Discovery of “Pronghorn” and “Lewis and Clark” Fields: Sweet-Spots within the Bakken Petroleum System Producing from the Sanish/Pronghorn Member NOT the Middle Bakken or Three Forks: AAPG Annual Convention and Exhibition, Long Beach, California, April 22-25, 2012, [Search and Discovery Article #110176 \(2015\)](#). Website accessed August 2017.

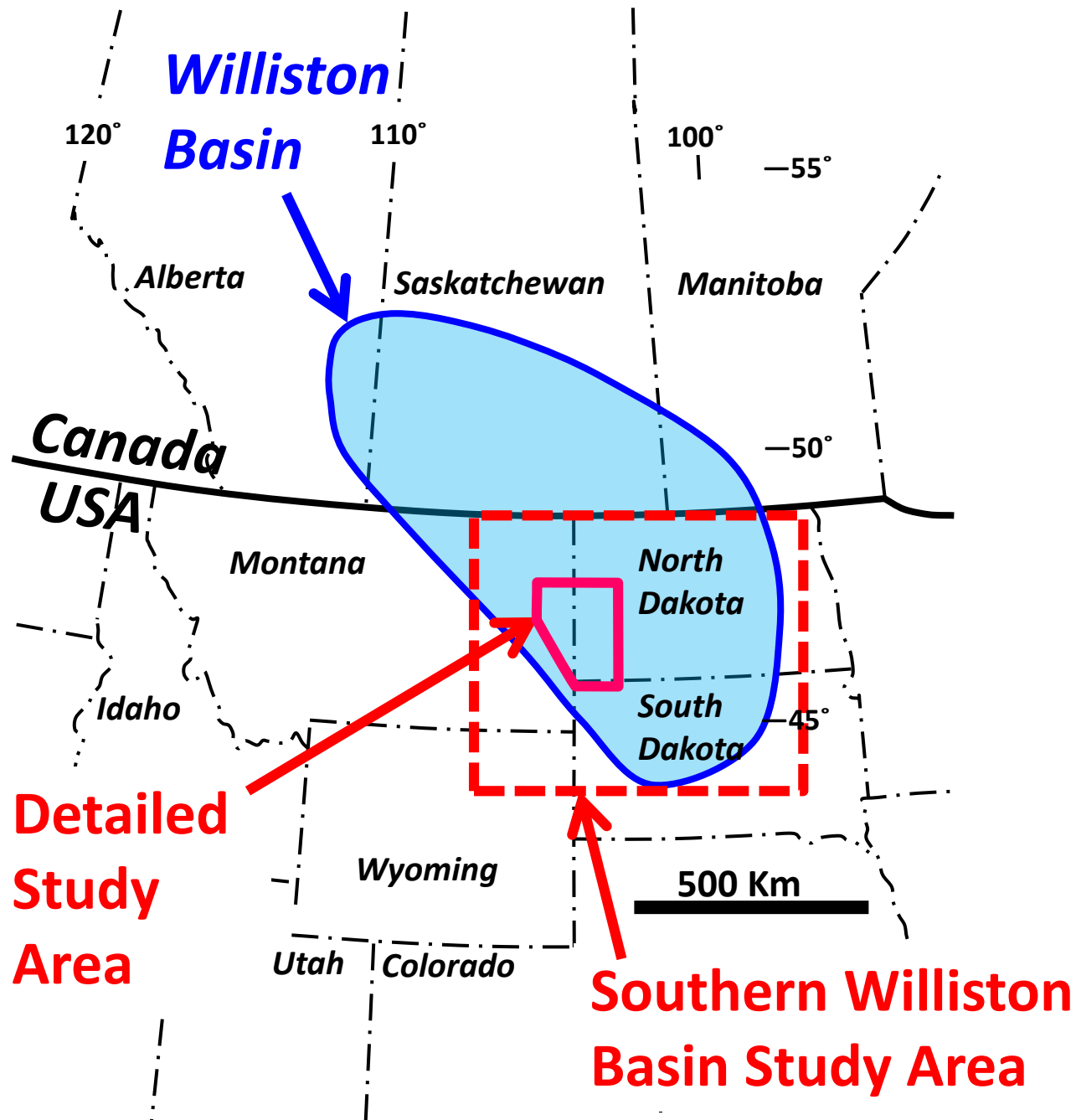
Smith T.M., and S.L. Dorobek, 1989, Dolomitization of the Devonian Jefferson Formation, South-Central Montana: The Mountain Geologist, v. 26/3, p. 81-96.

Wheeler, H.E., 1963, Post-Sauk and pre-Absaroka Paleozoic Stratigraphic Patterns in North America: American Association of Petroleum Geologists Bulletin, v. 47, p. 1497-1526.

Stratigraphic Framework for Basin-Margin, Sub-Unconformity Diagenesis Below the Acadian Unconformity in the Southern Williston Basin

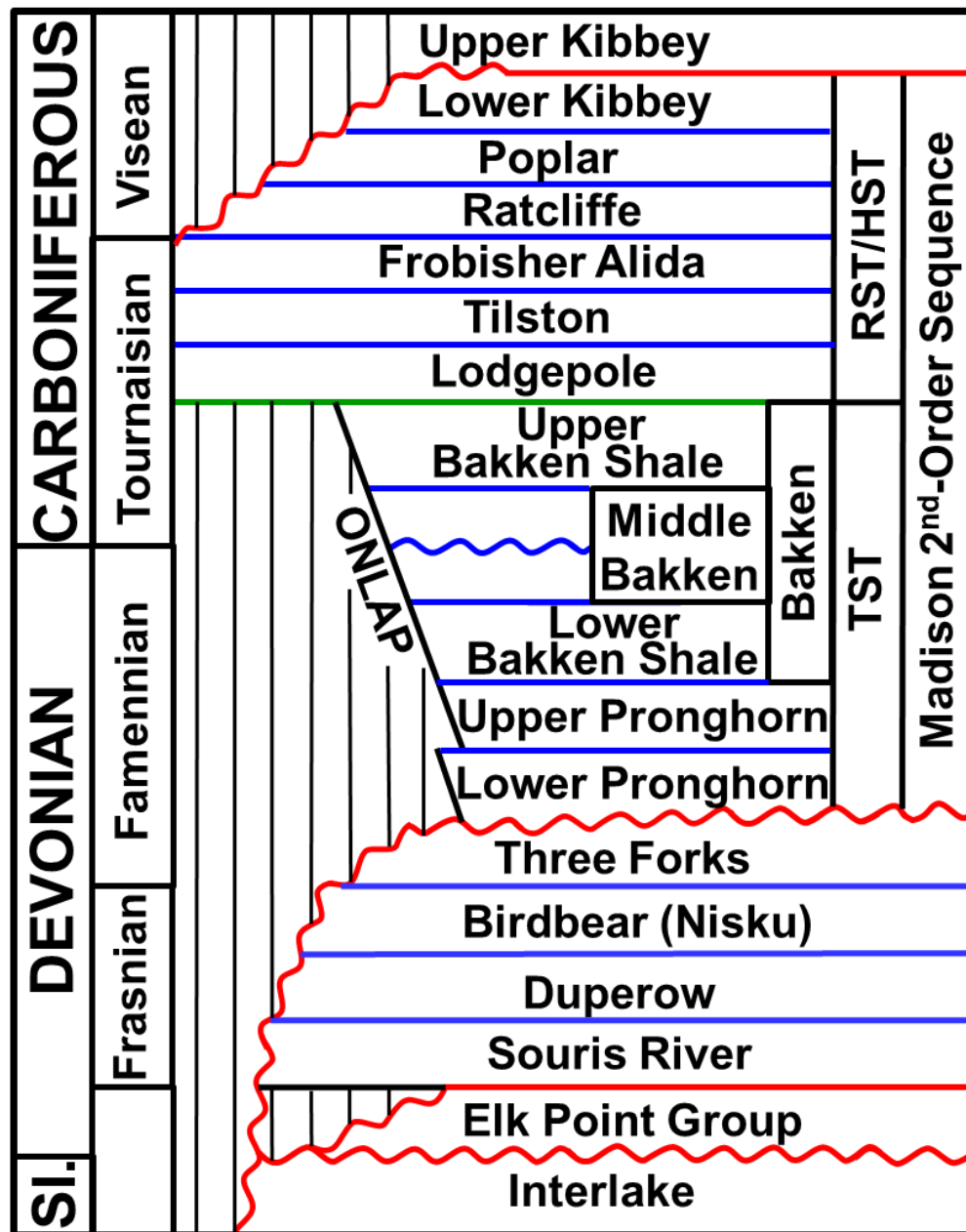
By
David M. Petty

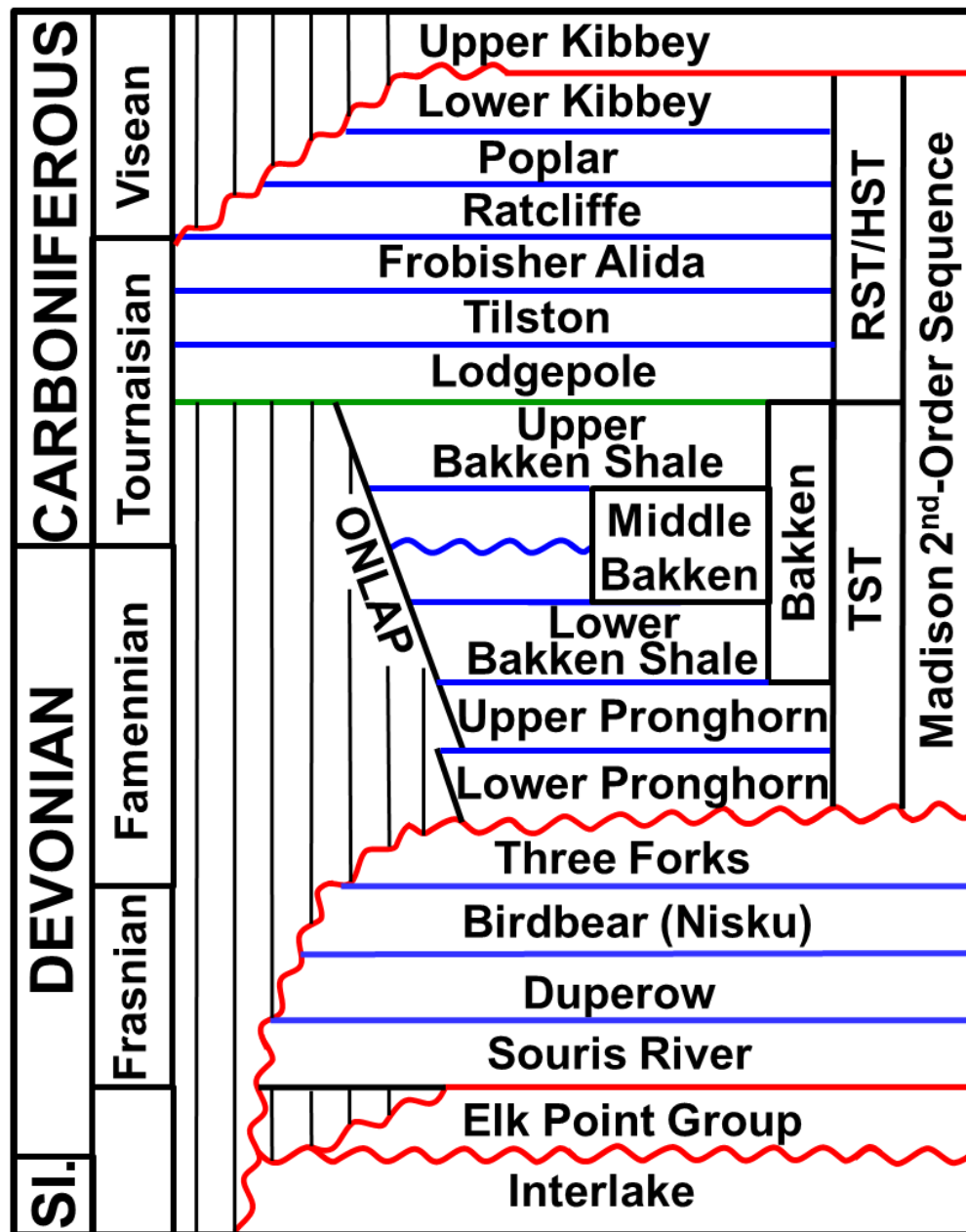




Introduction

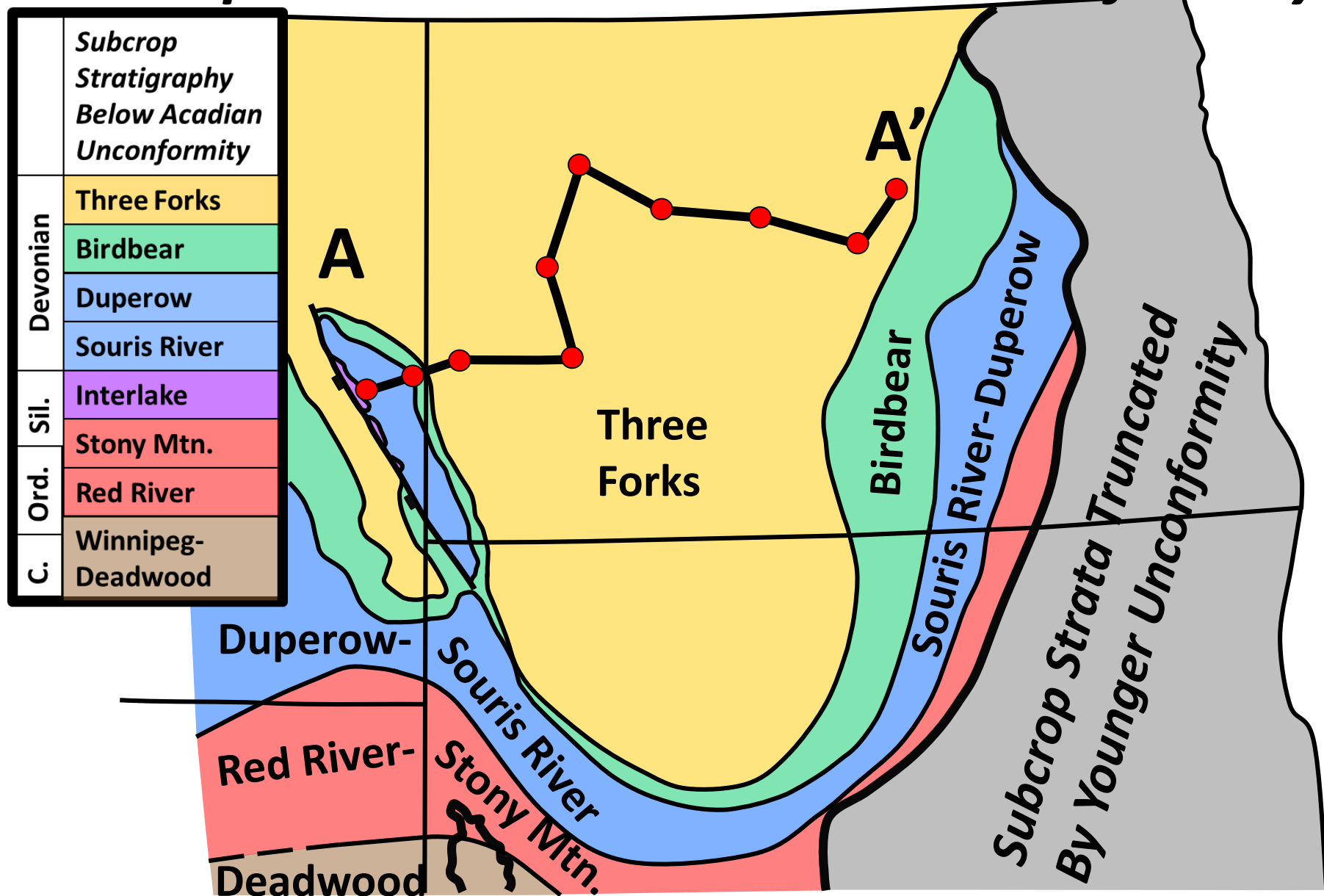
- **The “Acadian Discontinuity” is represented by a late-Devonian, angular unconformity (the “Acadian unconformity”) that lies on top of the Three Forks Formation in the central portions of the Williston basin, and on older Devonian through Cambrian strata in basin-margin areas (Wheeler, 1963, Boyd, 1997)**

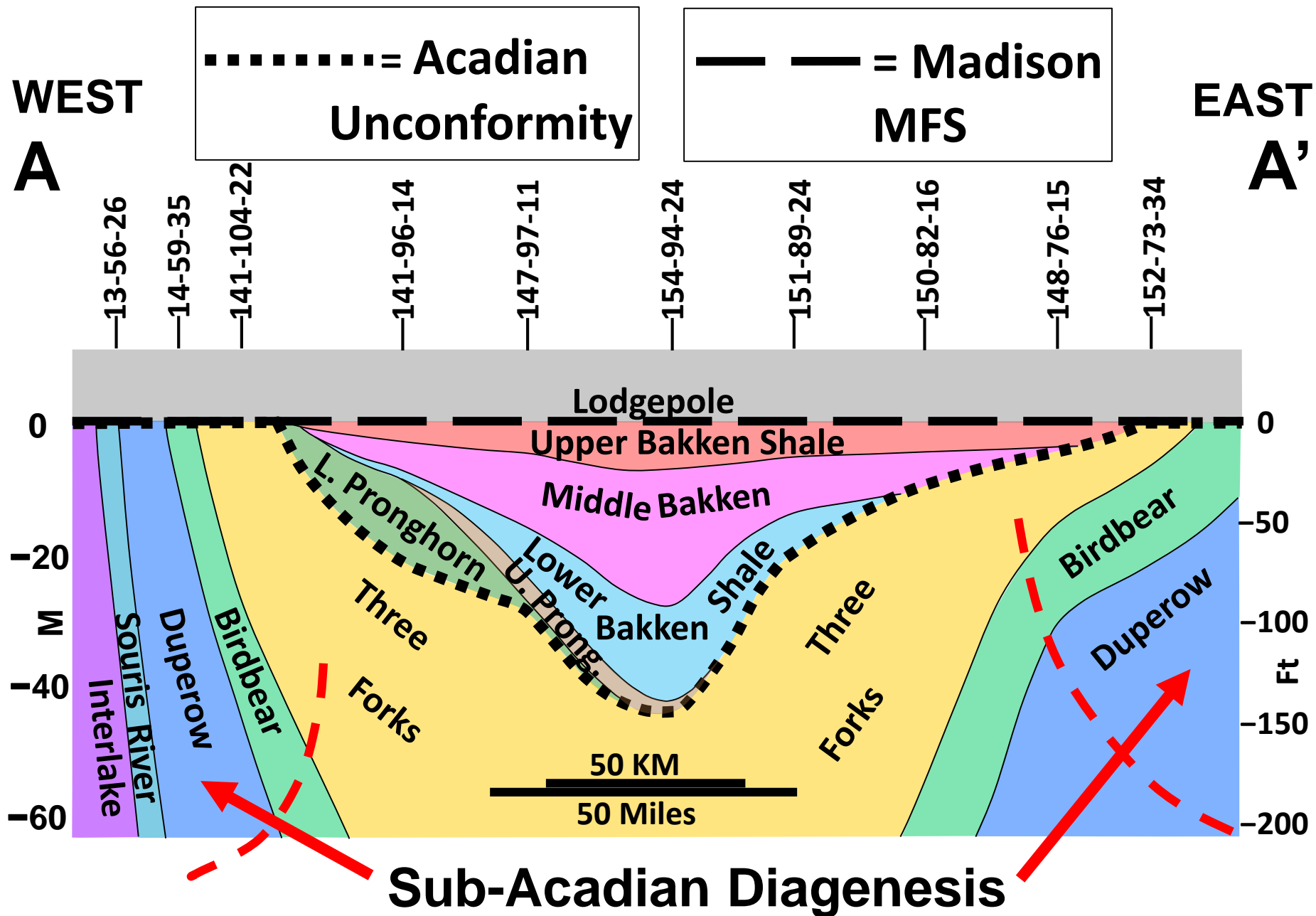




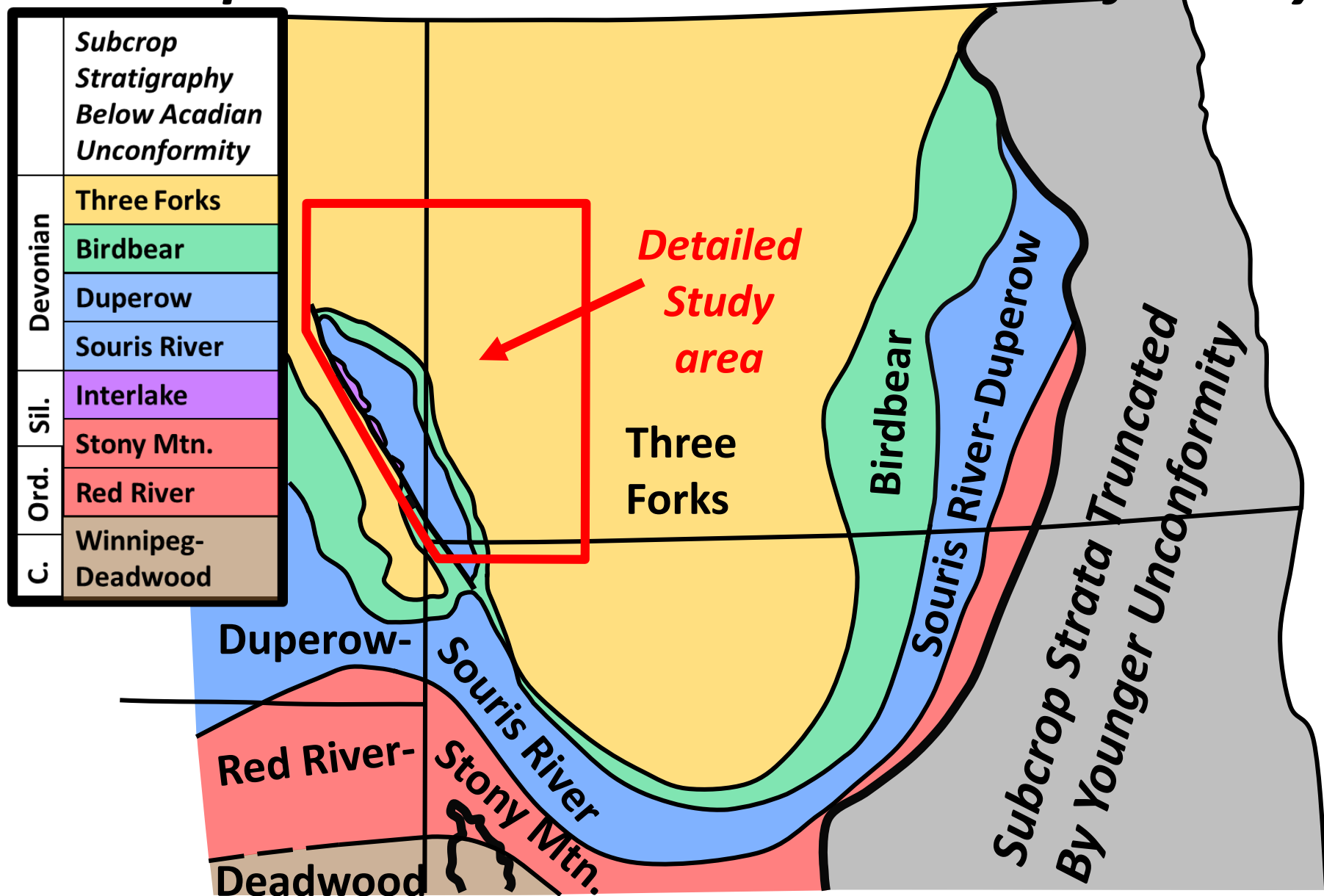
- 1st or 2nd-Order Boundary
- 2nd-Order MFS
- 3rd or 4th-Order Boundary

Subcrop Strata Below Acadian Unconformity





Subcrop Strata Below Acadian Unconformity

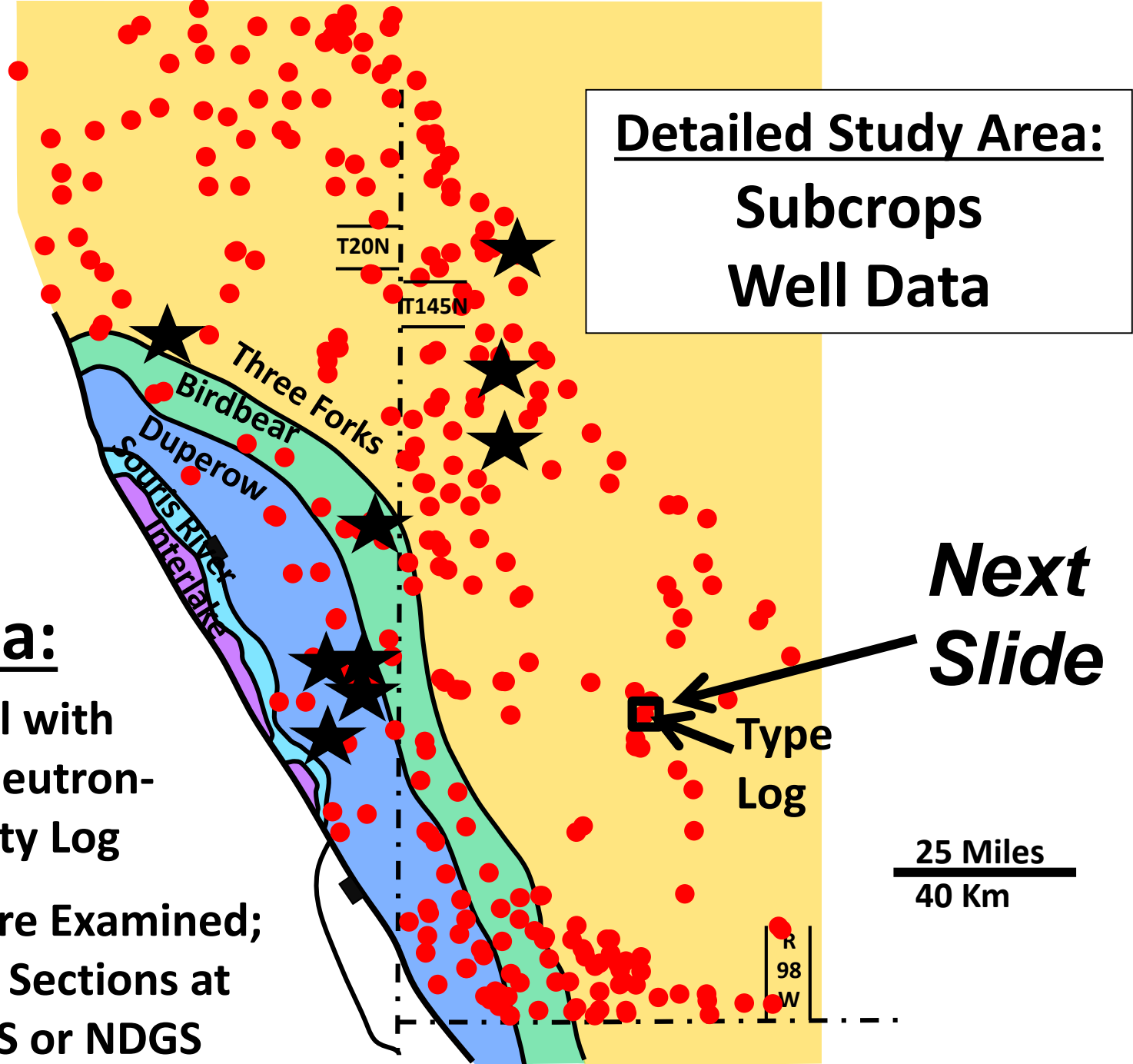


Detailed Study Area:
Subcrops
Well Data

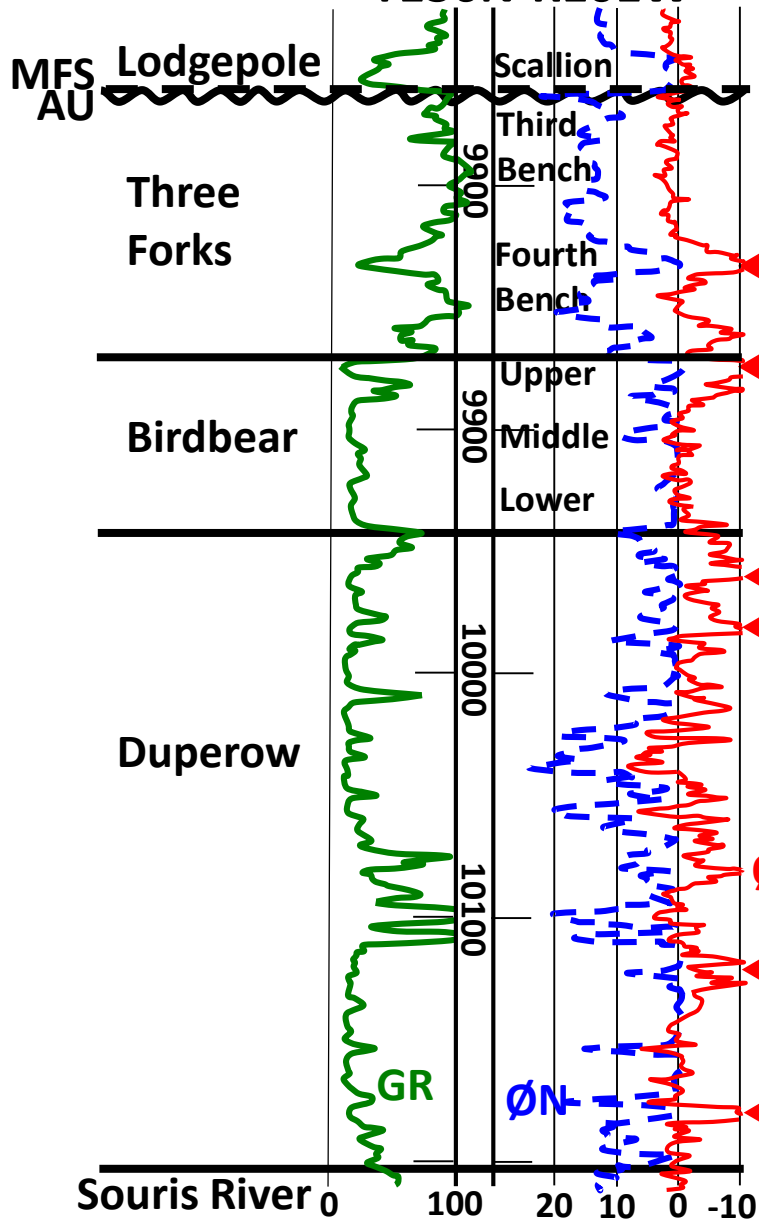
Data:

● = Well with
GR- Neutron-
Density Log

★ = Core Examined;
Thin Sections at
USGS or NDGS



NWNE Sec. 34,
T136N-R101W



Regional
Anhydrite
Beds

Mapped
Anhydrite Bed
Limits

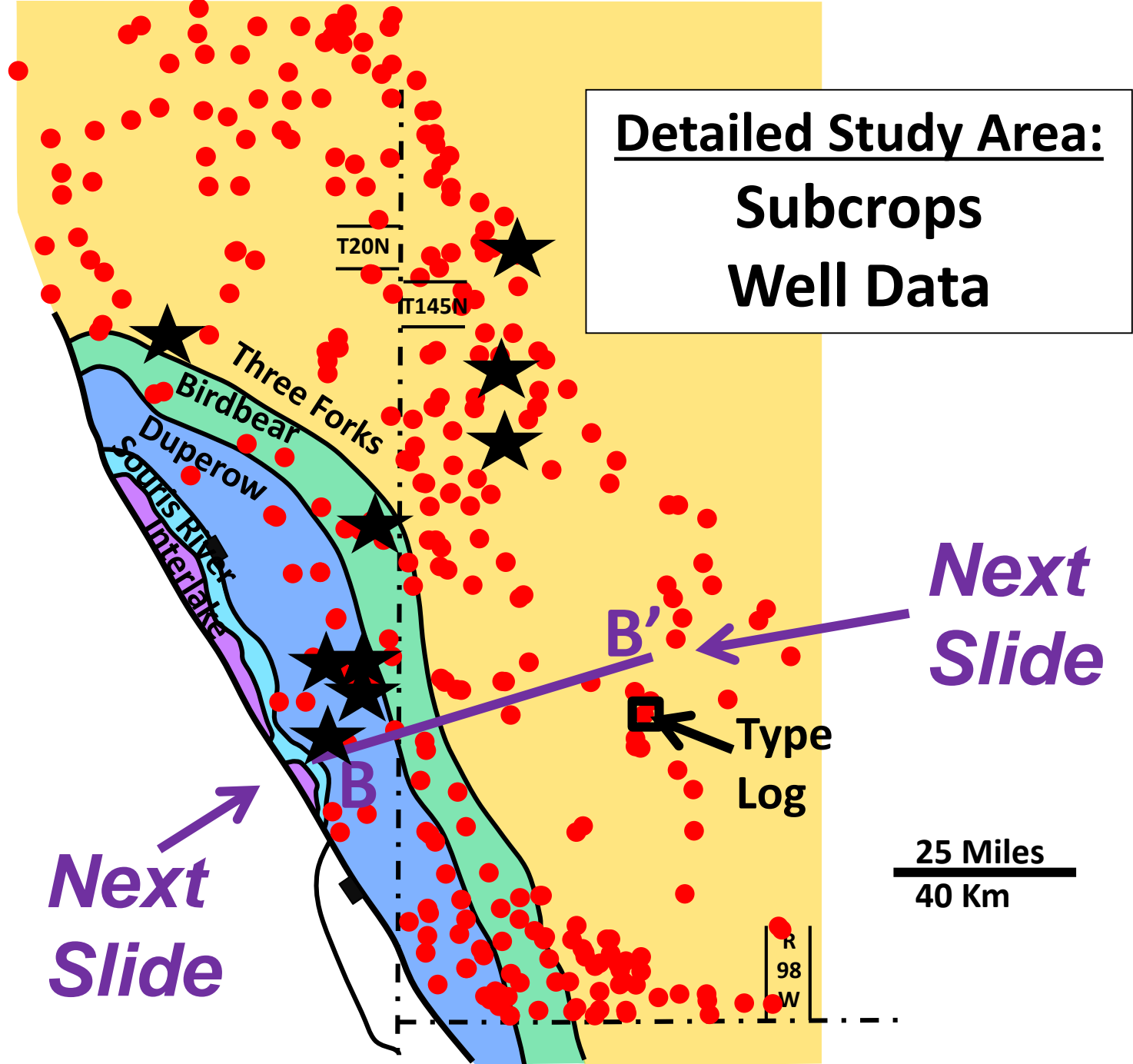
← Lower Three Forks

← Upper Birdbear

} ← Upper Duperow

} ← Lower Duperow

Detailed Study Area:
Subcrops
Well Data



B

Sec. 23,
T9N-
R59E

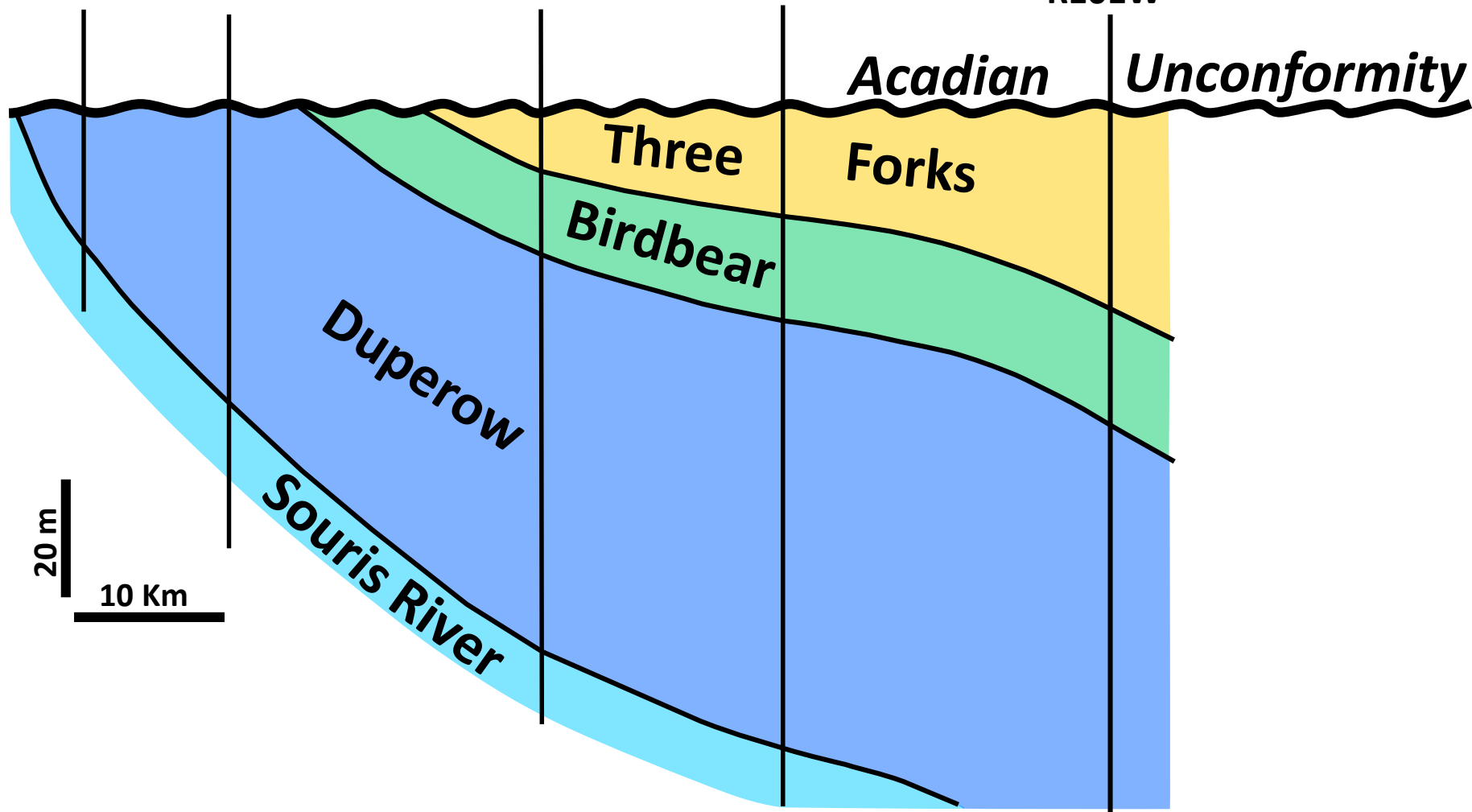
Sec. 9,
T10N-
R60E

Upper Devonian Cross Section

Sec. 9,
T136N-
R105W

Sec. 28,
T137N-
R103W

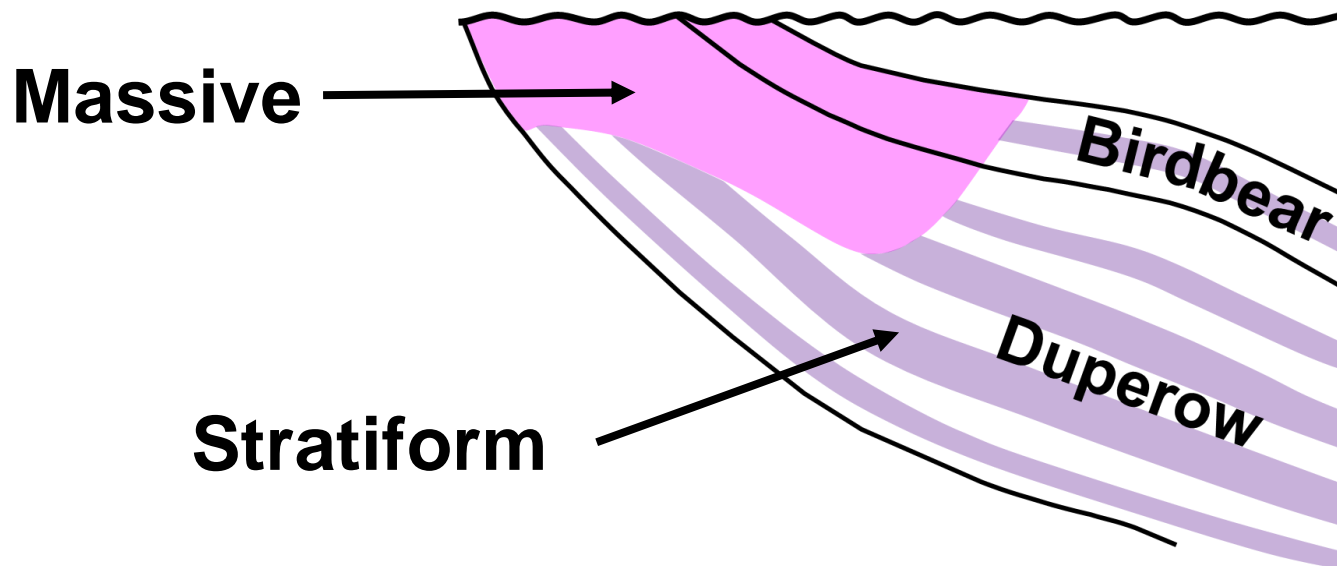
Sec. 34,
T136N-
R101W

B'

Diagenesis

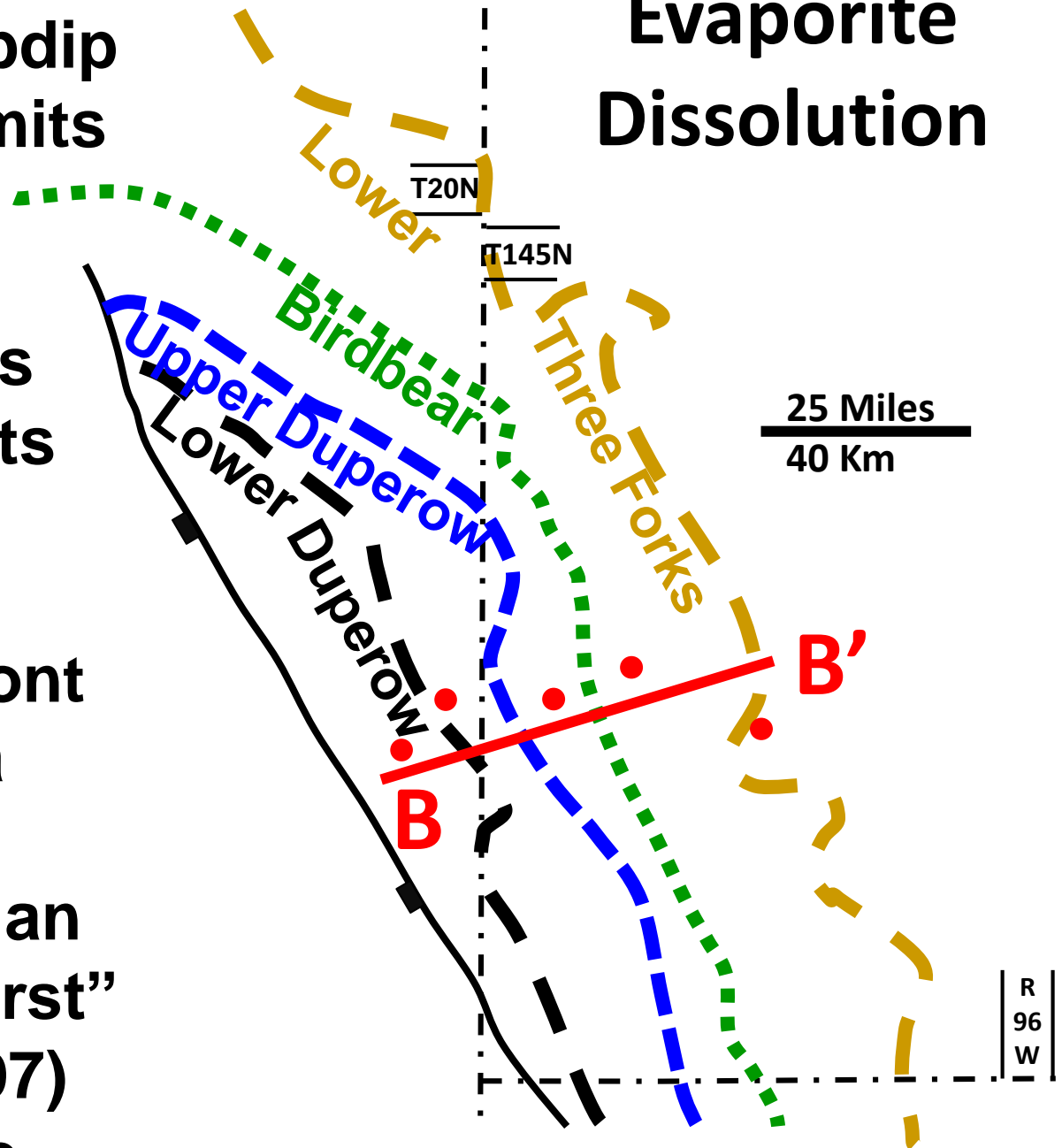
2 Major, Regional and *Mappable* diagenetic features will be discussed:

1. Evaporite dissolution
2. Massive dolomitization
 - “Massive” = lack of stratiform geometry



- Map shows updip dissolution limits for regionally continuous anhydrite beds
- Evaporite limits define a sub-unconformity dissolution front
- This defines a type of karst referred to as an “Evaporite Karst” (Johnson, 1997)
- B-B’ next slide

Evaporite Dissolution



B

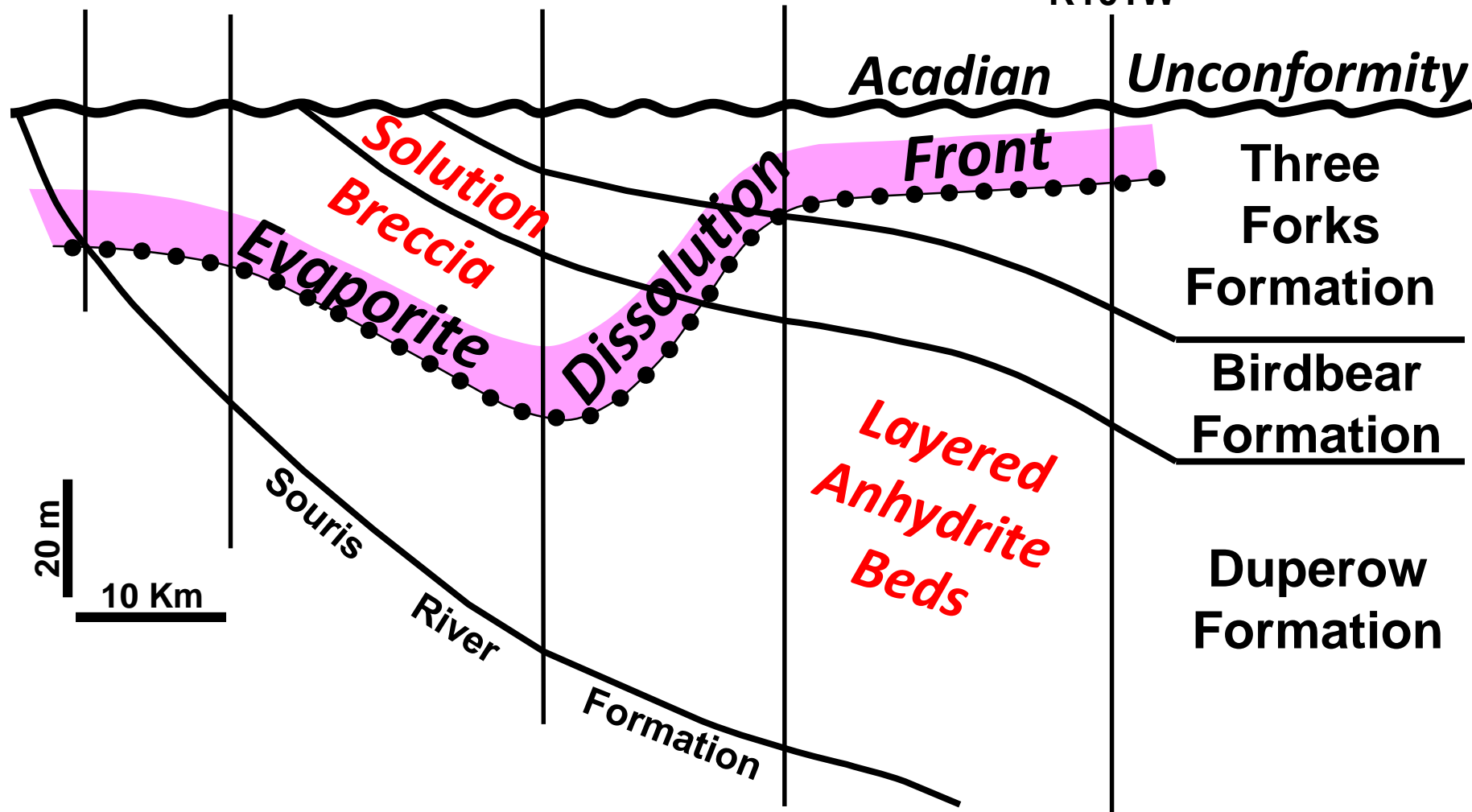
B'

Sec. 23, T9N-
R59E Sec. 9, T10N-
R60E

Sec. 9, T136N-
R105W

Sec. 28, T137N-
R103W

Sec. 34, T136N-
R101W



Map of Subcrop Massive Dolostone

Regionally,
massive
dolostone
(pink) occurs
in all subcrop
carbonates
below the
Acadian
unconformity

Duperow-

Red River-

Deadwood

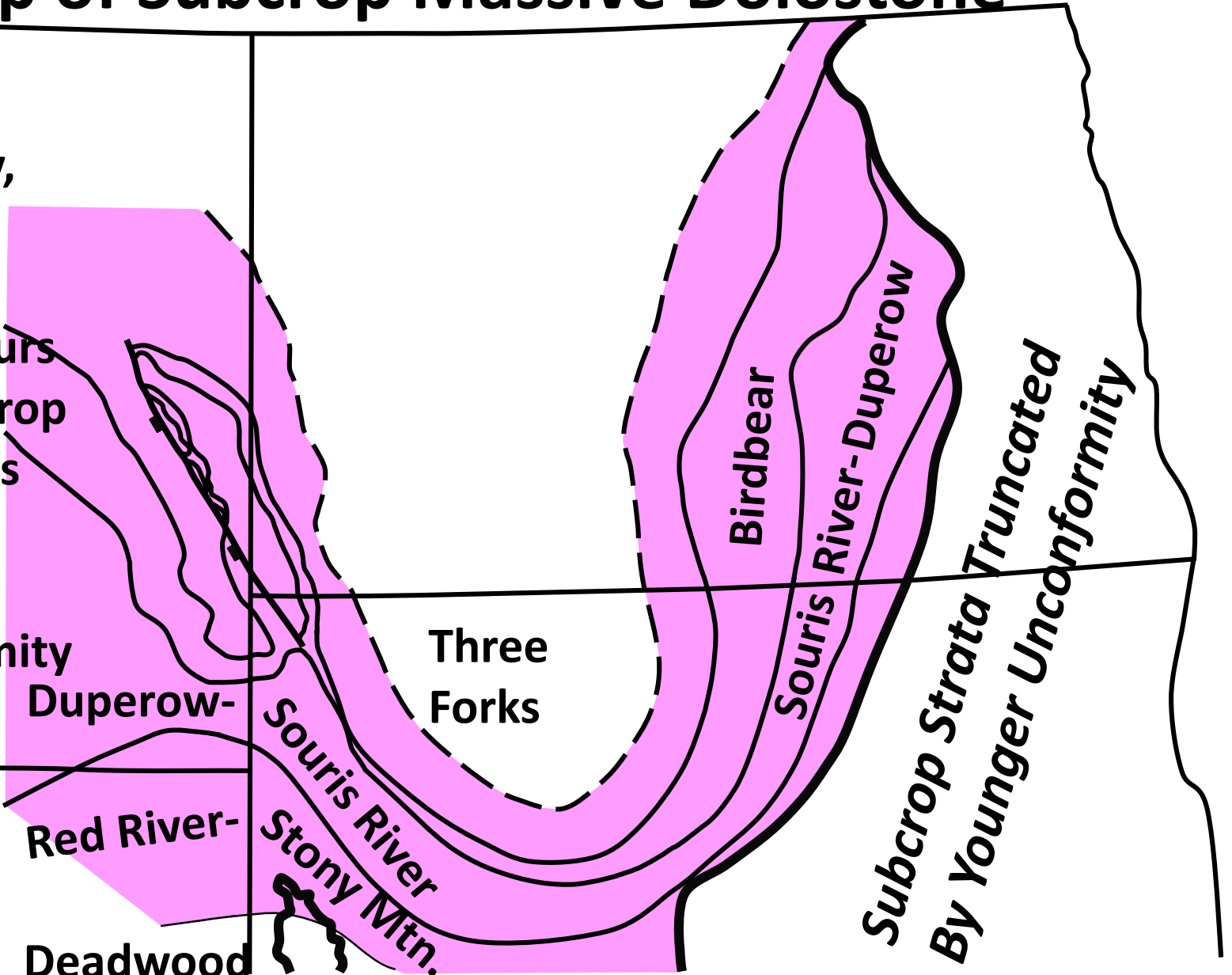
Souris River
Stony Mtn.

Three
Forks

Birdbear

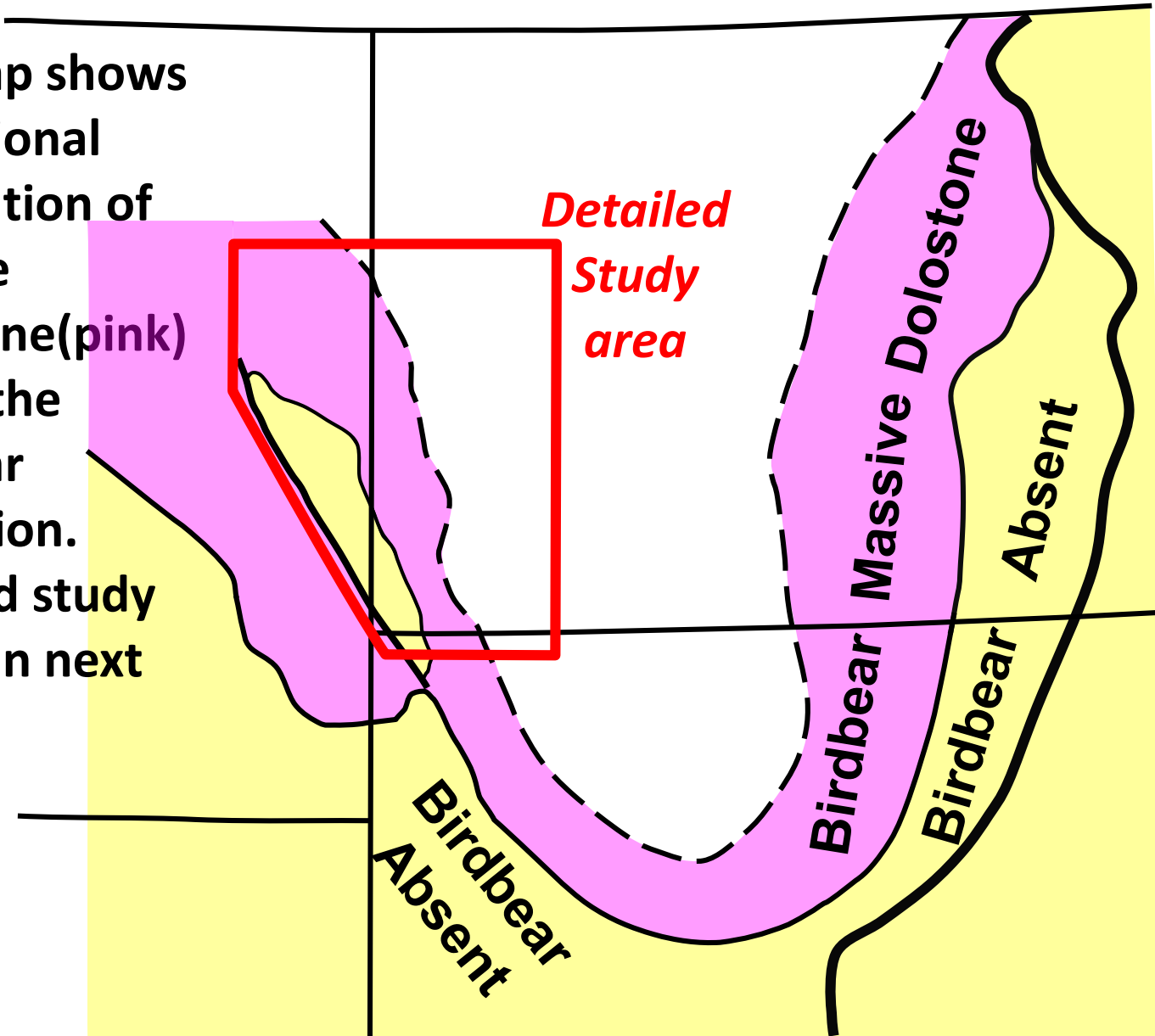
Souris River-Duperow

*Subcrop Strata Truncated
By Younger Unconformity*



Birdbear Massive Dolostone Distribution

This map shows the regional distribution of massive dolostone (pink) within the Birdbear Formation. Detailed study area is in next slide.



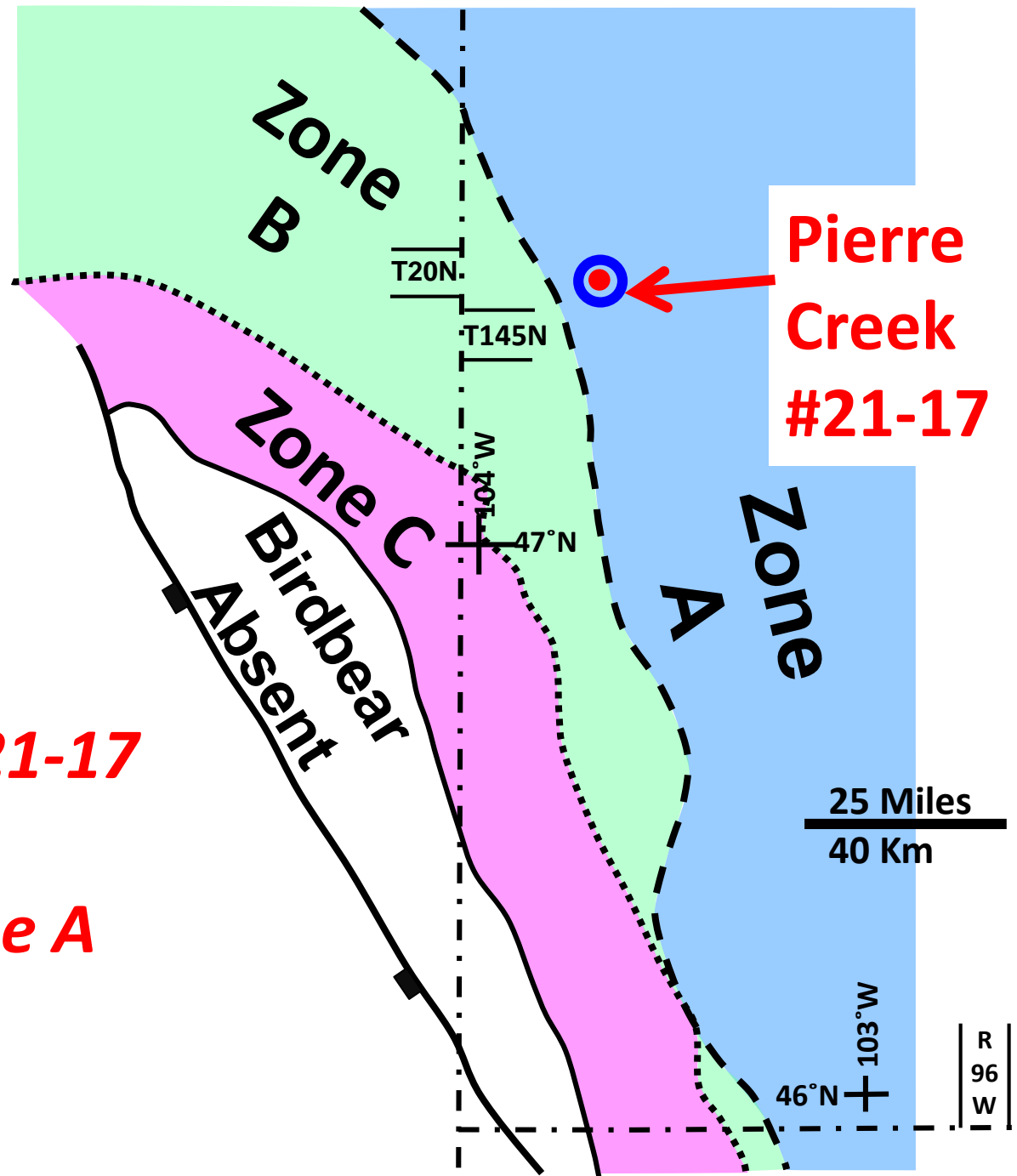
Birdbear Diagenetic Zones

- Lithology-based
- Mappable
- Zone A = layered anhydrite, dolostone and limestone
- Zone B = anhydrite beds and massive dolostone
- Zone C = Massive dolostone only

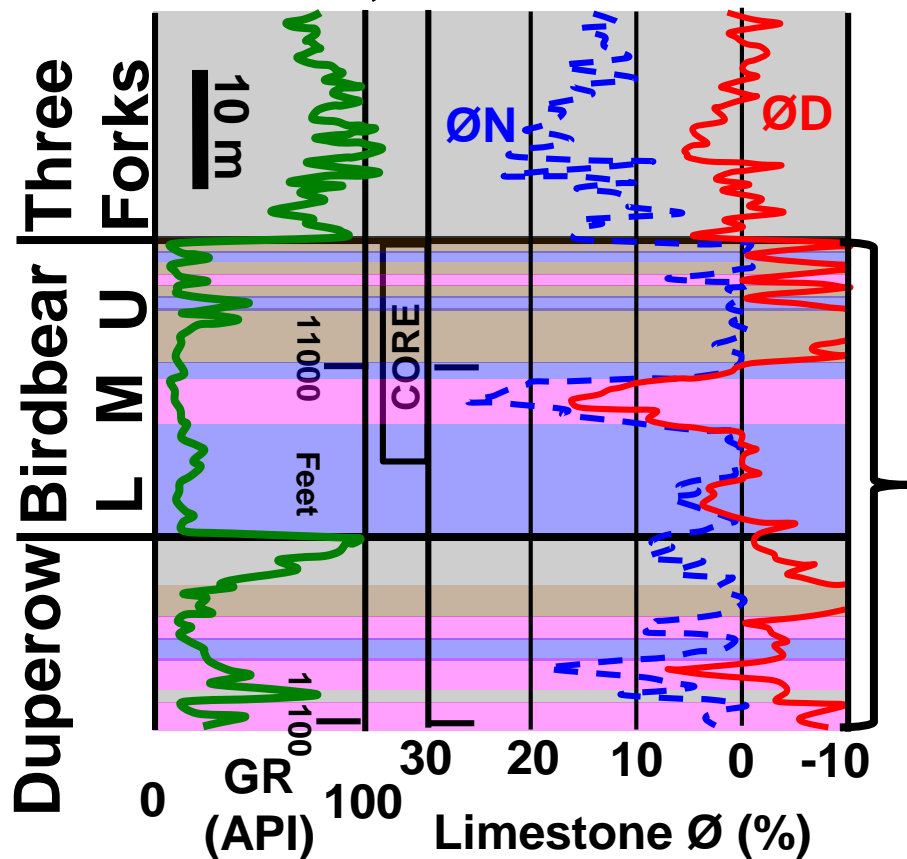


Birdbear
Diagenetic
Zones

Zone A
Pierre Creek #21-17
illustrates
Diagenetic Zone A

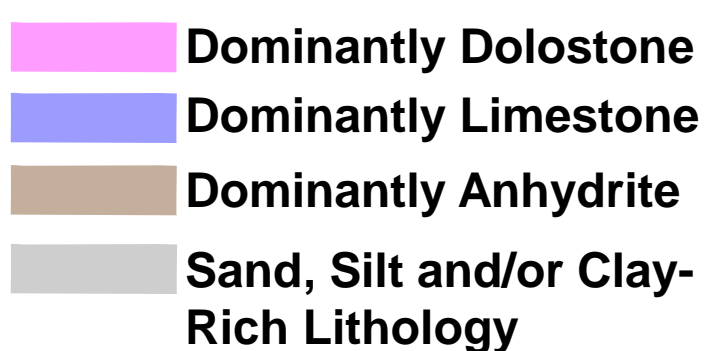


**Pierre Creek #21-17:
Sec. 17, T146N-R102W**



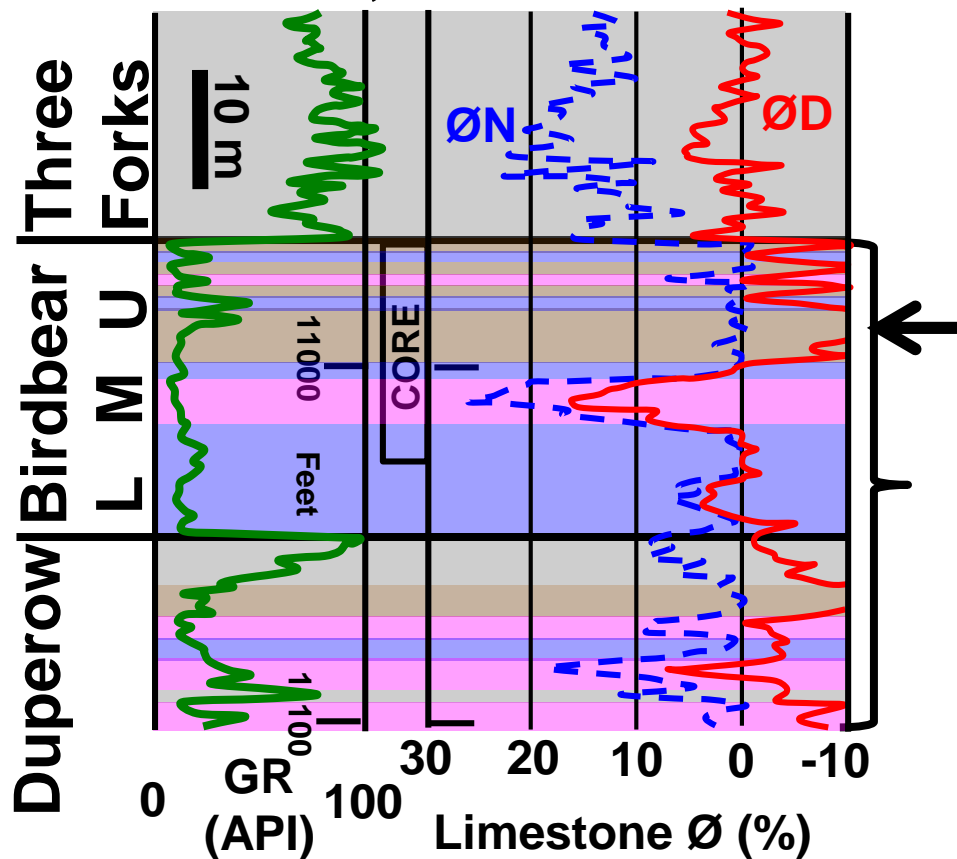
Diagenetic Zone A

Layered limestone, dolostone and anhydrite = downdip area unaffected or marginally affected by sub-unconformity diagenesis



Basic Lithology from GR-Neutron-Density logs tied to core data

**Pierre Creek #21-17:
Sec. 17, T146N-R102W**

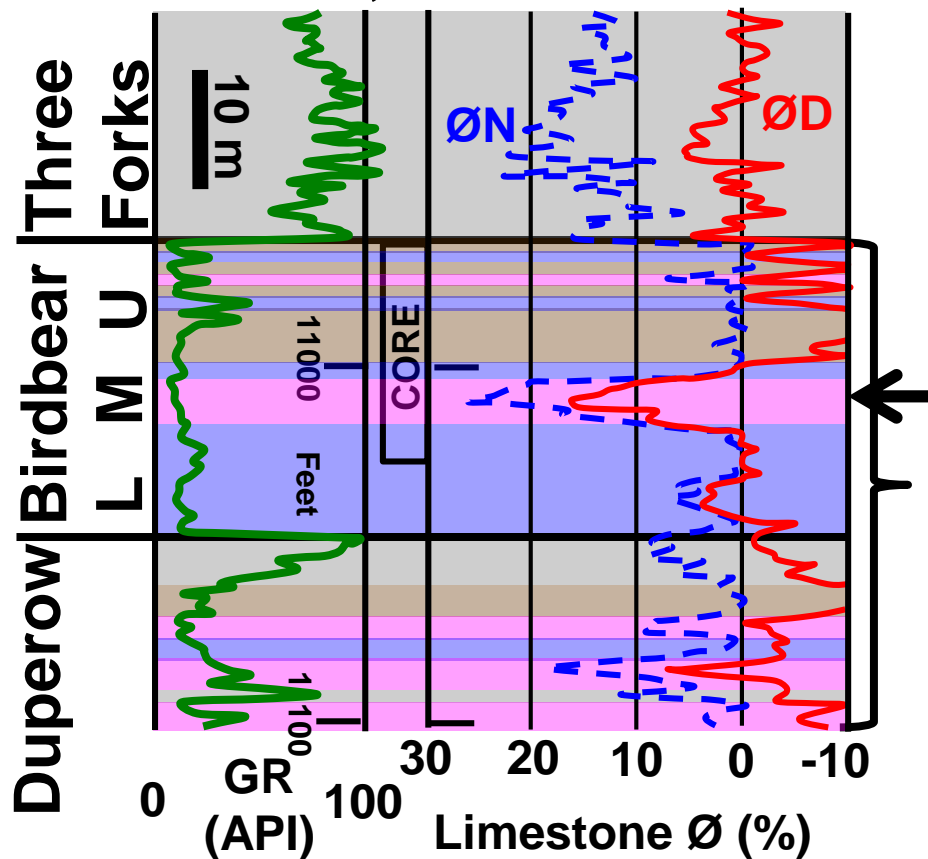


- Dominantly Dolostone
- Dominantly Limestone
- Dominantly Anhydrite
- Sand, Silt and/or Clay-Rich Lithology

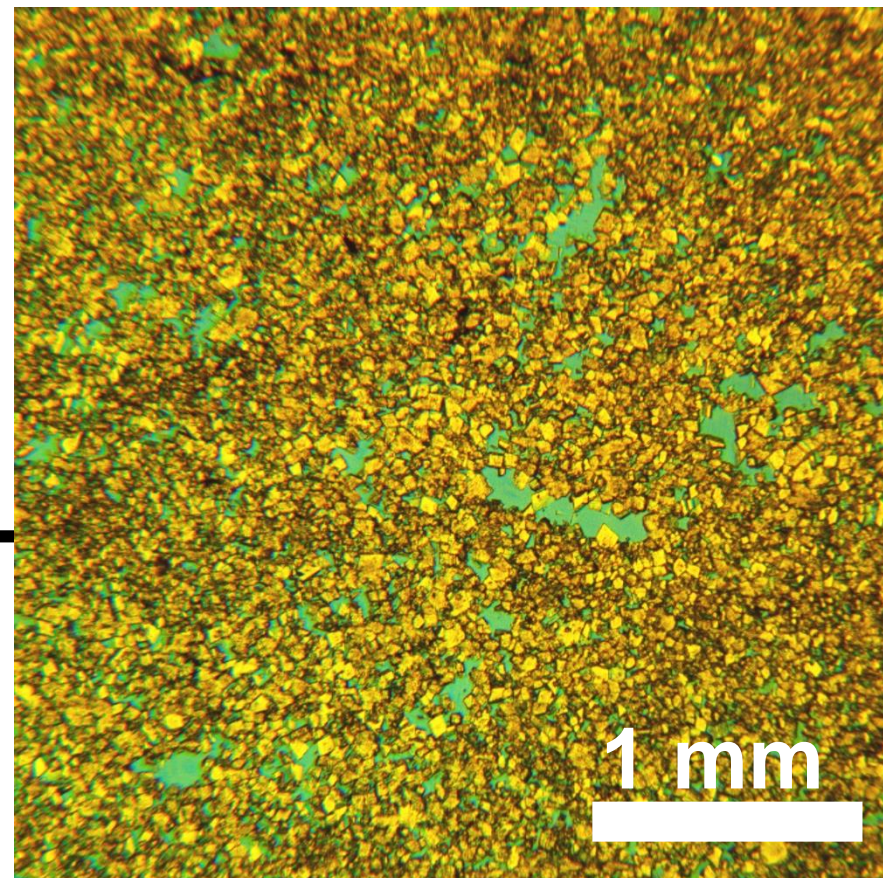
Birdbear Anhydrite:



**Pierre Creek #21-17:
Sec. 17, T146N-R102W**



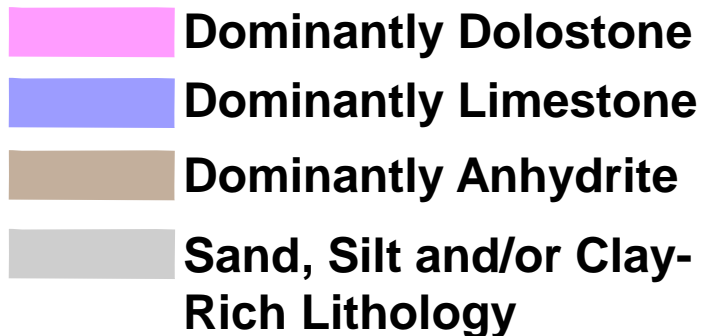
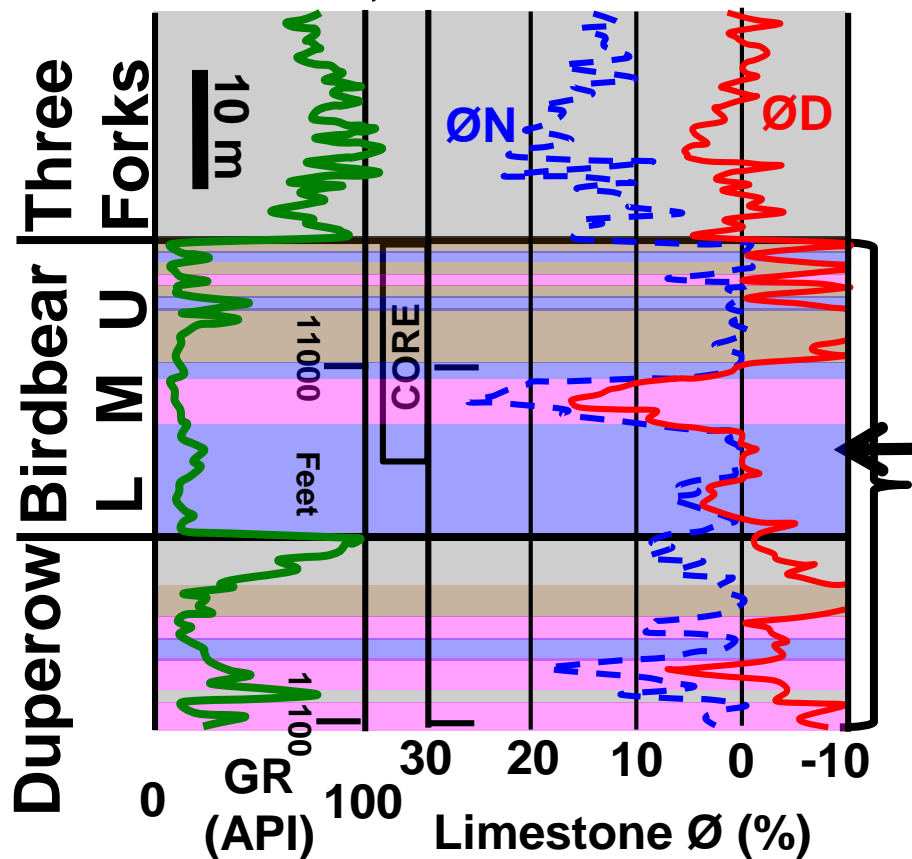
- Dominantly Dolostone
- Dominantly Limestone
- Dominantly Anhydrite
- Sand, Silt and/or Clay-Rich Lithology



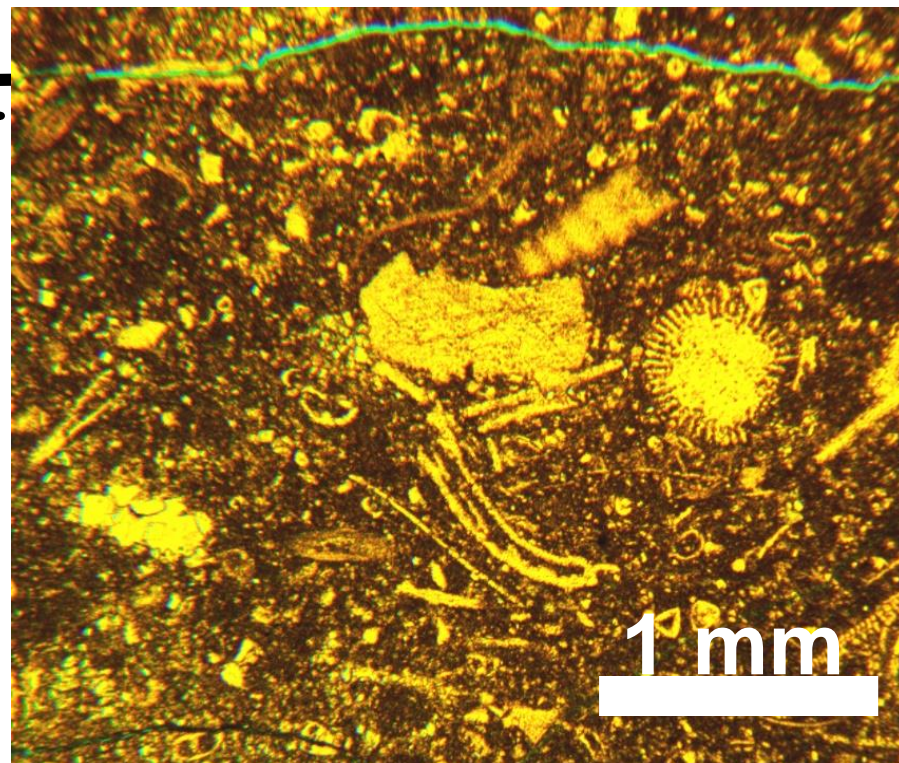
**11,007.5' core depth
Dolomitic Skeletal
Mudstone**

**Average crystal size = 30 μ
 $\text{Ø} = 20$; Ka = 71 md**

**Pierre Creek #21-17:
Sec. 17, T146N-R102W**

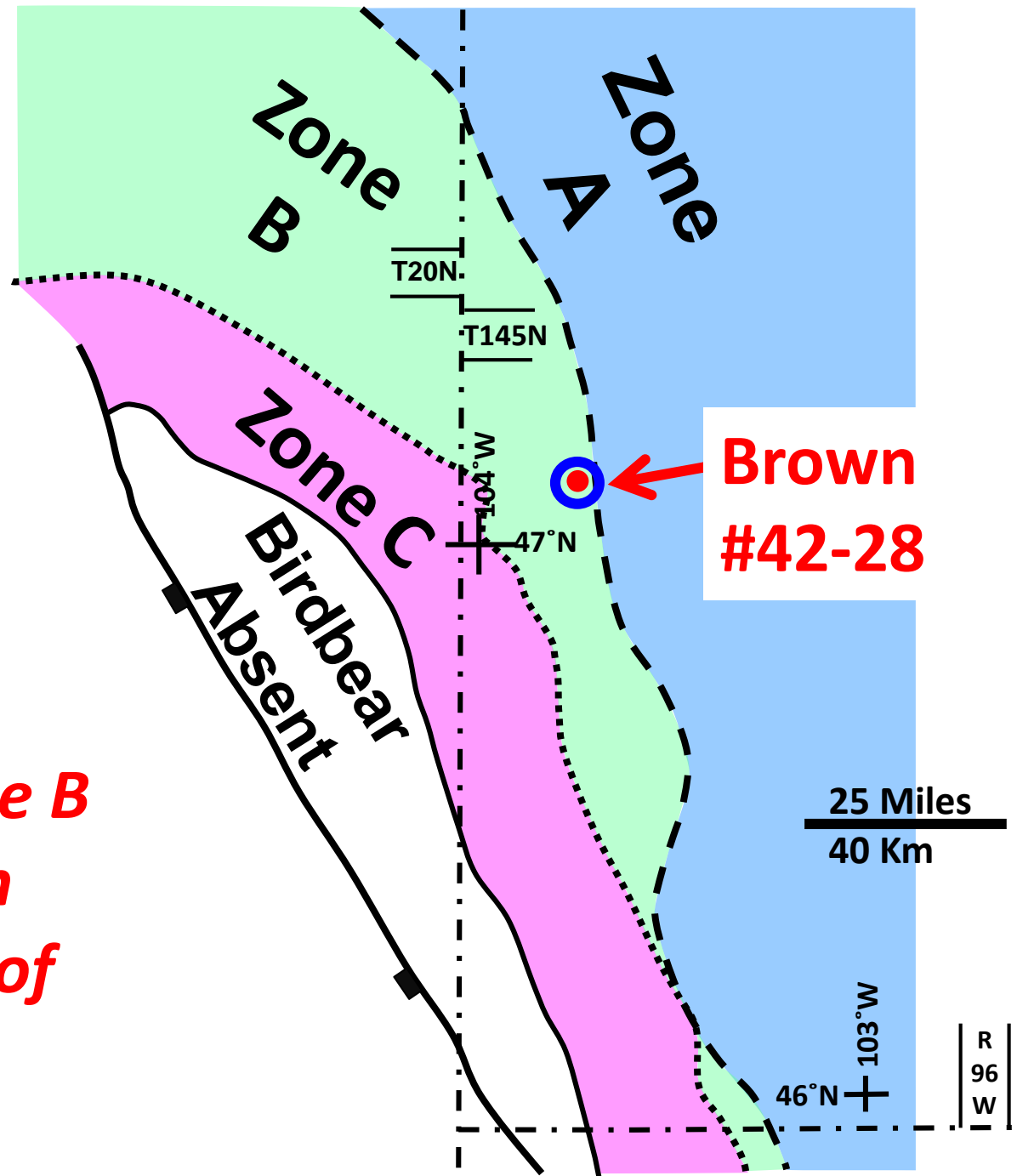


**11,019' core depth
Limestone
Skeletal wackestone**



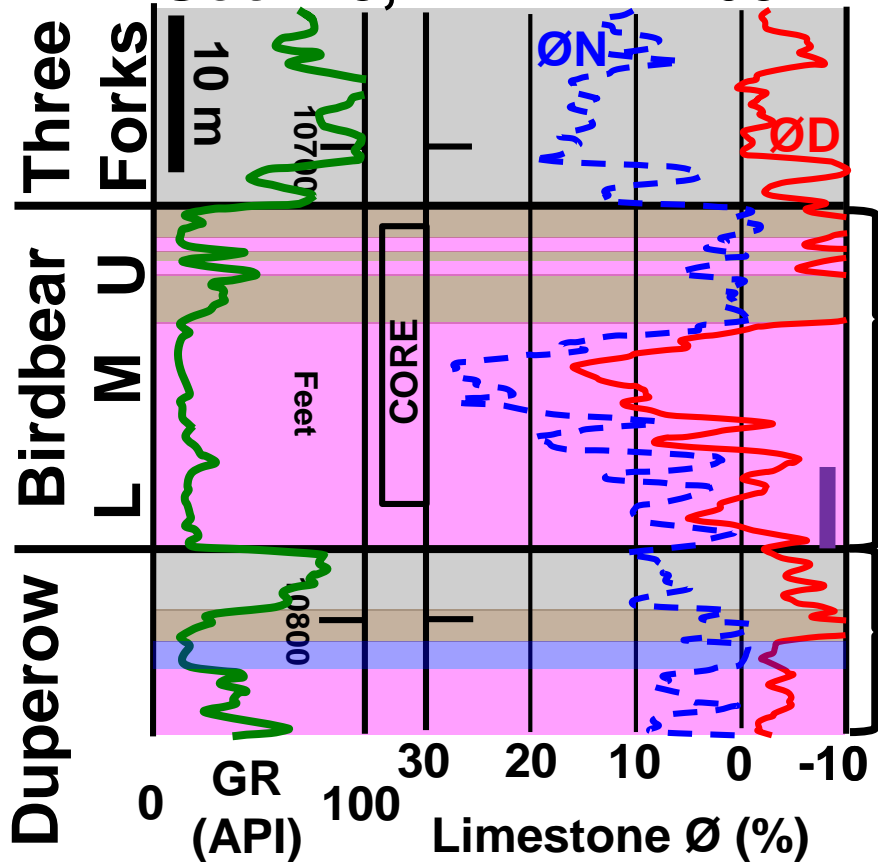
Birdbear Diagenetic Zones

*Brown #42-28
illustrates
Diagenetic Zone B
Note: well is on
downdip edge of
Zone B*



Brown #42-28

Sec. 28, T142N-R103W



Middle Well

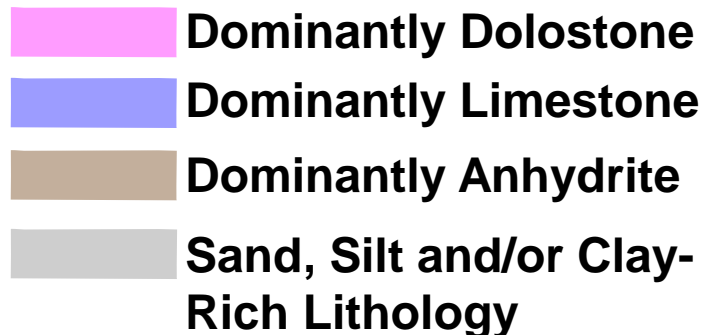
Diagenetic Zone B

Layered anhydrite
and dolostone

*Calcareous Dolostone with
patchy dolomite distribution*

Diagenetic Zone A

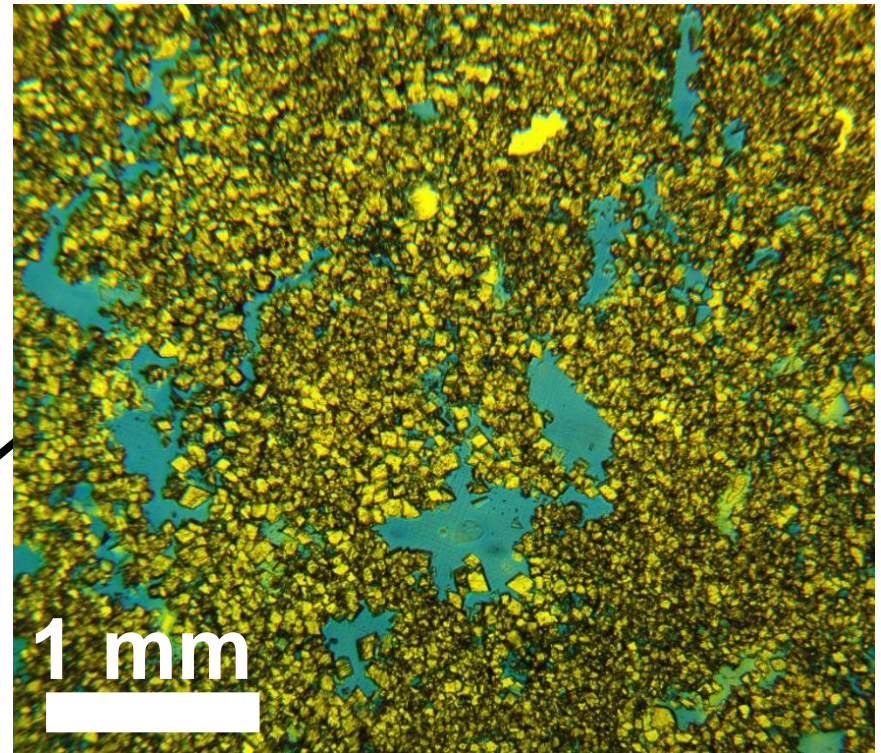
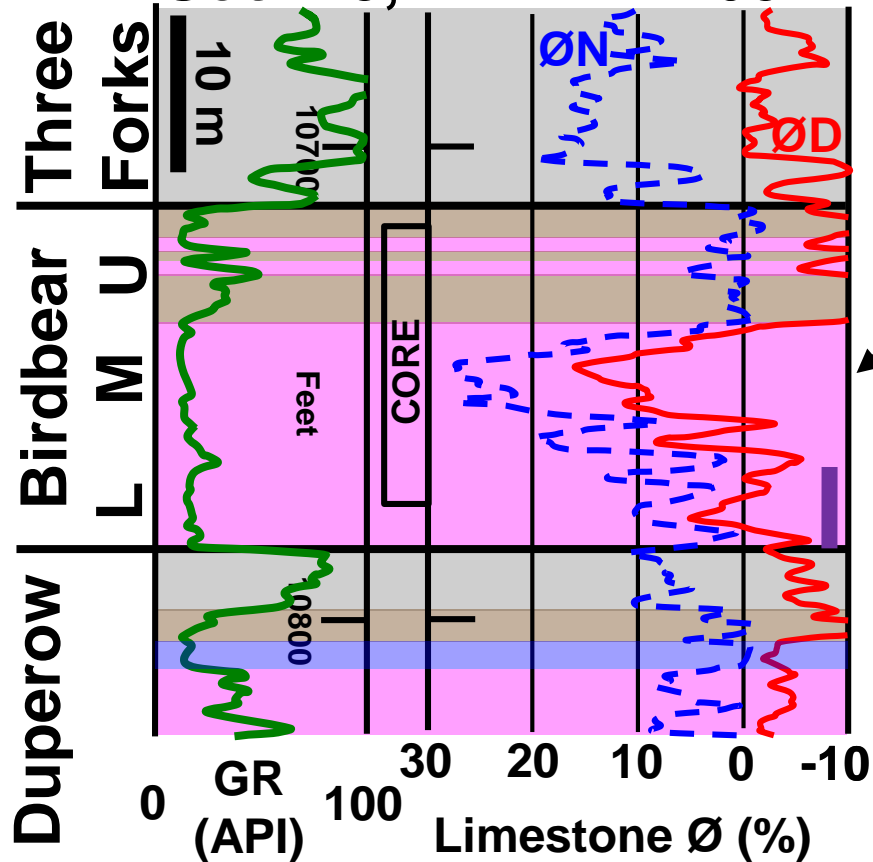
B
A



***Note: well lies near
A/B transition in the
Birdbear Formation***

Brown #42-28

Sec. 28, T142N-R103W



10,730' core depth

Dolomitic

Mudstone/wackestone

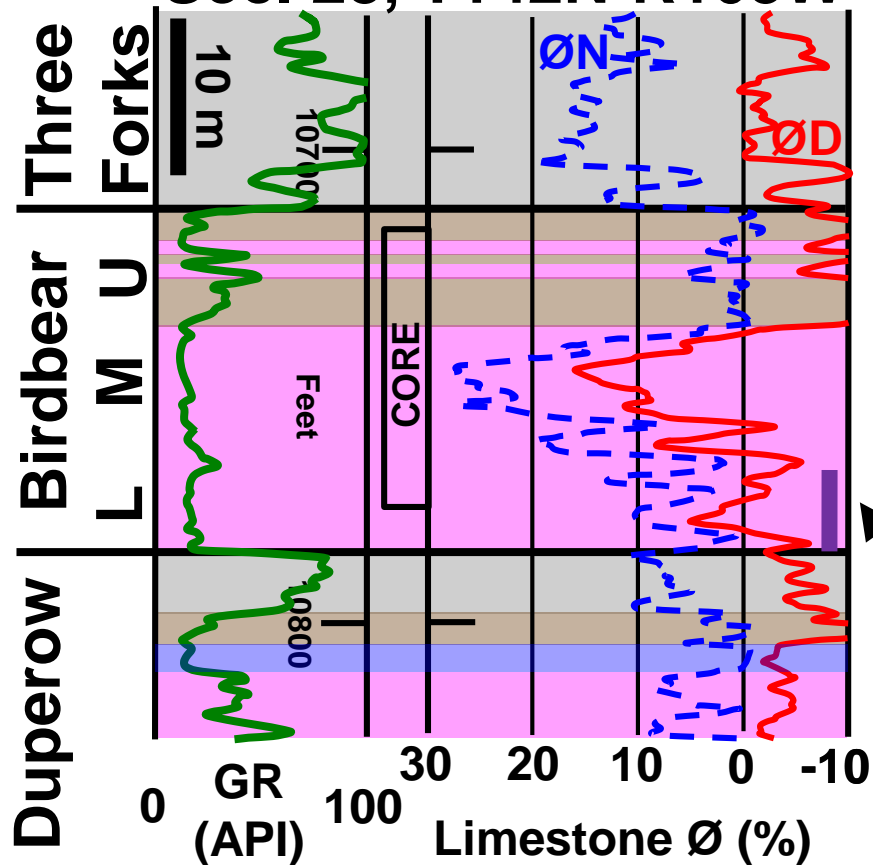
Average crystal size = 50 μ

$\text{Ø} = 22$; $\text{Ka} = 261$ md

- Dominantly Dolostone
- Dominantly Limestone
- Dominantly Anhydrite
- Sand, Silt and/or Clay-Rich Lithology

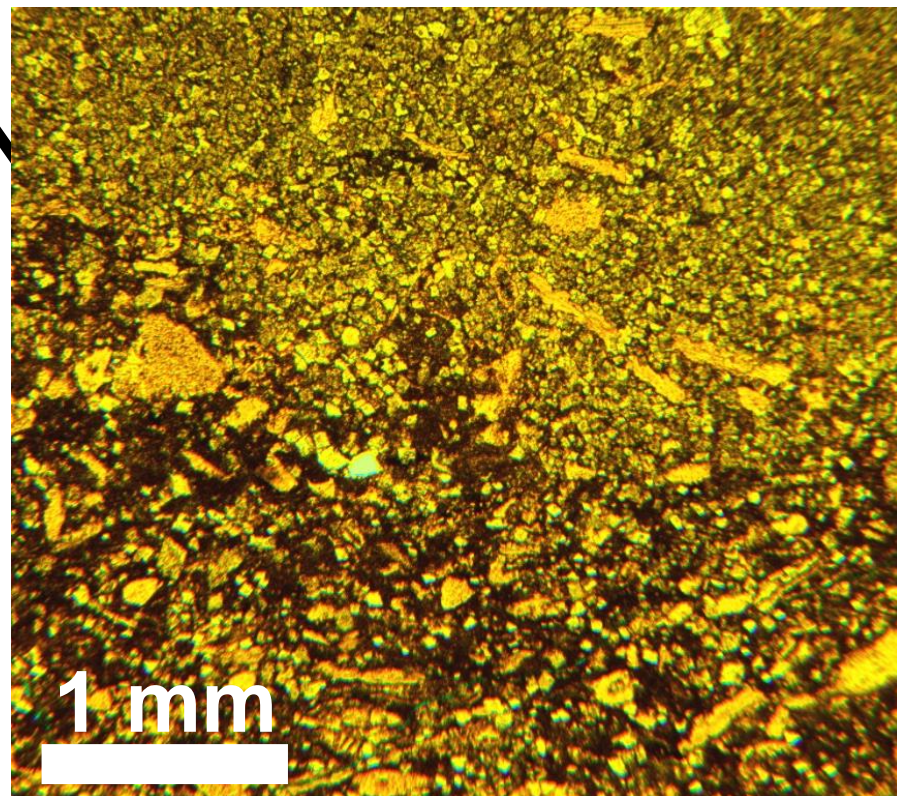
Brown #42-28

Sec. 28, T142N-R103W



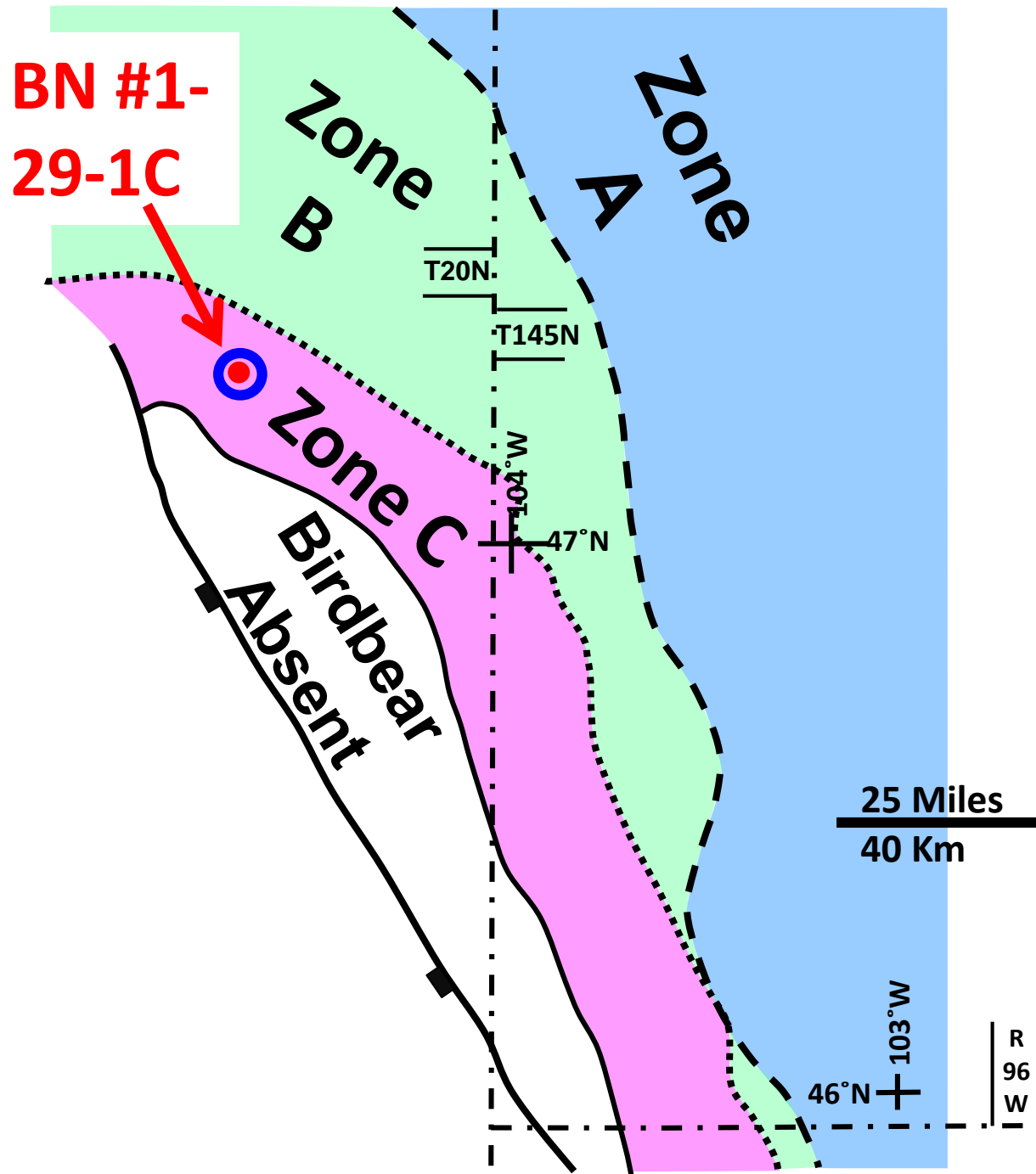
**Calcareous Dolostone at
Zone A/B Transition; Note
“Patchy” Dolomite
Distribution**

**10,752' core depth
Average crystal size = 25 μ**

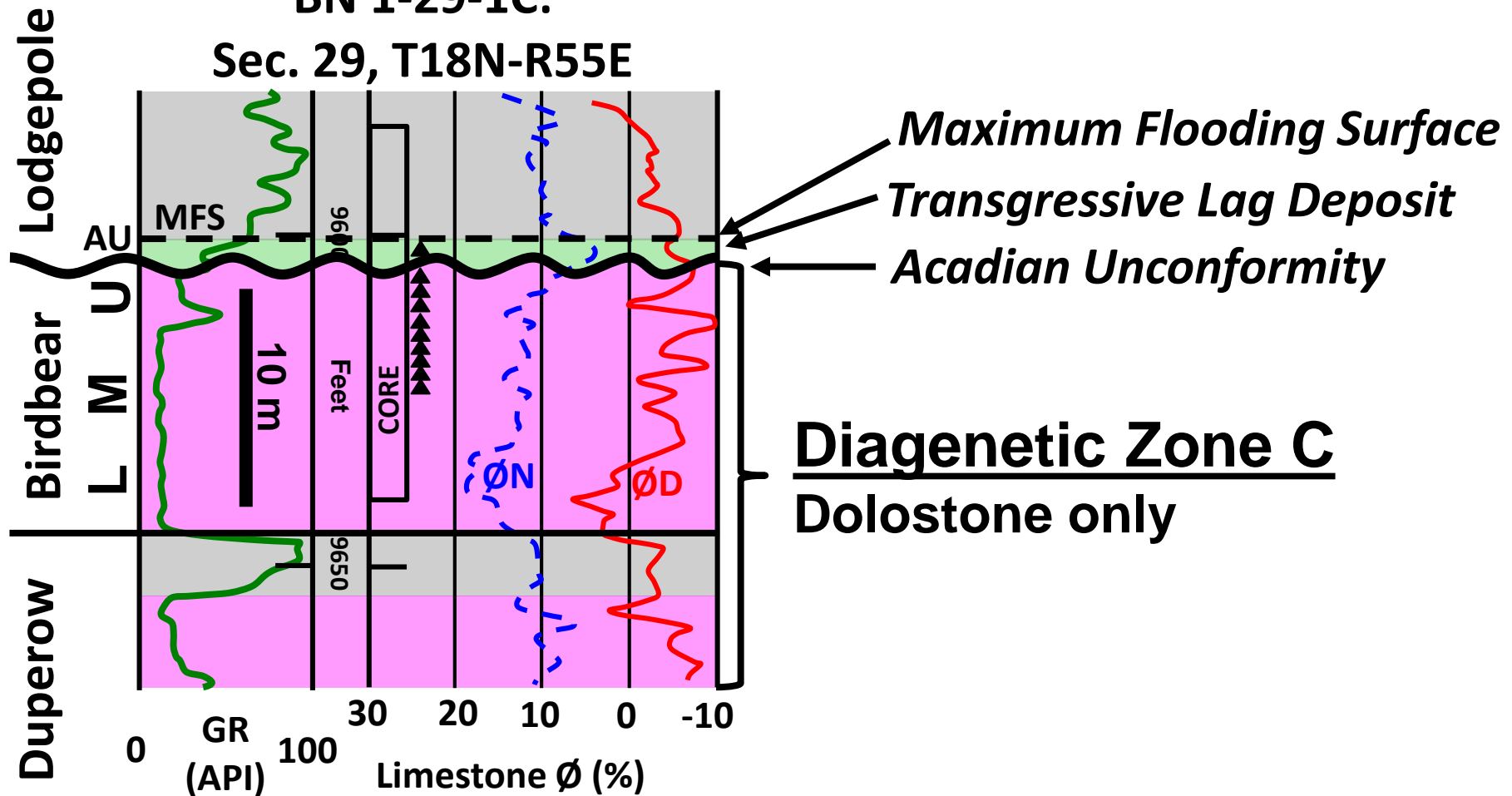


Birdbear
Diagenetic
Zones

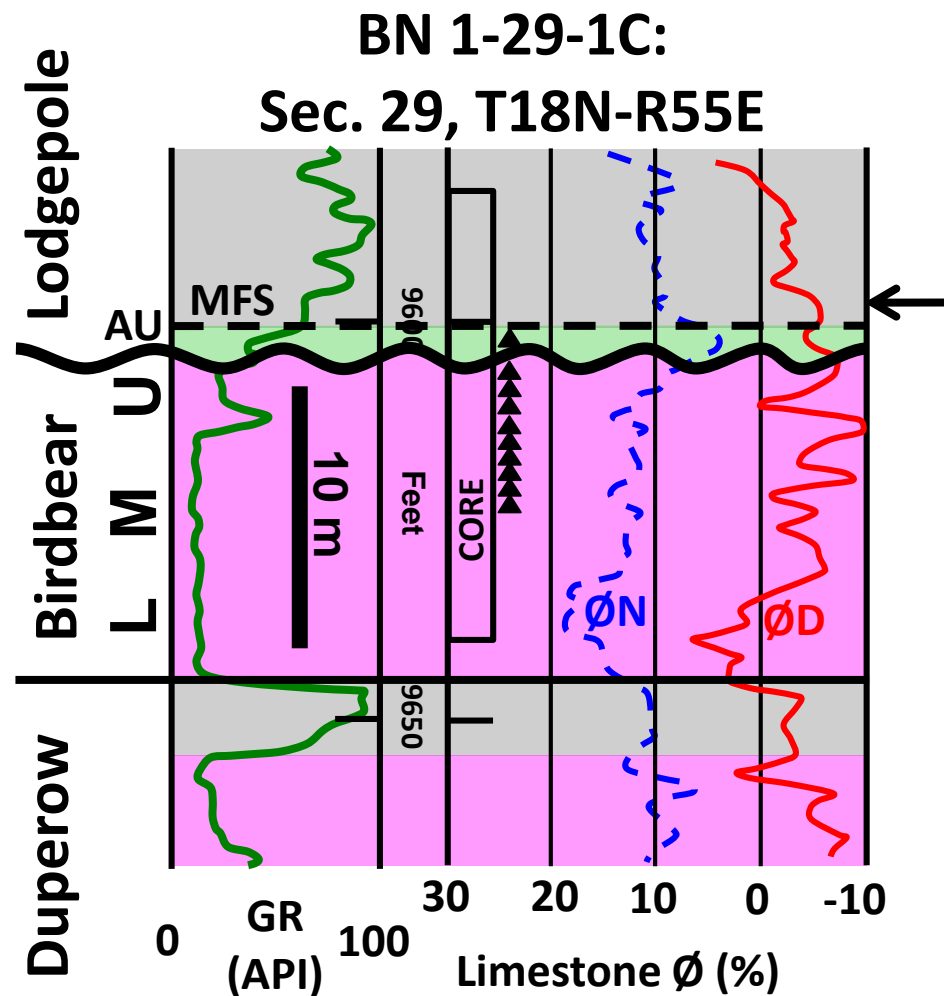
BN #1-29-1C
illustrates
Diagenetic
Zone C



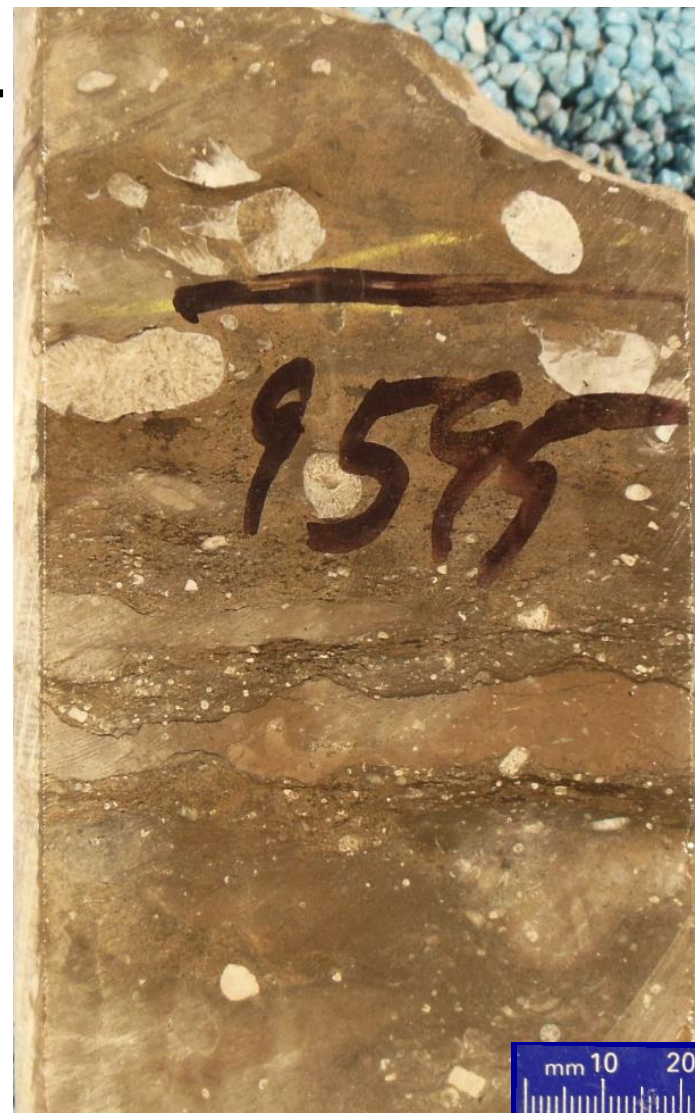
**BN 1-29-1C:
Sec. 29, T18N-R55E**



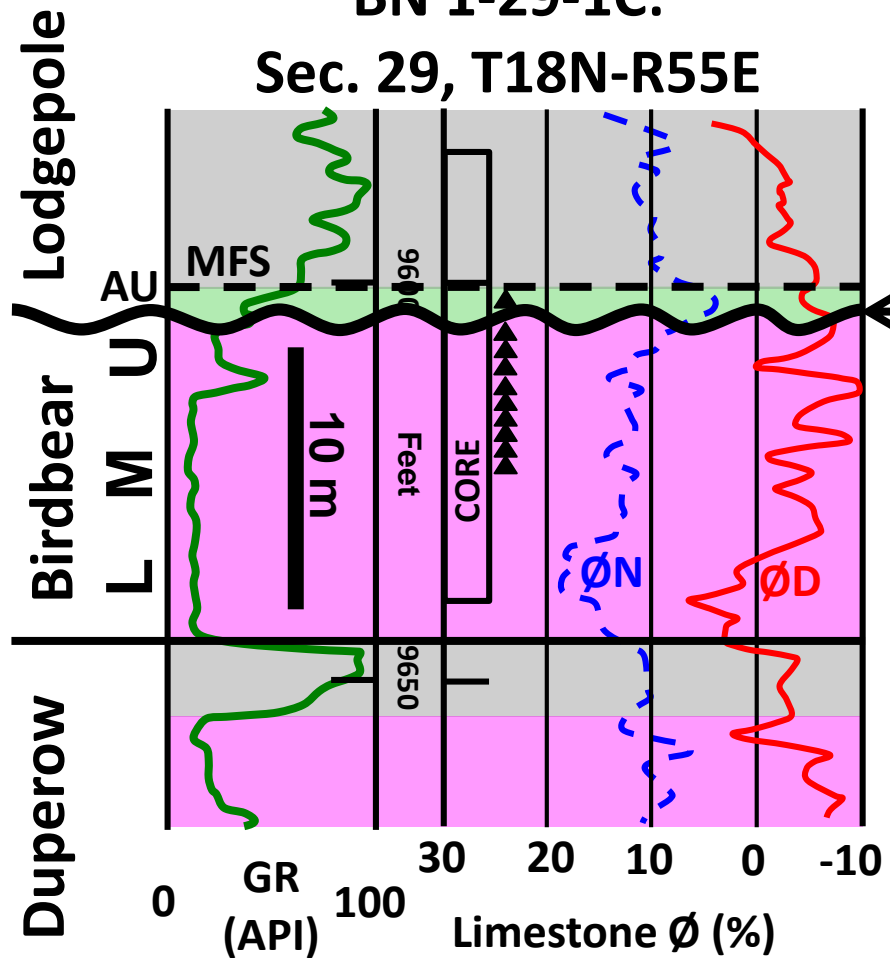
- Dominantly Dolostone
- Skeletal-Fragmental Limestone/Dolostone
- Sand, Silt and/or Clay-Rich Lithology
- ▲ = Breccia



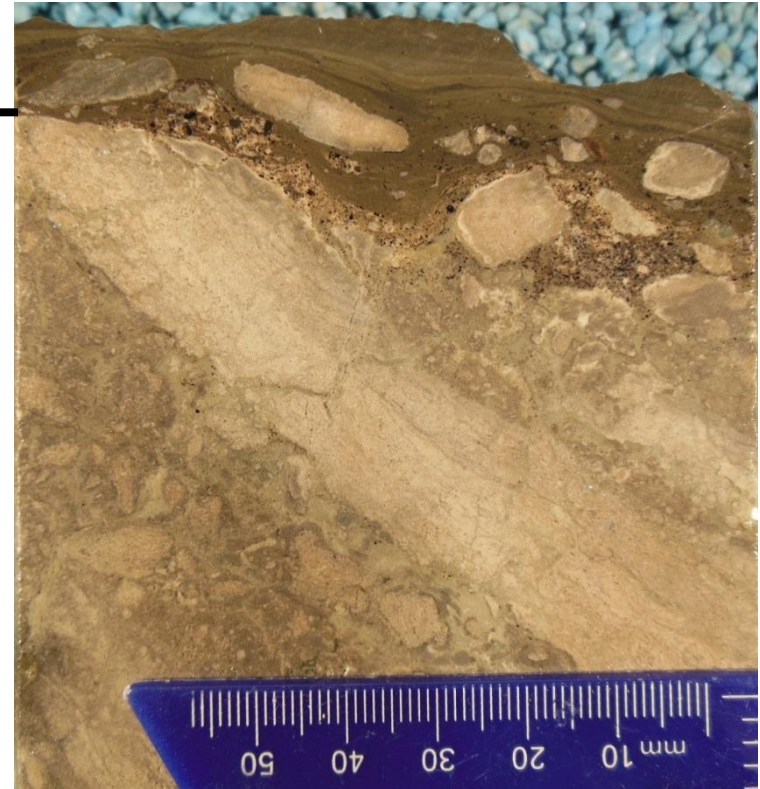
Basal Lodgepole:
shaly, coral-crinoid
limestone



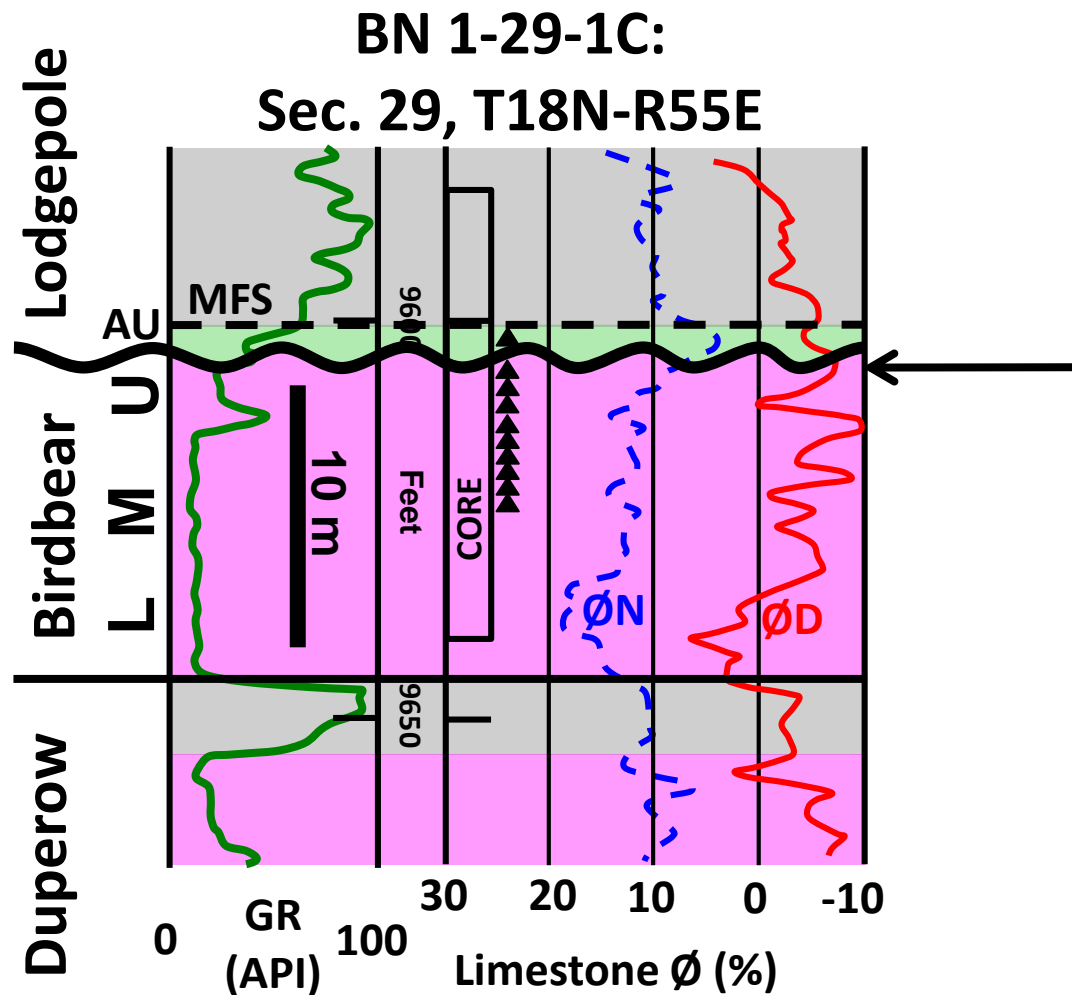
**BN 1-29-1C:
Sec. 29, T18N-R55E**



Acadian Unconformity



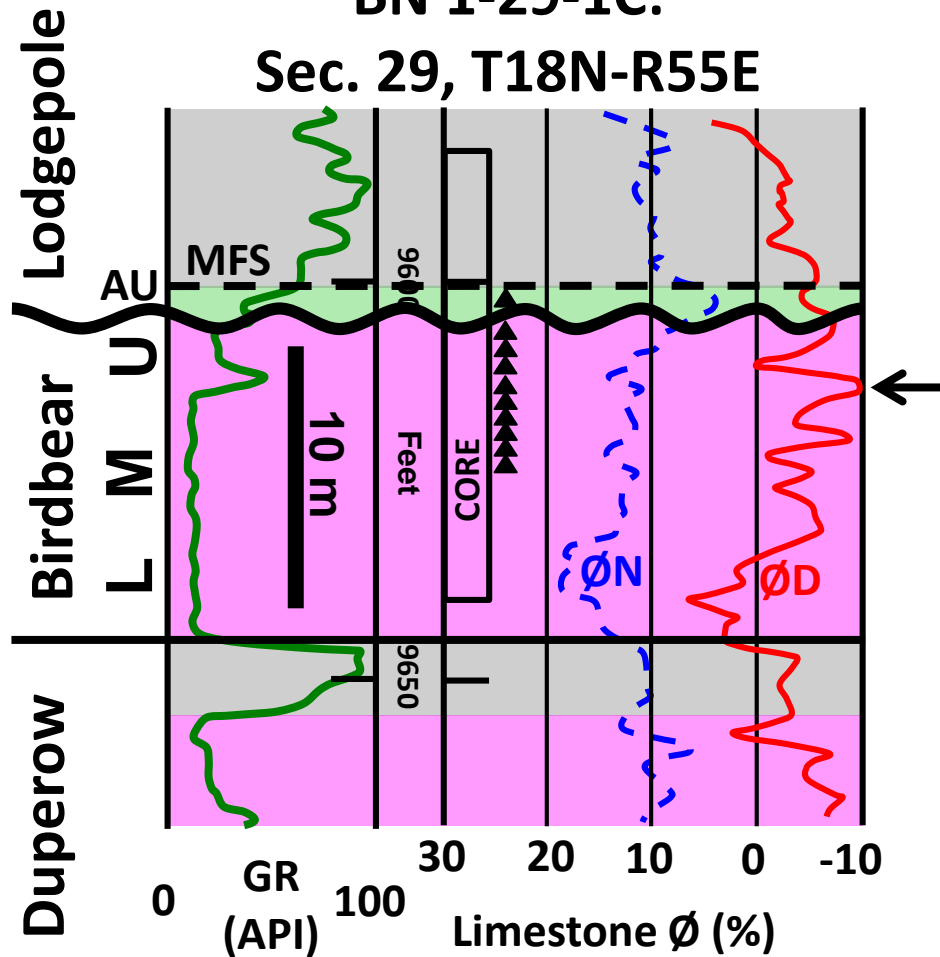
- Dominantly Dolostone
- Skeletal-Fragmental Limestone/Dolostone ▲ = Breccia
- Sand, Silt and/or Clay-Rich Lithology



Birdbear solution breccia (dolostone)



**BN 1-29-1C:
Sec. 29, T18N-R55E**

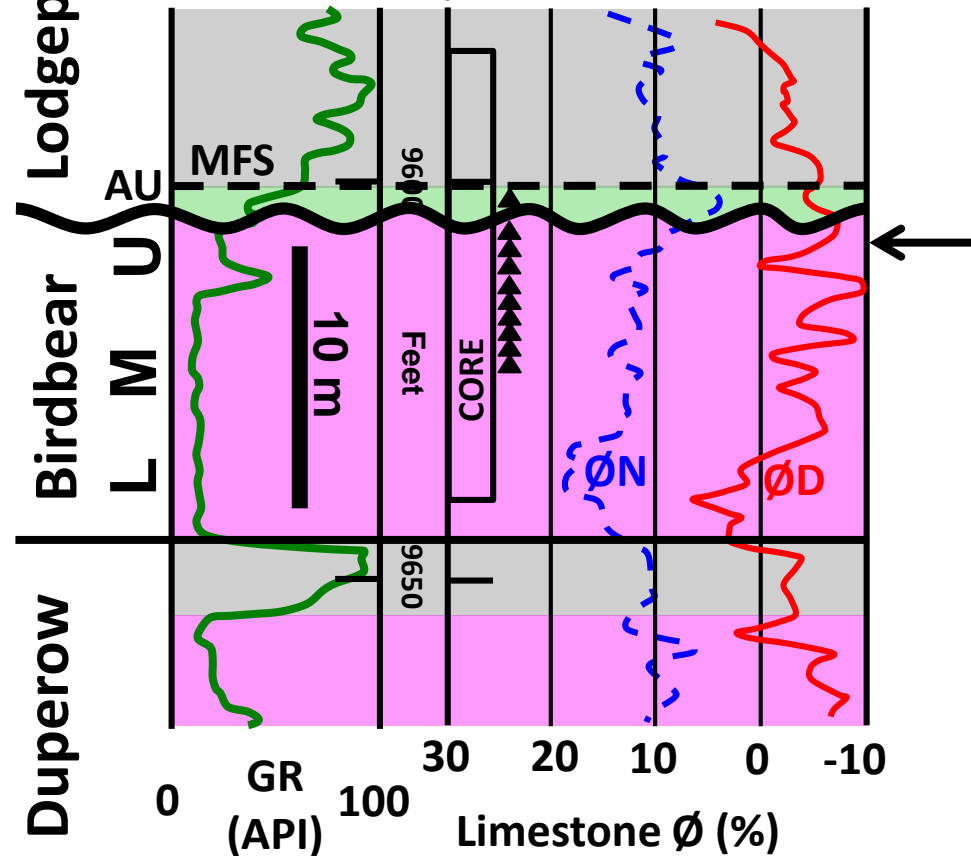


**Anhydrite cement in solution
breccia (anhydrite occurs
throughout Birdbear)**



- Dominantly Dolostone
- Skeletal-Fragmental Limestone/Dolostone
- Sand, Silt and/or Clay-Rich Lithology
- ▲ = Breccia

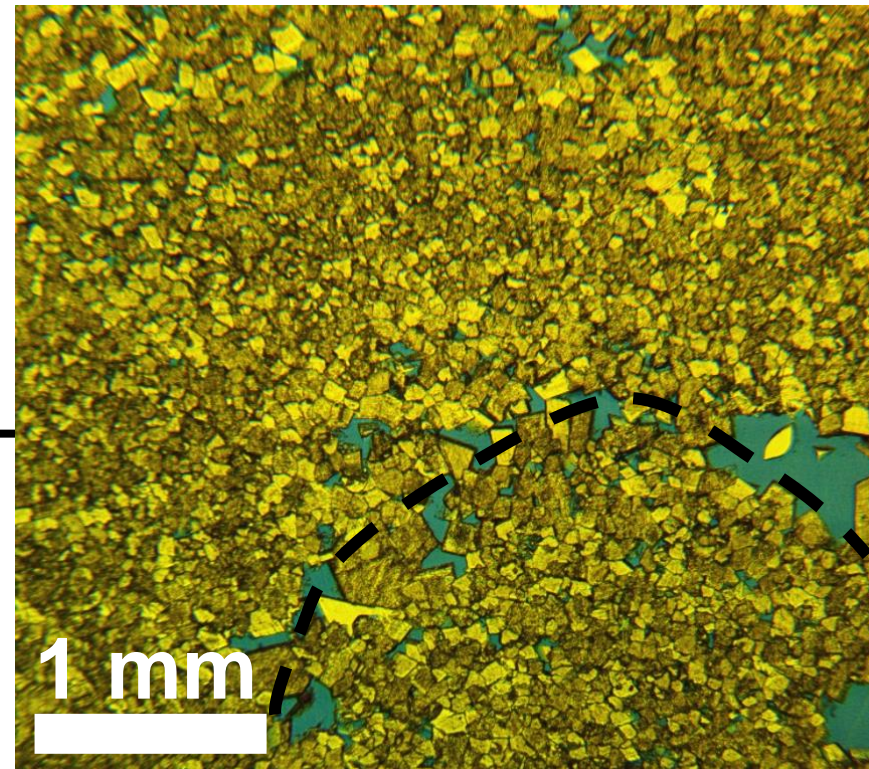
Duperow | Birdbear / Lodgepole



**Skeletal-Fragmental
Limestone/Dolostone**

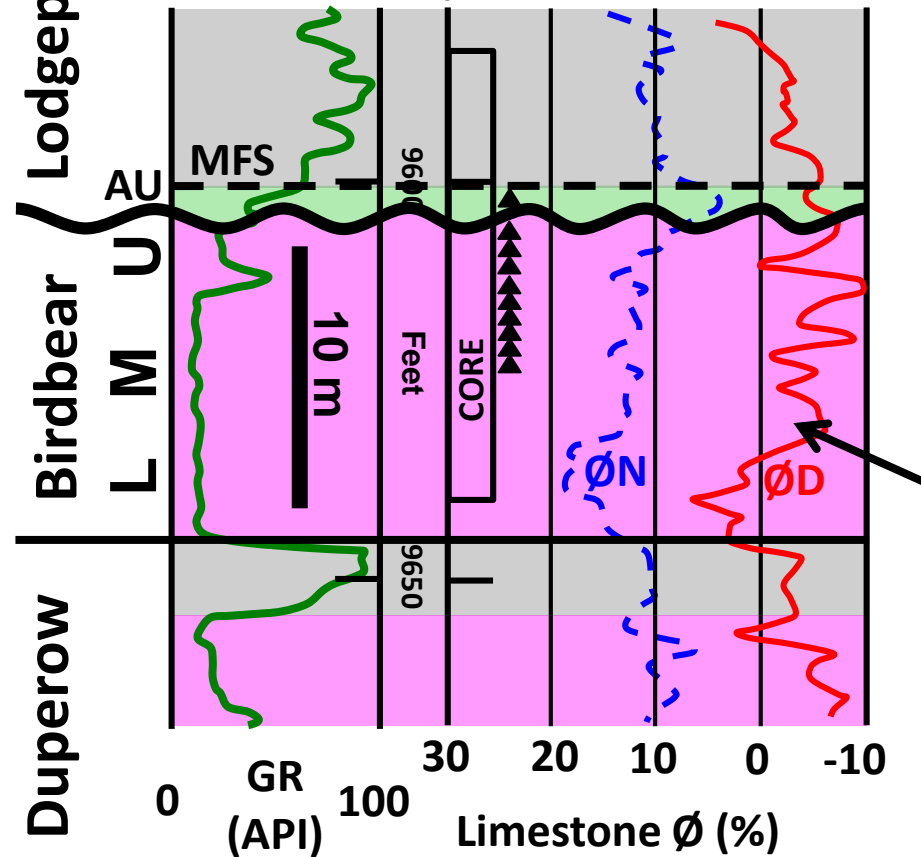
▲ = Breccia

Sand, Silt and/or Clay-Rich Lithology



9607.5' core depth
Breccia with fabric-
destructive dolomite; note
breccia fragment outlined in
lower-right corner of photo
Average crystal size = 100 μ
Thin Section porosity = 8%

Duperow | **Birdbear** / **Lodgepole**

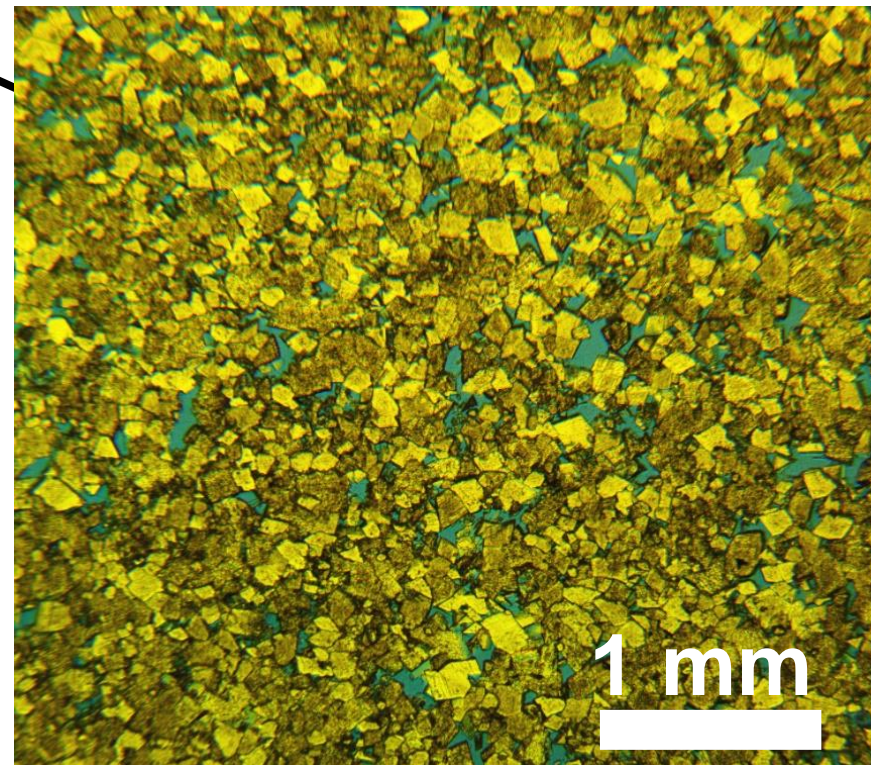


Dominantly Dolostone

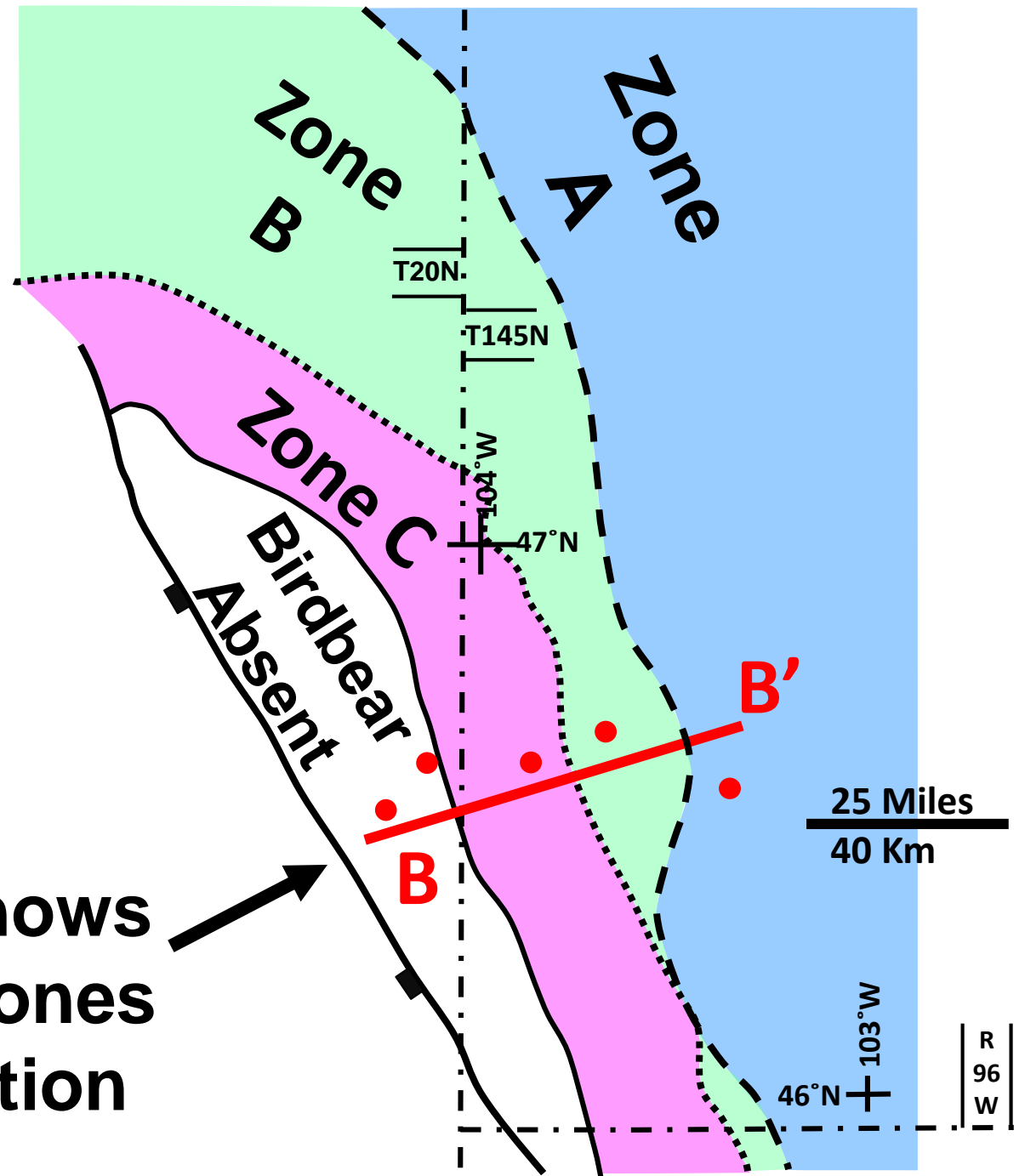
Skeletal-Fragmental Limestone/Dolostone ▲ = Breccia

Sand, Silt and/or Clay-Rich Lithology

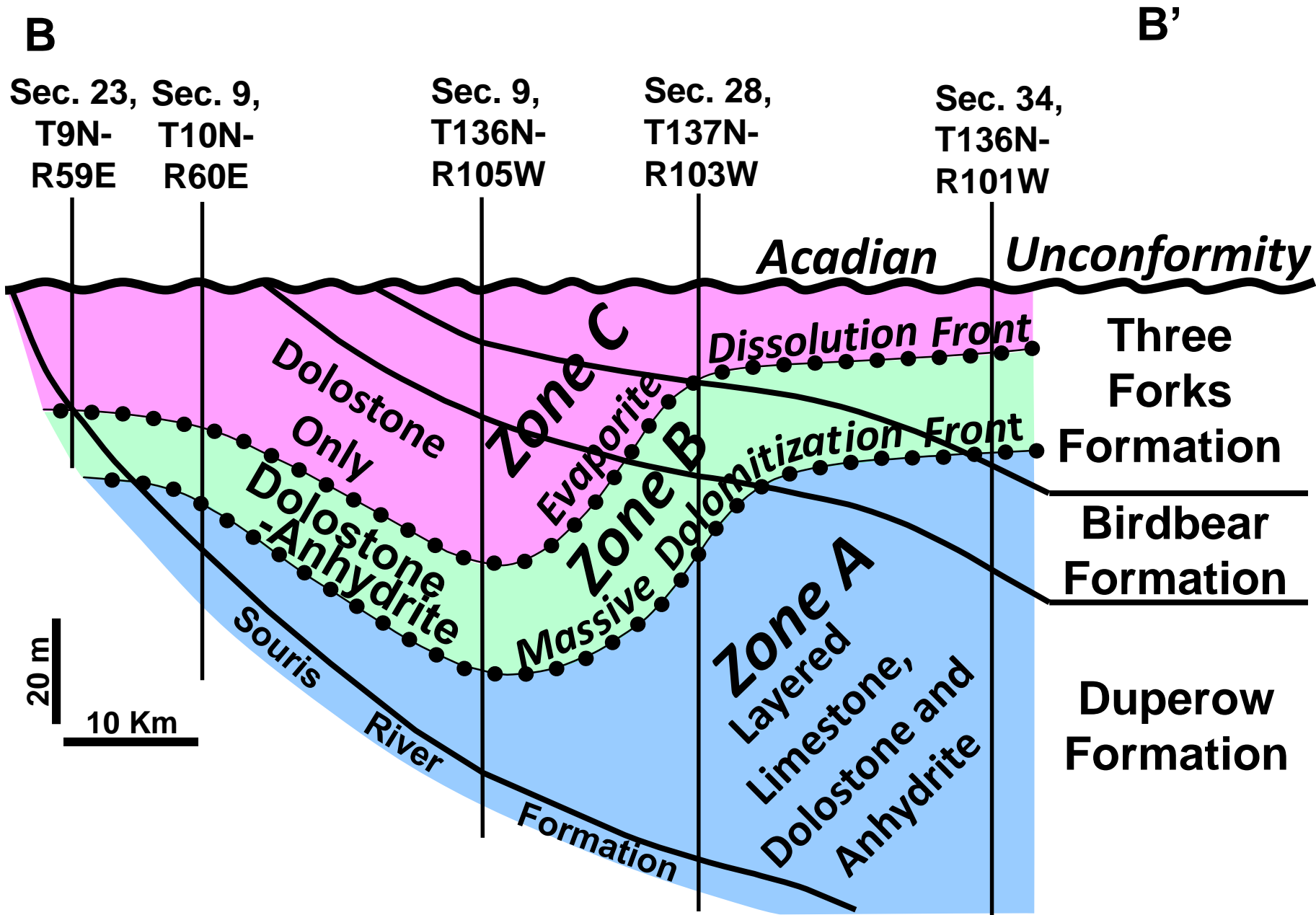
9634.5' core depth
Average crystal size = 80 μ
Thin Section porosity = 10%



Birdbear Diagenetic Zones



Next Slide shows
Diagenetic Zones
in Cross Section



Note Scale; fronts are sub-parallel to Acadian Unconformity

B

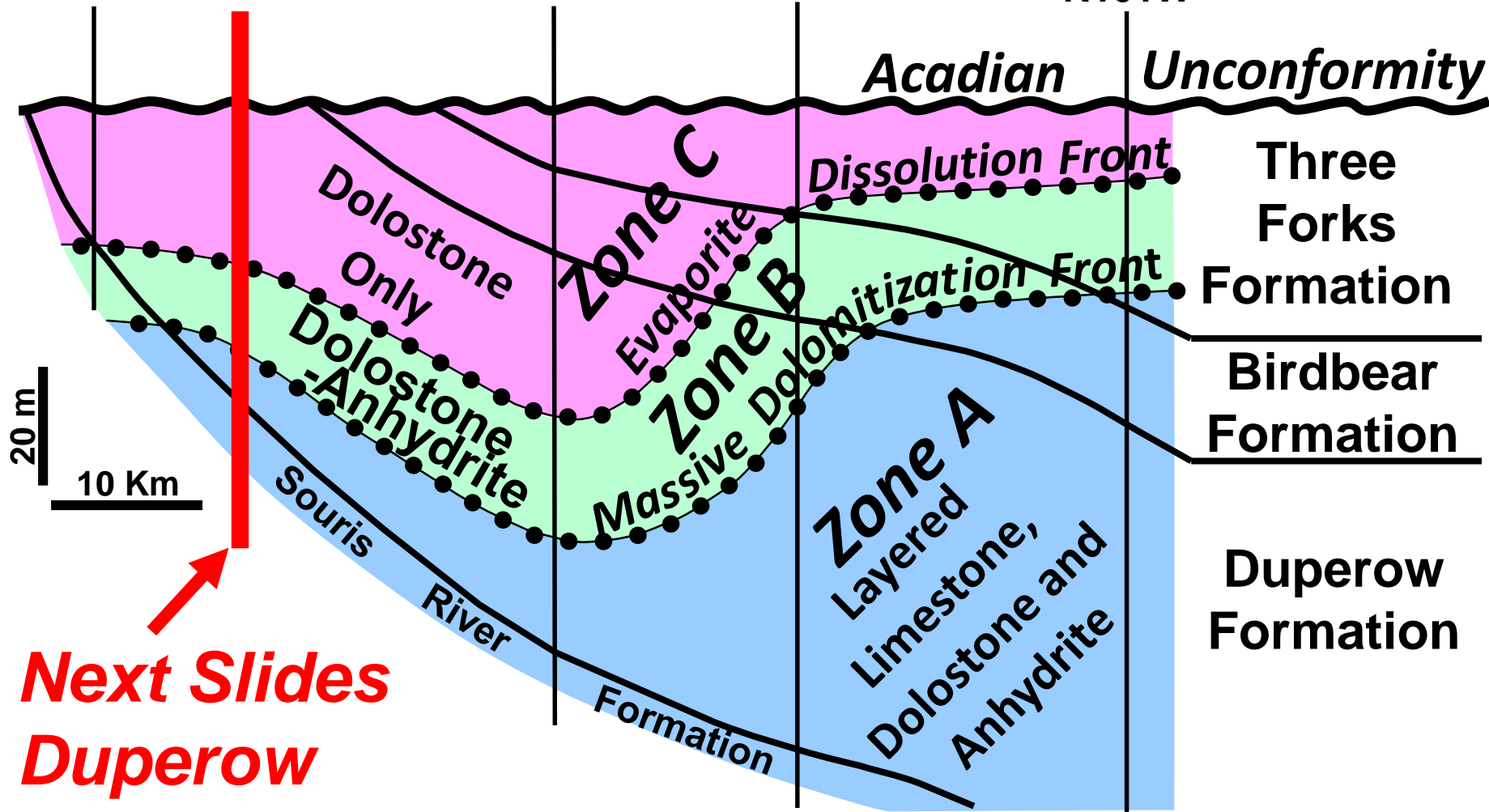
Sec. 23, Sec. 9,
T9N- T10N-
R59E R60E

Sec. 9,
T136N-
R105W

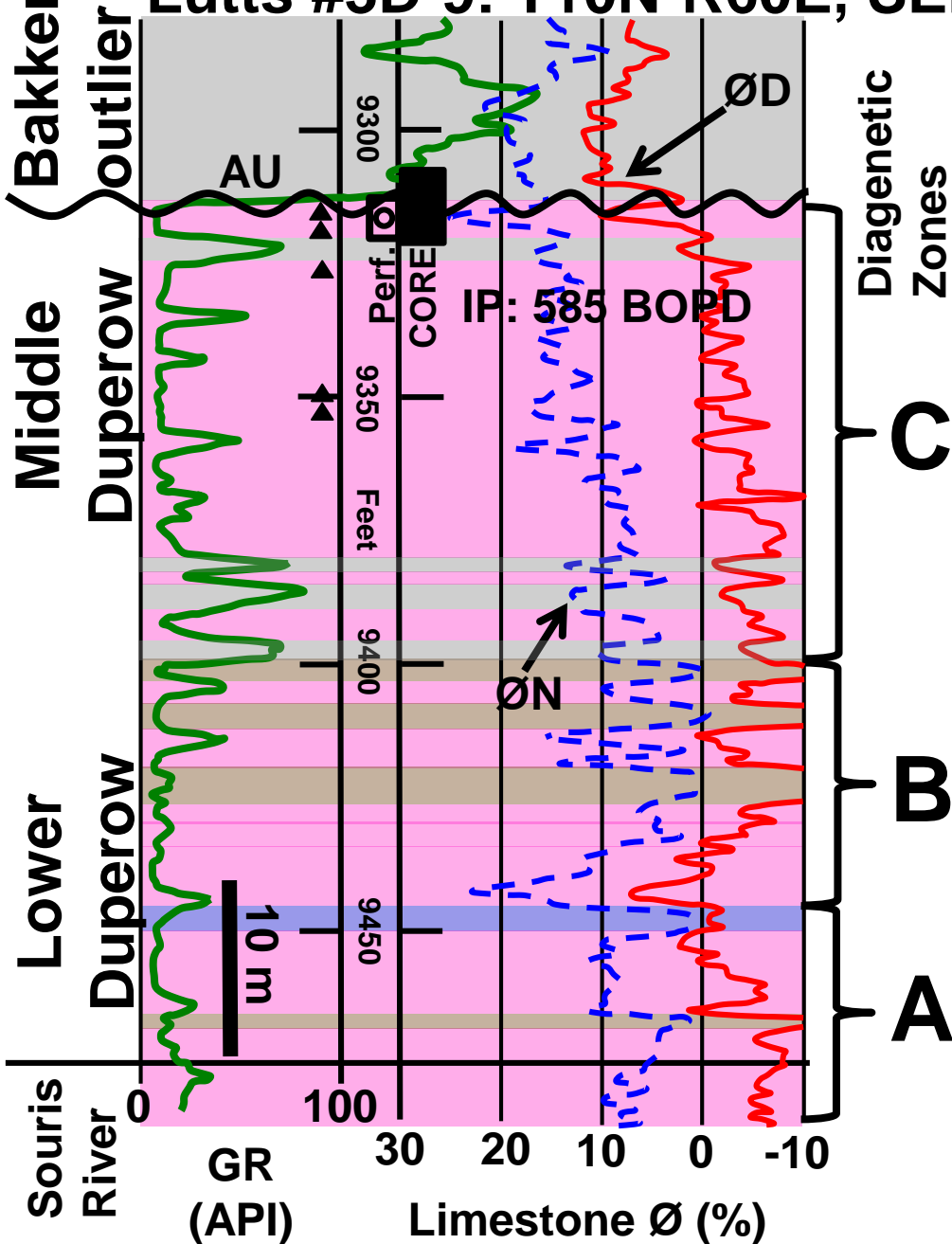
Sec. 28,
T137N-
R103W

Sec. 34,
T136N-
R101W

B'



Lutts #3D-9: T10N-R60E, SENE Sec. 9; Montana



Michels #1

#A1 Stark
SESE Sec. 15 T10N-R60E
#2-9 Lutts
NENE Sec. 9 T10N-R60E
Michels #1
SESW Sec. 6 T10N-R60E

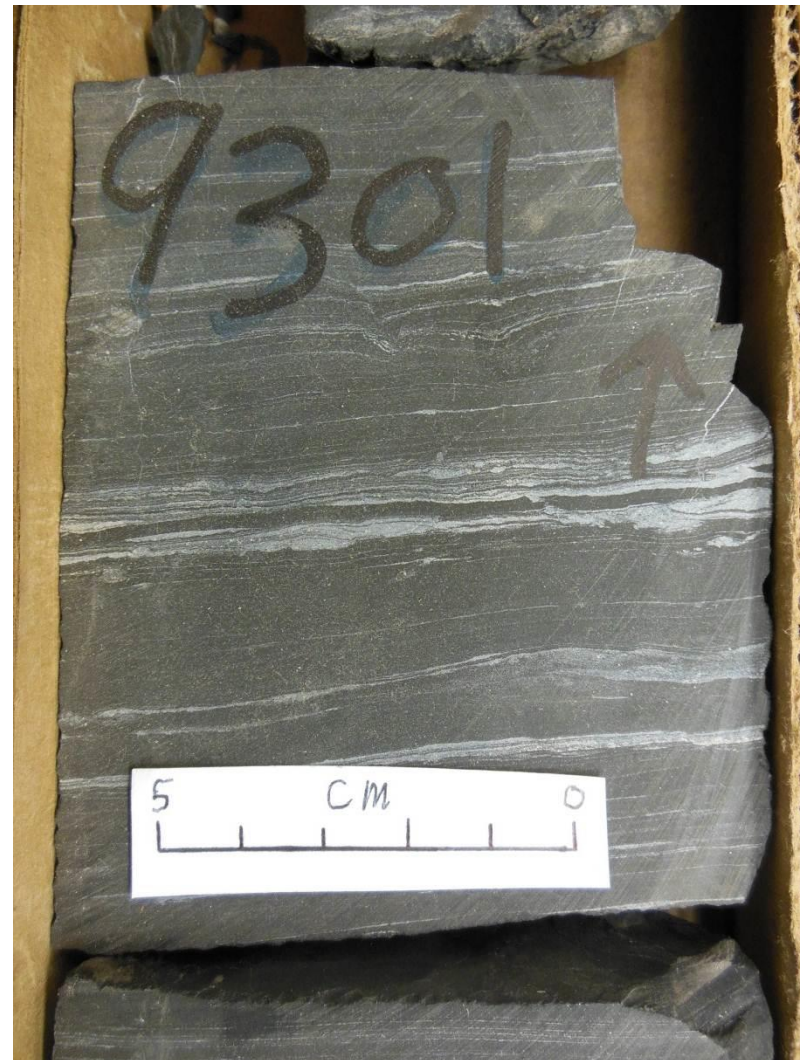
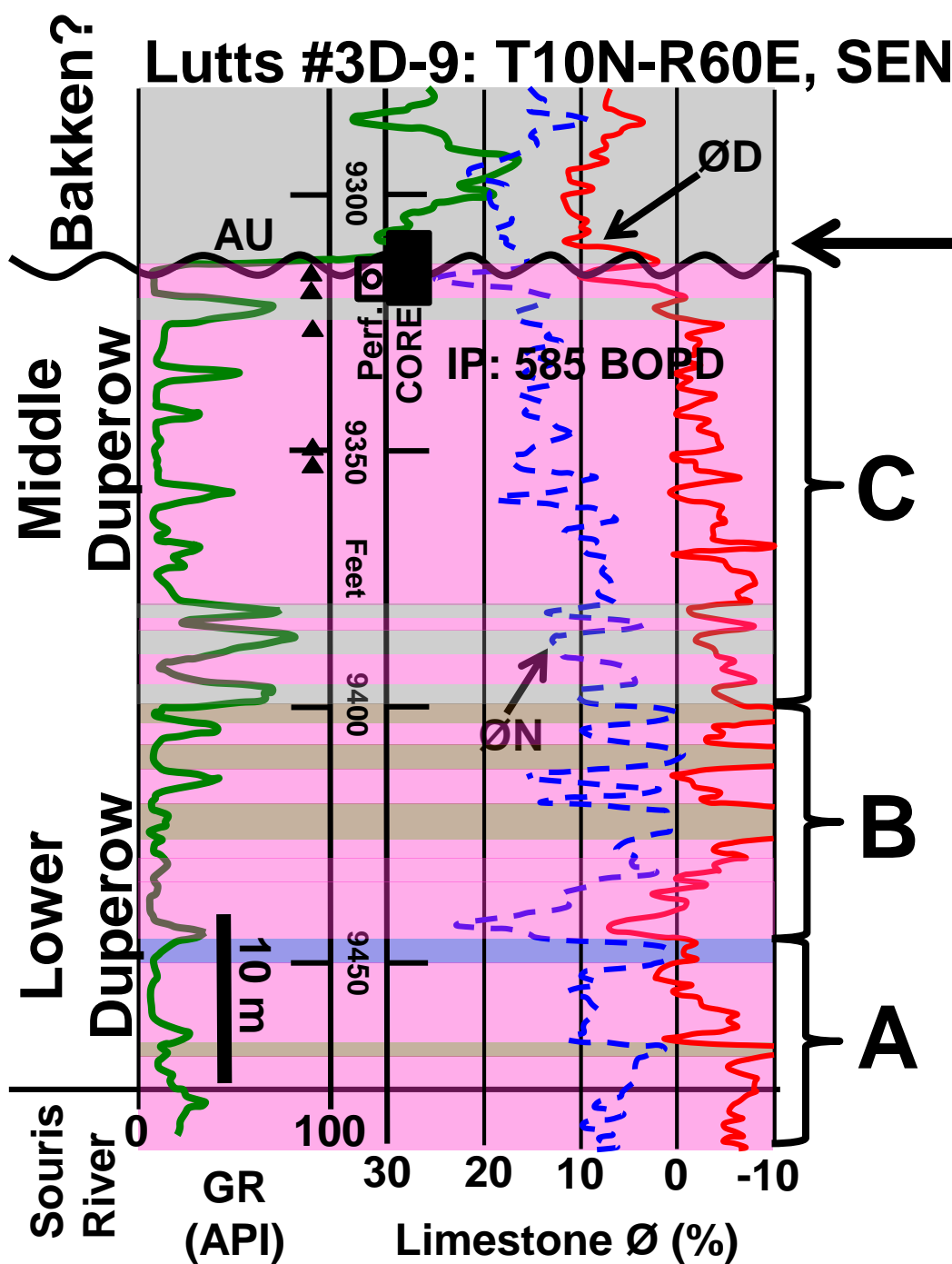
Stratigraphic position of cores in nearby wells

Basic Lithology

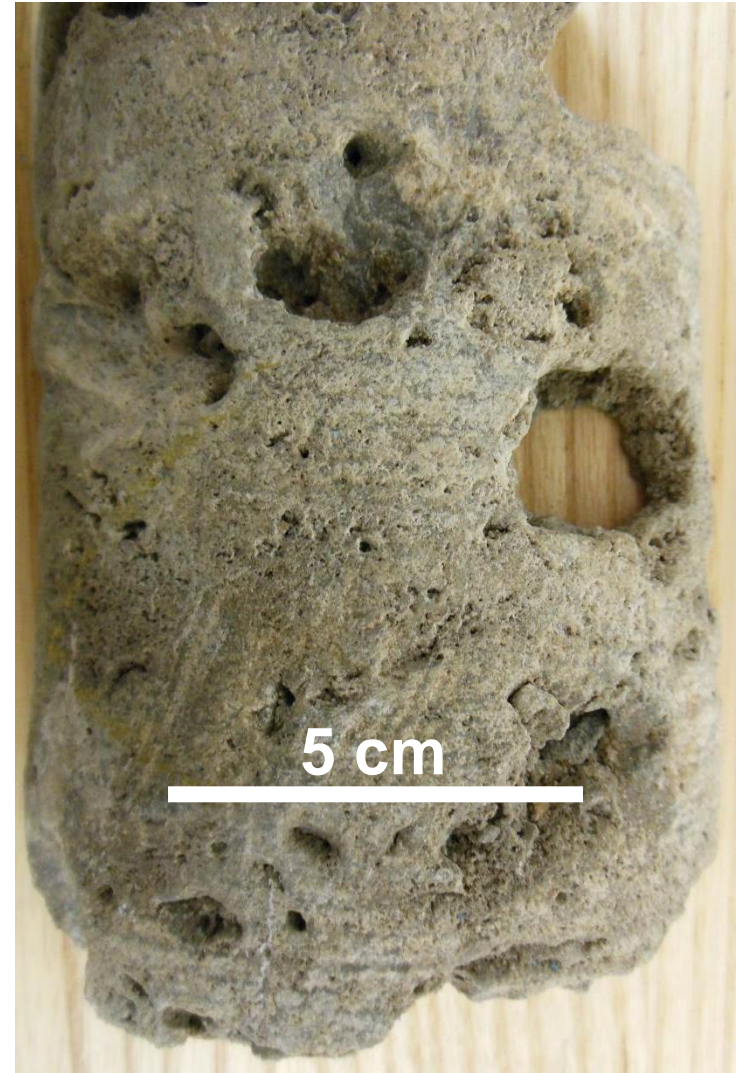
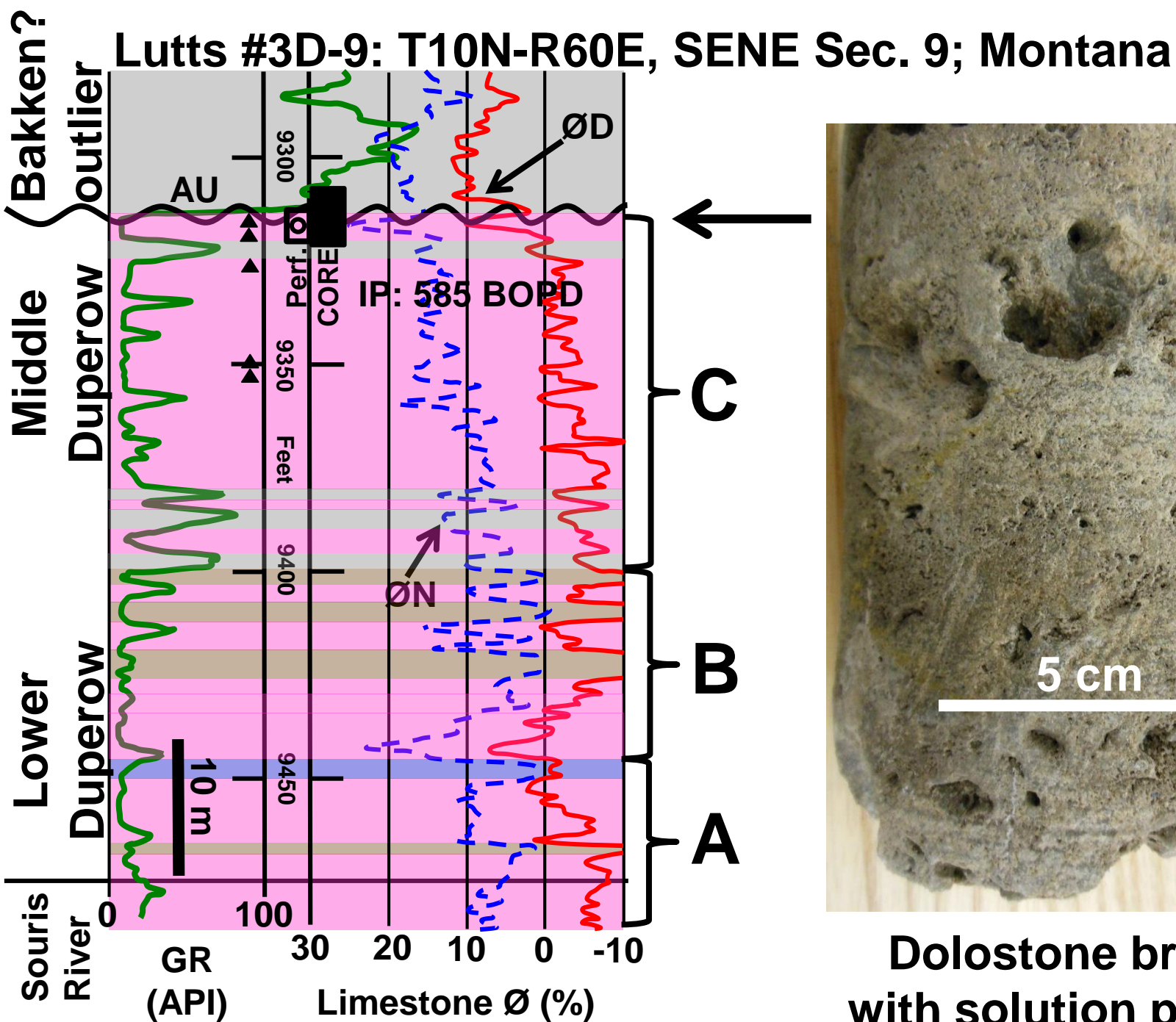
- Dominantly Dolostone
- Dominantly Limestone
- Dominantly Anhydrite
- Sand, Silt and/or Clay-Rich Lithology

▲ = Breccia in core

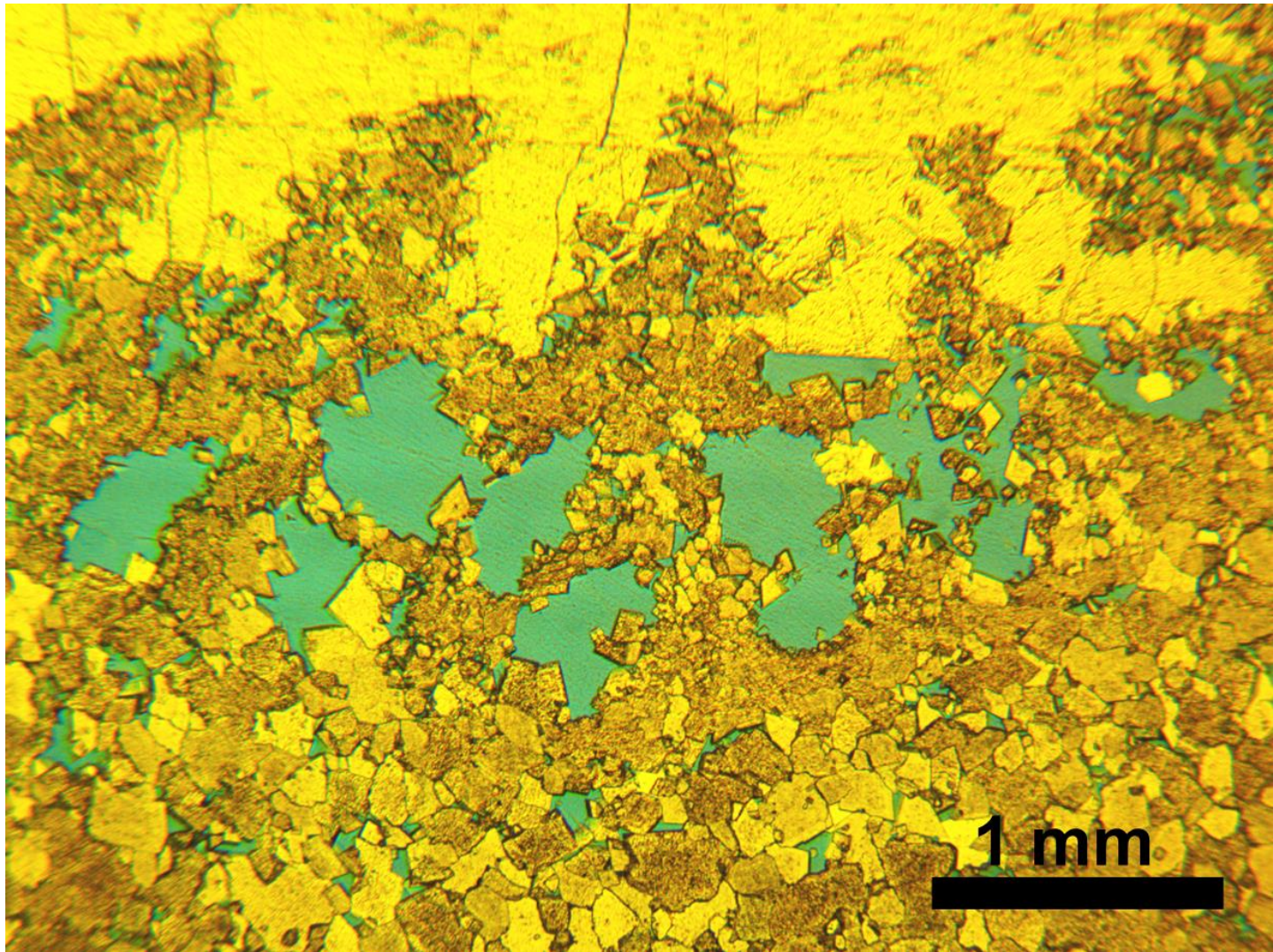
Lutts #3D-9: T10N-R60E, SENE Sec. 9; Montana



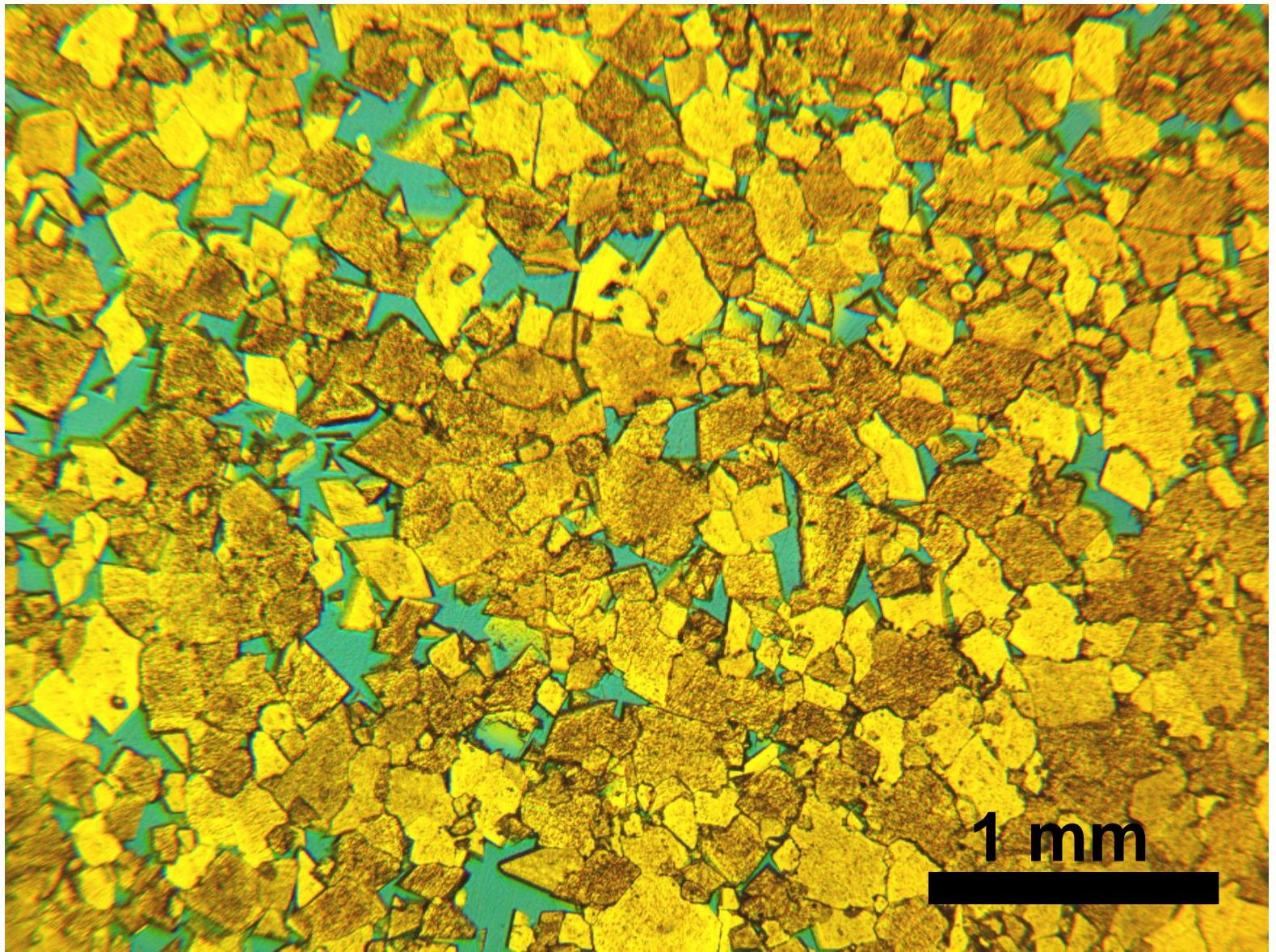
Laminated Shale
Upper Bakken (?) Outlier



**Dolostone breccia
with solution porosity**



**Lutts #2-9_9271.5; 1.0 foot below unconformity.
Coral replaced by fabric-destructive dolomite;
anhydrite fills coral center**



Lutts #2-9_9271.5_175 μ dolomite

Previous USA Studies

Massive Dolomitization

- Rarely discussed in literature
 - Loeffler (1982) suggested mixing zone model for Birdbear massive dolostone in North Dakota
 - Smith and Dorobek (1989) inferred massive dolomitization for Birdbear-Duperow in central Montana caused by Mg brines generated during deposition of overlying Three Forks evaporite
- *Model discussed here differs from all previous models*

Stratigraphic Constraints on Timing for Basin-Margin Diagenesis in the Birdbear and Duperow

1. Massive dolomitization occurred before Lodgepole deposition
 - No regional dolomitization in basal Lodgepole
2. Massive dolomitization and sub-karst dissolution occurred after most or all unconformity development:
 - *Massive dolostone and evaporite dissolution are sub-unconformity features;* they occur only in proximity to sub-unconformity areas
 - *Massive dolostone front and evaporite dissolution front are sub-parallel to the unconformity;* they are sub-parallel to existing (final) unconformity
 - *Top-down diagenesis below unconformity*

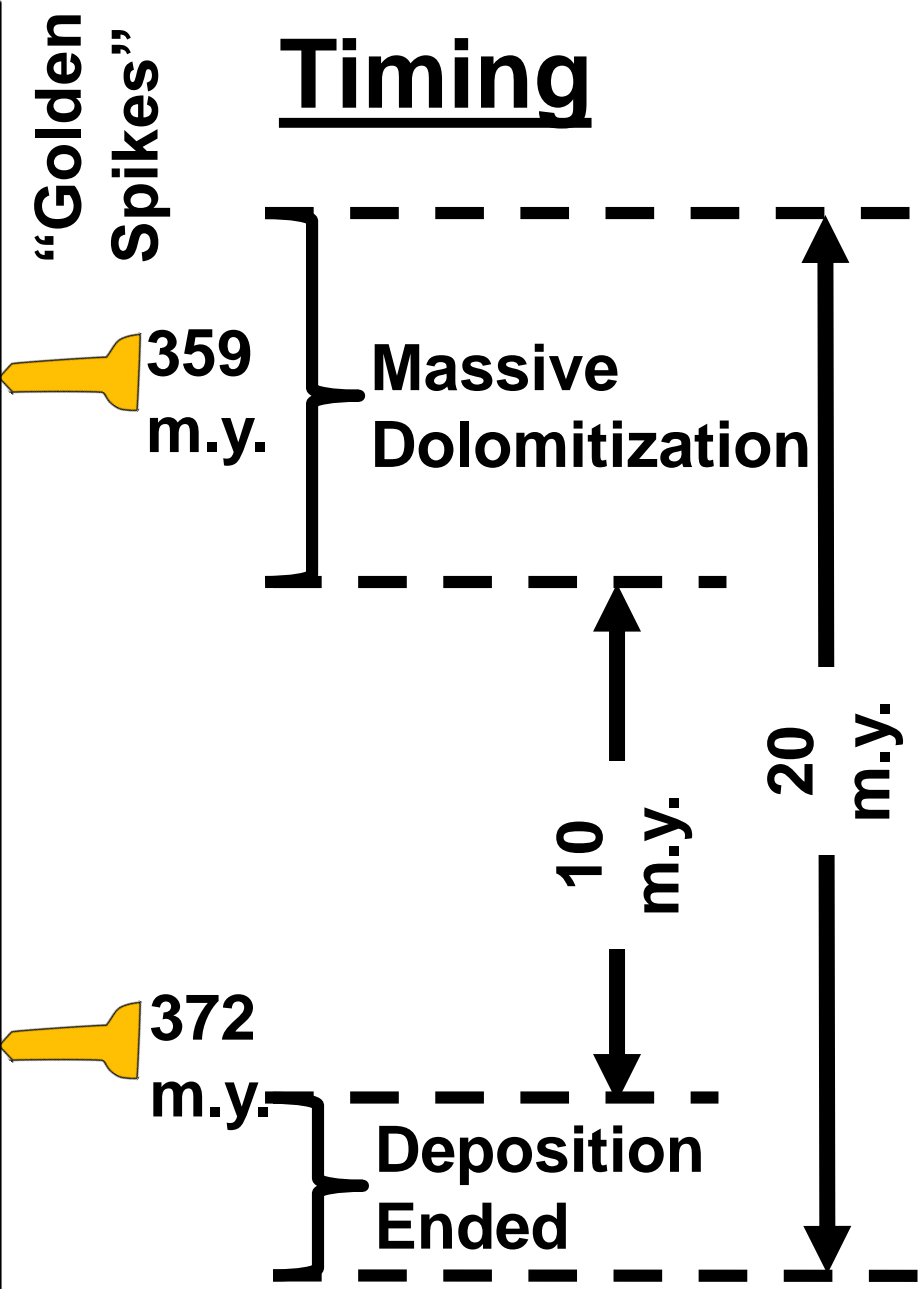
Additional Stratigraphic Observations:

- **L. Pronghorn sandy dolostone in southwest North Dakota is a detrital deposit derived from erosion on Cedar Creek anticline (Skinner et al., 2015)**
- **L. Pronghorn sandy dolostone mineralogy is very similar to upper Three Forks sandy dolostone mineralogy**

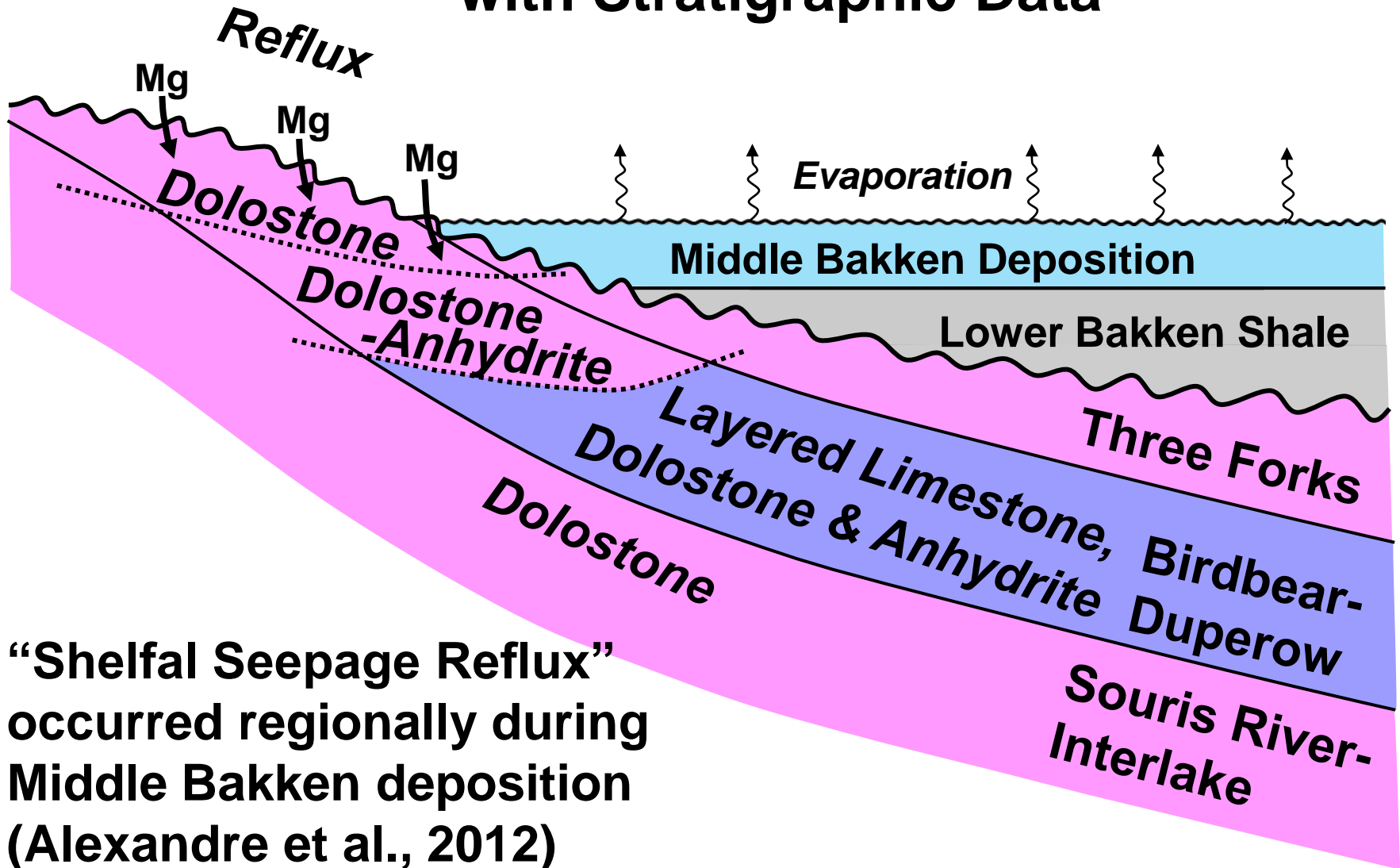
Stratigraphic Interpretations:

- **Lower Pronghorn deposition represents initial exposure of Cedar Creek anticline**
- **Most of the observed sub-unconformity Birdbear-Duperow diagenesis occurred after L. Pronghorn deposition (i.e., after removal of Upper Three Forks on Cedar Creek anticline)**
- **Most sub-unconformity diagenesis probably occurred during Bakken deposition**

CARBON.		Conodont Zones	Age Ma	Formation/Member	
DEVONIAN	Tournaisian	<i>U. crenulata</i>	354	Lodgepole	
		<i>L. crenulata</i>			
		<i>sandbergi</i>	356	Upper Bakken Shale	
		<i>duplicata</i>			
		<i>sulcata</i>	358	Middle Bakken	
	Famennian	<i>praesulcata</i>	360	Lower Bakken Shale	
		<i>expansa</i>			
		<i>postera</i>	362	? Prong- ? horn	
		<i>trachytera</i>	364		
		<i>marginifera</i>	366		
		<i>rhomboidea</i>	368	Three Forks	
		<i>crepida</i>			
		<i>triangularis</i>	370	No Data	
	Frasnian	<i>linguiformis</i>	372		
		<i>rhenana</i>	374	Birdbear	
			376	Duperow	



Dolomitization During Middle Bakken Deposition is Consistent with Stratigraphic Data



“Shelfal Seepage Reflux”
occurred regionally during
Middle Bakken deposition
(Alexandre et al., 2012)

Conclusions

1. Regional evaporite dissolution formed below an “evaporite karst” that developed during development of the Acadian Unconformity

- **Mappable**
- **Dissolution up to 50 m below unconformity**

Conclusions

2. Stratigraphic constraints indicate that massive dolostone in the Birdbear and Duperow formations formed as a result of one or more diagenetic events that occurred after unconformity development and before Lodgepole deposition

- Dolomitization front sub-parallel to unconformity**
- Massive dolostone up to 70 m below unconformity**
- Dolomitization occurred during Bakken deposition**
- Speculative conclusion that massive dolomitization occurred during Middle Bakken dolomitization**
- Dolomitization 10-20 million years after deposition is responsible for “massive” characteristics**

END

Acknowledgements:

- **Cores were examined at the USGS and NDGS**
- **Thin sections collected from these studies are on file at these organizations**
- **NDIC web data base indispensable**