

Evolution of Large-Scale Topography and River Drainage Direction from Mantle Flow Models*

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Search and Discovery Article #41769 (2016)**

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Abstract

Flow within the mantle imposes a deformation of the Earth's surface called dynamic topography. However, quantifying this topography in space and time remains a challenge. Over the last few years, we have developed dynamic Earth models that progressively assimilate cutting-edge tectonic reconstructions with continuously closing plates and plate deformation. These models read-in plate velocities and subduction zone location and geometry defined in million-year increments for the last 230 Myr. The approach allows us to simultaneously model global mantle flow and large-scale lithospheric deformation, and in turn estimate the evolution of both dynamic and isostatic topography back to the Jurassic. In the South Atlantic domain, we show that long-wavelength topographic features, including the anomalously high elevations of southern Africa and East Brazil, could result from whole mantle flow, rather than asthenospheric flow, as previously suggested. Comparing model tectonic subsidence along the South American passive margin with that estimated from boreholes, we attribute the post-rift subsidence of the East Brazil Rift System since the Eocene to the motion of South America over ancient subducted slabs. This interplay between plate motion and mantle flow, combined with the development of flat slab subduction in Peru, also explains the reversal of the drainage direction of the Amazon River and the drying of the Pebas System since the Miocene. On a smaller scale, we show that deformation within the Central Andes imposes an inboard migration of the subduction zone relative to the South American plate that could have resulted in the migration of the depocenter of the Bolivian Chaco foreland basin and associated drainage

reorganization. In South East Asia, we use simpler models without lithospheric deformation to attribute missing sedimentary sections of latest Cretaceous to Paleocene age across Sundaland to dynamic uplift resulting from a ~10 Myr subduction hiatus along the Sunda active margin. In the Arctic, we show that subduction-driven dynamic topography can explain the rapid Middle to Late Jurassic subsidence of the Slave Craton and North Slope of Alaska and the vertical motions of the Barents Sea region, characterized by subsidence between ~170–50 Ma followed by uplift since 50 Ma. Together, these results illustrate the important role of mantle flow on the evolution of large-scale topography, making dynamic Earth models powerful for frontier hydrocarbon exploration.

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Melbourne
September 2015

Evolution of Large-Scale Topography and River Drainage Direction from Mantle Flow Models

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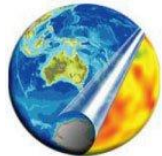
*R.D. Müller^a, M. Gurnis^b, S. Zahirovic^a, Grace Shephard^c, D.J. Bower^b
M. Seton^a, S. Williams^a & J. Skogseid^d*

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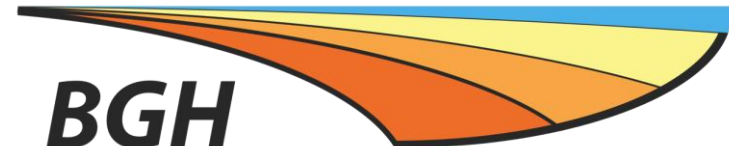
^c University of Oslo – ^d Statoil ASA



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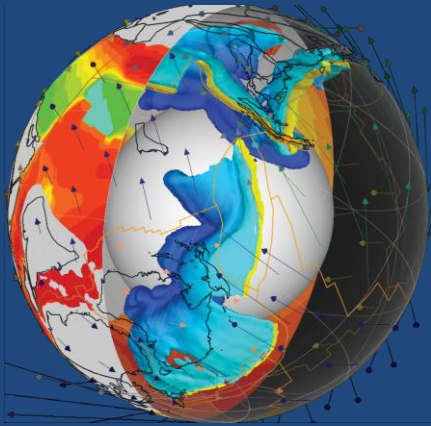


EarthBYTE

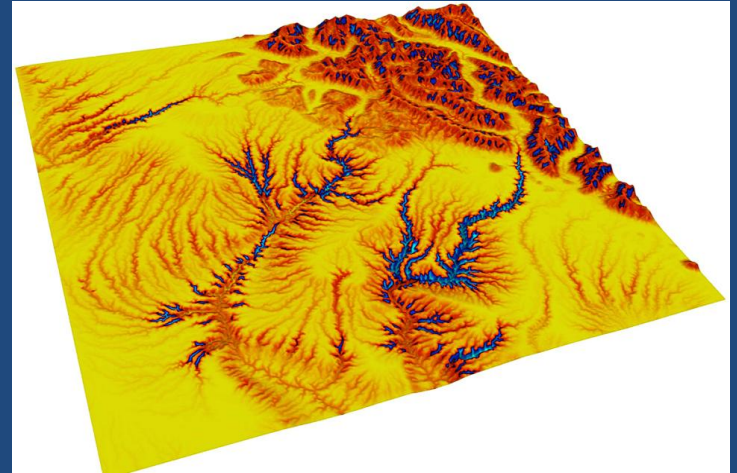


Basin GENESIS Hub

- 5-year industry linkage project with research focus on Australia, PNG and the Atlantic



Modelling effects of crustal- and mantle-scale processes on basin evolution



Modelling dynamics of surface topography, erosion, sedimentation

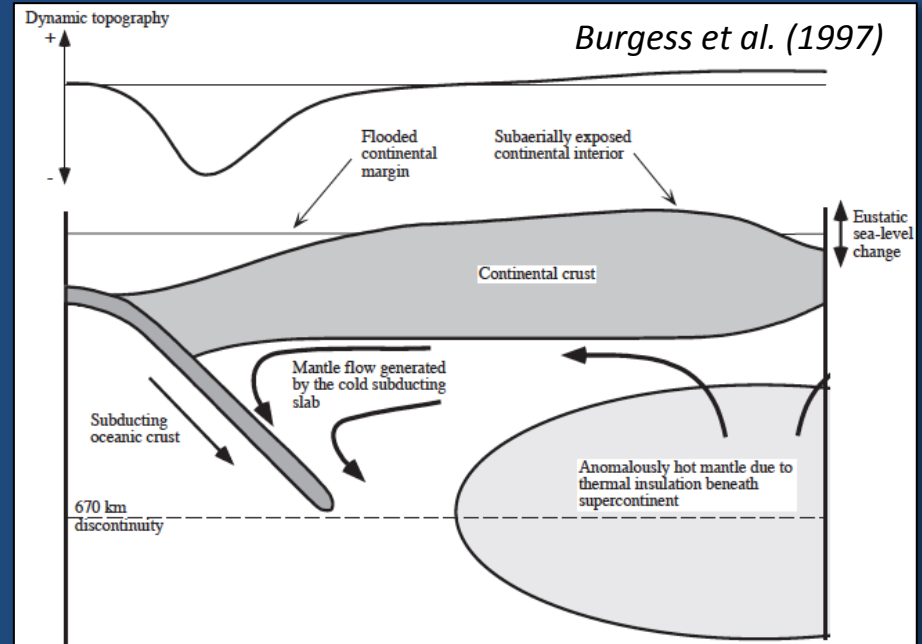
Isostatic (“tectonic”) and dynamic topography



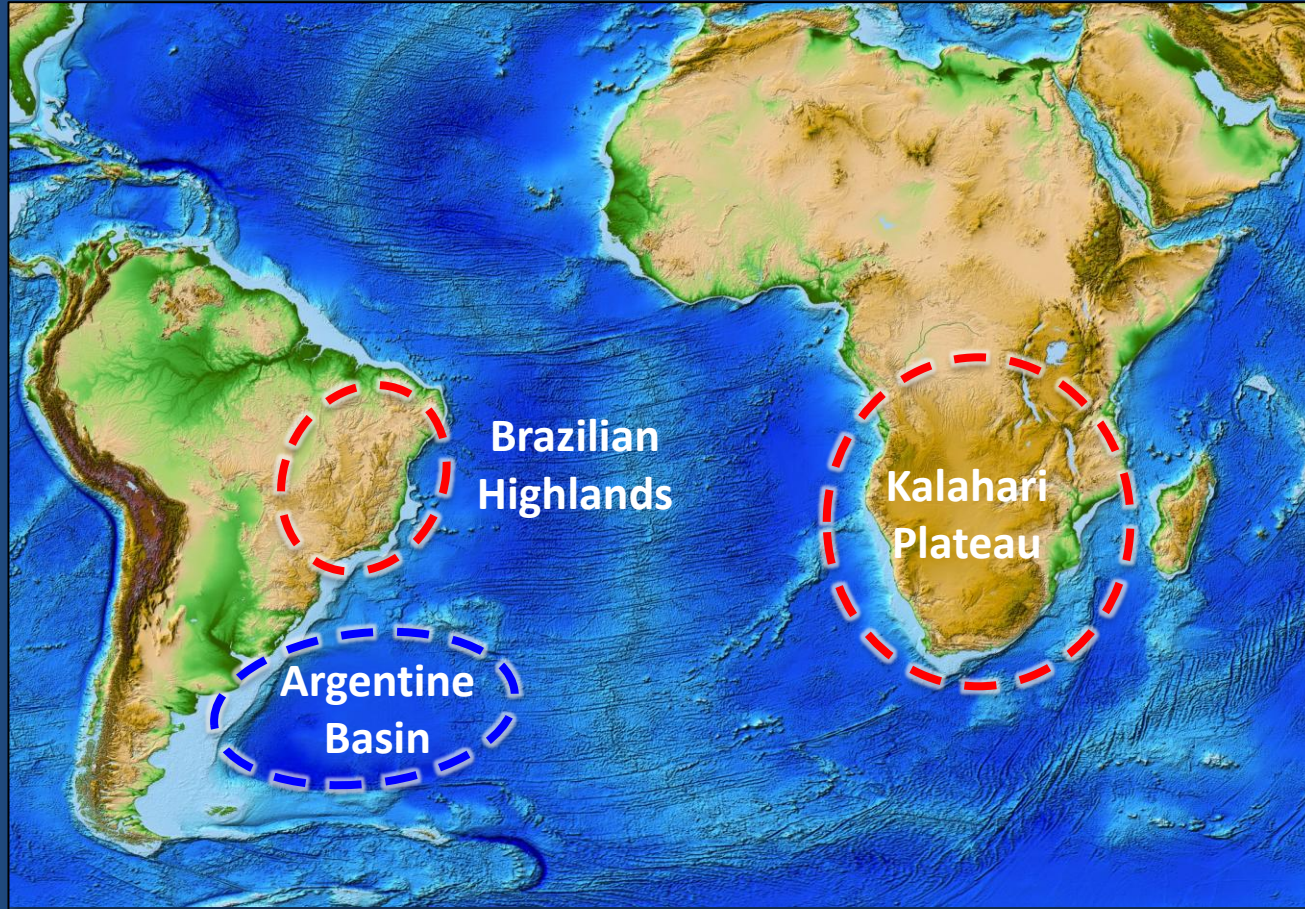
Many other processes affect topography, including:

- Thermal diffusion that affects the thickness of the lithosphere
- Mantle flow

- Most of the topography at the Earth's surface is *isostatically compensated*
- This topography is related to the thickness of the continental crust and lithosphere



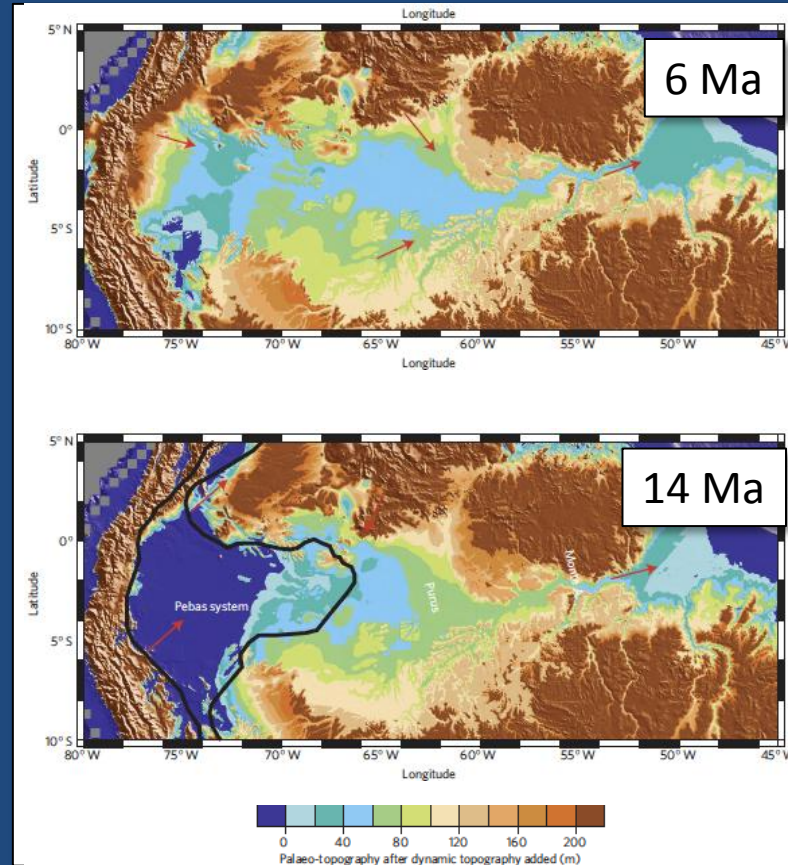
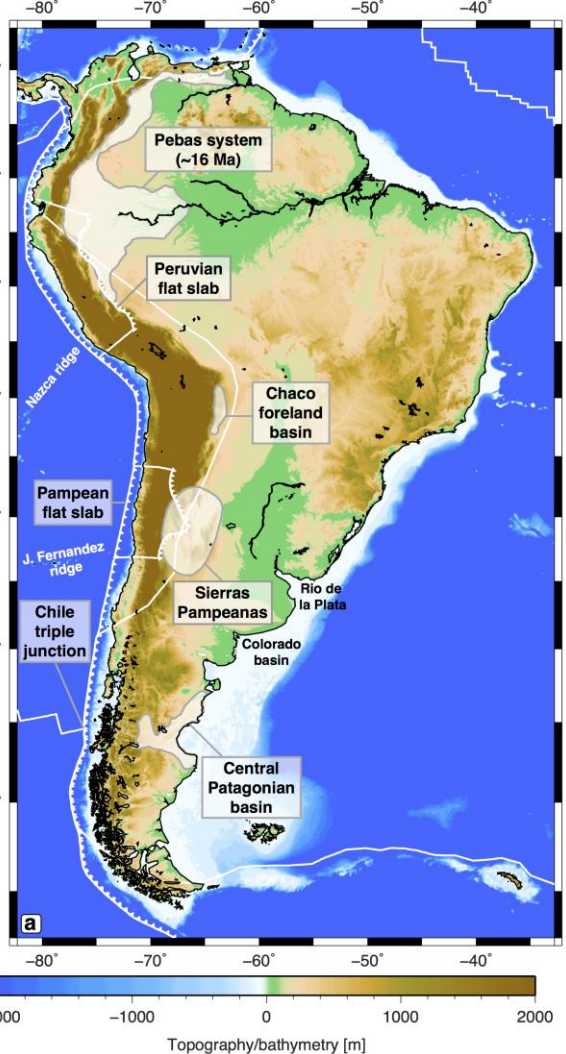
Topographic asymmetry of the South Atlantic domain



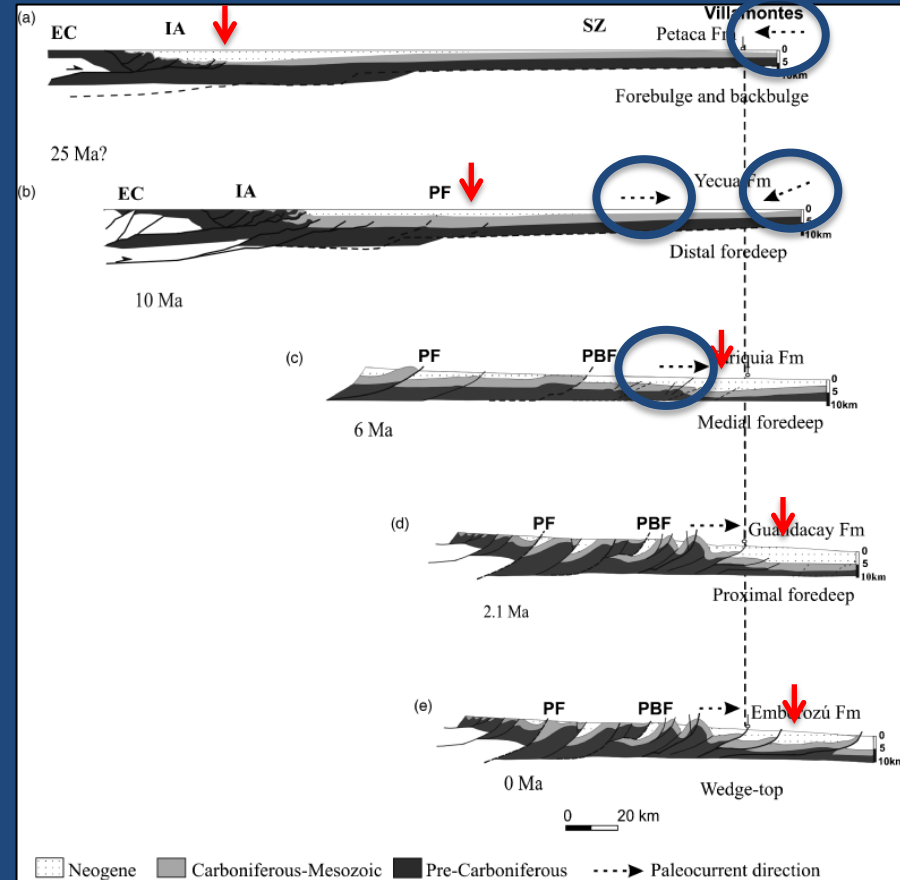
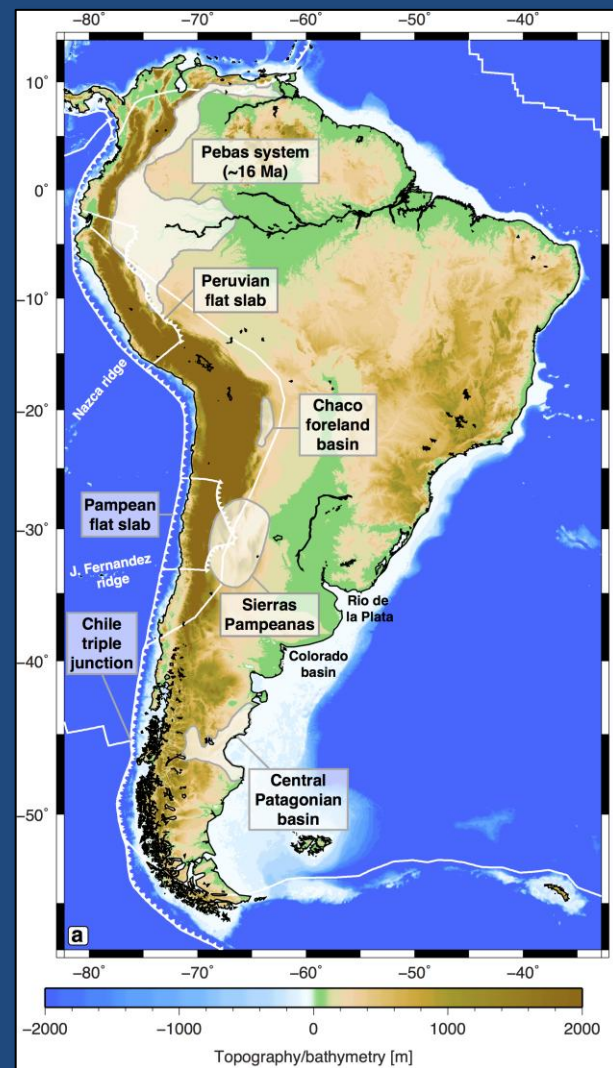
Anomalousy
high
low
regions

Pebas system and reversal of the drainage direction of the Amazon River

*Shephard
et al.
(2010)*

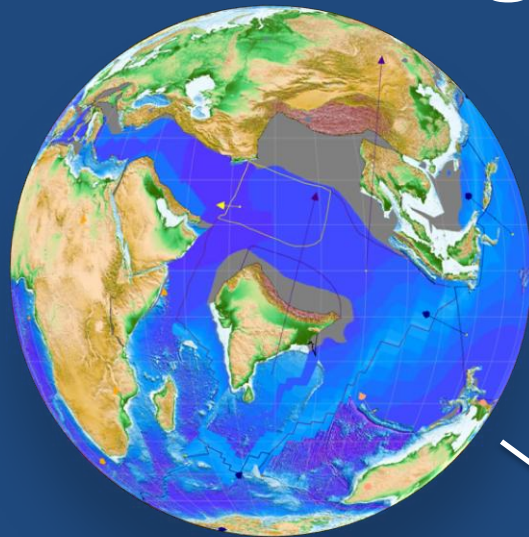


Migration of the depocentre of the Chaco Basin



Uba
et al.
(2006)

Forward global mantle flow models



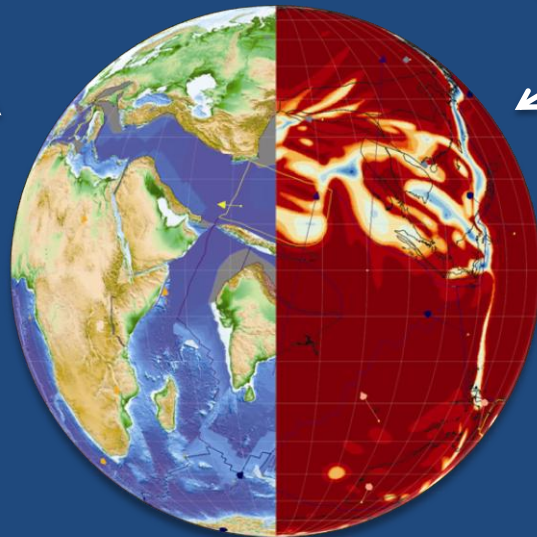
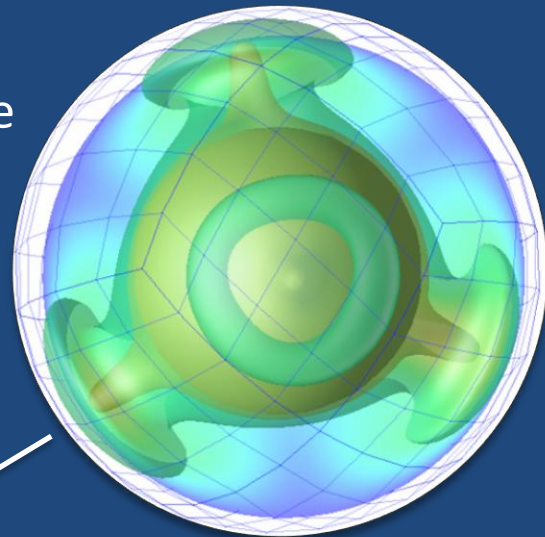
GPlates

Global reconstructions back to 230 Ma with continuously closing plate polygons (velocities; plate boundaries)



CitcomS

- 3D spherical finite element mantle convection code
- Modified for data assimilation



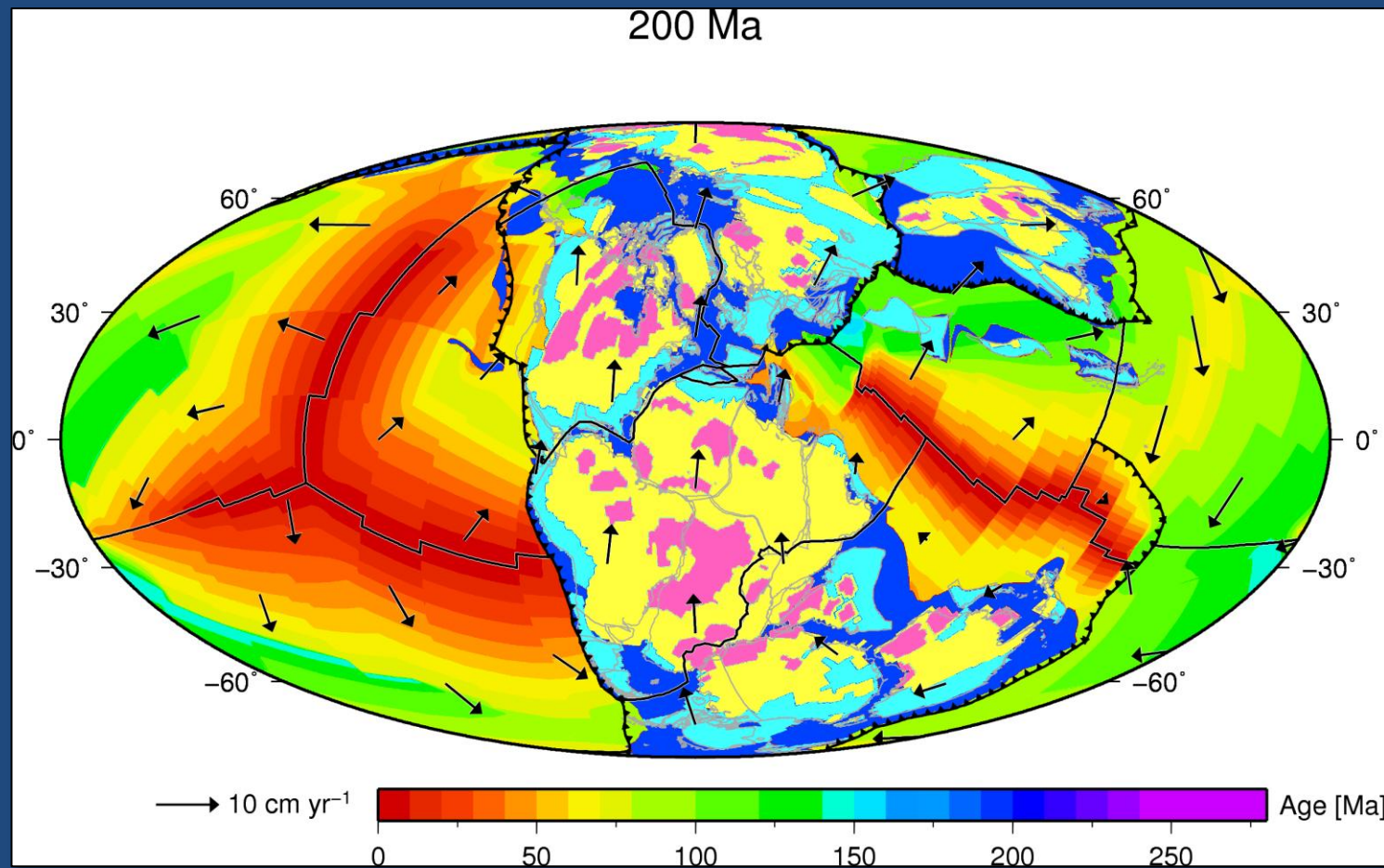
13 x 10⁶ points
~ 40,000 CPU hours/model

Forward global mantle flow models

*Seton
et al.
(2012)*

*Gurnis
et al.
(2012)*

*Artemieva
(2006)*

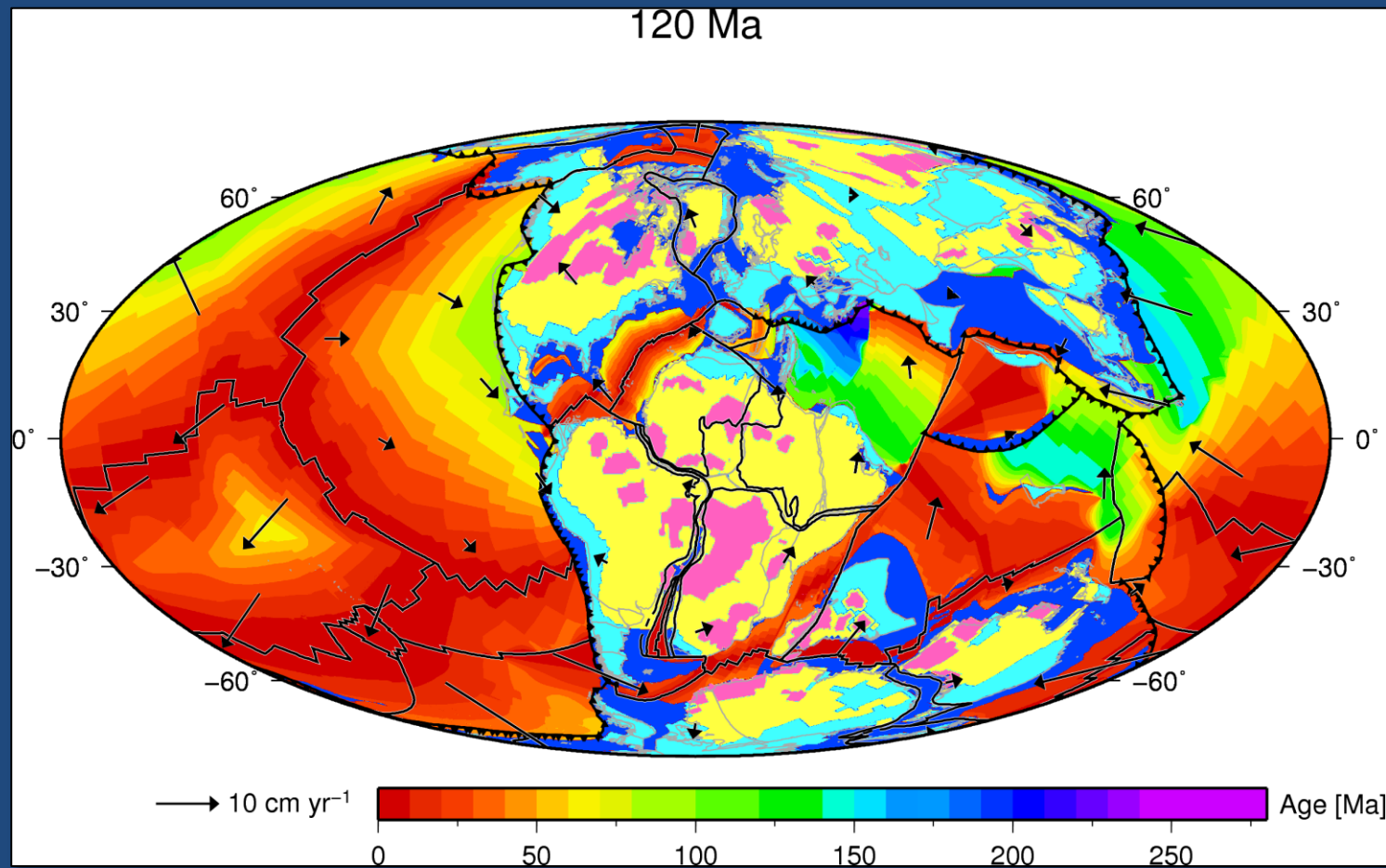


Forward global mantle flow models

*Seton
et al.
(2012)*

*Gurnis
et al.
(2012)*

*Artemieva
(2006)*

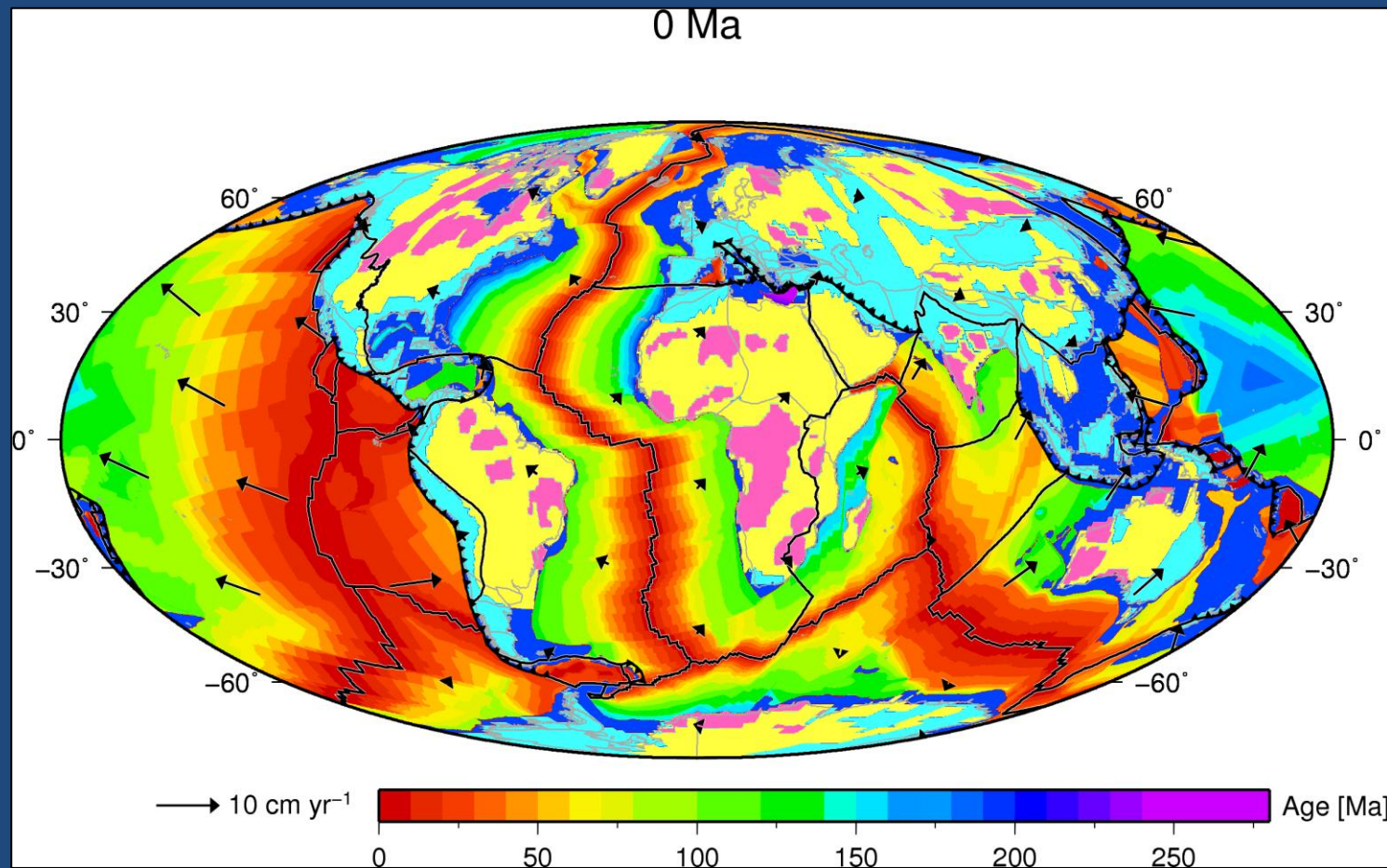


Forward global mantle flow models

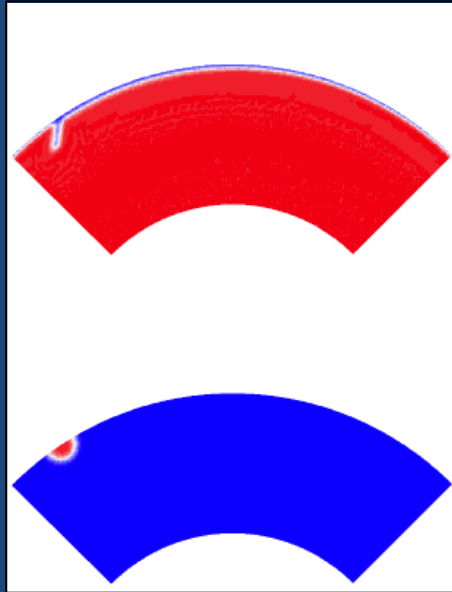
*Seton
et al.
(2012)*

*Gurnis
et al.
(2012)*

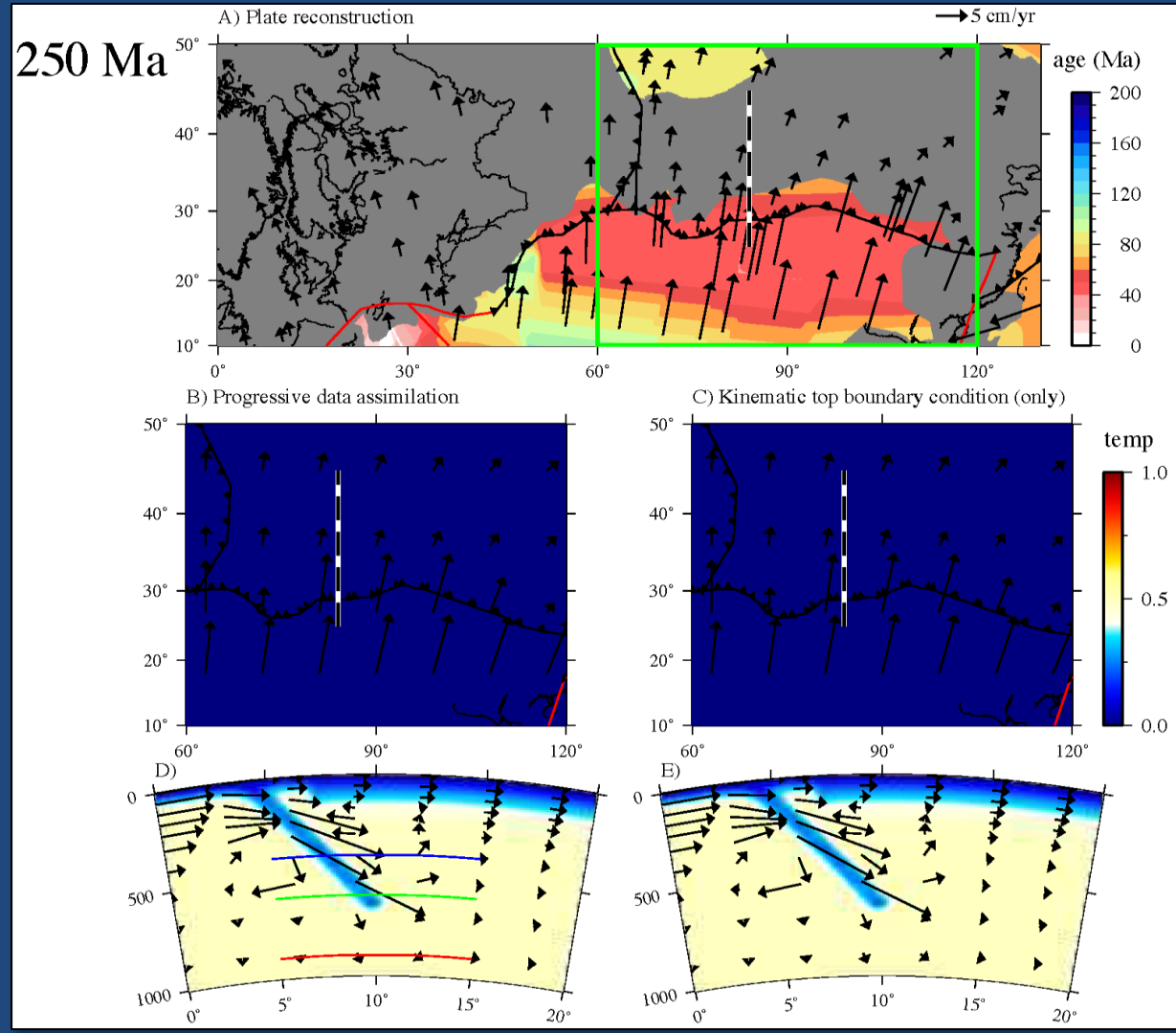
*Artemieva
(2006)*



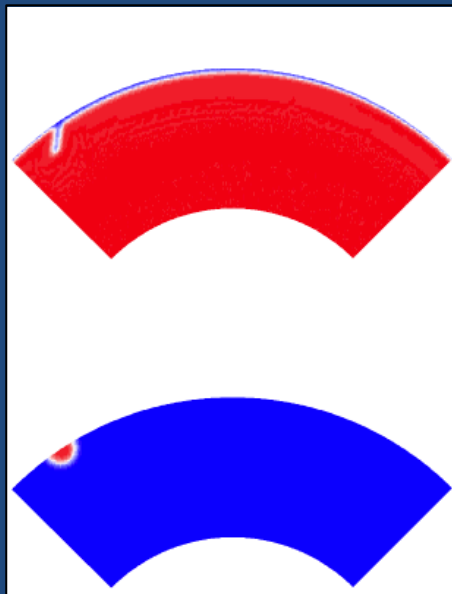
Progressive data assimilation



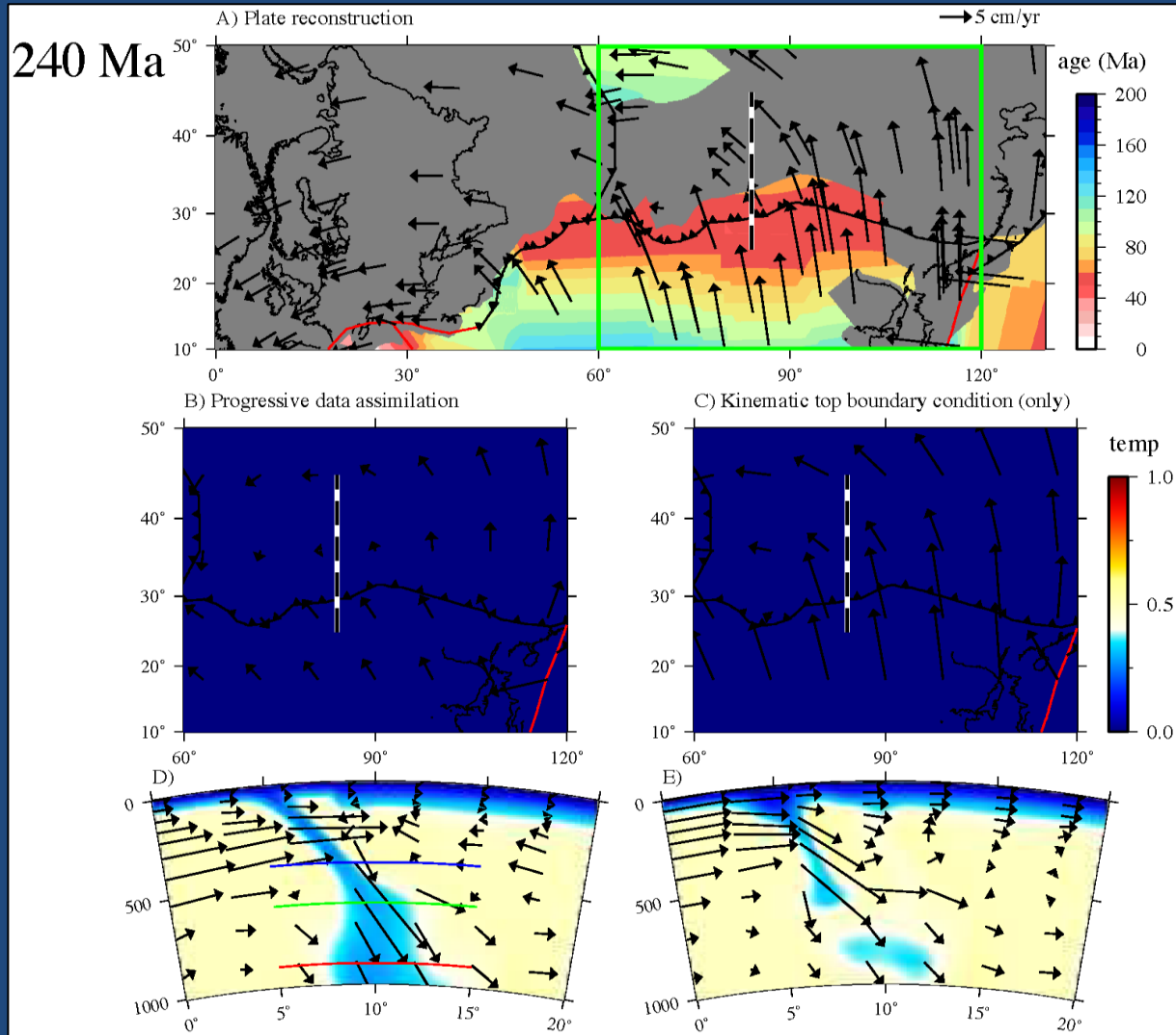
*Bower et al.
(2015)*



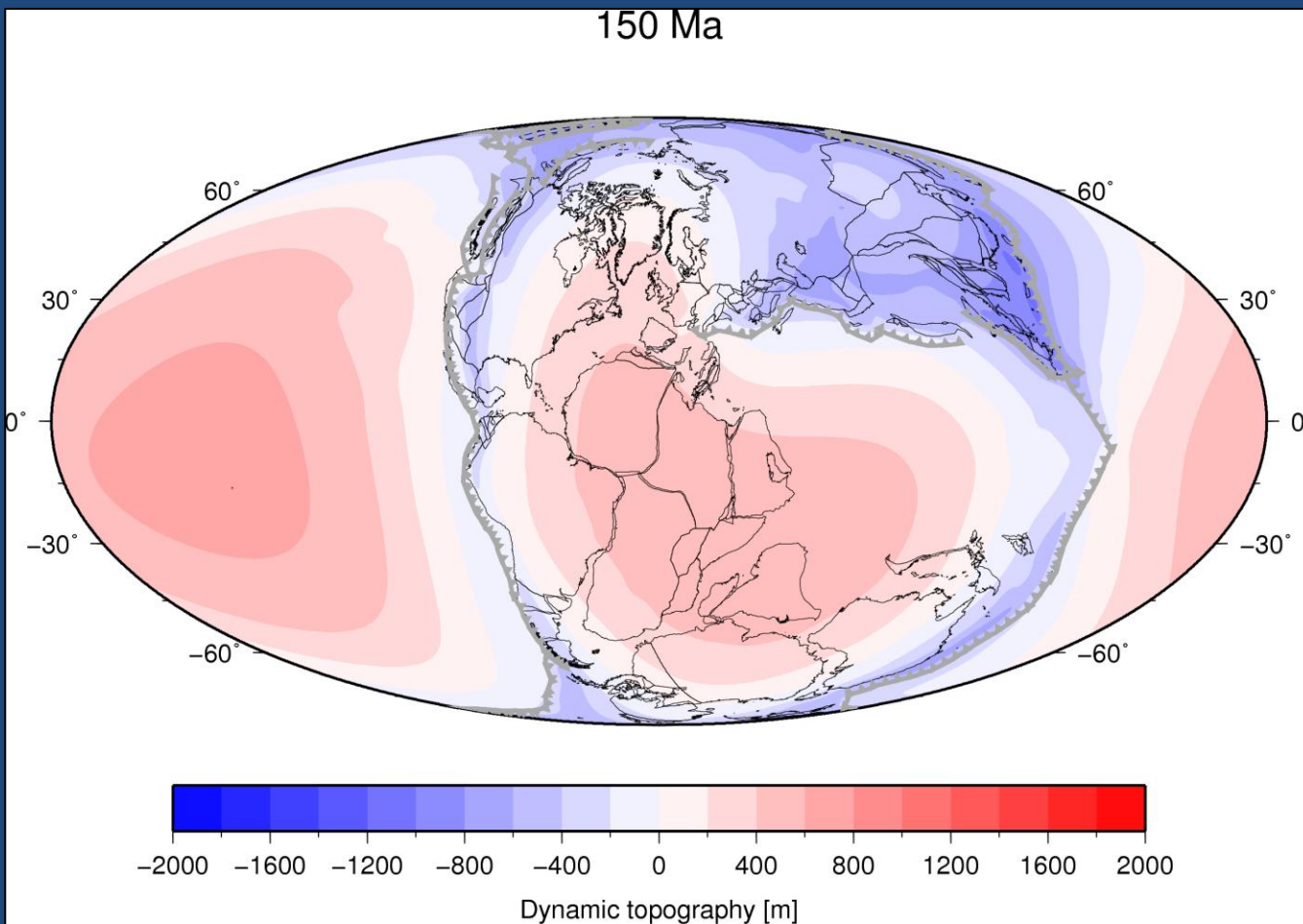
Progressive data assimilation



Bower et al.
(2015)



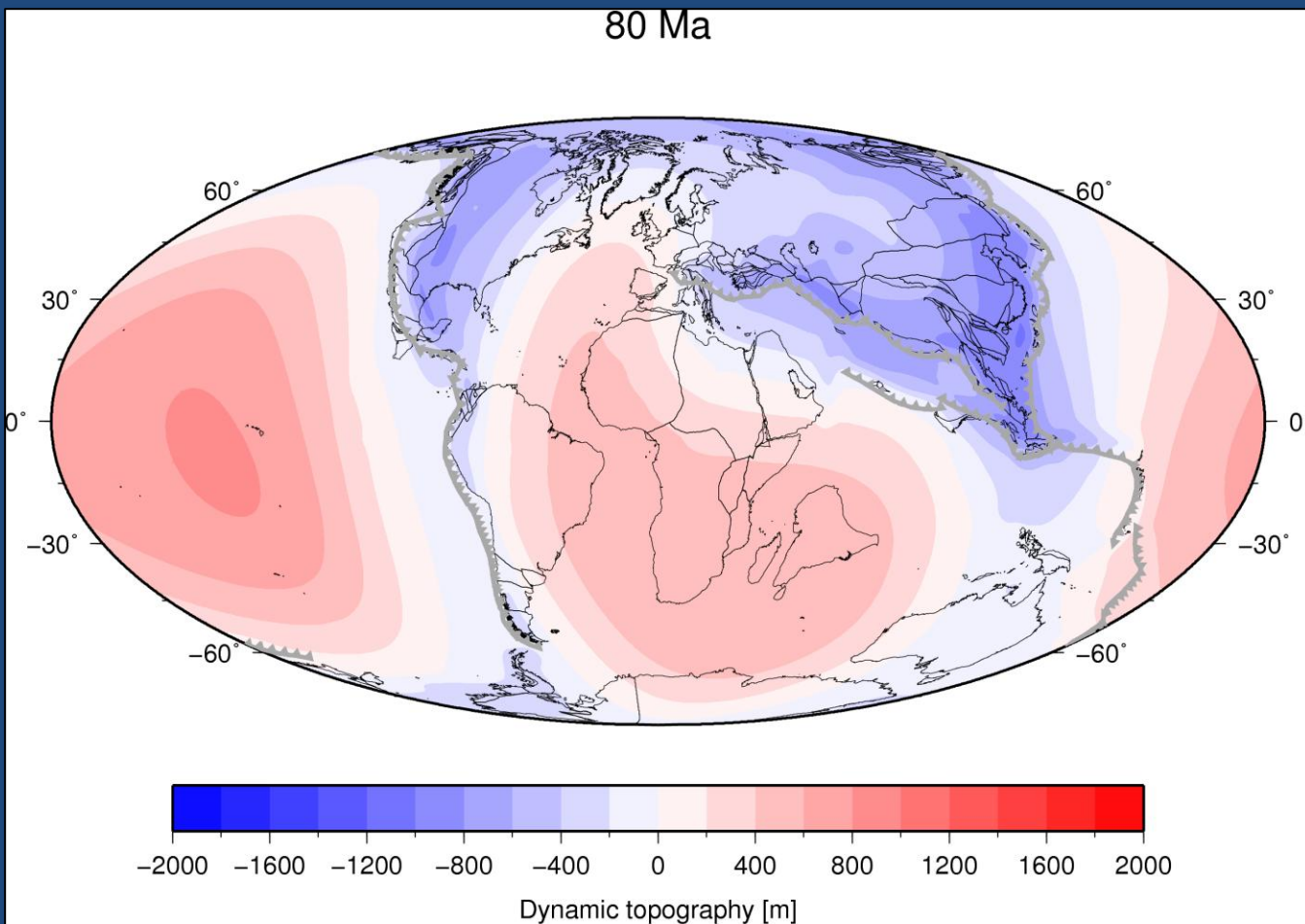
Evolution of global dynamic topography



Dynamic topography

- at the surface,
- from flow deeper than 350 km;
- free-slip boundary condition

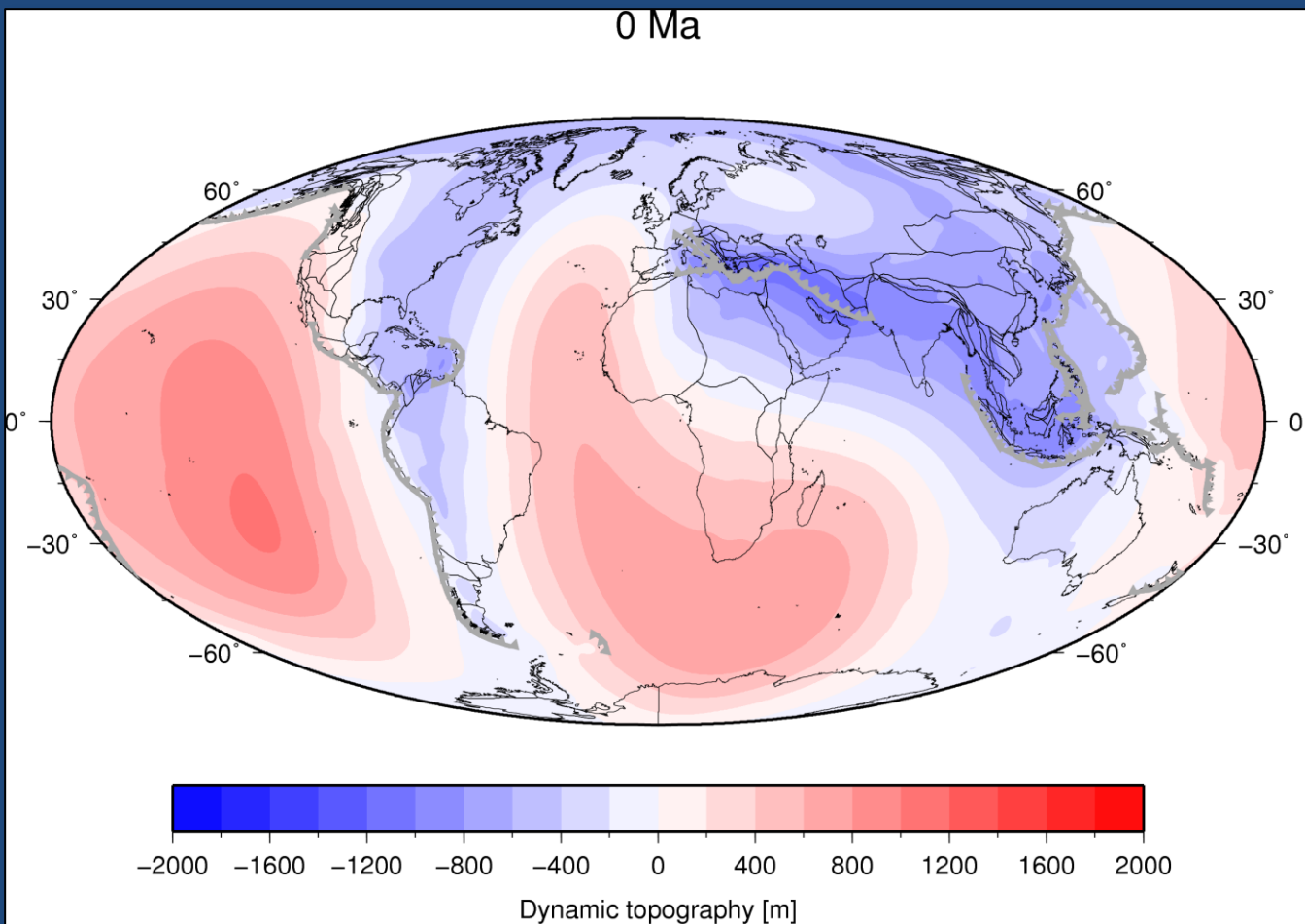
Evolution of global dynamic topography



Dynamic topography

- at the surface,
- from flow deeper than 350 km;
- free-slip boundary condition

Evolution of global dynamic topography



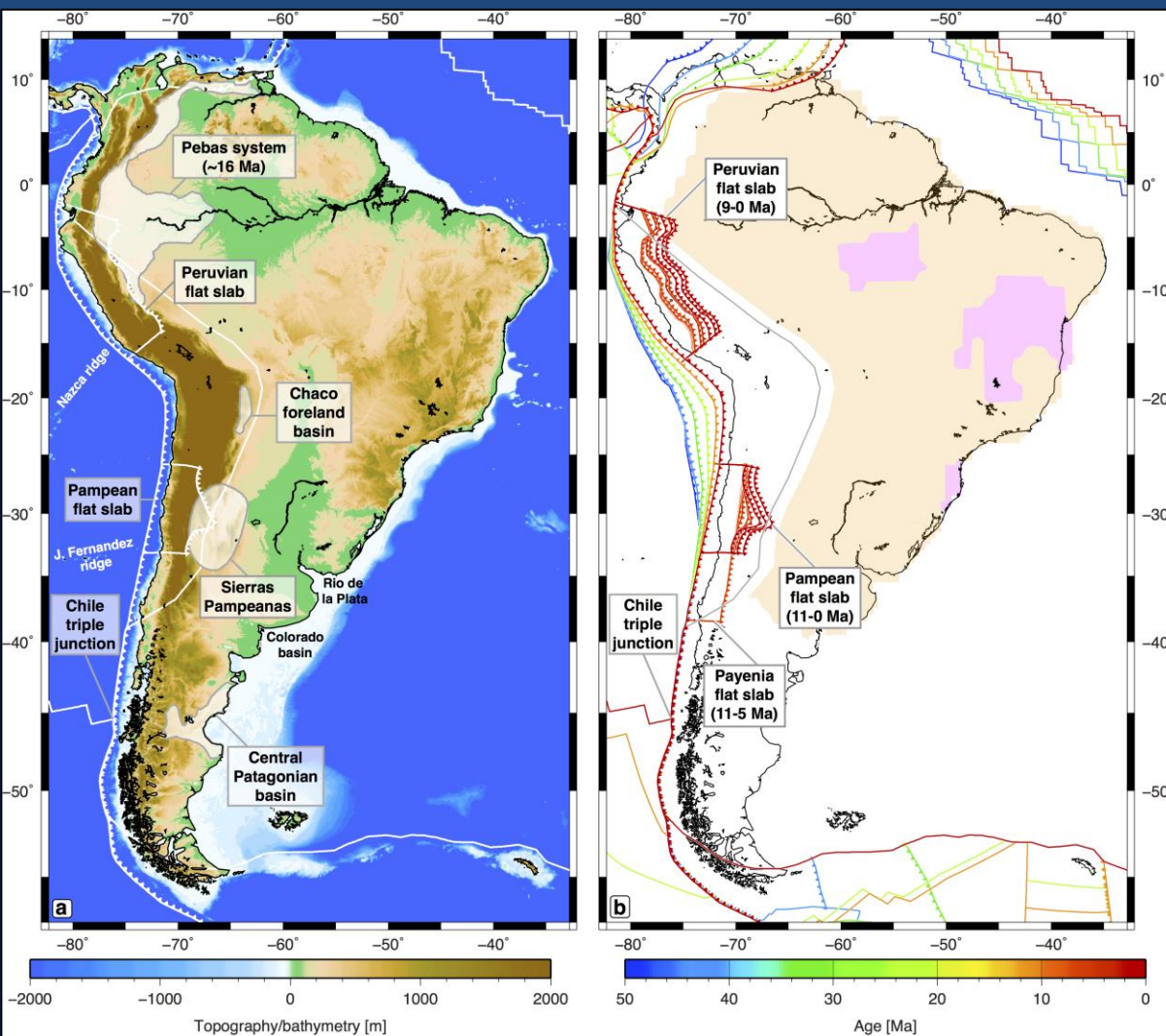
Dynamic topography

- at the surface,
- from flow deeper than 350 km;
- free-slip boundary condition

South American subduction history

- Andean deformation
(*Arriagada et al., 2008*)
- Flat-slab subduction
(*e.g. Ramos & Folguera, 2009*)

Flament et al. (2015)

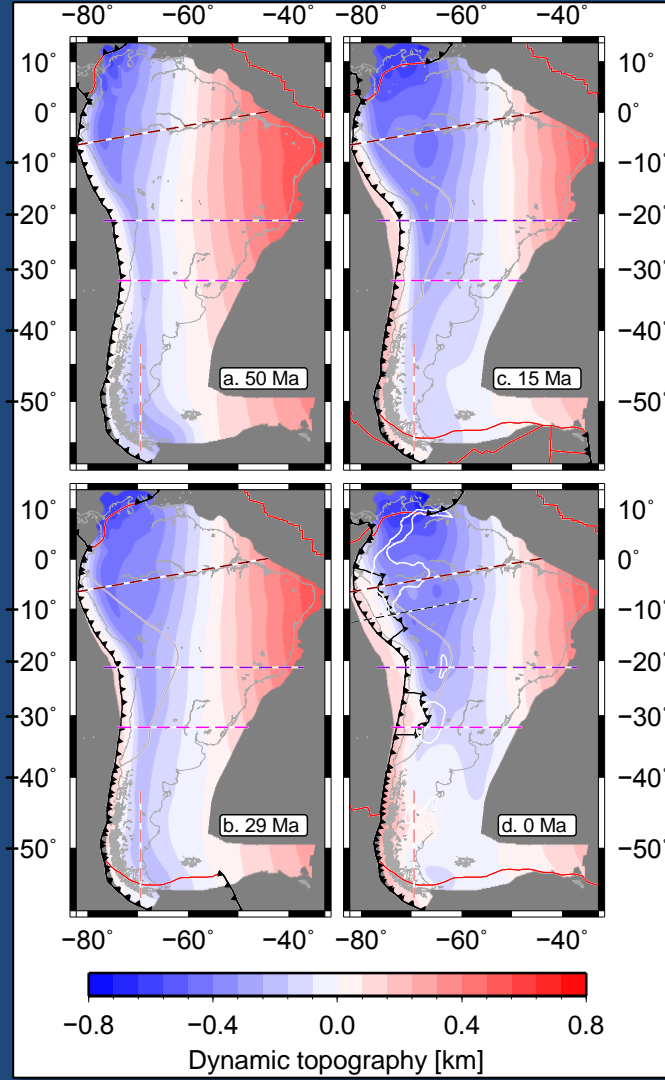


Predicted South American dynamic topography

Eastward propagation of a dynamic topography low

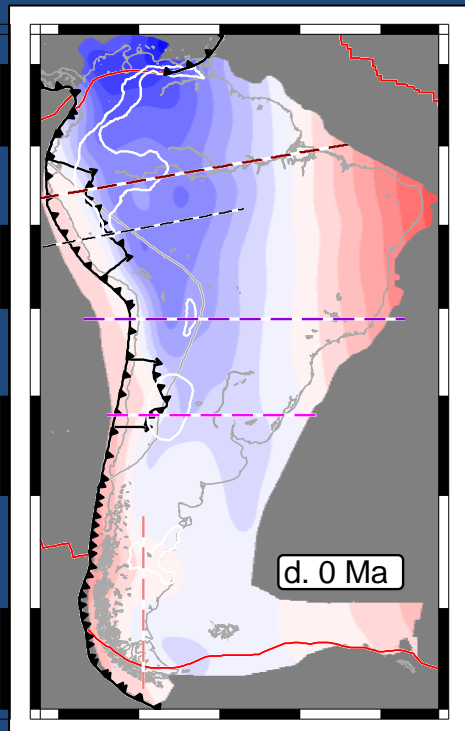
Effects of:

- Uplift associated with flat slabs
- Trench migration under the Central Andes



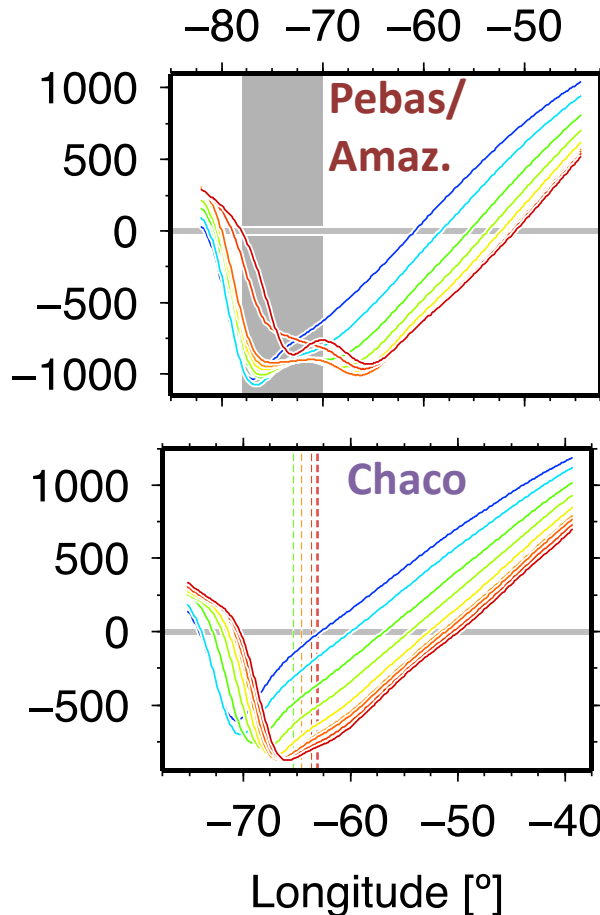
Flament et al. (EPSL, 2015)

Predicted South American dynamic topography

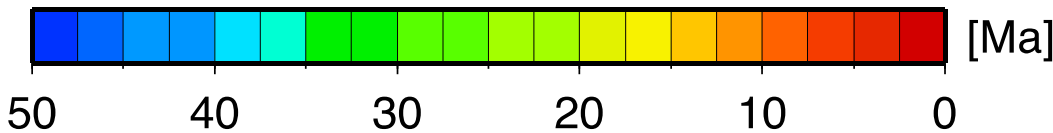
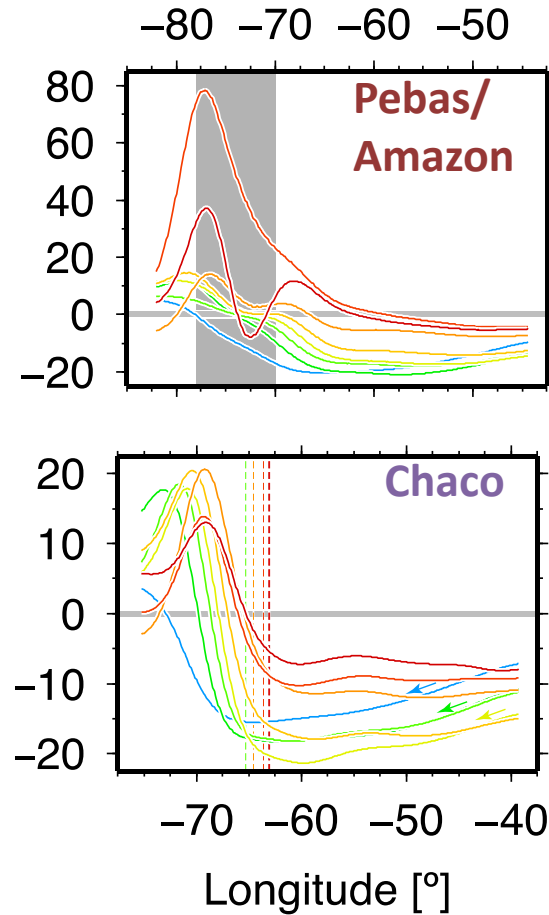


Flament et al. (EPSL, 2015)

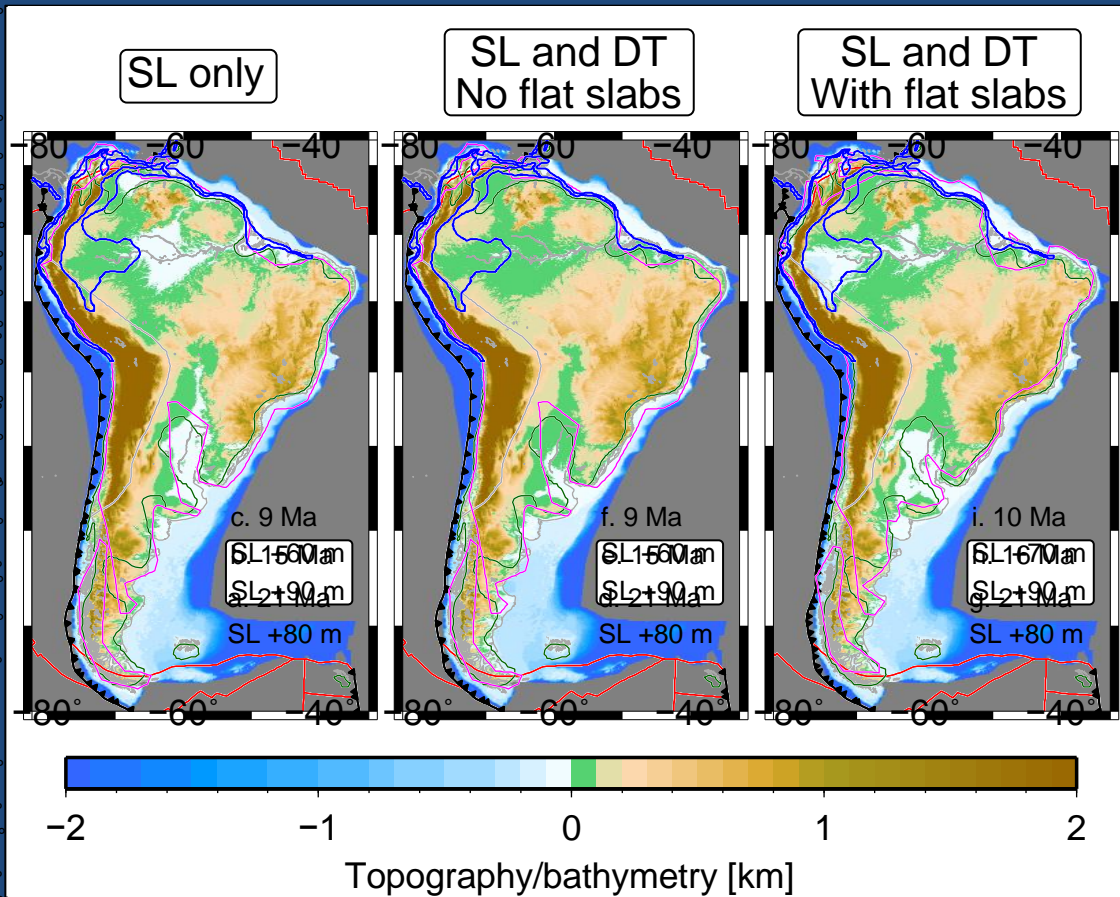
Dynamic Topography [m]



Rate of dynamic topography change [m Myr⁻¹]



Predicted and “observed” Miocene South American flooding



Simple model
paleogeography:

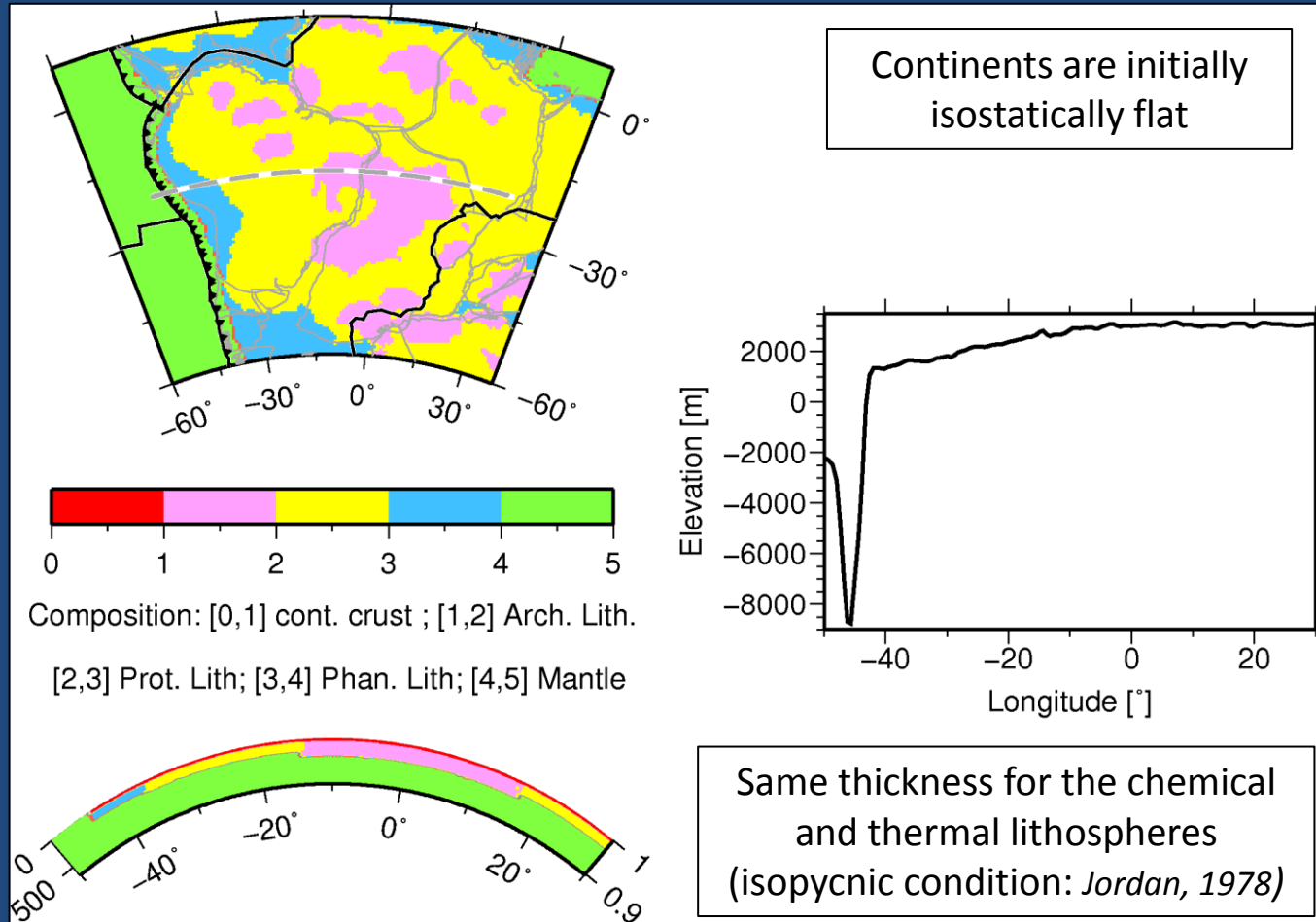
- long-term sea level change (*Haq et al., 1987*)
- changes in dynamic topography

Paleoshorelines (*Smith et al., 2004; Golonka, 2009; Wesselingh et al., 2010*)

→ Both sea level change and flat slabs are required to match paleoshorelines

Flament et al. (EPSL, 2015)

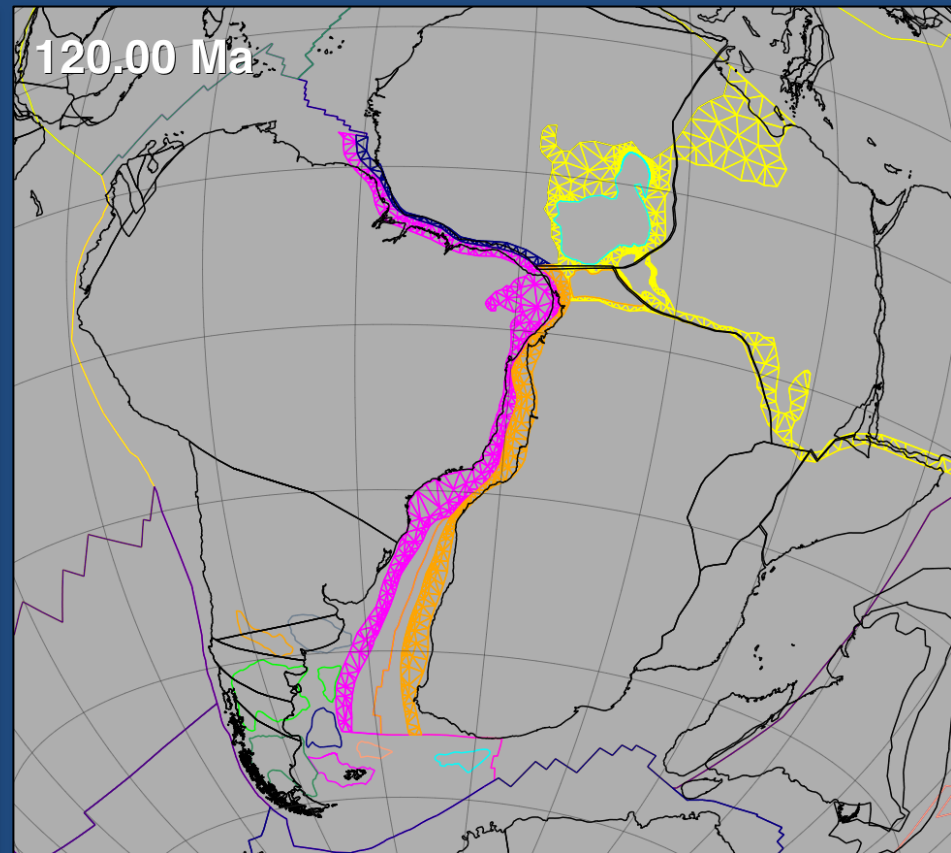
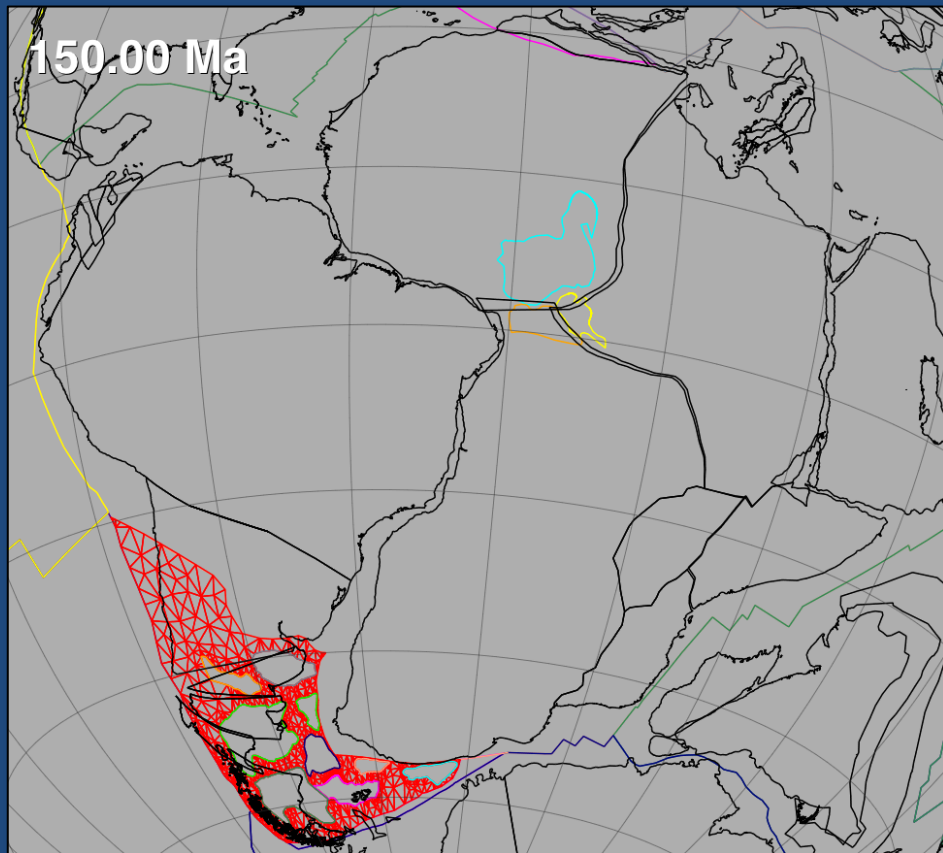
Modelling isostatic and dynamic topography



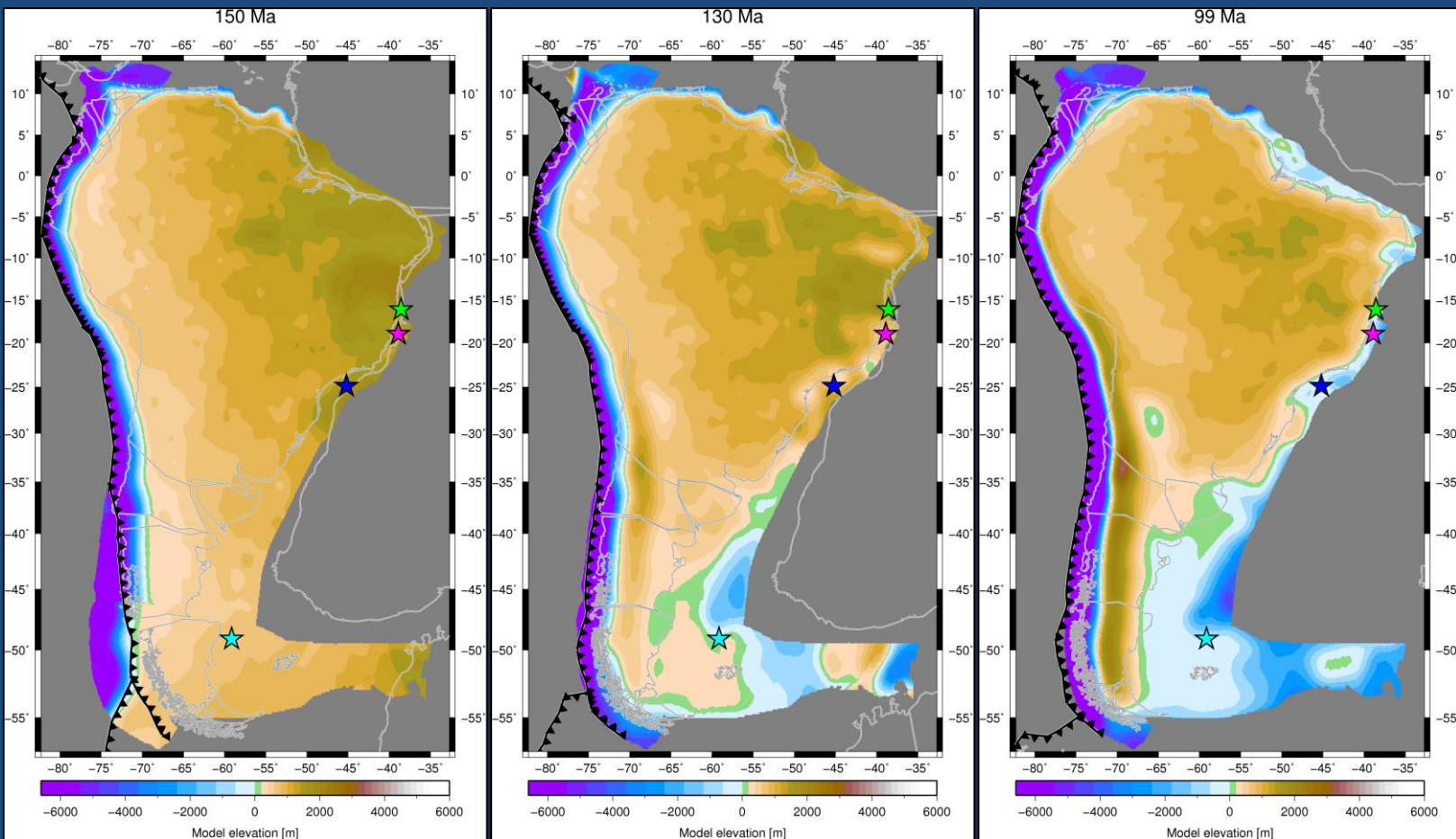
*Flament
et al.
(2014)*



Reconstruction with deformation



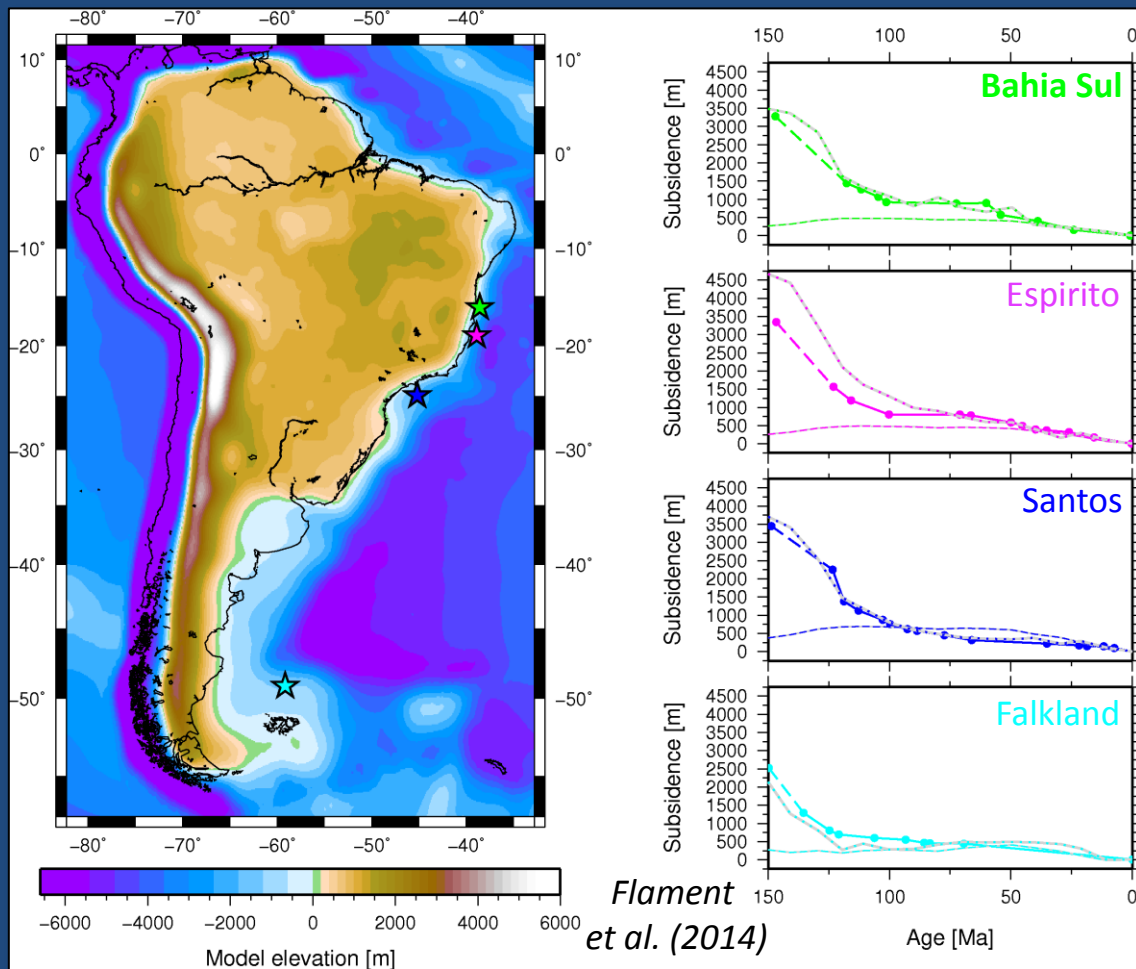
Evolution of South American *total model elevation*



Bahia Sul Basin
Espirito Santo Basin
Santos Basin
Chang et al. (1992)

North Falkland Basin
Jones et al. (2004)

Predicted and “observed” tectonic subsidence



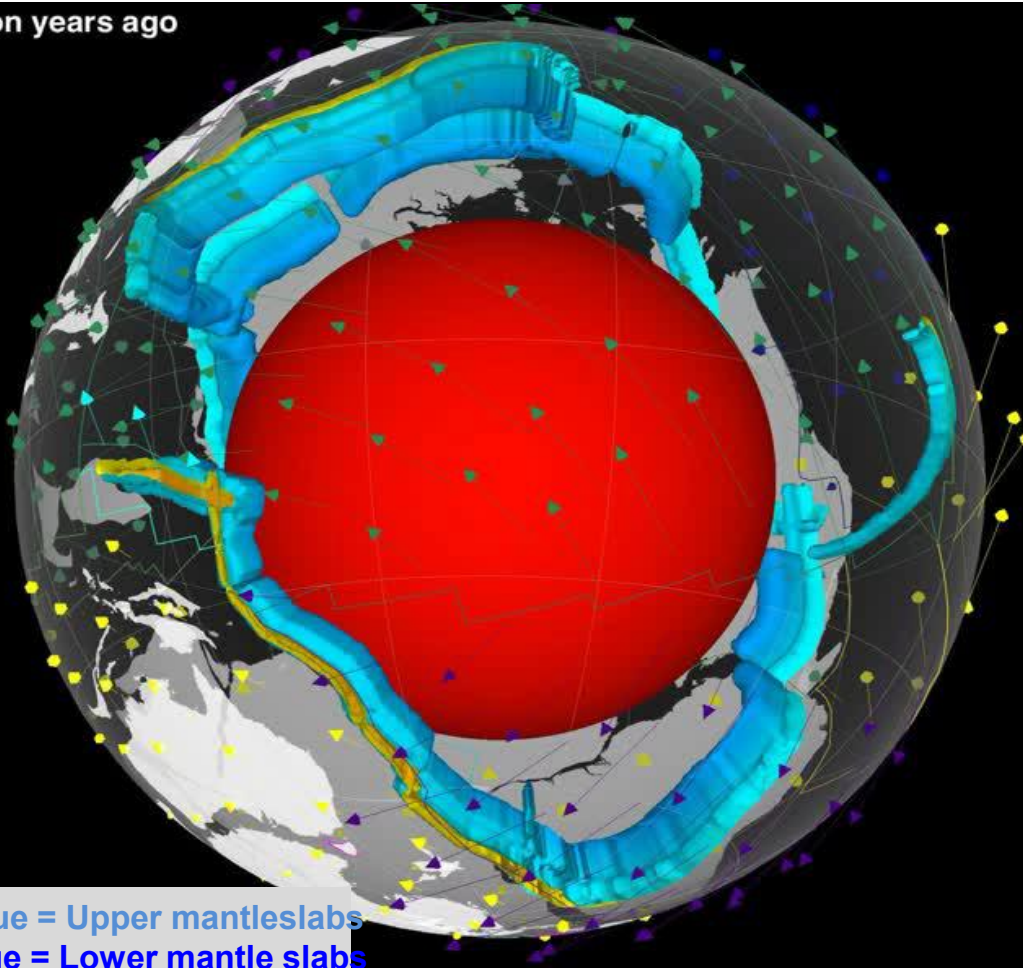
Chang et al. (1992)

- Rift dynamics captured to first order: rift phase followed by thermal subsidence: essentially the model of *McKenzie (1978)* in 3D, time-dependent

Jones et al. (2004)

- Accelerated tectonic subsidence since ~50 Ma corresponds to predicted dynamic subsidence

230 million years ago

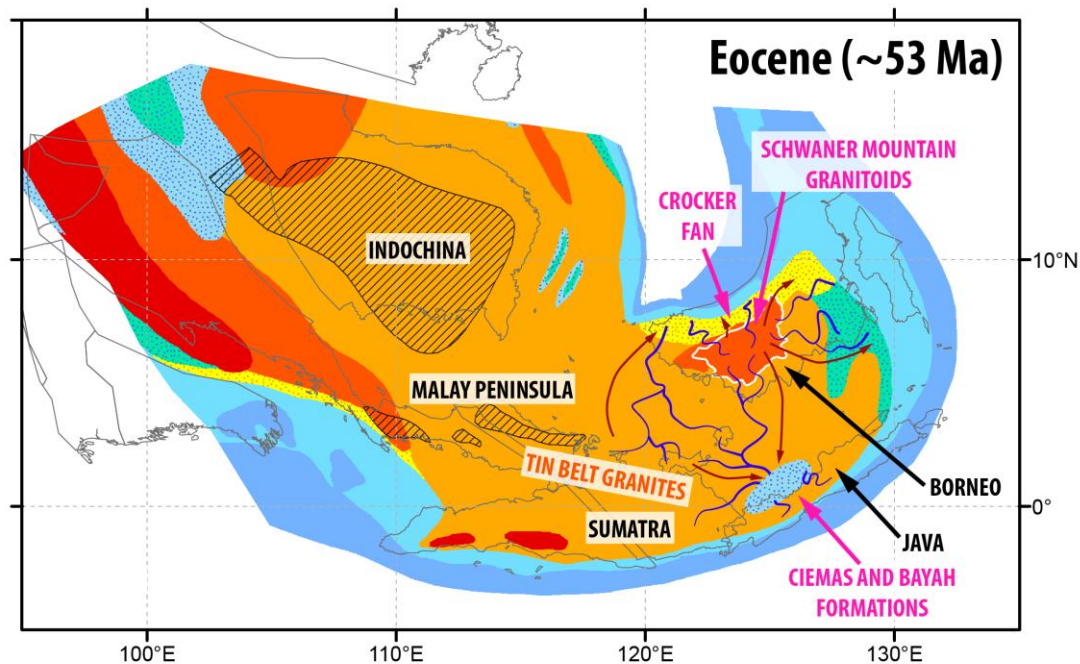


Light blue = Upper mantle slabs
Dark blue = Lower mantle slabs

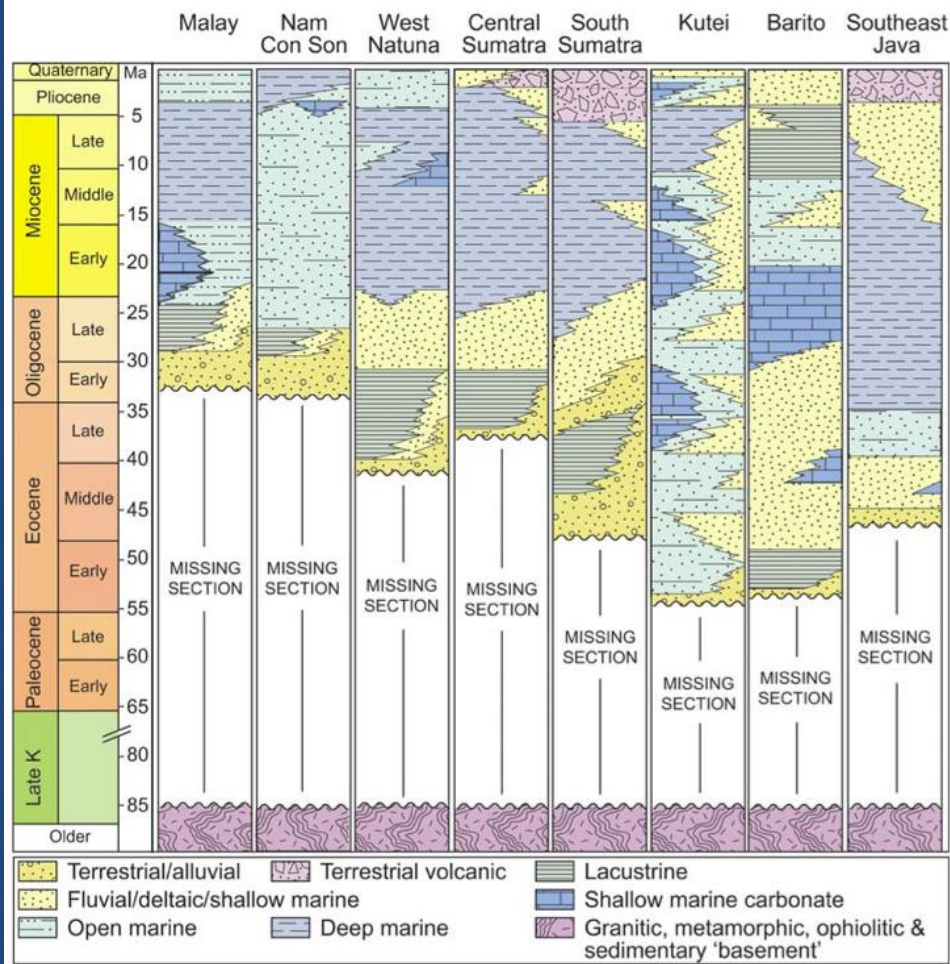
- Dynamic topography of SE Asia
- A complex, deforming region profoundly affected by subduction through time

Presenter's notes: Geodynamic simulation---We can then use these plate velocities to drive mantle convection in a global spherical mantle shell, allowing slabs to sink and interact in a more realistic manner. This allows us to track subducted material through time and then account for it at present day – as well as test alternative tectonic scenarios when the surface geology is vague. The first key here is the colours. Red represents upwelling hot mantle, and blue represents sinking slab material. Notice that New Guinea and SE Asia are overriding this 'slab burial ground', resulting in strong dynamic subsidence and flooding of these regions since ~30 Ma.

Late Cretaceous to Eocene Sundaland unconformity



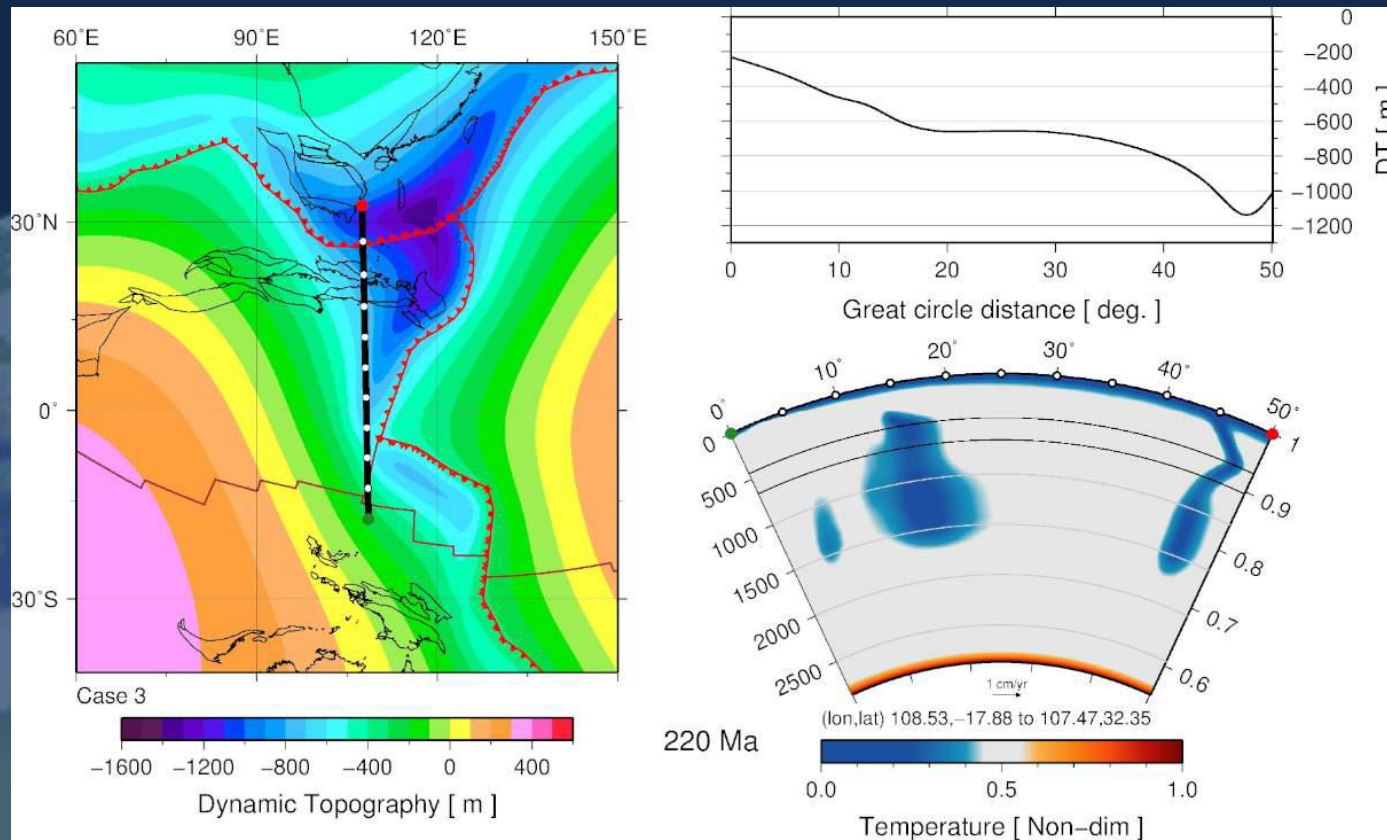
- Missing sedimentary sections across wide region Late K to Eocene
- Regional unconformity
- Unlikely to be from “tectonic” topography or from flexure alone



Regional unconformity, Clements et al. (2011)

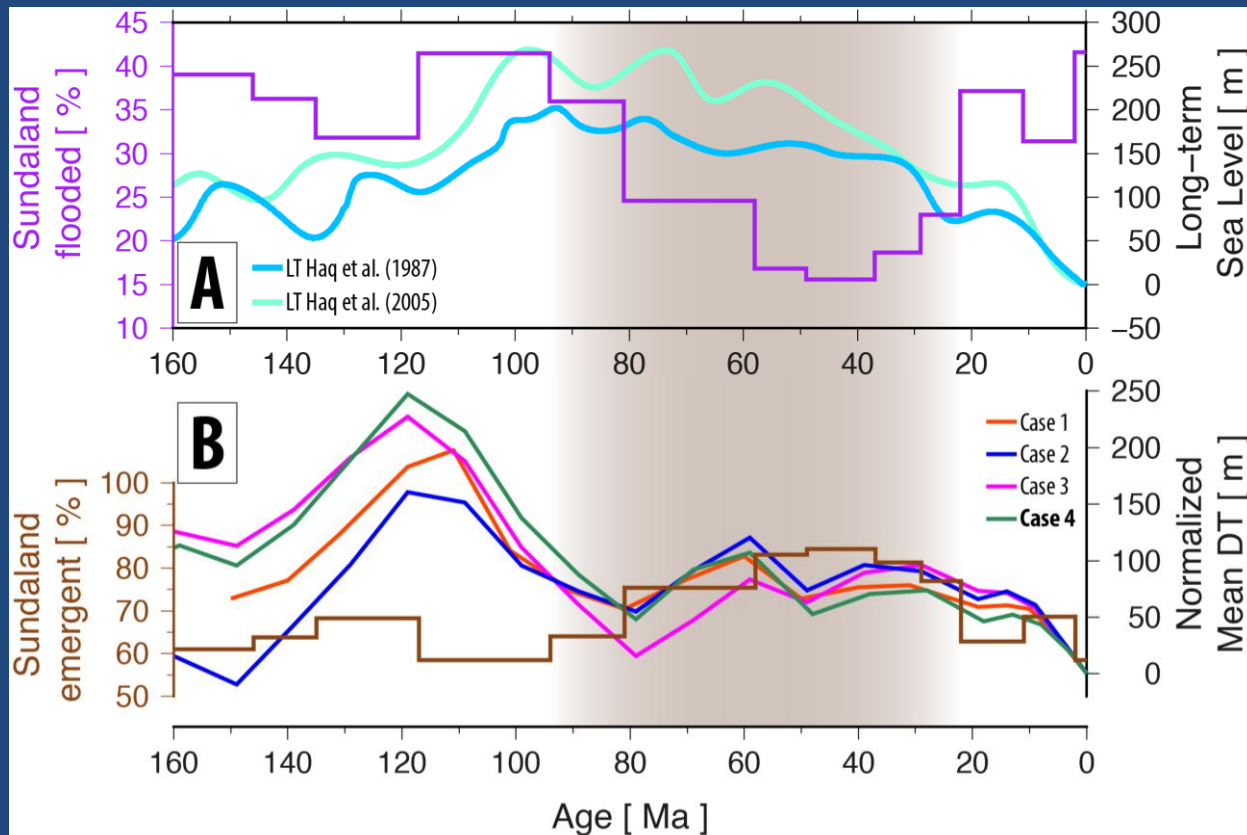
Presenter's notes: Sundaland Case Study---Presented here is a case study of Sundaland emergence and subsequent flooding, which we argue is controlled largely by dynamic topography, which is topographic responses to mantle convection. There is an enigmatic Late Cretaceous to Paleocene regional unconformity, and Clements et al. (2011) proposed that the accreted Gondwana fragments may have choked a subduction zone, resulting in dynamic uplift due to the absence of subduction. The emergence and flooding pattern is also supported by paleogeography.

Geodynamic model of dynamic topography



Presenter's notes: Dynamic Topography: We can then interrogate the geodynamic model, both vertical slices through the mantle, but also the surface dynamic topography. Here is a slice across Sumatra, showing the sinking slab material.

Dynamic topography and flooding/emergence

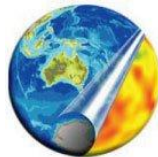


Presenter's notes: Results---What we find is that the eustatic sea level trends are anti-correlated to the flooding since ~30 Ma, where the falling sea level curve would suggest a tendency toward emergence. Instead, the pattern of emergence, resulting from a hiatus in subduction, and subsequent flooding since ~30 Ma from re-established subduction, suggests that dynamic topography had a dominant control on the flooding of Sundaland. We need to look at New Guinea in this context as well, as it seems it could provide the mechanism for widespread contemporaneous flooding and carbonate deposition since ~30 Ma (Oligocene times).

Conclusions

- Ability to jointly investigate mantle flow and crustal deformation over 100's and 1000's of km and 10's of Myr
- Direct comparison between predicted and observed tectonic subsidence/uplift
- **Exploration Geodynamics:** An emerging exploration tool

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