3-D Surface-Wave Estimation and Separation Using an Iterative Closed-Loop Approach*

Tomohide Ishiyama¹ and Gerrit Blacquiere²

Search and Discovery Article #41898 (2016)**
Posted September 26, 2016

*Adapted from oral presentation given at AAPG GEO-2016, 12th Middle East Geosciences Conference & Exhibition, Manama, Bahrain, March 7-10, 2016

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Abstract

Surface waves in seismic data are often dominant and mask primaries in land or shallow-water environments. Separating them from the primaries is of great importance either for removing them as noise for reservoir imaging and characterization, or for considering them as signal for near-surface characterization. However, their complex properties, such as dispersion, multi-modality and spatial variability, make the surface-wave separation significantly challenging in processing. To address the challenges, we introduced a method of 3-D surface-wave estimation and separation using an iterative closed-loop approach.

The closed loop contains a relatively simple forward model of surface waves and adaptive subtraction of the forward-modelled surface waves from the observed surface waves, making it possible to evaluate the residual between them. In this approach, the surface-wave model is parameterized by the frequency-dependent slowness and source properties for each surface-wave mode. The optimal model parameters are estimated in an iterative way such that the residual is minimized and, consequently, the approach solves the inverse problem.

We applied this method to several data sets to demonstrate its virtues, such as real 3D geophone/hydrophone seismic data onshore/offshore Abu Dhabi where ground-roll/mud-roll is significantly dominant in land/shallow-water environments. Through the examples, we observed that the method successfully estimates and separates out the surface waves from the seismic data to consequently obtain the subsurface signals. The method provides a better result than a conventional slowness/velocity-based filtering method which cannot handle both surface waves and subsurface signals overlapping each other. We also observed its wide range of applicability to under-sampled, asymmetrically sampled, irregularly sampled and blended seismic data. This suggests the possibility of relaxing requirements for seismic survey parameters in terms of surface-wave separation and, therefore, offers flexibility as well as potential effort reduction with respect to seismic surveys. It should be noted that recent advances in acquisition, such as point receivers and a large amount of stations, make the method more effective because of the improved spatial sampling of surface waves without negative array effects.

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References Cited

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Exhibition: 8 – 10 March 2016

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3-D Surface-wave estimation and separation using an iterative closed-loop approach

Tomohide Ishiyama, Inpex Corporation and Gerrit Blacquiere, Delft University of Technology















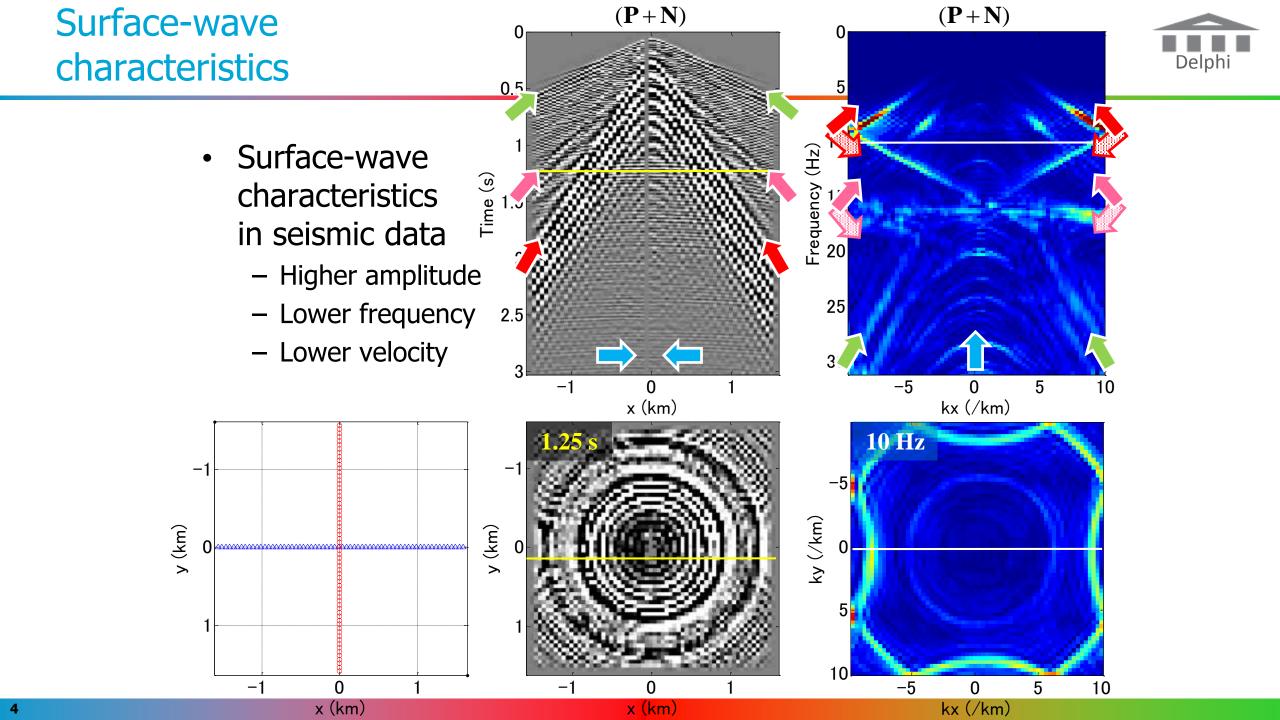
Outline

- Introduction
- Theory and method
- Real data examples
- Conclusions and remarks



Outline

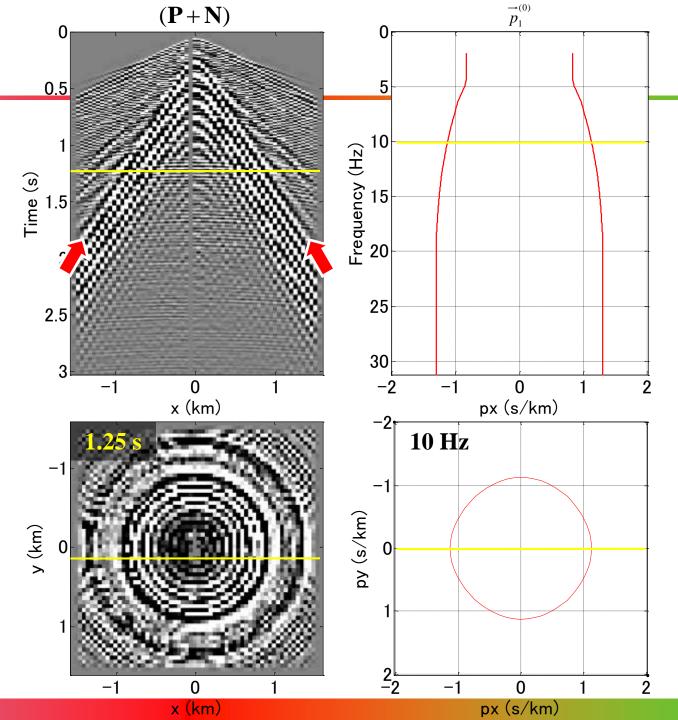
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Dispersive

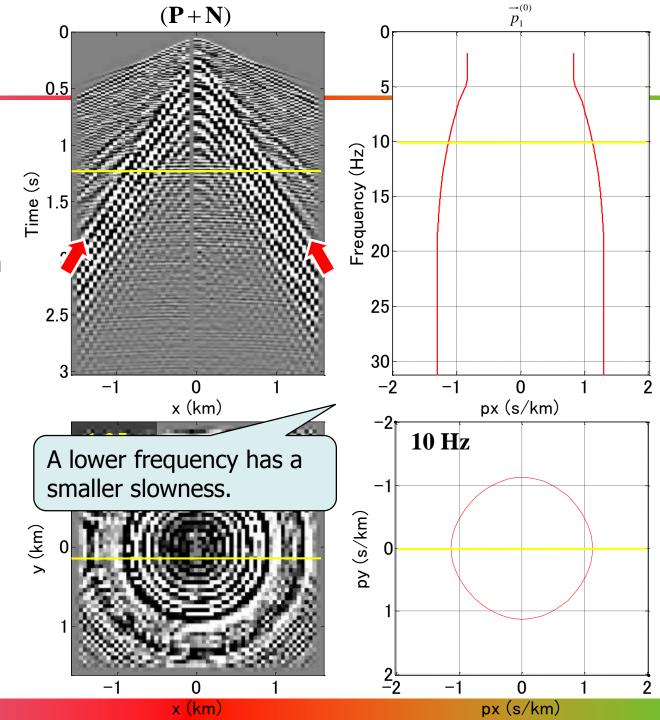
Different
 frequencies
 propagate with
 different
 velocities.





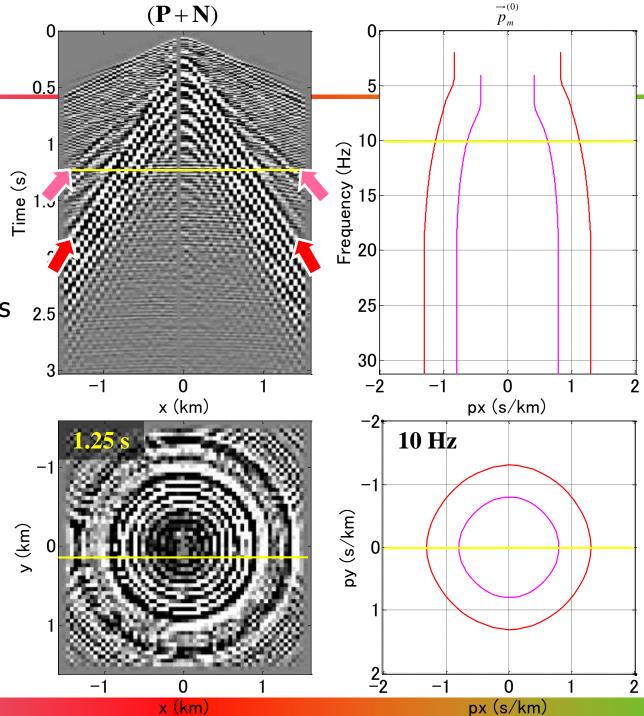
Dispersive

Different
 frequencies
 propagate with
 different
 velocities.





- Dispersive
- Multi-modal
 - Each frequency propagates with several velocities simultaneously.

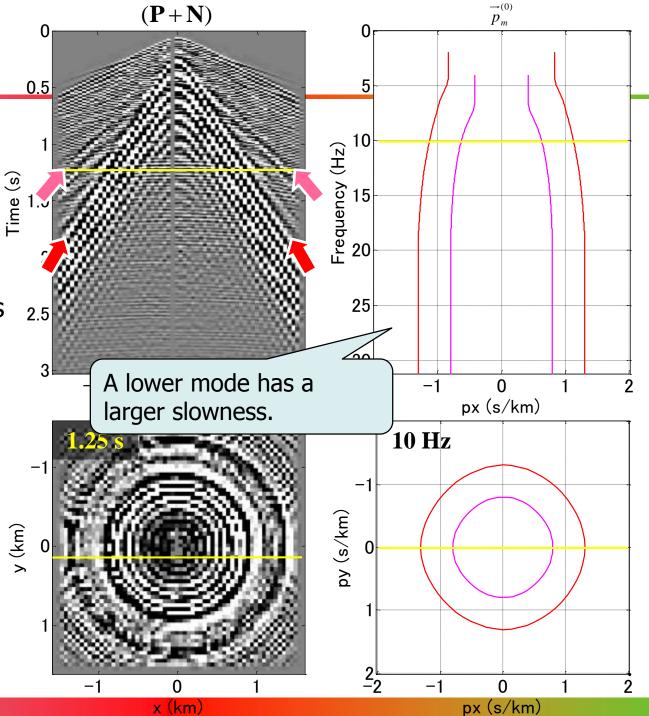




Dispersive

Multi-modal

Each frequency propagates with several velocities simultaneously.



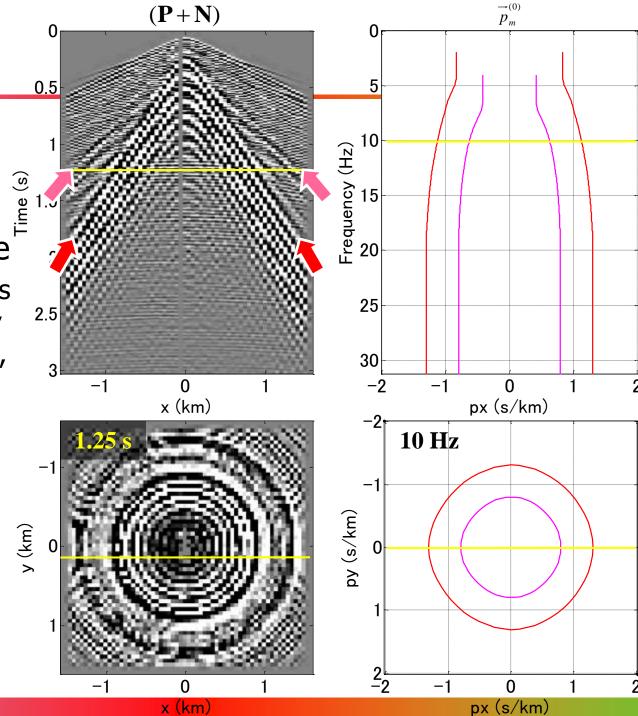


Dispersive

Multi-modal

Spatially variable

 These properties change spatially in a shot record, and one shot to another.

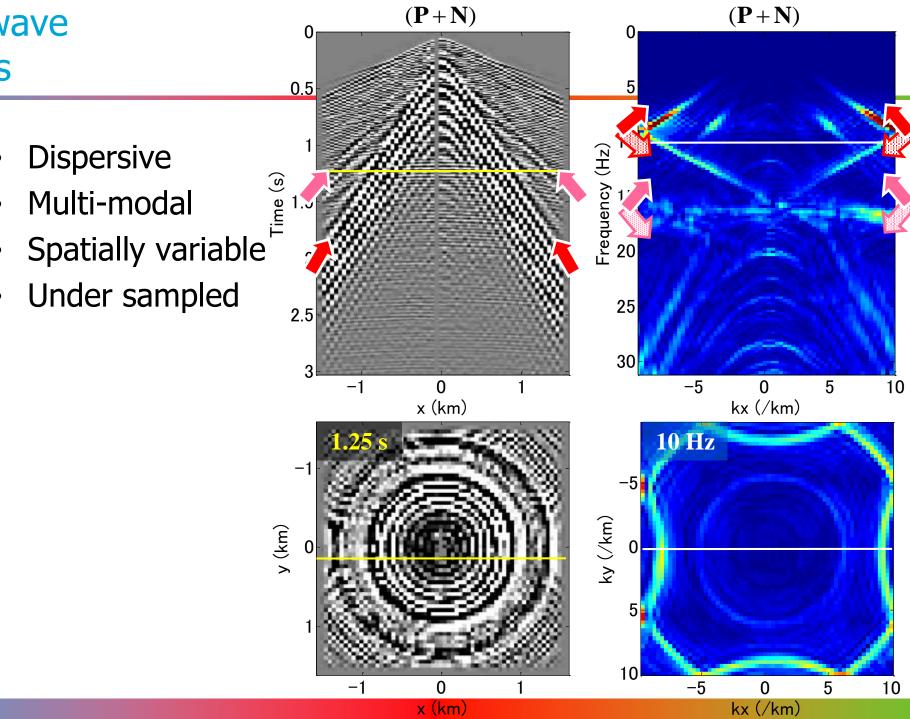




 $(\mathbf{P} + \mathbf{N})$ Delphi

- Dispersive

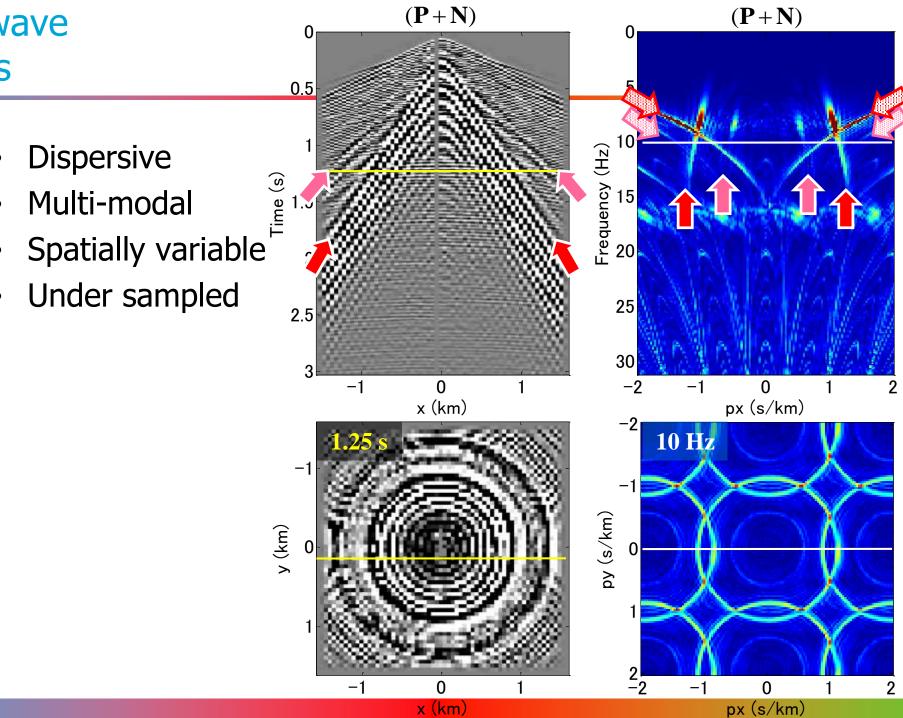
- Under sampled







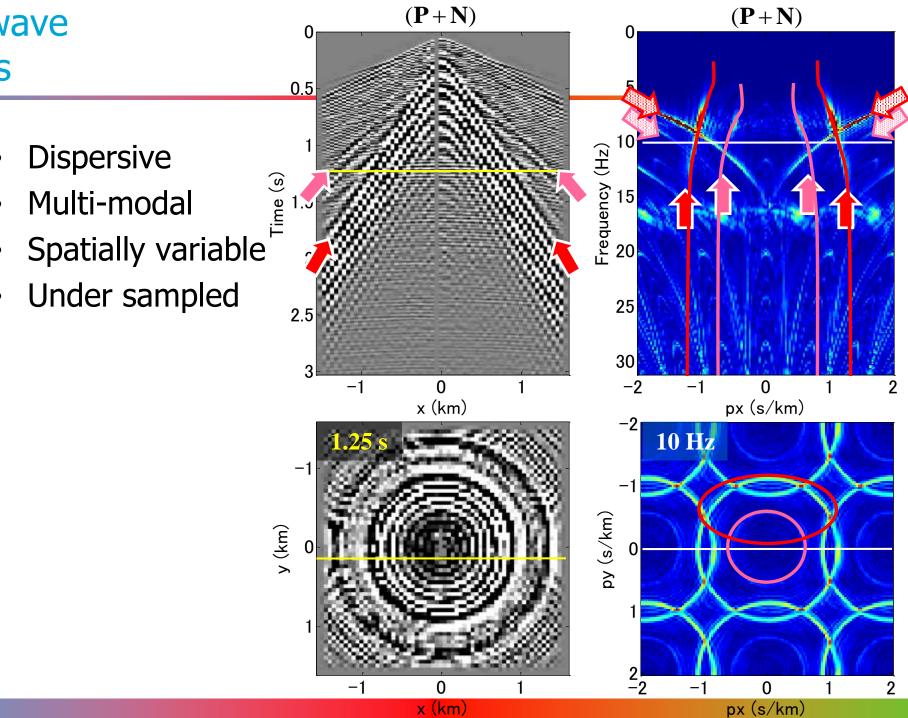
- Dispersive



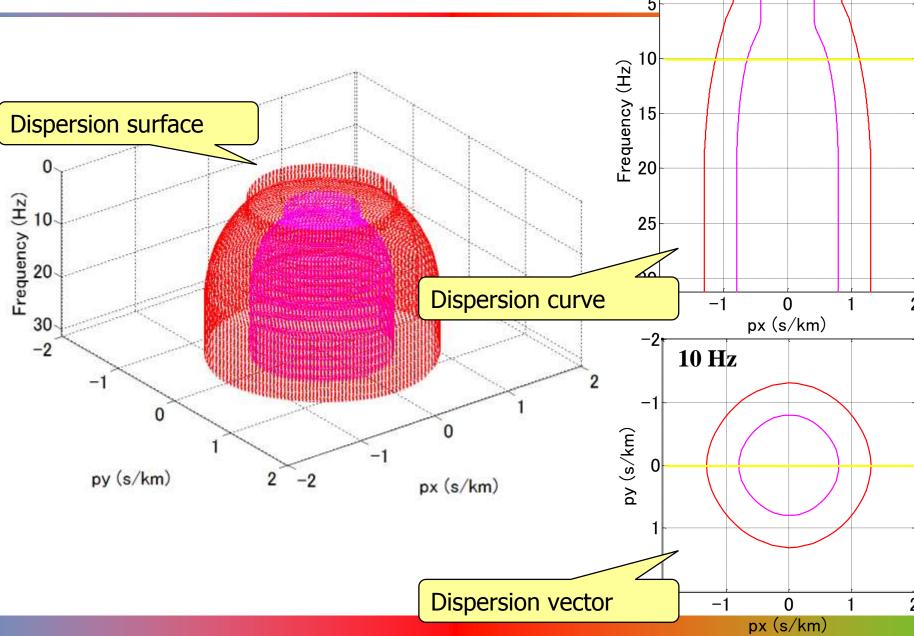


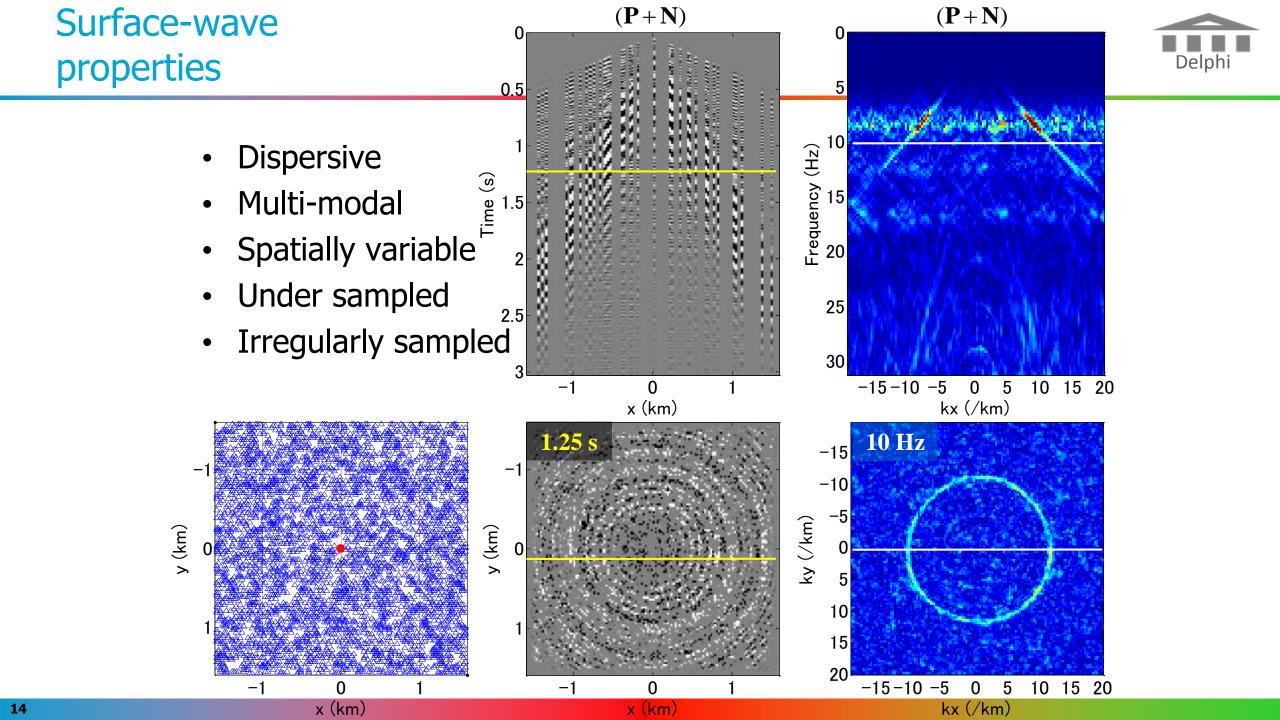


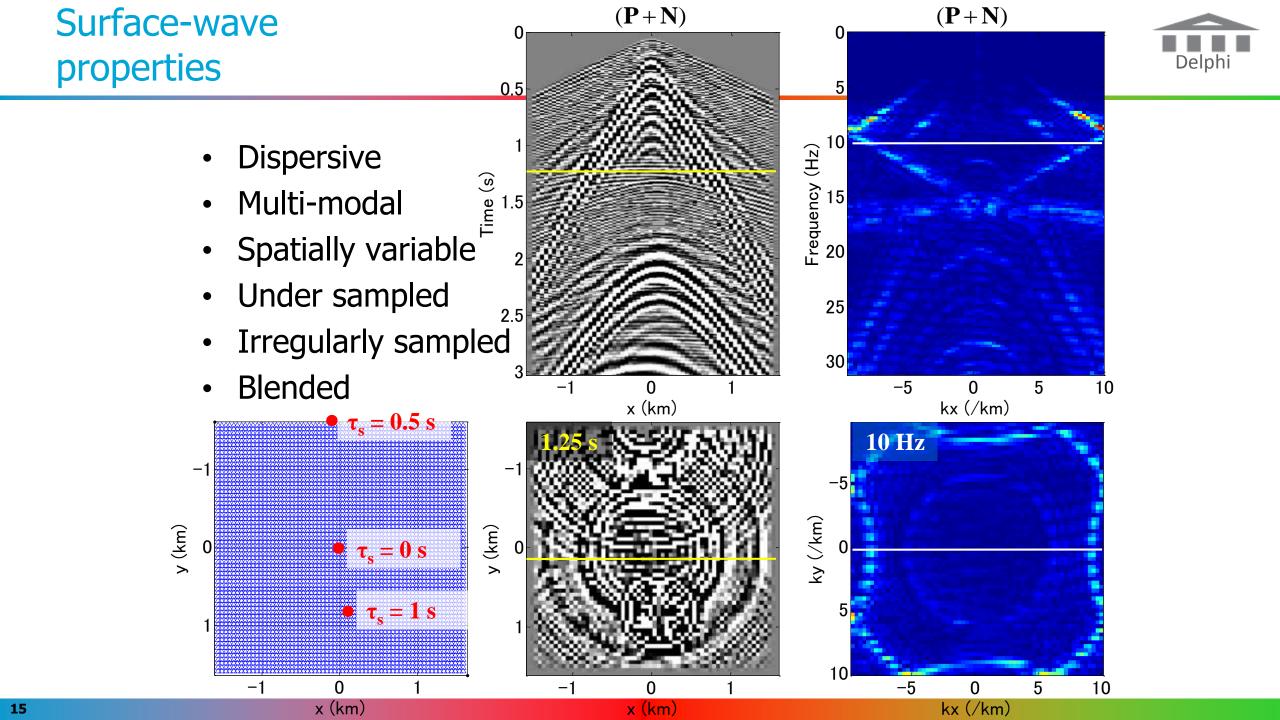
- Dispersive











Surface-wave separation



Surface waves

- traditionally treated as noise, masking primaries.
- today regarded as signal, for near-surface characterization.

Estimating them and separating them out from seismic data is important for both applications.

$$\mathbf{P}_{tot} = \mathbf{P} + \mathbf{N}$$

N: Surface waves

P : Subsurface signals

(P+N) : Seismic data

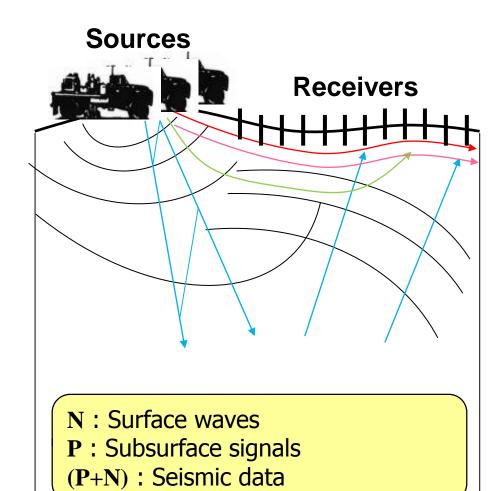


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Forward model



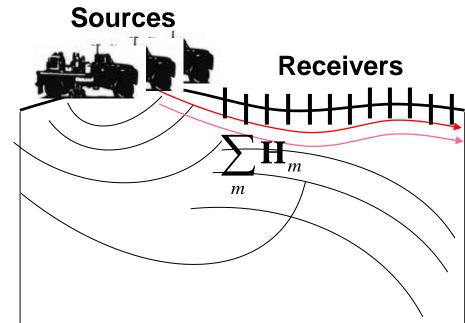


$$\mathbf{P}_{tot} = \mathbf{P} + \mathbf{N}$$

Forward model







S: Source properties

H: Horizontal propagation operator

X: Earth propagation operator

$$\mathbf{N} = \sum_{m} \mathbf{N}_{m}$$

$$\mathbf{N}_m = \mathbf{H}_m \mathbf{S}_m$$

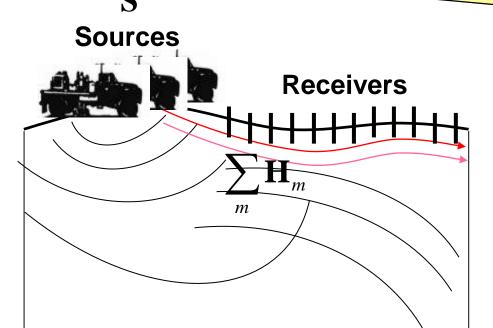
$$H_{m}(\vec{r},\omega) = \left(\frac{e^{-\vec{\lambda}_{m}(\omega)\cdot\vec{r}}}{\sqrt{|\vec{r}|}}\right)e^{-j\omega\vec{p}_{m}(\omega)\cdot\vec{r}}$$

Forward m

Multi-modes



m : the number of a surface-wave mode.



S: Source properties

H: Horizontal propagation operator

X: Earth propagation operator

$$\mathbf{N} = \sum_{m} \mathbf{N}_{m}$$

N_m dispersion vector

$$H_{m}(\vec{r},\omega) = (\frac{e^{-\vec{\lambda}_{m}(\omega)\cdot\vec{r}}}{\sqrt{|\vec{r}|}})e^{-j\omega\vec{p}_{m}(\omega)\cdot\vec{r}}$$

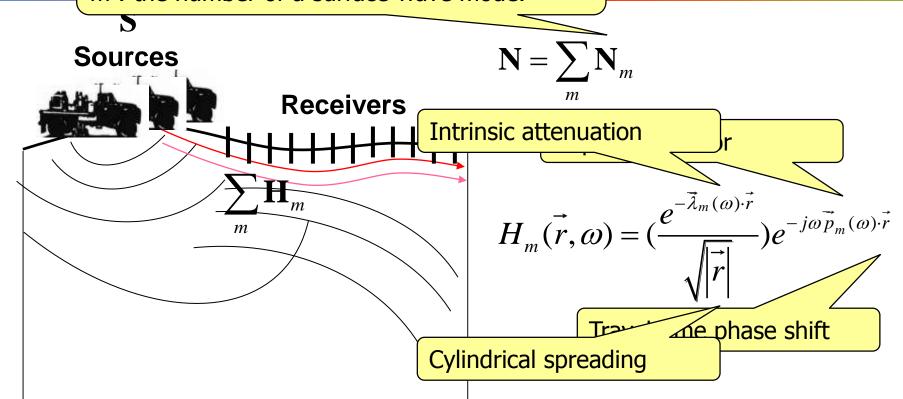
Travel-time phase shift

Forward m

Multi-modes



m : the number of a surface-wave mode.



S: Source properties

H: Horizontal propagation operator

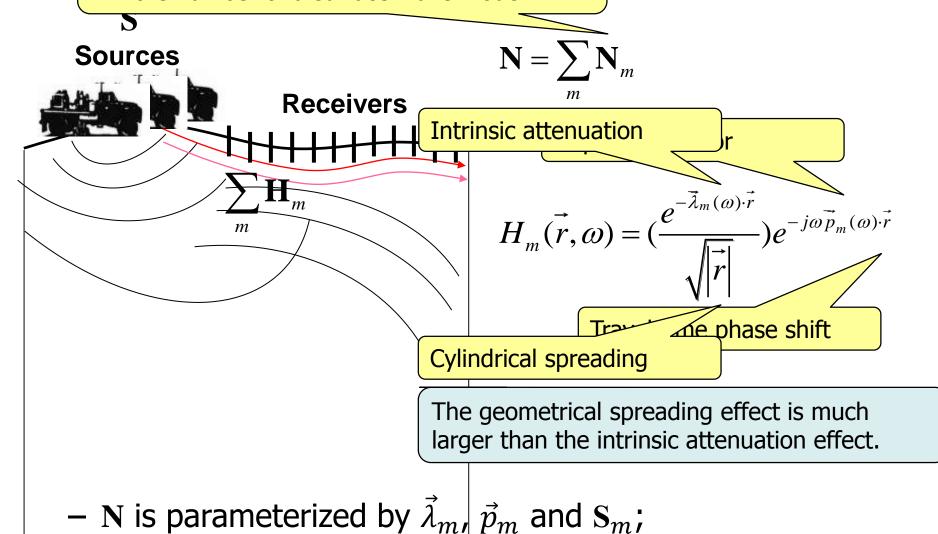
X: Earth propagation operator

Forward n

Multi-modes



m : the number of a surface-wave mode.



or \vec{p}_m and S_m with the assumption that $\vec{\lambda}_m$ is zero.

22

Inversion



$$\mathbf{J}: \mathbf{A}_m; \mathbf{H}_m, \mathbf{S}_m \xrightarrow{\mathbf{S}.t.} <\mathbf{P}>$$

$$\mathbf{J} = \sum_{\omega} \|\mathbf{P} + \Delta \mathbf{N}\|^{2}$$
$$= \sum_{\omega} \|(\mathbf{P} + \mathbf{N}) - \langle \mathbf{N} \rangle\|^{2}$$

$$= \sum_{\omega} \left\| (\mathbf{P} + \mathbf{N}) - \sum_{m} \mathbf{A}_{m} \mathbf{H}_{m} \mathbf{S}_{m} \right\|^{2}$$

<N> : Surface-wave estimate

<P>: Subsurface-signal estimate

- The parameters as well as the adaptive filter are solved such that the residual $P + \Delta N$ is minimized.
- The minimization scheme works on < N > only, and P remains untouched by a signal-protecting scheme.
- In the ideal situation, the resulting residual P + ΔN closely corresponds to P.

Inversion



$$\mathbf{J}: \mathbf{A}_m; \mathbf{H}_m, \mathbf{S}_m \xrightarrow{\mathbf{S.t.}} <\mathbf{P}>$$

$$\mathbf{J} = \sum_{\omega} \|\mathbf{P} + \Delta \mathbf{N}\|^{2}$$

$$= \sum_{\omega} \|(\mathbf{P} + \mathbf{N}) - \langle \mathbf{N} \rangle\|^{2}$$
A. Adaptive for

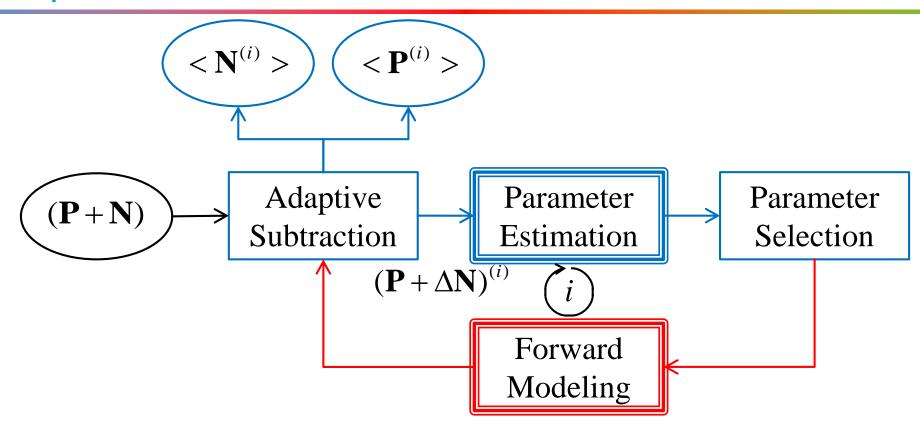
 $= \sum_{\omega} \| (\mathbf{P} + \mathbf{N}) - \langle \mathbf{N} \rangle \|^{2}$ minimized. $= \sum_{\omega} \| (\mathbf{P} + \mathbf{N}) - \sum_{m} \mathbf{A}_{m} \mathbf{H}_{m} \mathbf{S}_{m} \|^{2}$ The minimization scheme works on $\langle \mathbf{N} \rangle$ only, and \mathbf{R} remains untouched by a

 The parameters as well as the adaptive filter are solved such that the residual $P + \Delta N$ is

P remains untouched by a Parameters rotecting scheme.

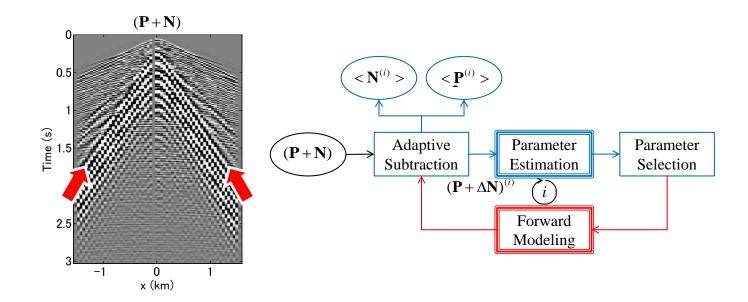
> In the ideal situation, the resulting residual $P + \Delta N$ closely corresponds to P.



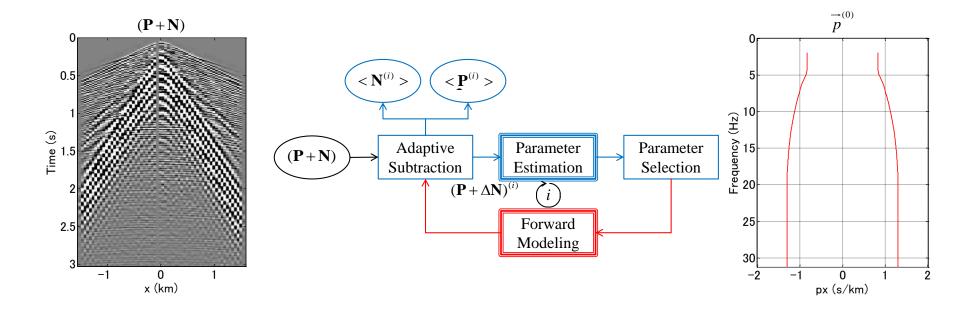


– For each loop, the modal parameters $\vec{p}_m^{(i)}$, $\Delta S_m^{(i)}$ as well as the modal adaptive filter $A_m^{(i)}$ (m=1,2,...) are estimated.

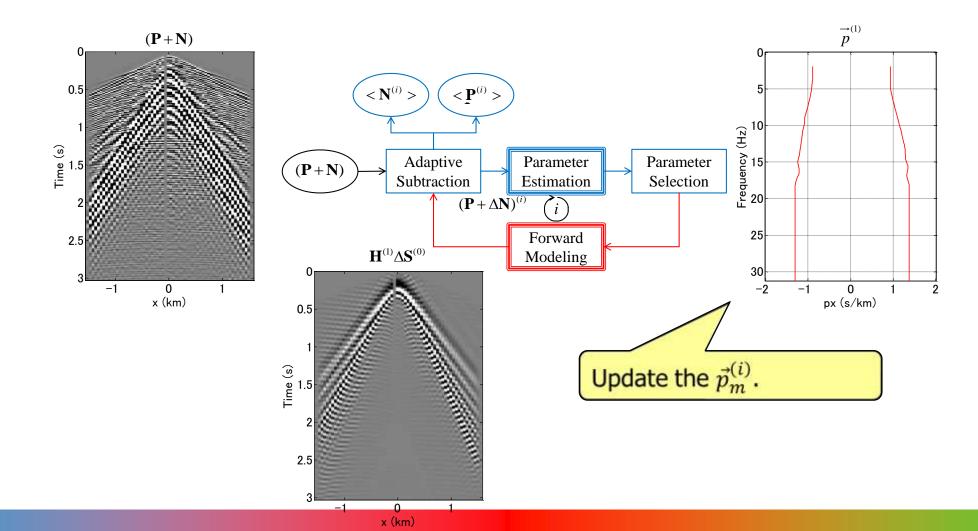




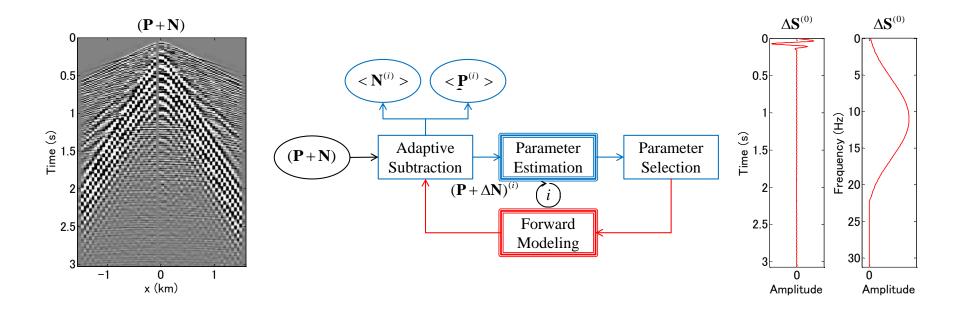




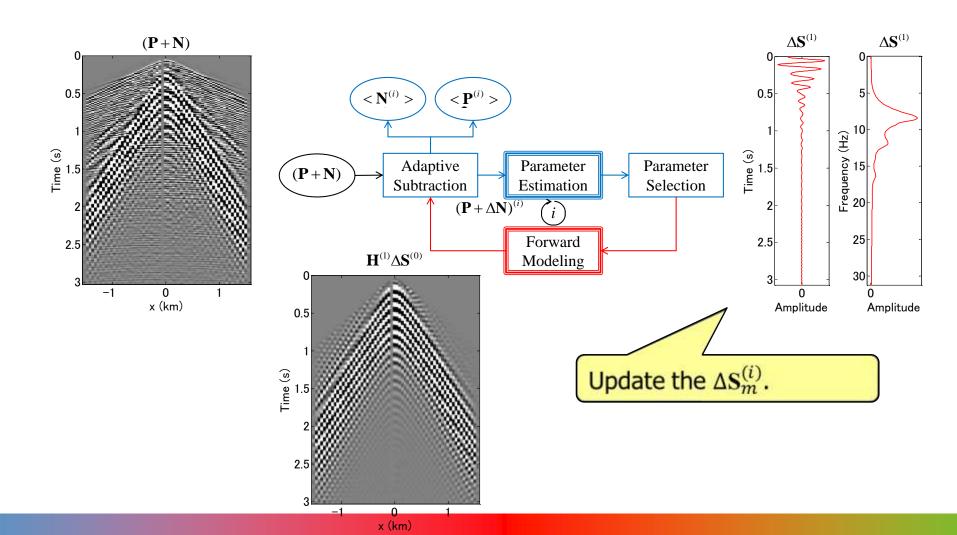




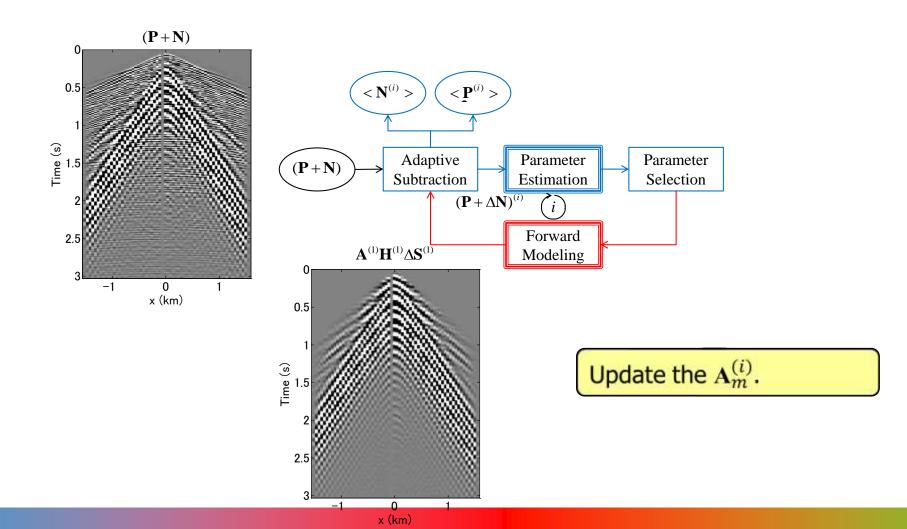




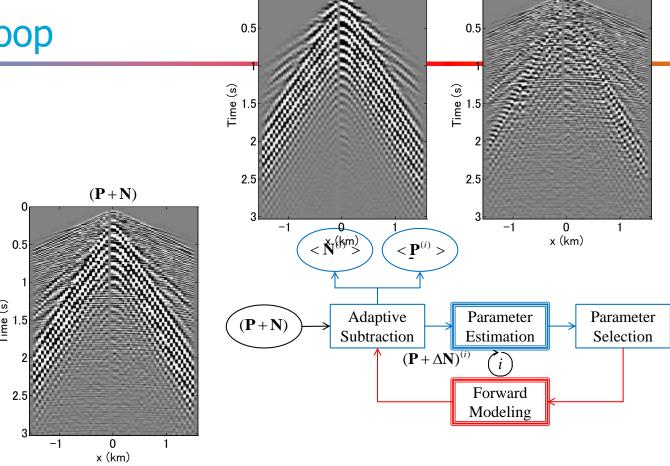








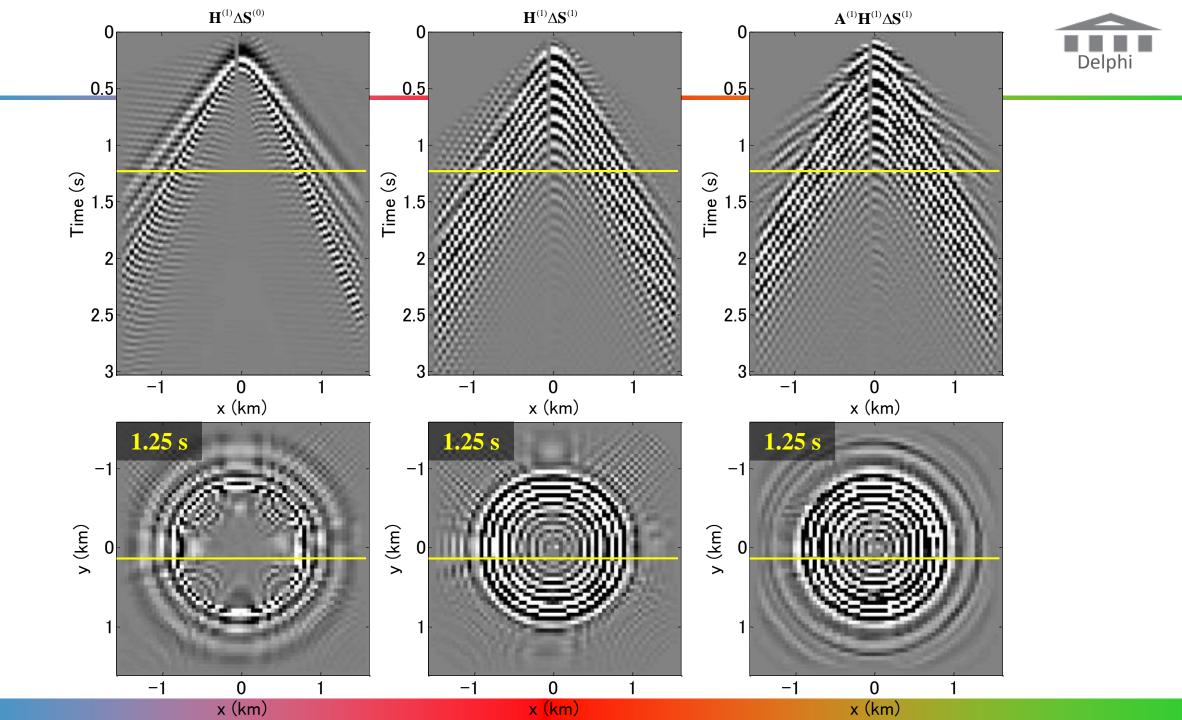


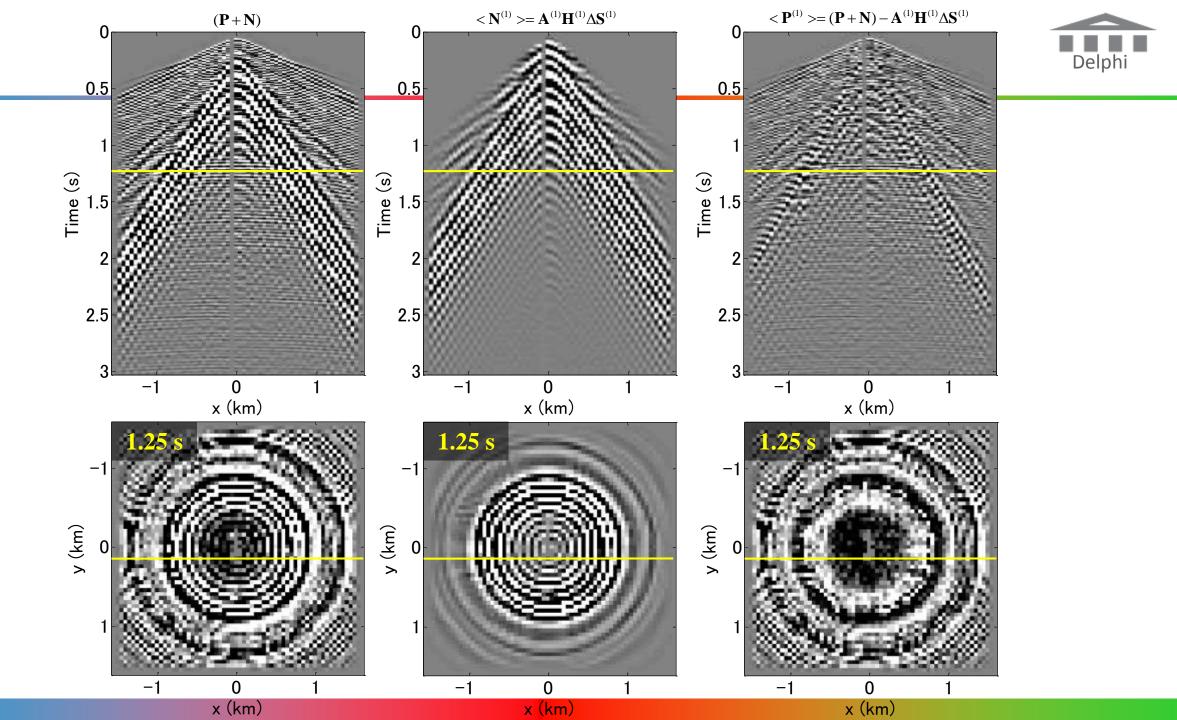


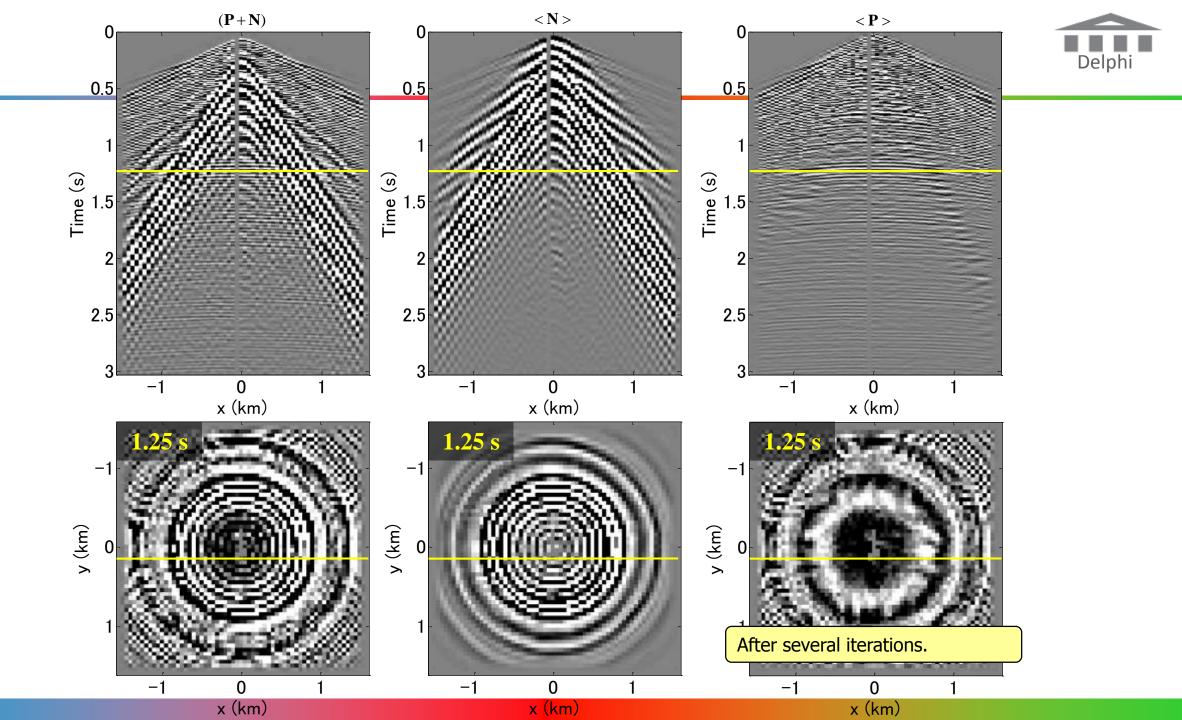
 $<\boldsymbol{N}^{(1)}>$

Go to the next loop, etc., etc...

 $< \mathbf{P}^{(1)} > = (\mathbf{P} + \Delta \mathbf{N})^{(2)}$









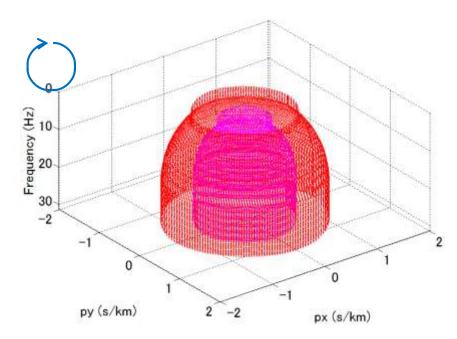


- Slowness/velocity-based filtering methods (e.g. Yilmaz, 2001)
- Data-driven, data-adaptive and model-based method (AROGA; Le Meur et al., 2008, 2010)
- Data-driven, data-adaptive and model-based method using an iterative closed loop (SWES+; Ishiyama et al., 2014, 2015)
- Near-surface model-based method (SWAMI; Strobbia *et al.*, 2010, 2011)
- FWI-based methods (e.g. Ernst, 2013)

Methods of surface-wave separation

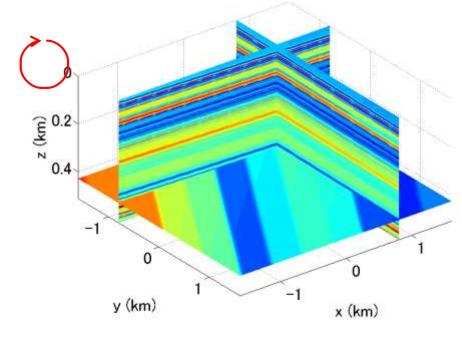






- Simple (and robust) model
- Parameterized by surface-wave properties: few parameters
- Data-adaptive
- Computationally affordable

FWI



- Complex (and sensitive) model
- Parameterized by near-surface properties: many parameters
- Computationally expensive



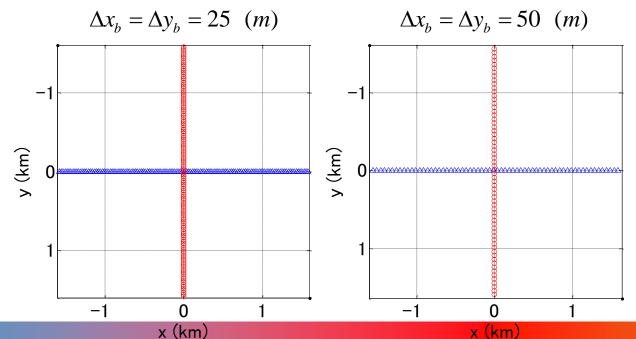
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Real data examples

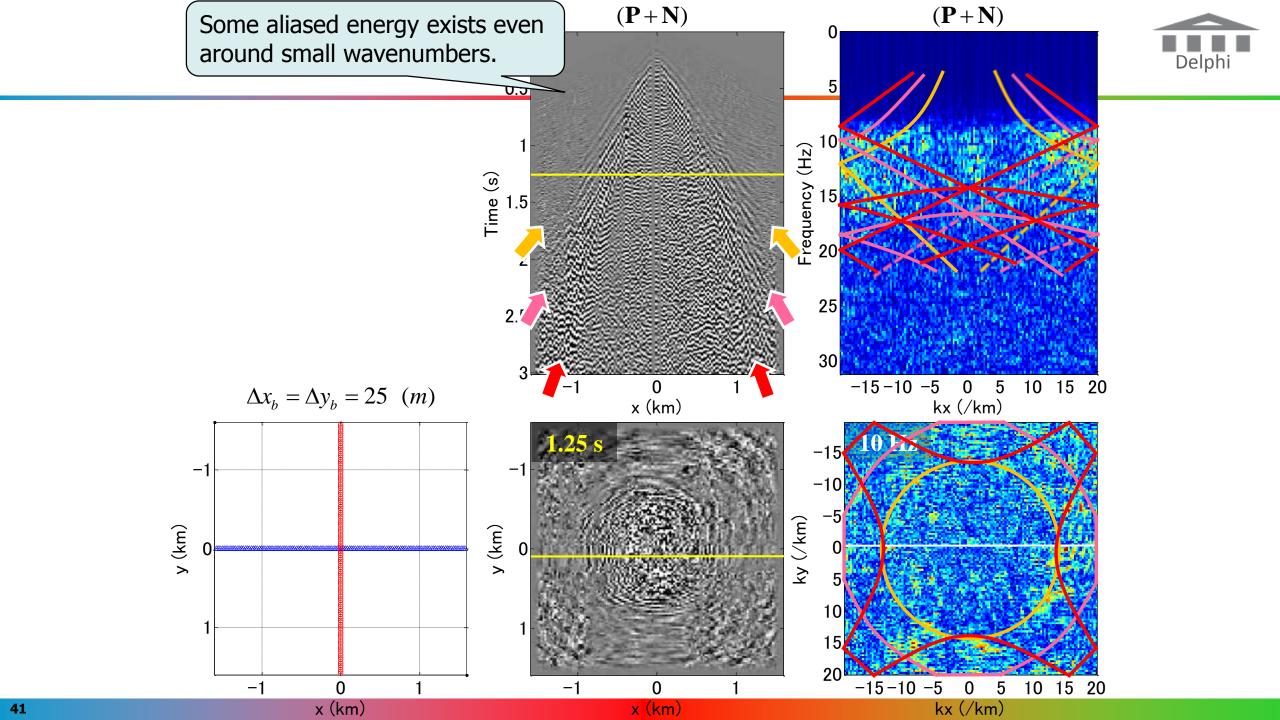


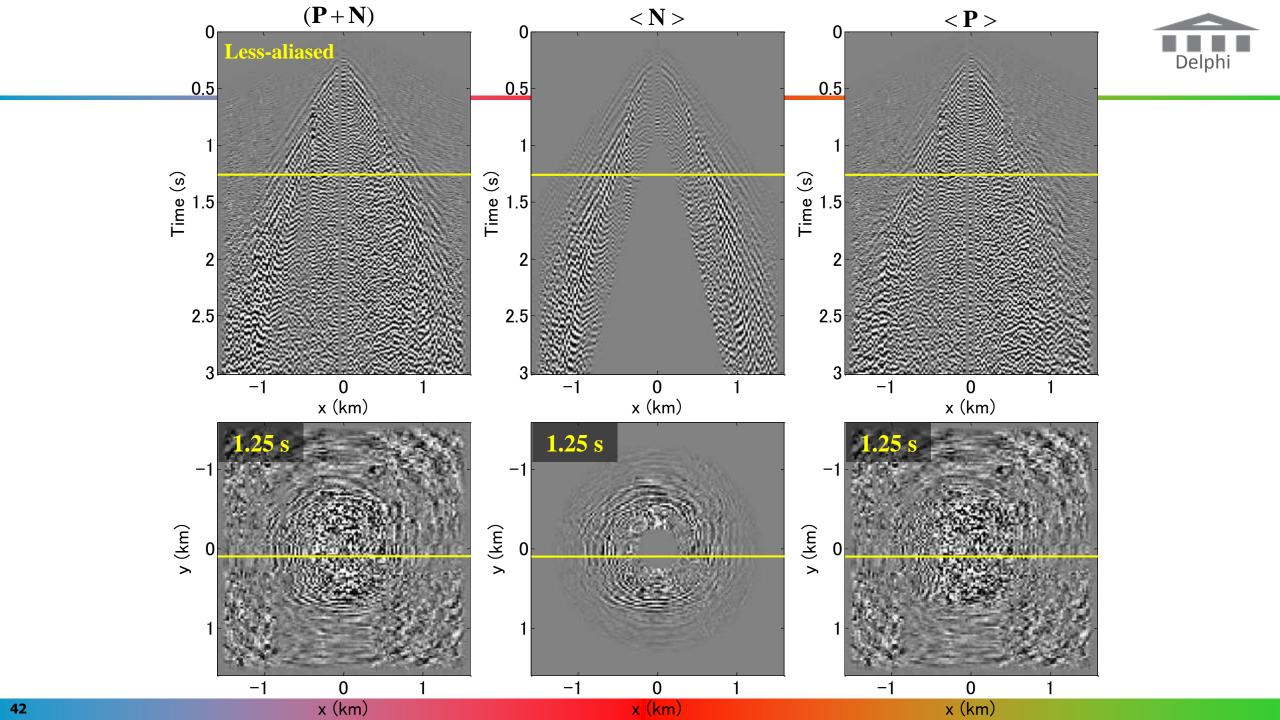
- 3D land data acquired onshore Abu Dhabi
 - (1) Less-aliased vertical geophone data
 - (2) More-aliased vertical geophone data

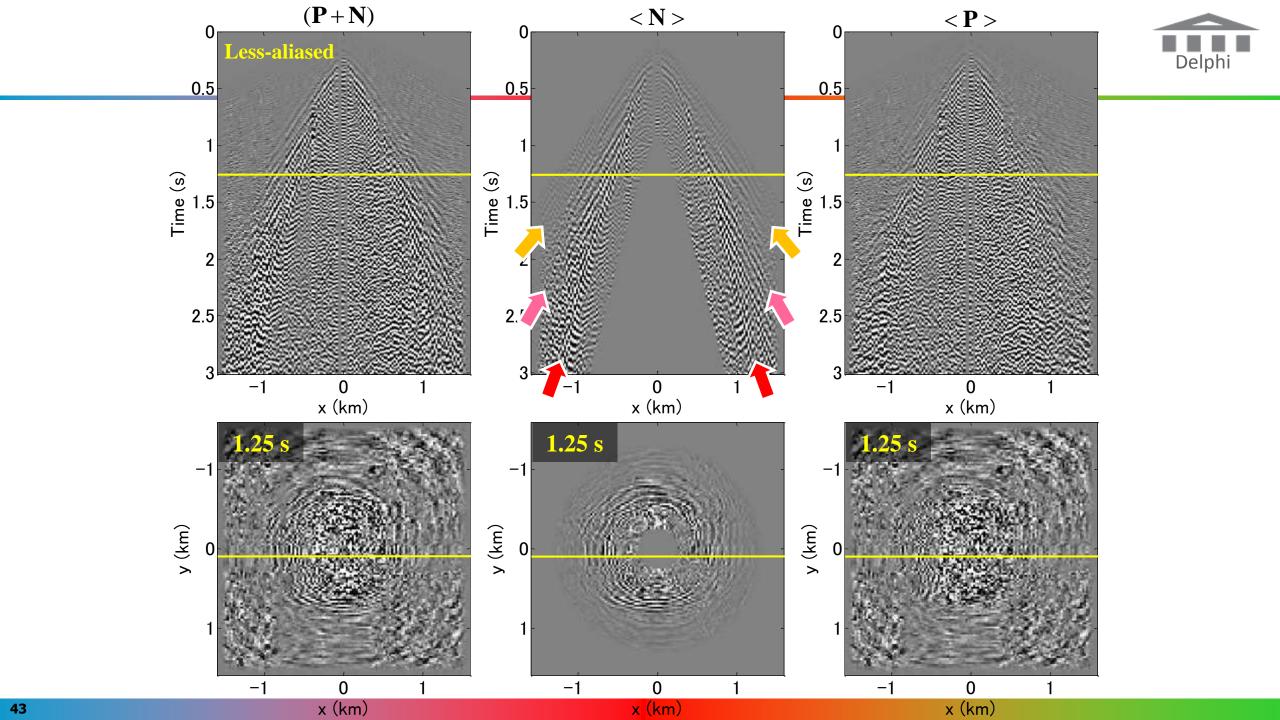


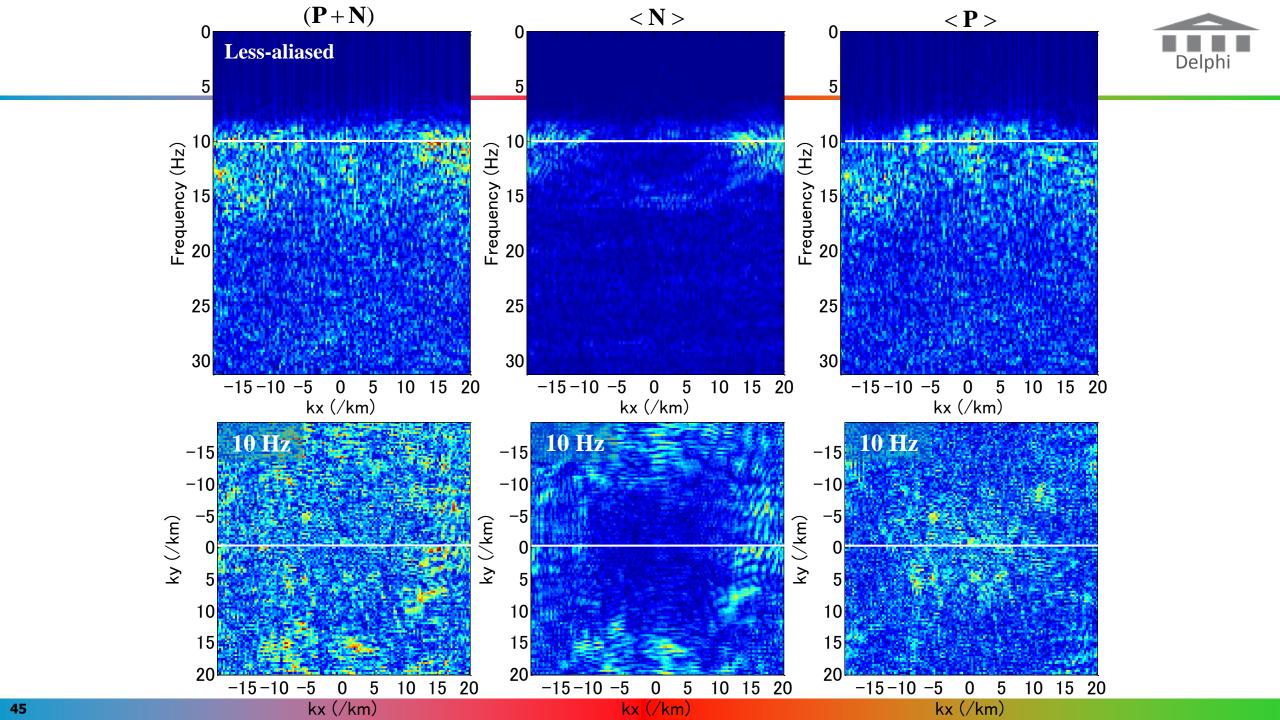
Receiver: Arrayed geophones

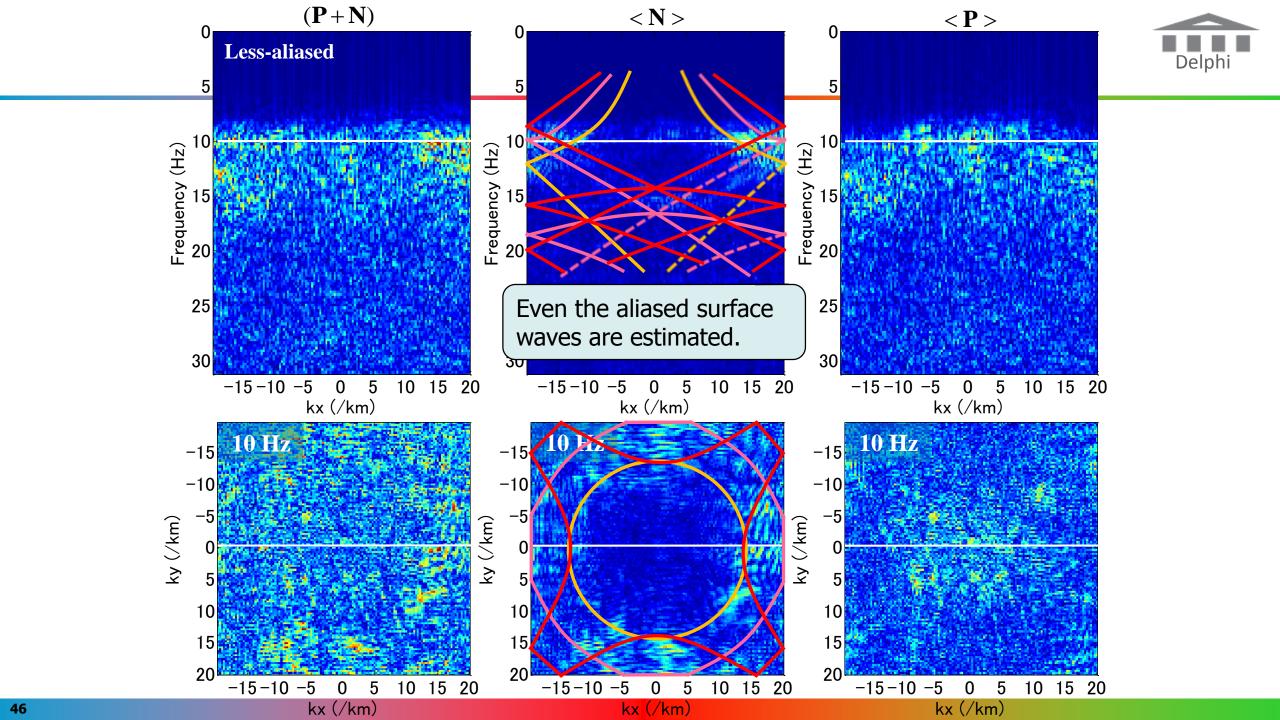
Source: Vibroseis with 6 Hz to 80 Hz linear sweep

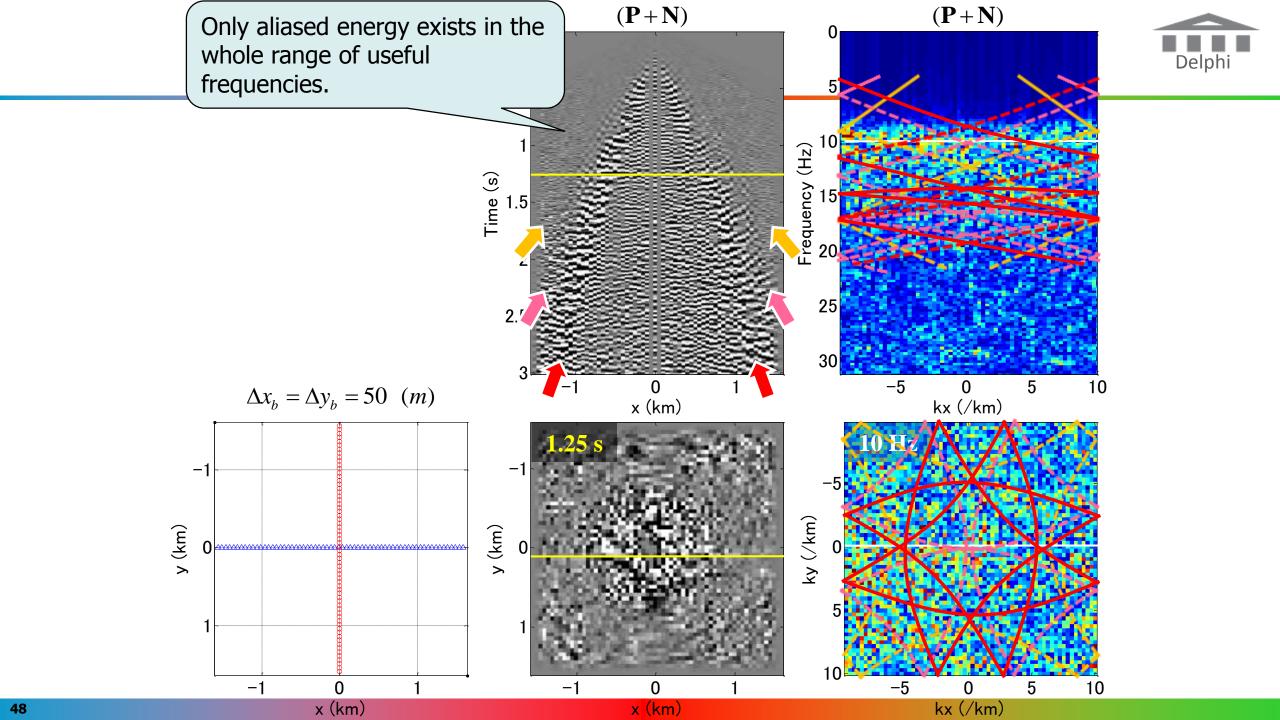


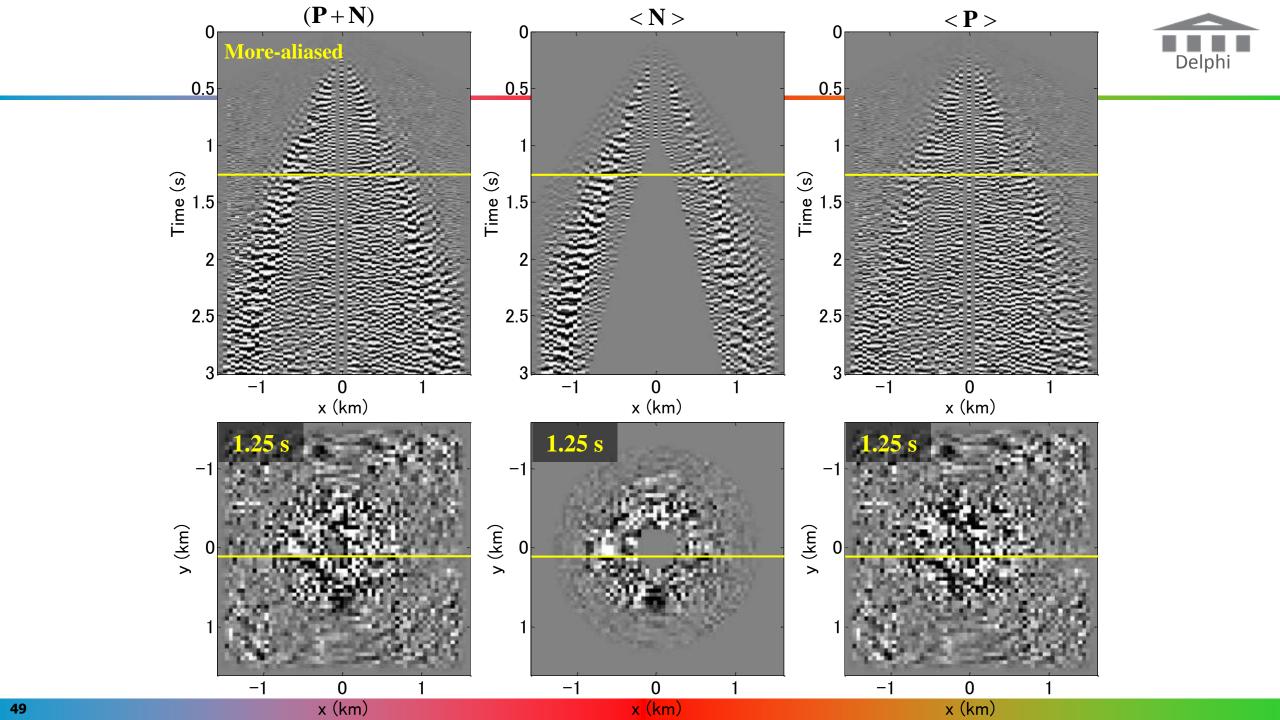


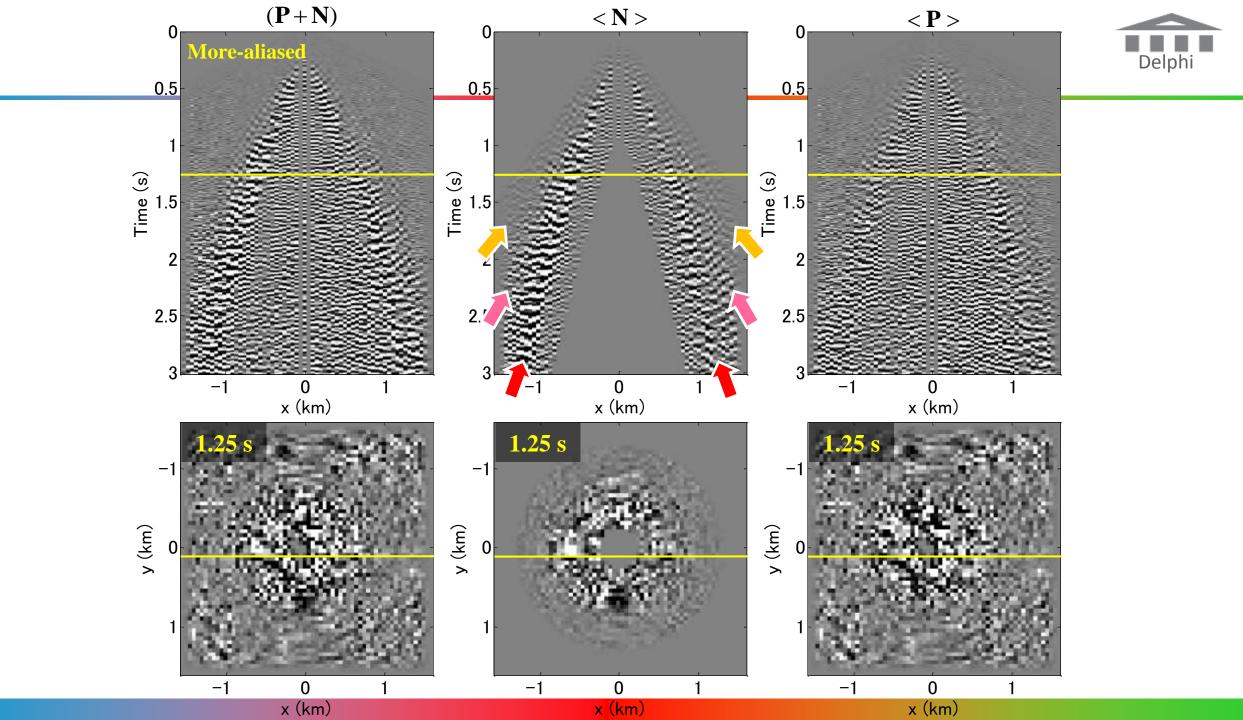


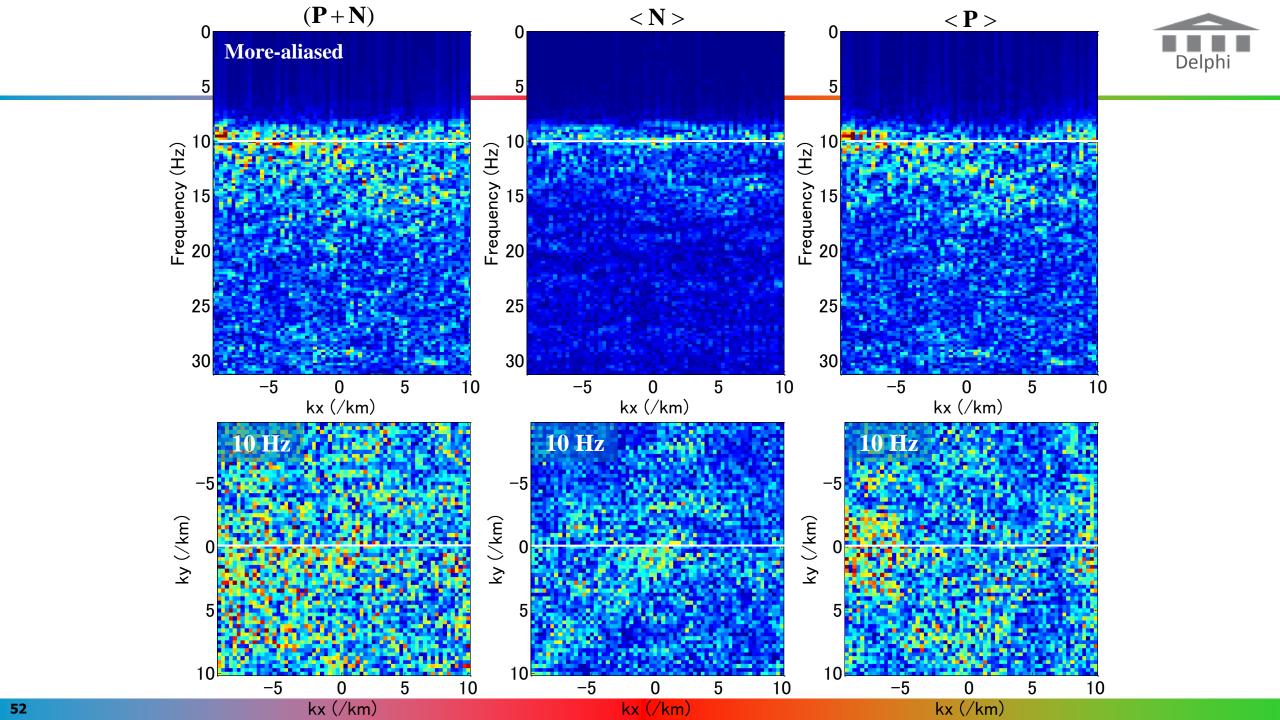


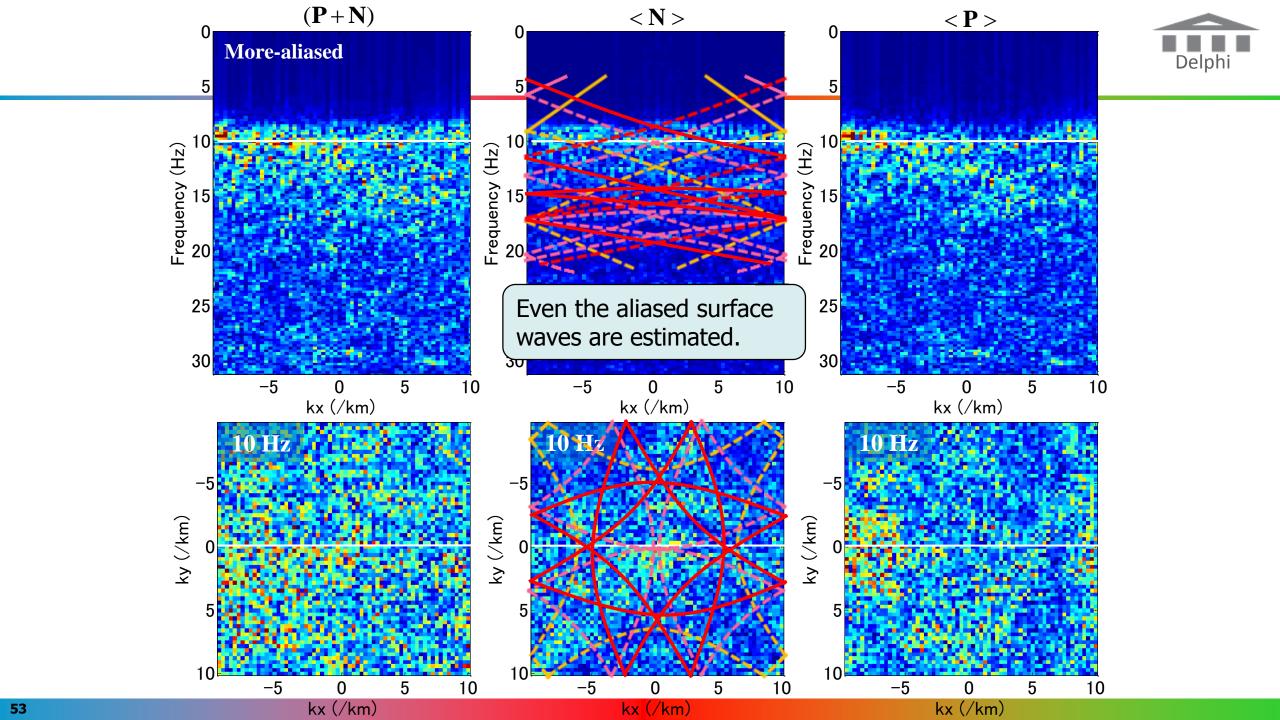








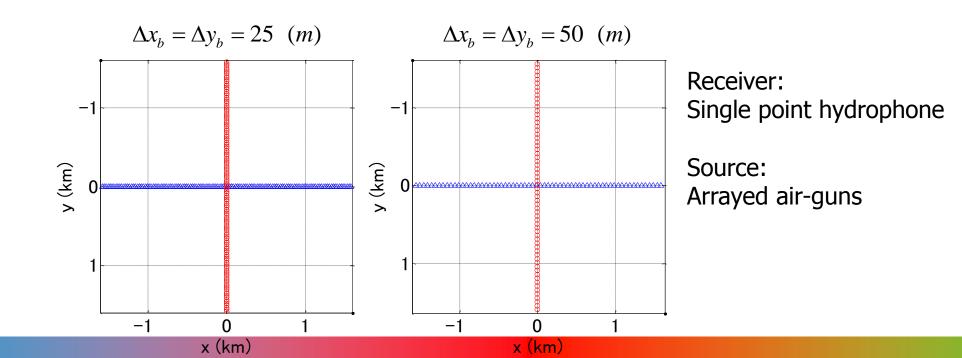


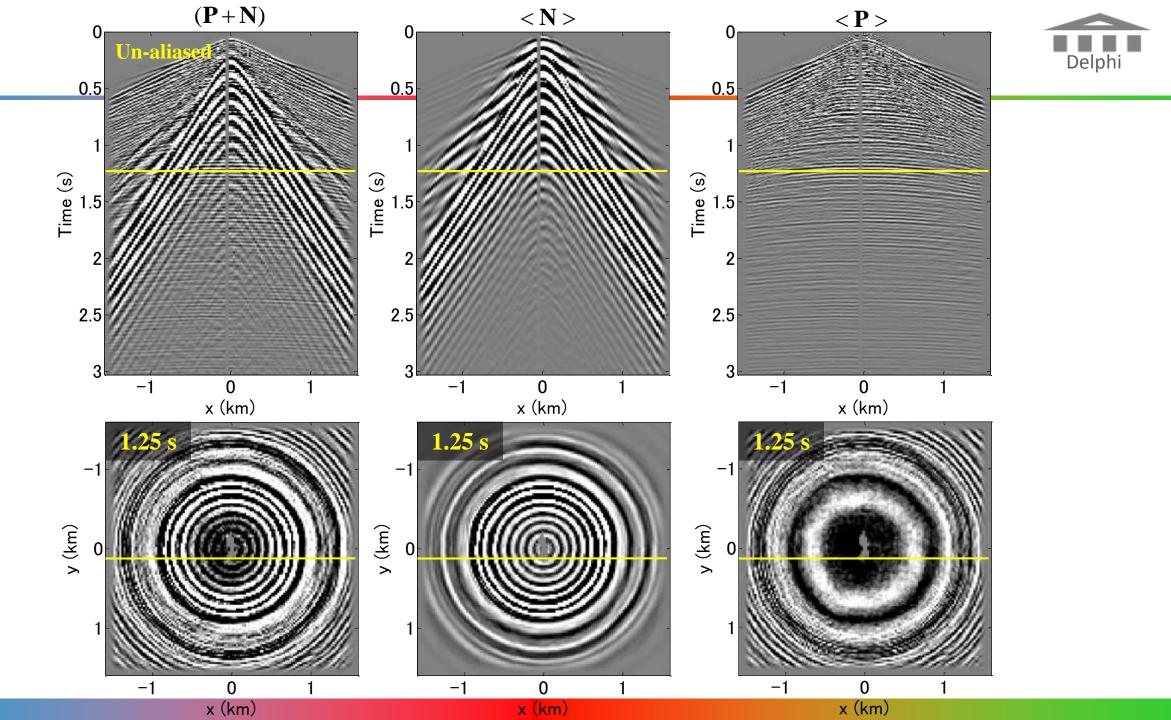


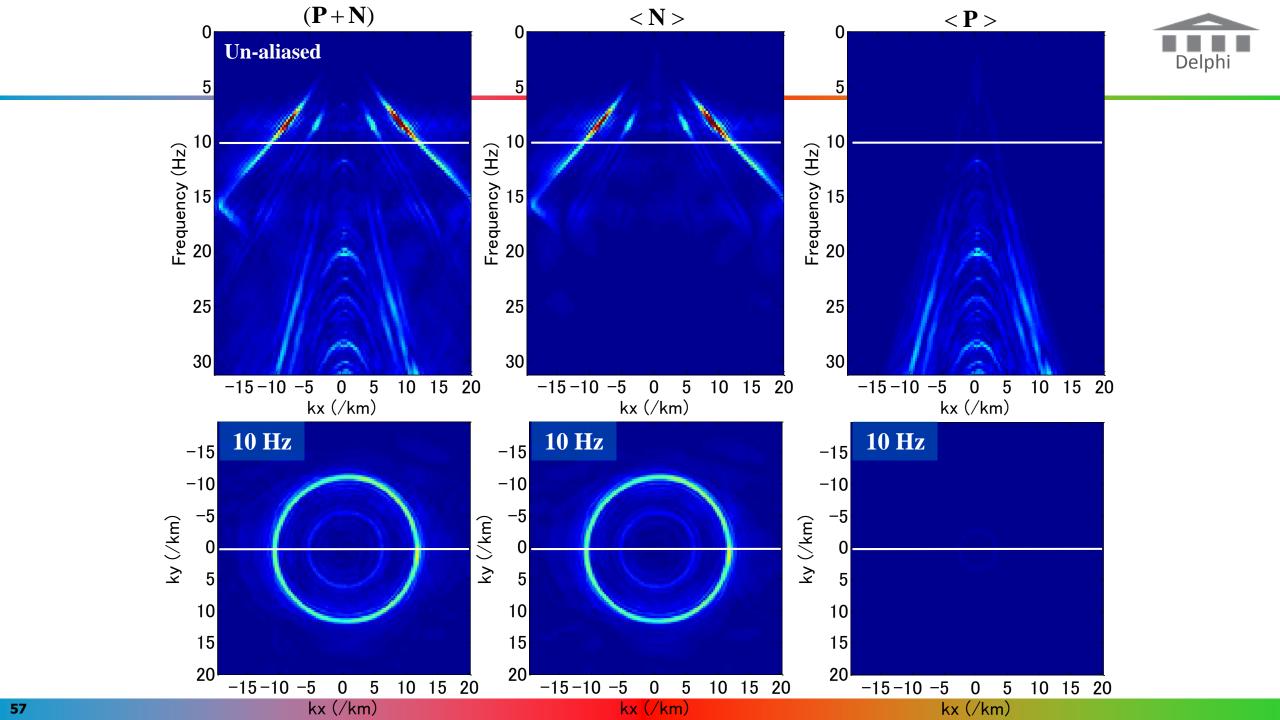
Real data examples

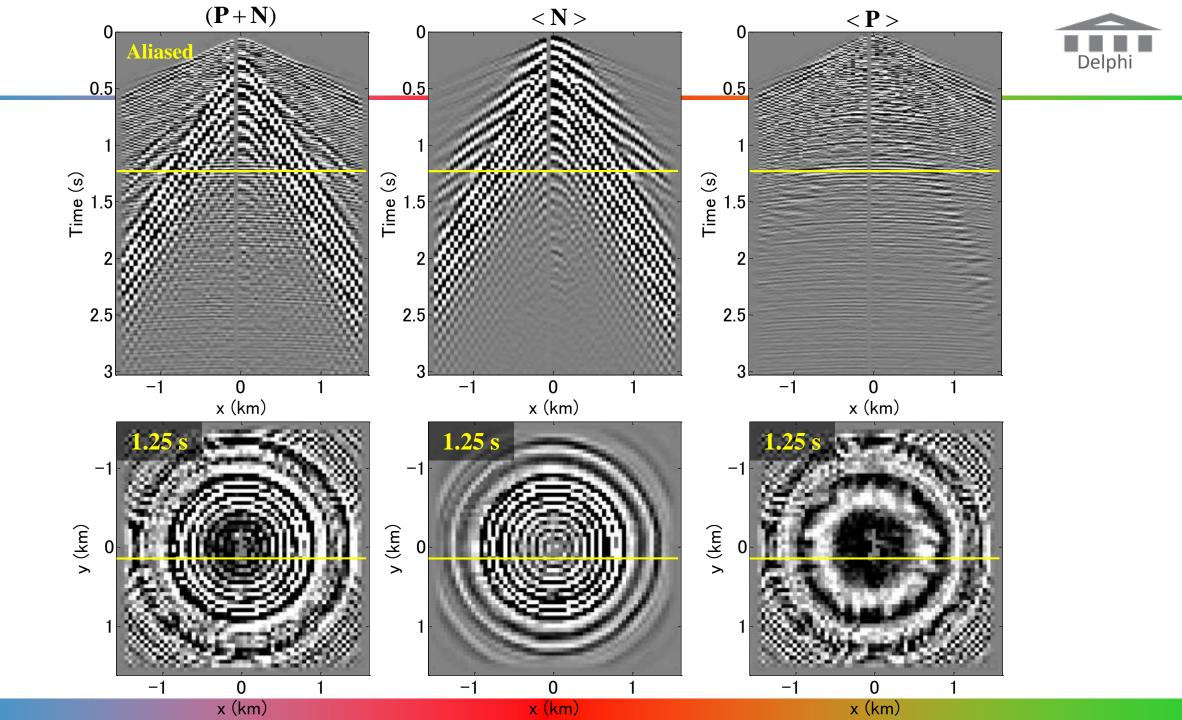


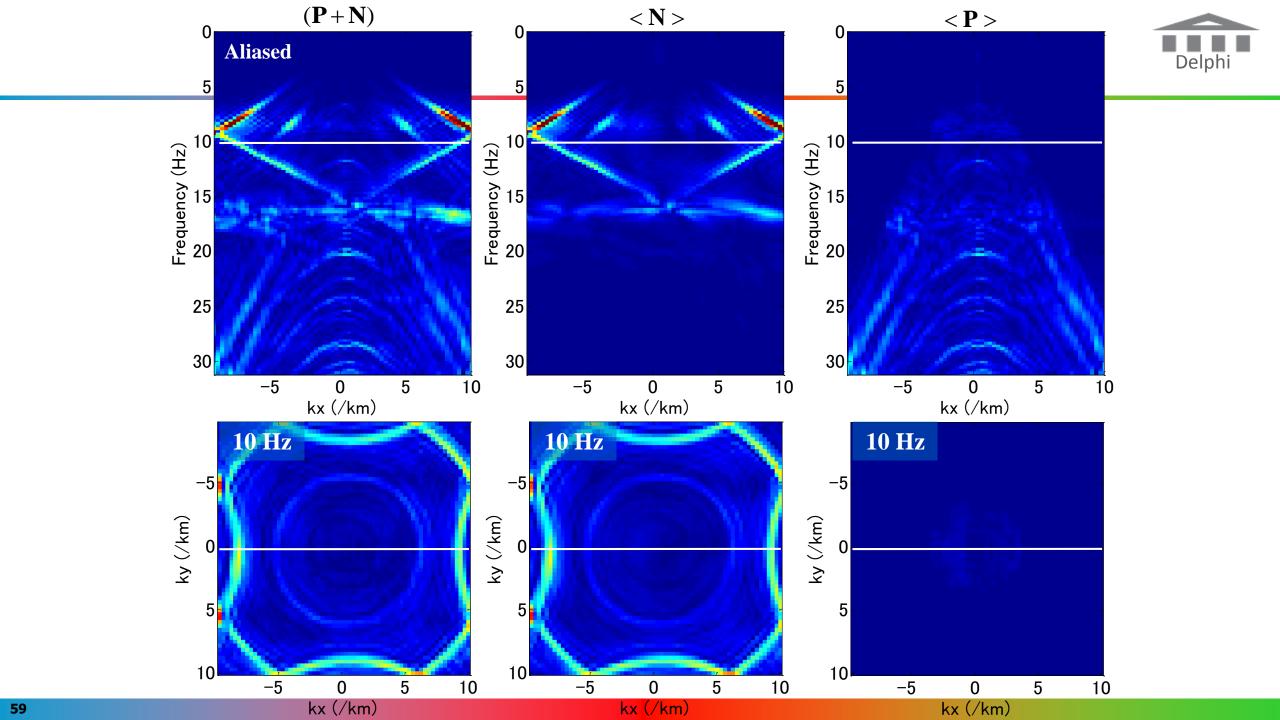
- 3D OBC data acquired offshore Abu Dhabi
 - (1) Un-aliased hydrophone data
 - (2) Aliased hydrophone data
 - (2)' The above data with a conventional filtering method

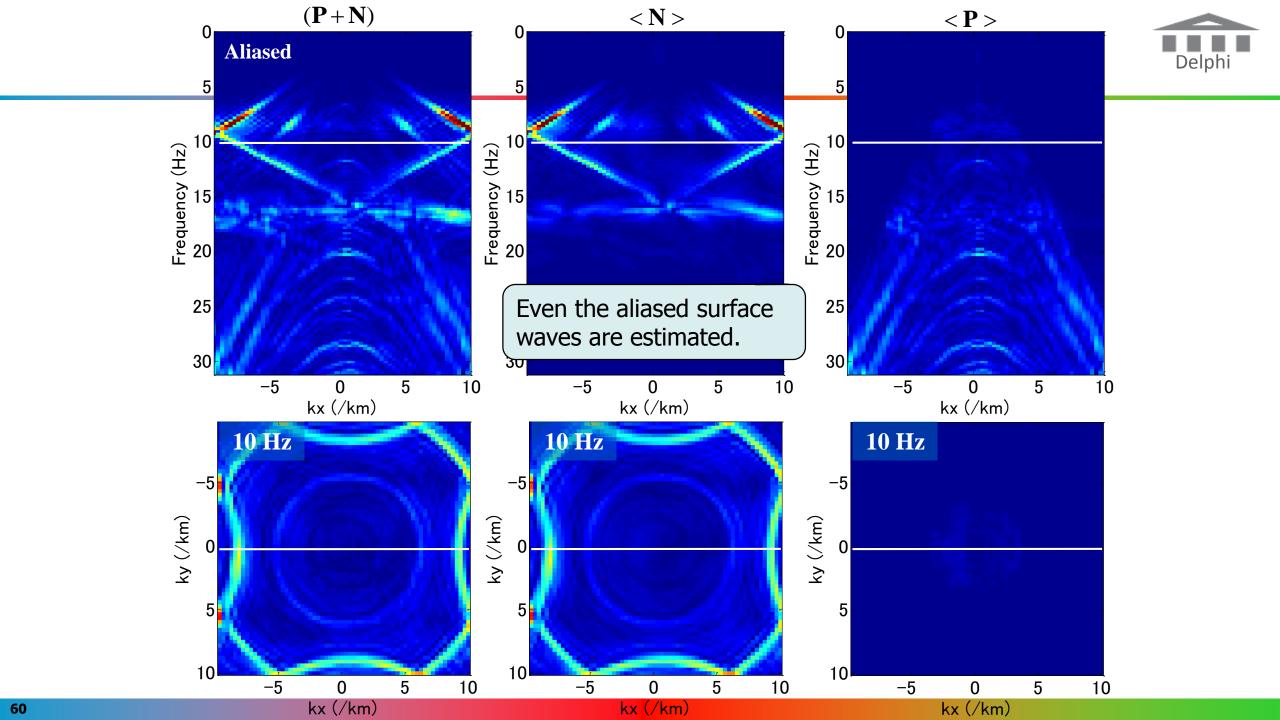


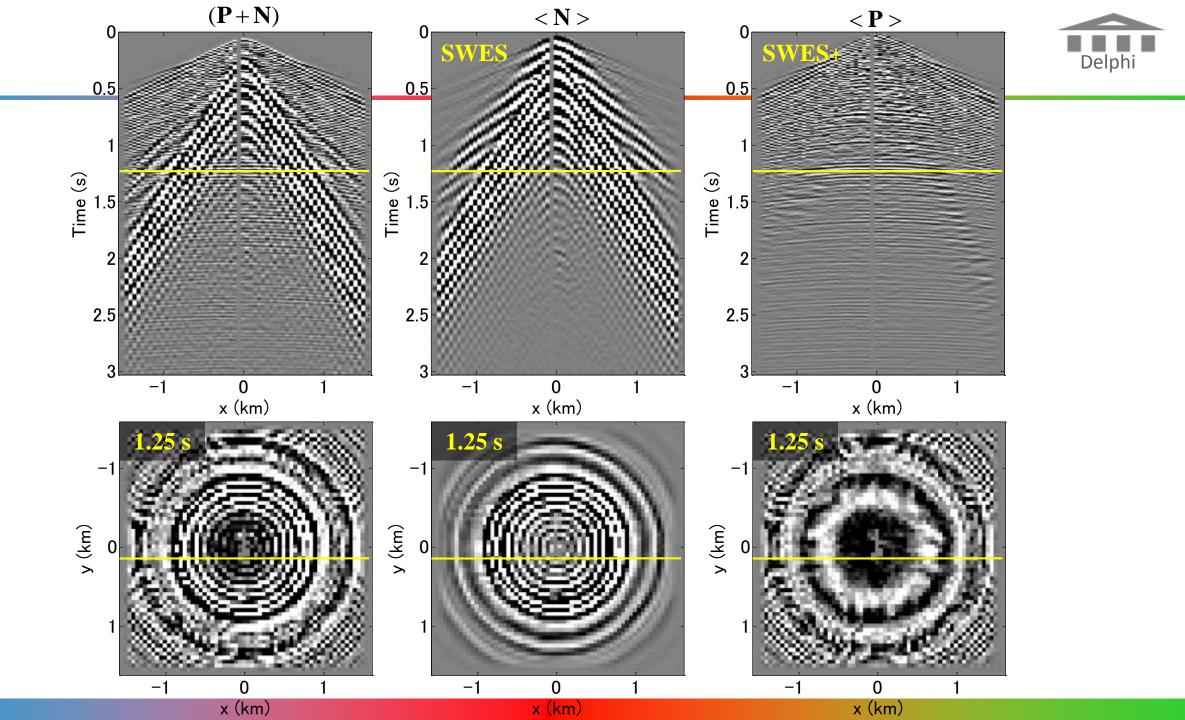


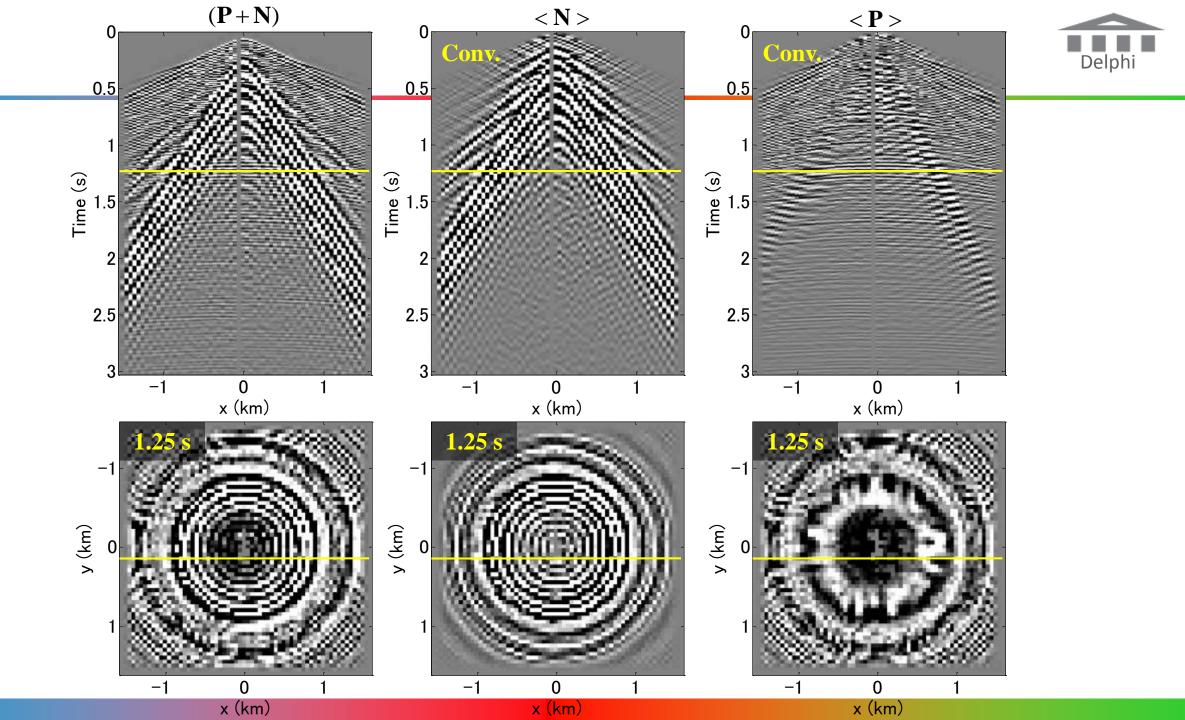


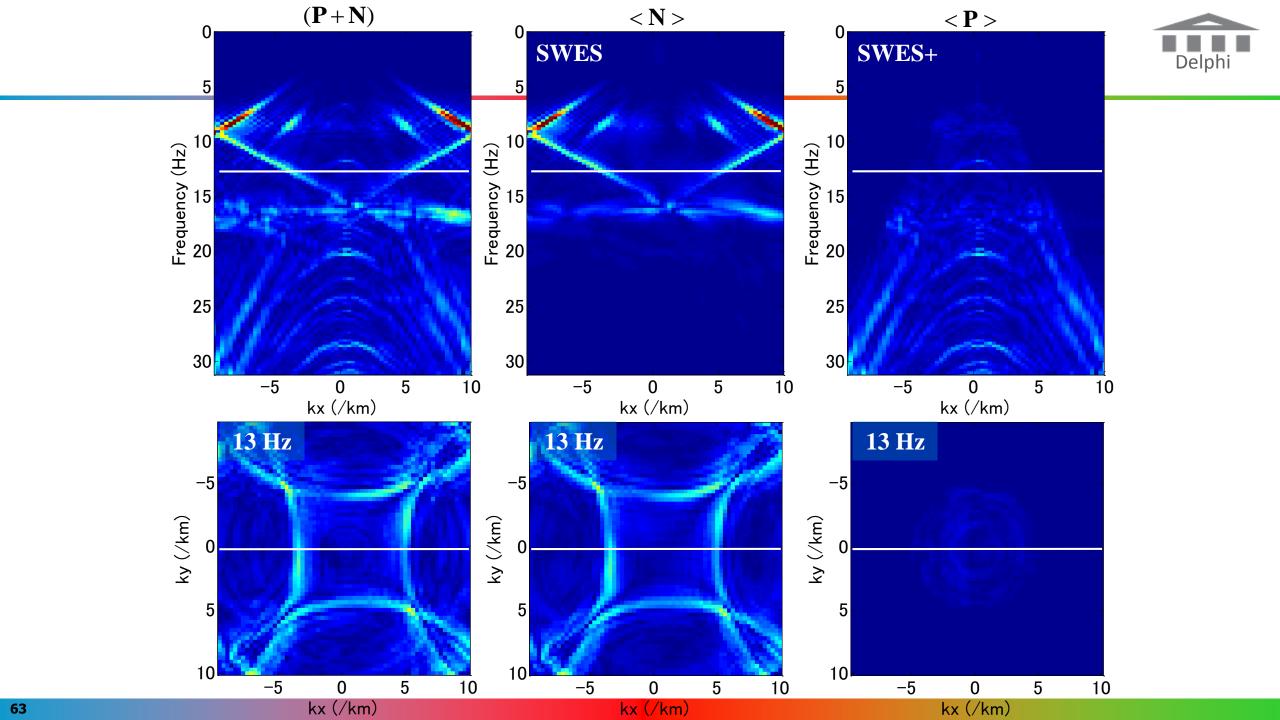


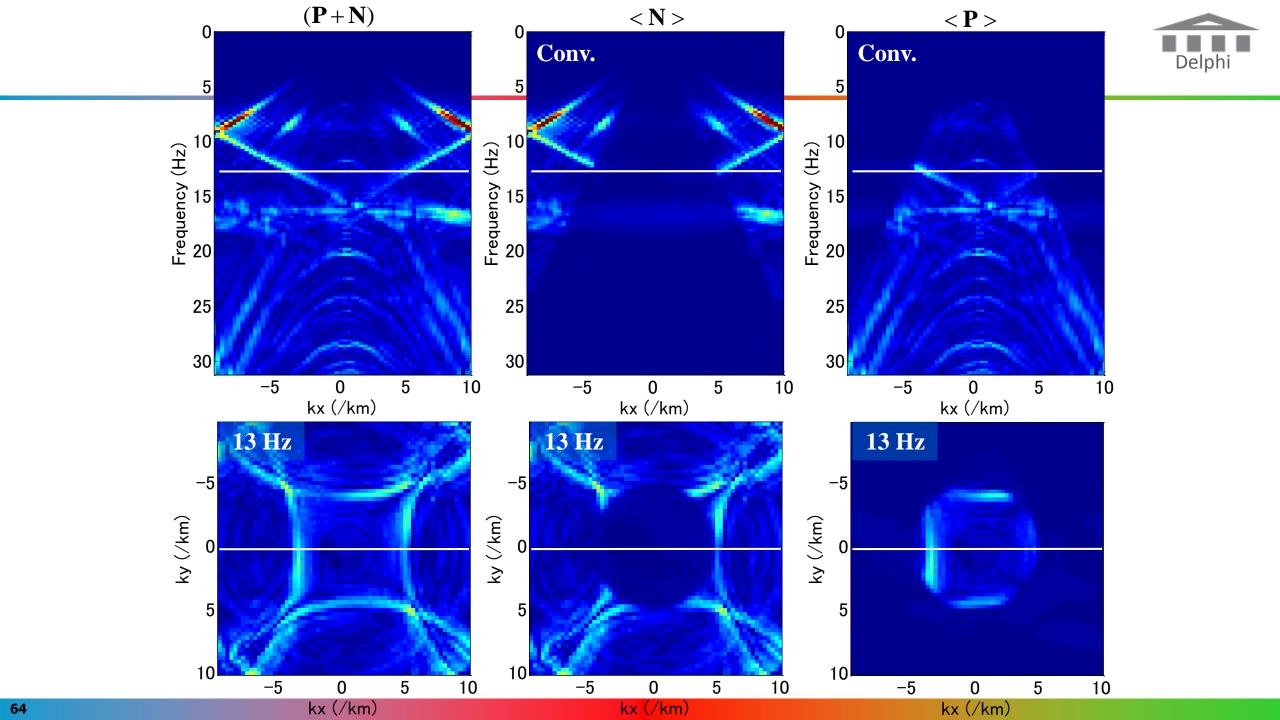


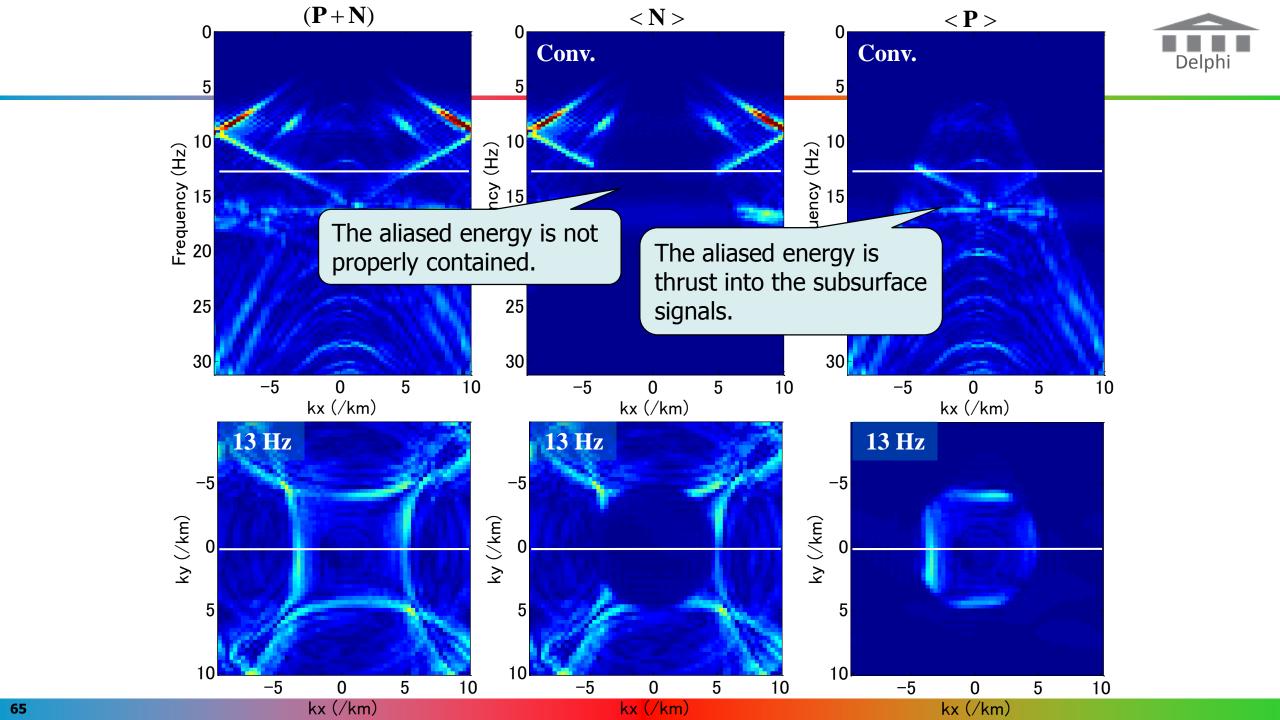








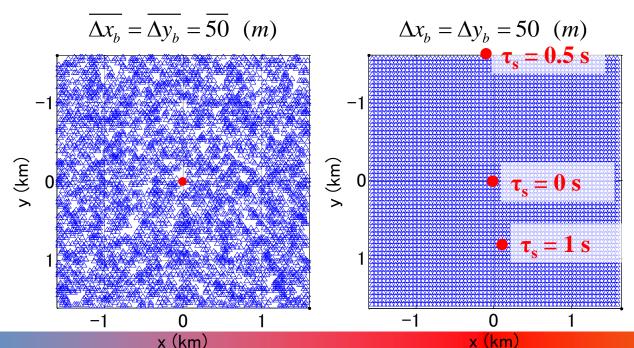




Real data examples

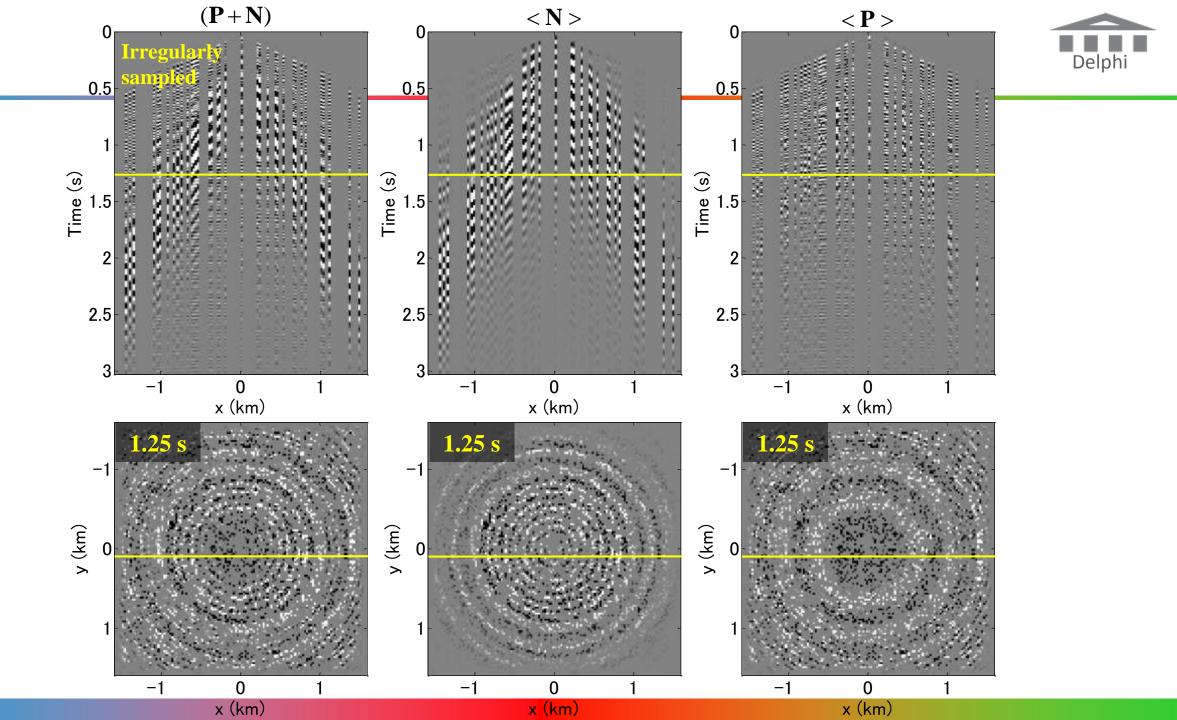


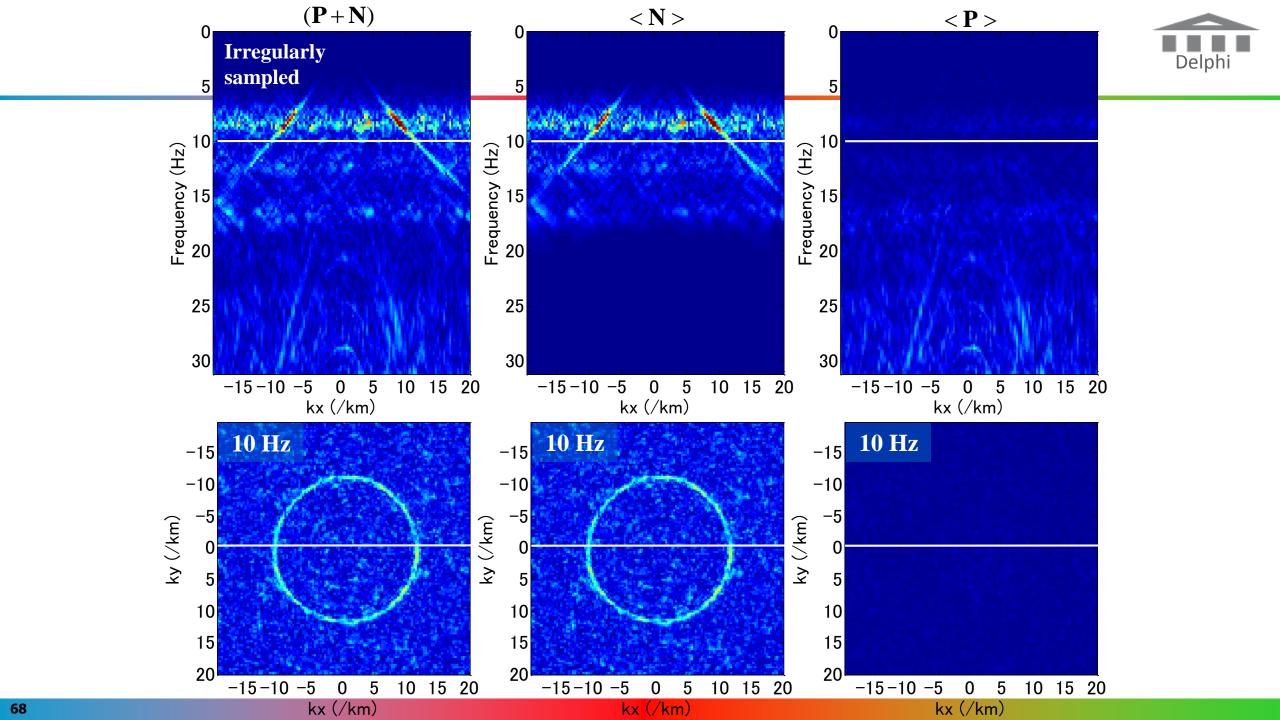
- 3D OBC data acquired offshore Abu Dhabi
 - (1) Irregularly sampled hydrophone data
 - (2) Blended hydrophone data

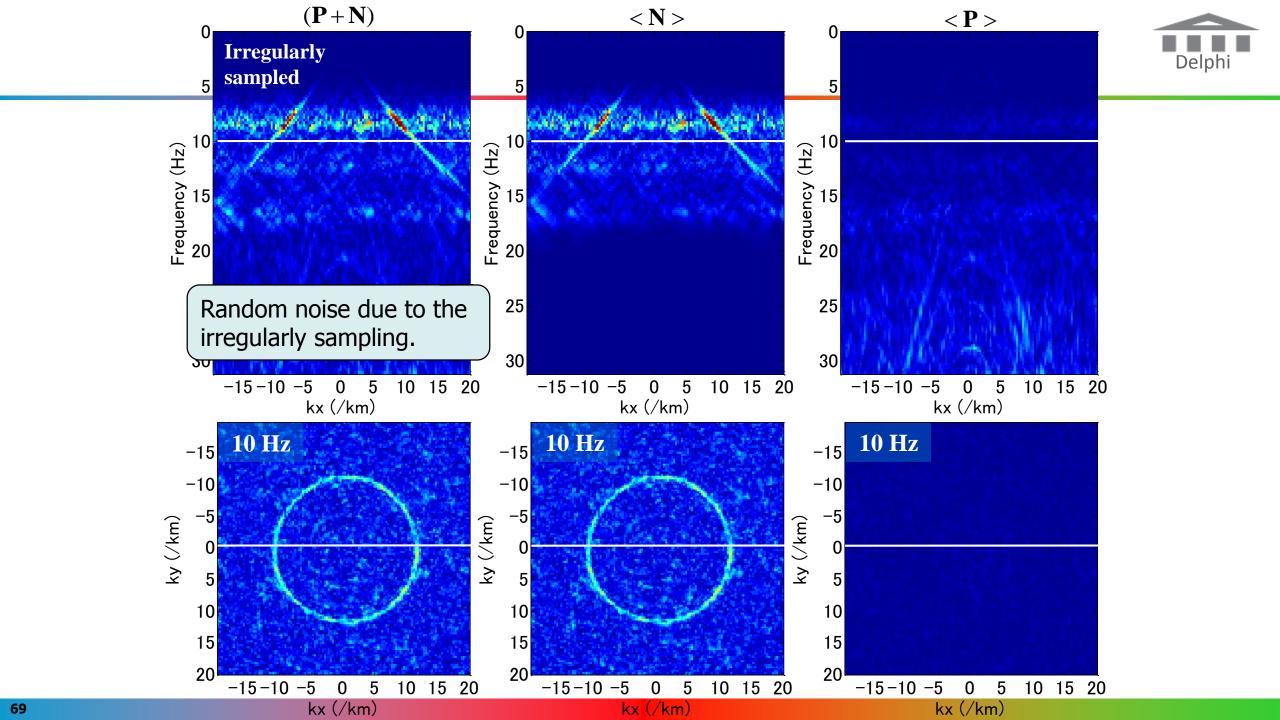


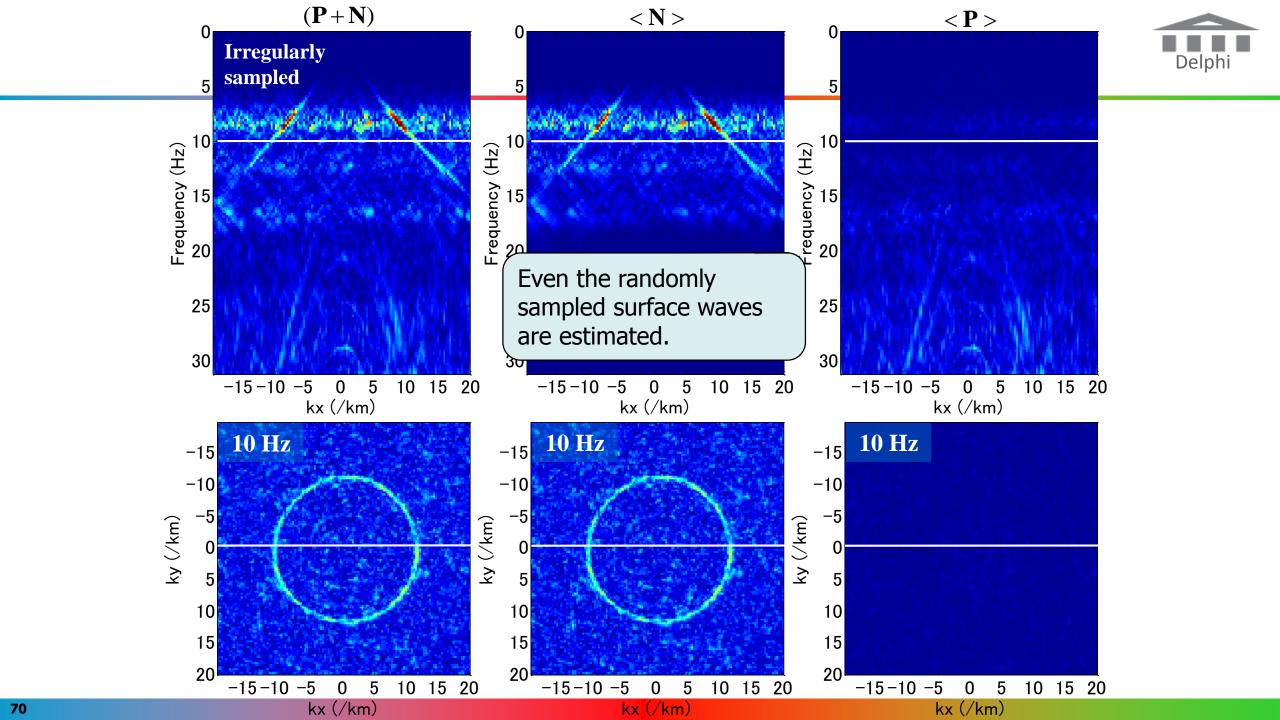
Receiver:
Single point hydrophone

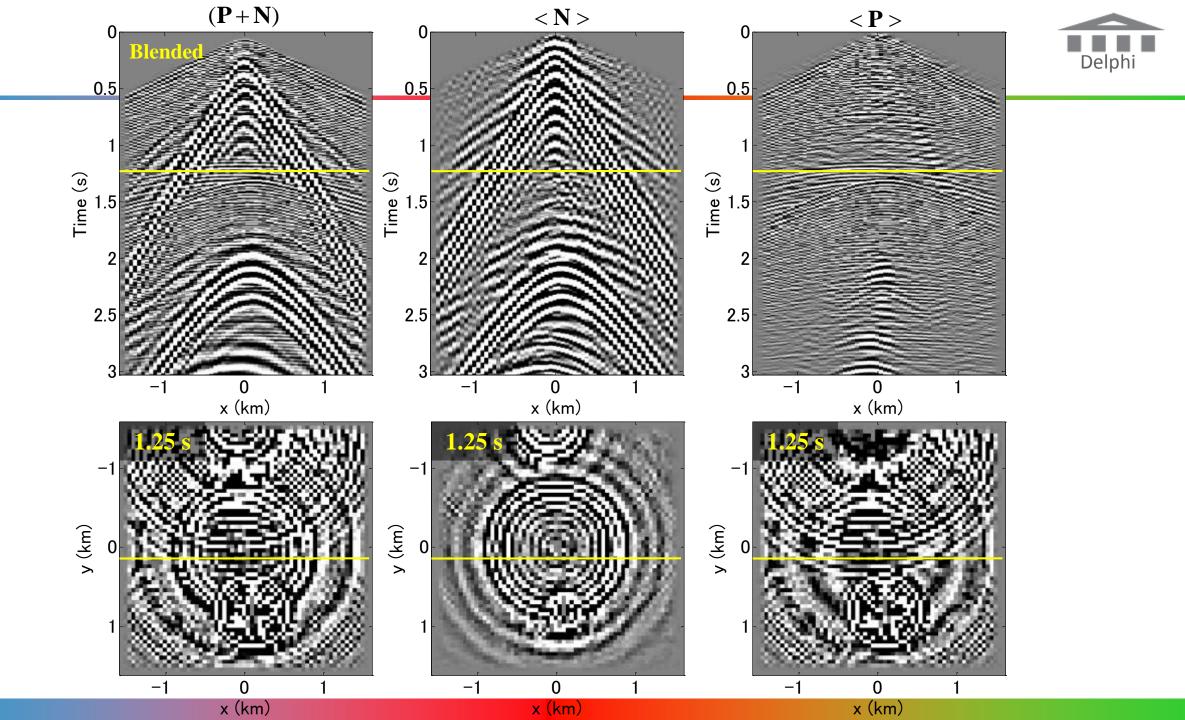
Source: Arrayed air-guns as a single point source (1) as a blended source (2)

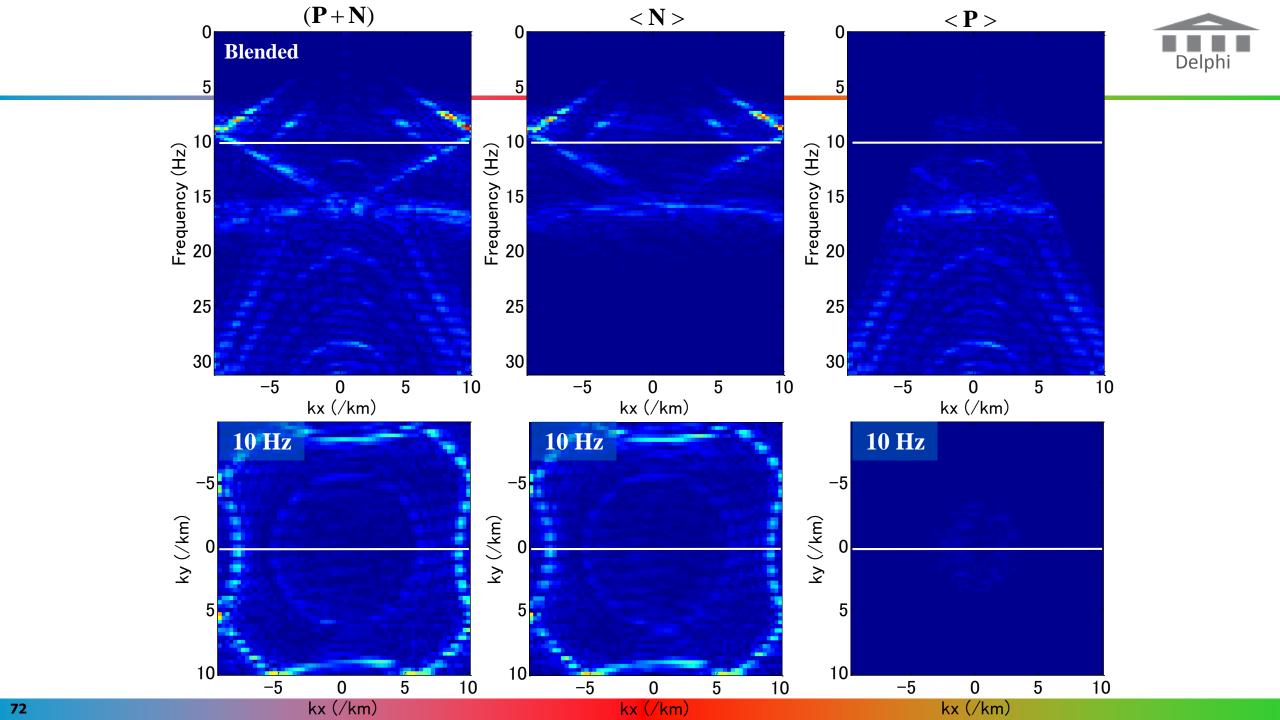


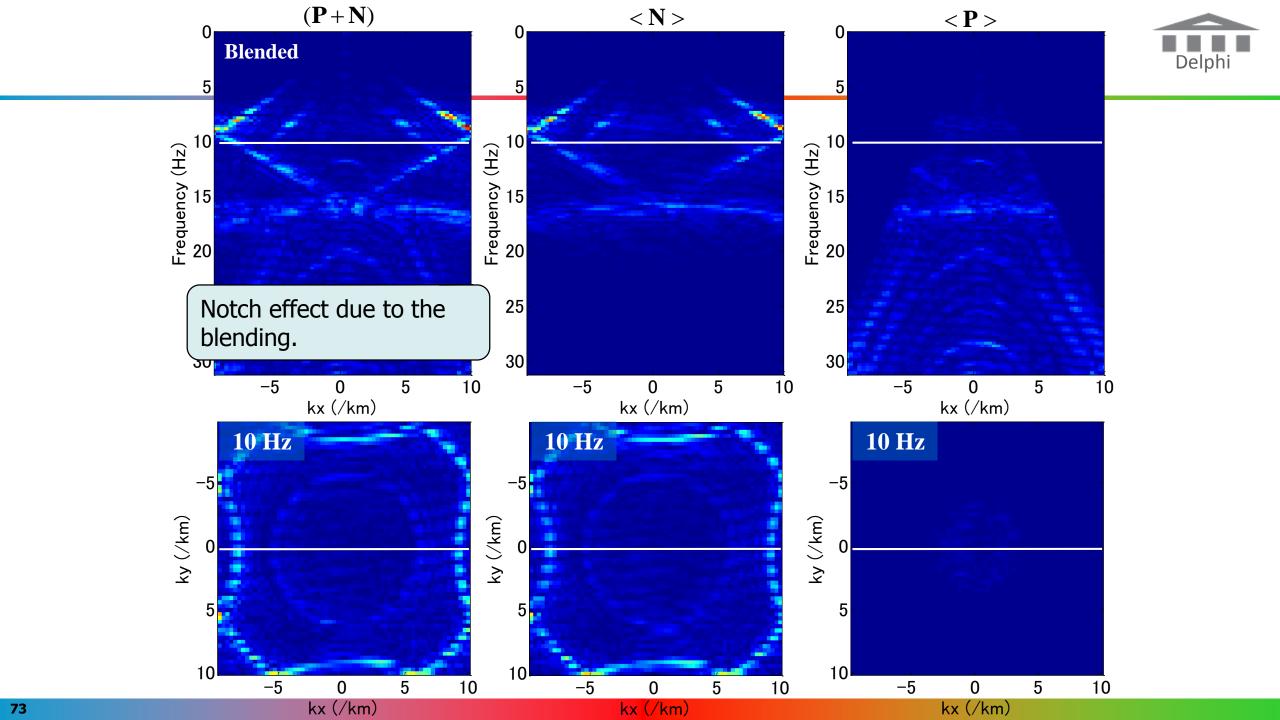


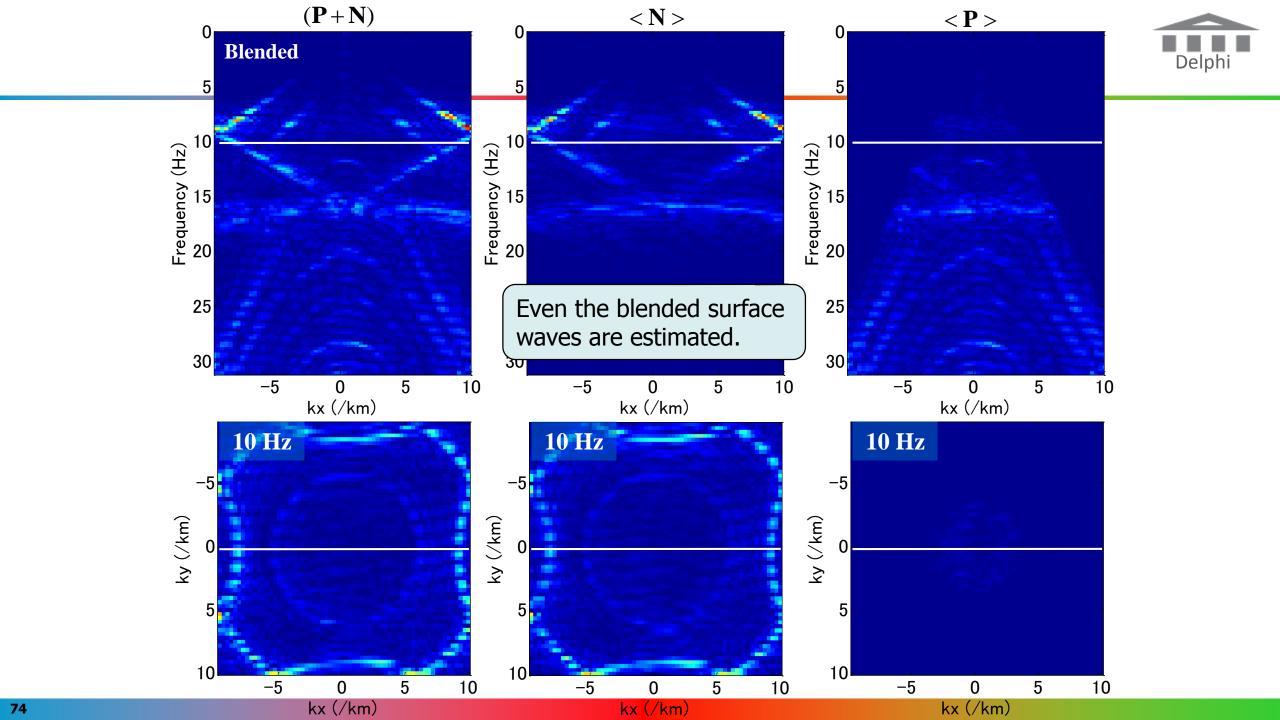














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Conclusions and remarks



- The proposed method addresses the surface-wave properties, i.e., dispersion and multi-modes.
- The method is data-driven and data-adaptive, automatically taking into account physical phenomena such as spatial variation, attenuation, anisotropy, etc.
- The method can be applied in any geometry domain, i.e., in any gathers.
 It can be applied to under-sampled, asymmetrically sampled, irregularly sampled and blended seismic data.





• This suggests the possibility of relaxing the spatial sampling interval, encourages random sampling / blending, and offers flexibility with respect to acquisition geometry.

Acknowledgements



- ADNOC R&D OSC / PI
- INPEX / JODCO
- DELPHI / TUD



Thank you.



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