Anomalous Thermal Indicators from the Authigenic Minerals in Upper Paleozoic Strata of the Michigan Basin*

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Abstract

Indications of anomalous paleo-temperatures exist in strata of the Michigan Basin, USA. The thermal history of the basin was investigated through identification and analysis of authigenic components in two Upper Paleozoic strata, the Devonian Dundee Formation and Mississippian Marshall Sandstone. Formation conditions for diagenetic phases in both units were estimated through a variety of means and compared to conditions expected for a simple burial model. Authigenic clays identified in the Marshall via X-ray diffraction are expected to have formed at temperatures above what would have been produced by burial alone. Stable isotope and fluid inclusion analysis indicate the same for secondary carbonates in the Dundee. K-Ar age dating of Marshall authigenic illite provide ages of approximately 280 Ma. This age may reflect an illite formation period in the Permian or may be an integrated age representing multiple periods of illitization. Dolomite in the Dundee and illite in the Marshall are likely to have formed under hydrothermal conditions. The results of this study are part of a larger body of evidence that suggests episodic and extensive hydrothermal events have influenced the thermal history of the Michigan Basin. These events are due to the reactivation of faults in the Proterozoic rift that underlies the basin and maybe related to thermal convection within the igneous rift body or Appalachian tectonic events.

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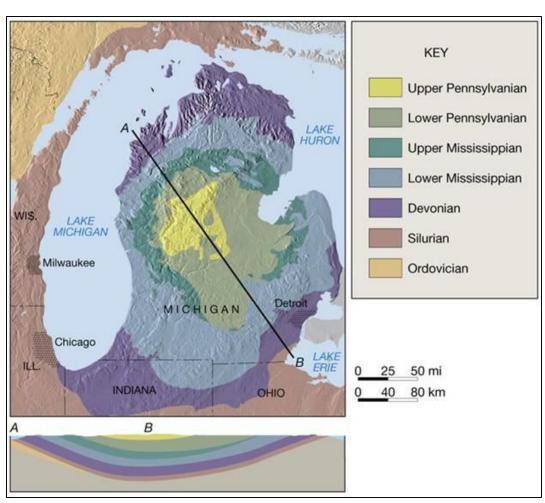


Importance of Thermal History

- To understand the evolution of an intracratonic sedimentary basin you must understand its thermal history
- The thermal history is an important control on the occurrence oil and gas resources in said basin.
- Unconventional resource exploitation depends upon in situ thermal alteration

Intro: Michigan Basin

- Circular intracratonic sedimentary basin.
- Up to 5km of sedimentary rock
- Most strata Paleozoic

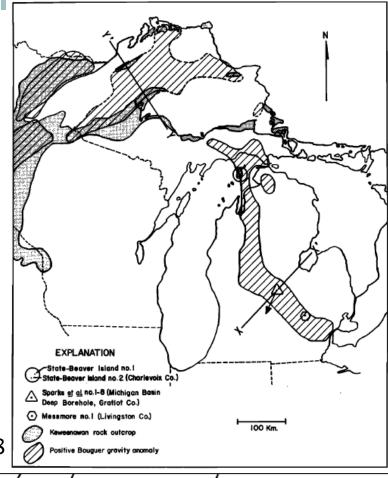


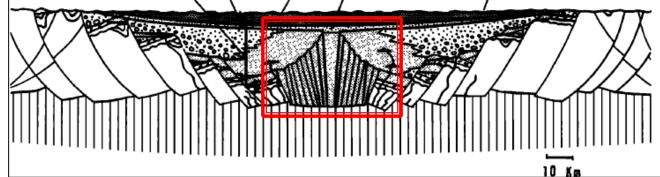
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Intro: Mid-Michigan Rift (MMR)

- Linear gravimetric and magnetic anomaly.
- Proterozoic Mid-Continent Rift System (Hinze et al., 1975).
- Related basement faults and fractures can extend into the overlying sedimentary strata (Prouty, 1988).

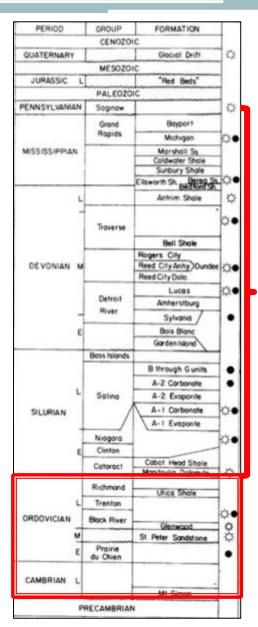
Fowler and Kuenzi, 1978





Intro: Michigan Basin Thermal History

- Michigan basin diagenetic minerals & organic material indicate temperatures significantly higher than presentday burial temperatures.
- Current geothermal gradient is 19-22° C/km (Vugrinovich, 1988)
- Organic material
 - Only Ordovician should be mature. (Nunn et al., 1984).
 - Mature source rocks and coal shallower (East et al., 2012).
- Diagenetic mineralization
 - St. Peter illite (Girard & Barnes, 1995)
 - Hydrothermal dolomitization (Prouty, 1988; Coniglio et al., 1994; Luczaj et al., 2006; Haeri-Ardakani et al., 2013; and others)
- Basin brine
 - Thermal event (Ma et al., 2009)
 - Mantle signature (Castro et al., 2009)



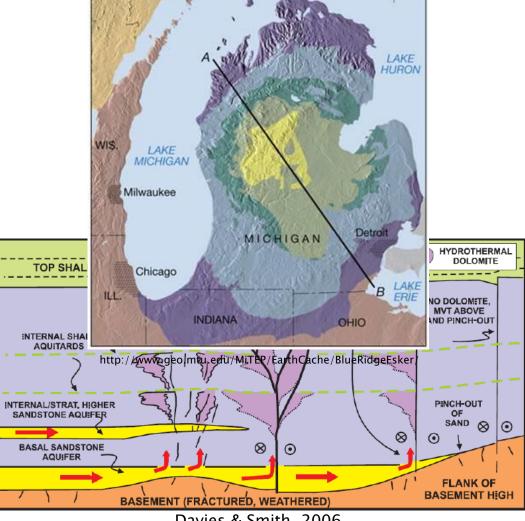
Intro: Models for Elevated Paleotemperatures

Burial

- 0.3–3 km add. sed.
- Late Paleozoic-Early Mesozoic
- Would produce basin shaped alteration

Hydrothermal alteration

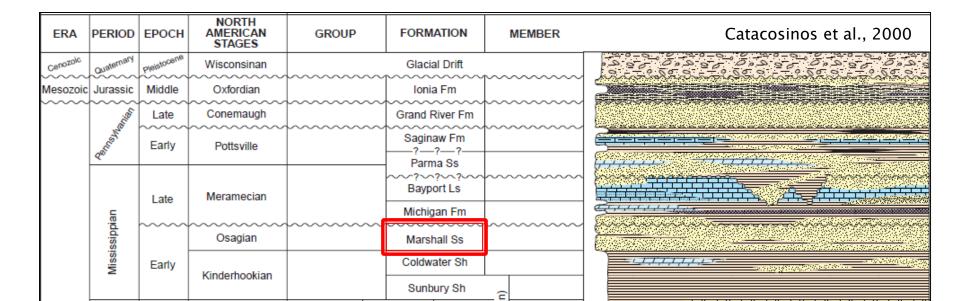
- Hot fluids migrating via basement rooted faults/fractures
- Shallow formations, deep brines
- Spatial relationship with structures
 - MMR
- Appalachian tectonic events



Davies & Smith, 2006

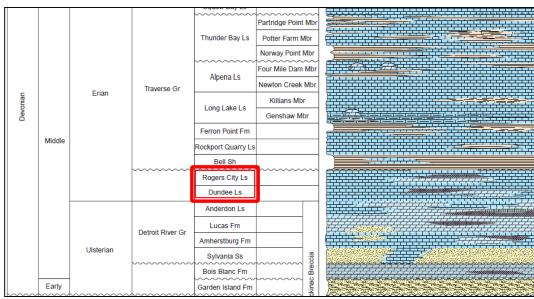
Intro: Marshall Sandstone

- Early Mississippian (~345 Ma) shallow marine sandstone (Harrell et al., 1991).
- Authigenic components indicate elevated paleotemperatures
 - Zacharias et al., 1993
- Why focus on the Marshall?
 - Promising preliminary work
 - · Illite cement in sandstone
 - Implications
 - Age data



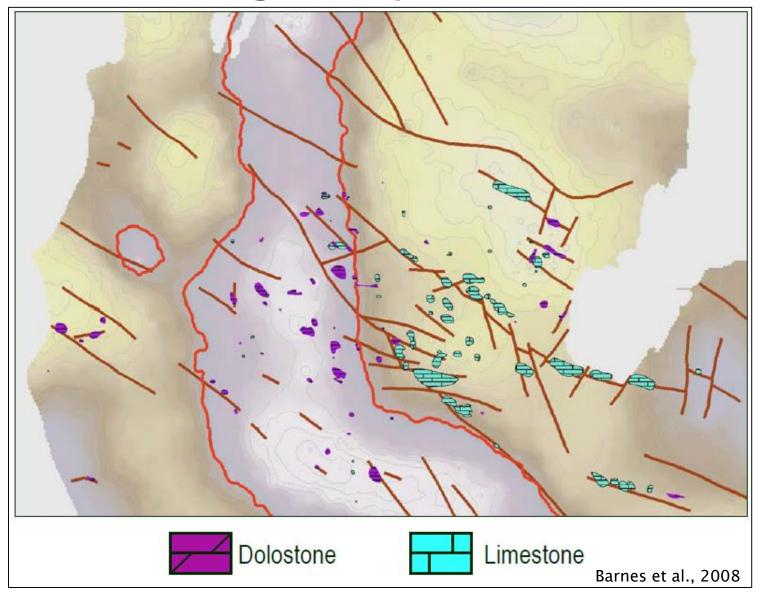
Intro: Dundee/Rogers City

- Middle Devonian carbonate mudstone to grainstone
 - Coastal sabkha to shallow marine platform environment. (Catacosinos et al. 2001).
 - Is locally dolomitized in many parts of the Michigan basin
- Secondary saddle dolomite indicate elevated formation temp.
 - 120–150° C at 1.2km (Luczaj et al., 2006)
 - Isotopic analysis also supports high temp. (Smith Williams, 2003)
- Why focus on the Dundee?
 - Evidence for elevated temperatures
 - · More complete analysis possible
 - Evidence of structural control of dolomitization (Barnes et al., 2008)



Catacosinos et al., 2000

Intro: Dundee/Rogers City



Distribution of dolomite and limestone oil fields in the Michigan basin

Research Questions

1. What authigenic minerals are found in the Marshall and Dundee Formations?

Petrographic Microscopy, XRD

2. What were the conditions of formation for these minerals?

Stable Isotopes, FIA, Literature

3. When did they form?

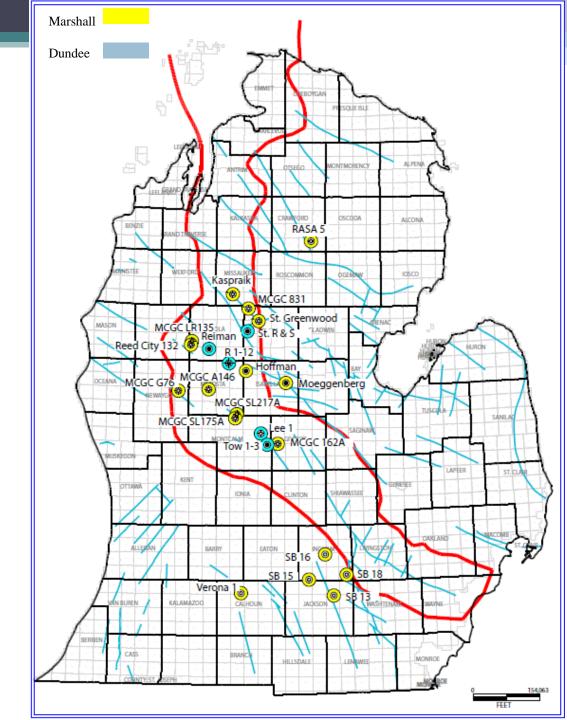
K-Ar Age Dating

- 4. How do the conditions of formation compare with a simple burial model?
- 5. Are anomalous paleo-temperature indicators spatially related to the rift and known structures?

Mapping

Results: Marshall Sampling

- 18 Marshall Wells
 - □ 14–477m
 - 15 petrographic thin sections
 - 51 XRD samples
 - 3 K-Ar age dating samples

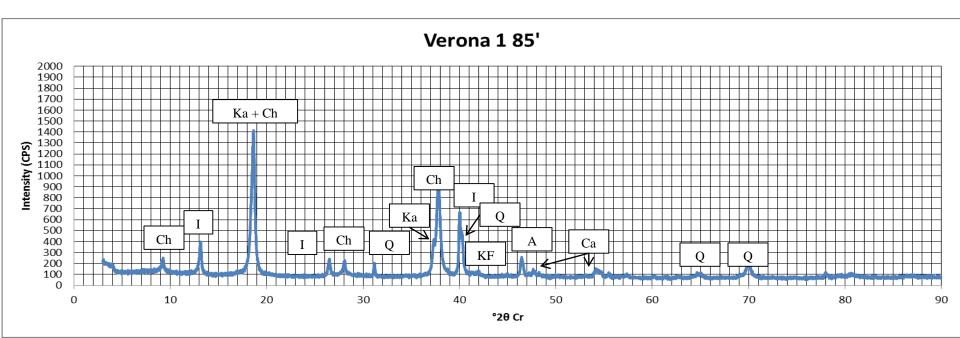


Results: Marshall Petrography

- 15 "clean" sandstone thin sections were examined
- Identified Authigenic Minerals:
 - Quartz
 - Overgrowths
 - K-feldspar
 - Overgrowths
 - Dissolved and replaced by clays
 - Carbonates
 - Zoned Ankerite Rhombs
 - Poikilotopic Calcite Cement
 - Clay
 - Grain rimming, pore filling, K-feldspar replacing
 - "Micaceous" birefringence and crystallinity
 - Moderate burial (>2-4km) diagenetic mineral assemblage

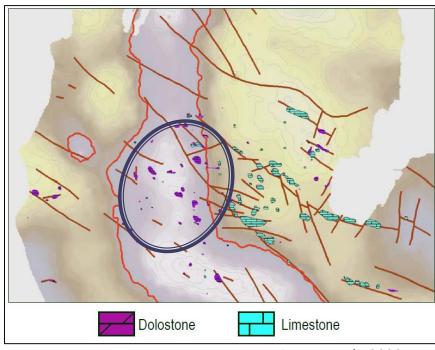
Results: Marshall XRD

- Neoformed minerals identified for 51 Marshall sandstone samples via XRD
 - All detrital particles (>2 μm) removed
- Clays were identified as illite, kaolinite, and chlorite
 - "Micaceous" clays identified as illite
 - Kaolinite replaced K-feldspar, illite replaced kaolinite
 - Common sequence in sandstone (Meunier & Velde, 2004)

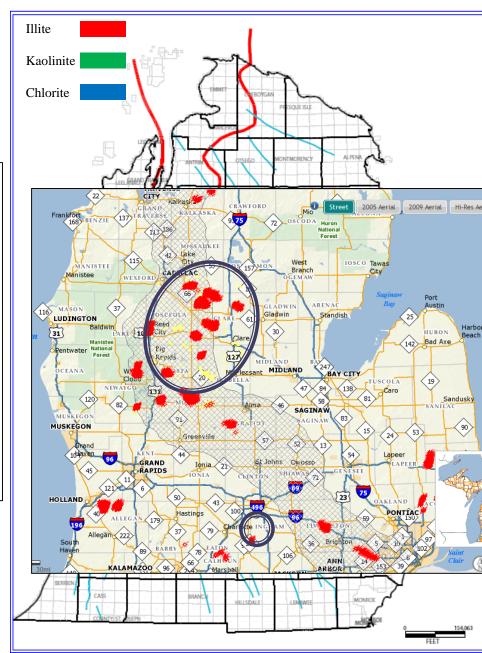


Depth (ft.)	Illite	Kaolinite	Chlorite	Depth (ft.)	Illite	Kaolinite	Chlorite		
	MGSC 831				REED CITY 132				
1432.5	X			1226	X	X	X		
1434	X			1232.3	X				
State Greenwood				RASA 5					
1462	X	X		520	X	X	X		
1466	X	X		531	X	X	X		
MCGC NH 162-A				Verona 1					
983.5	X	X		47	?	X			
1009.5	X	X	X	65	X	X	X		
1012	X	X	X	80	X	X	X		
1021.5	X	X	X	85 97	X	X	X		
	Hoffman				X	X	X		
1424	X	X		123.5	X	X	X		
1456	X		X		SB-18		,		
	Moeggenberg			389	X	X			
1273		X		415	X	X	X		
1274	X	X	X	472	X	X	X		
	MCGC SL-217-A			497	X	X	X		
1252	X	X		570	X	X	X		
1253	X	X	X	637	X	X	X		
MCGC A-146				SB-16					
1380	X			616.5	X	X	X		
1400	X	X			SB-15	<u> </u>			
	Kaspraik			321	X	X			
1520	X	X	?	327	X				
1521	X	X	?	353	X	X	X		
1523					SB-13	•	,		
1524	X		X	182	X	X	X		
1527	X	X		274	X	X	X		
1564	X	X		X=major x=minor					
	MCGC SL-175-A	•		74 1110	., 91 / 11111191				
1304	X	X	X						
1353.5	X	X	X	• Illite- 18/18					
	MCGC G-76	1		Illite-	18/18				
1139	X	X							
1161	X	X	X						
1248	X NGGGAN 12.5	X	 Kaolinite – 17/18 						
	MCGC LR-13-5	1		ιασιπ					
1250	X	X							
1279	X	X			1.4.1.0				
				 Chlori 	te- 14/18				
				J					

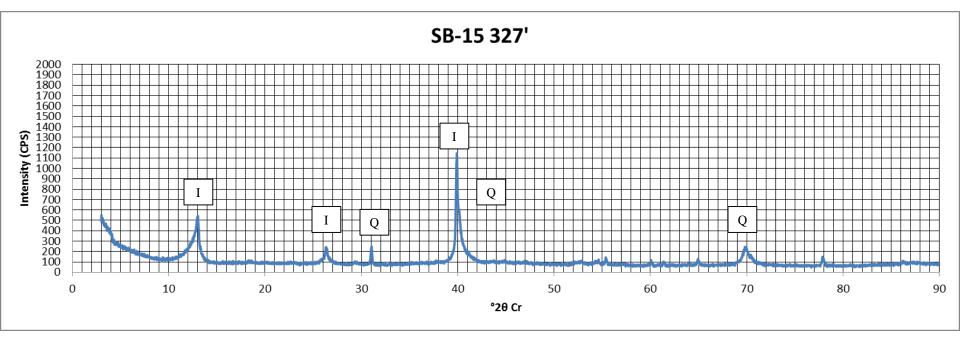
Results: Illite Distribution



Barnes et al., 2008



Results: Marshall XRD

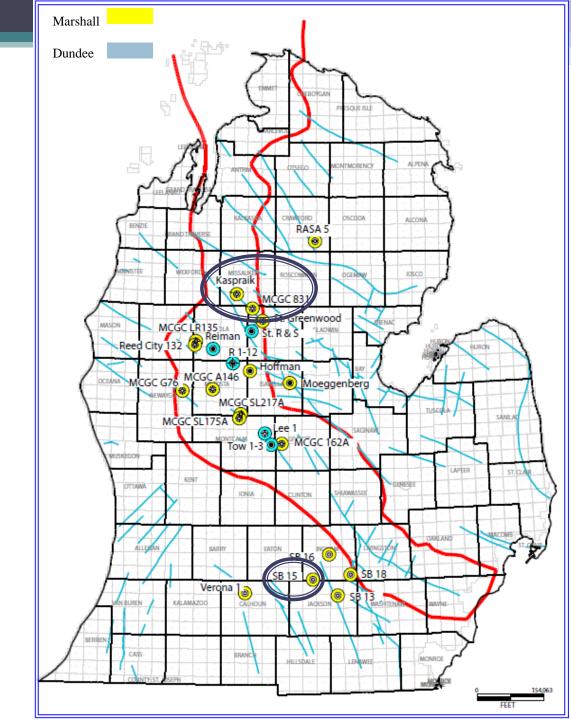


• Illite dominated XRD sample used for K-Ar age dating

Results: K-Ar Age Dating

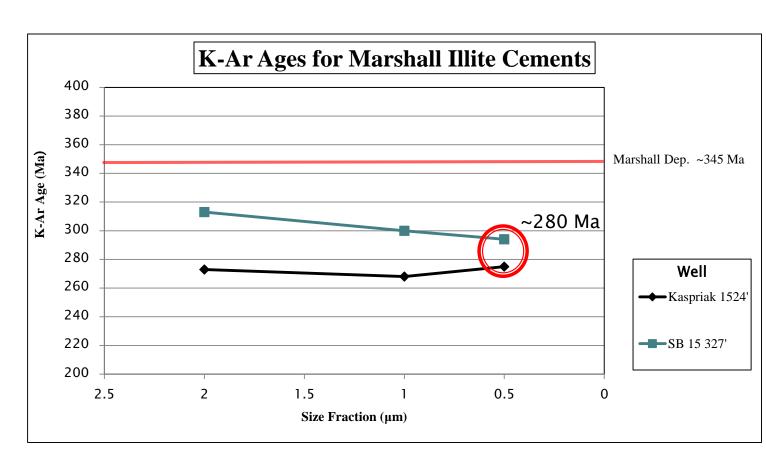
3 Samples

- Kaspraik 1524'
- MGSC 831 1432.5'
- SB 15 327'



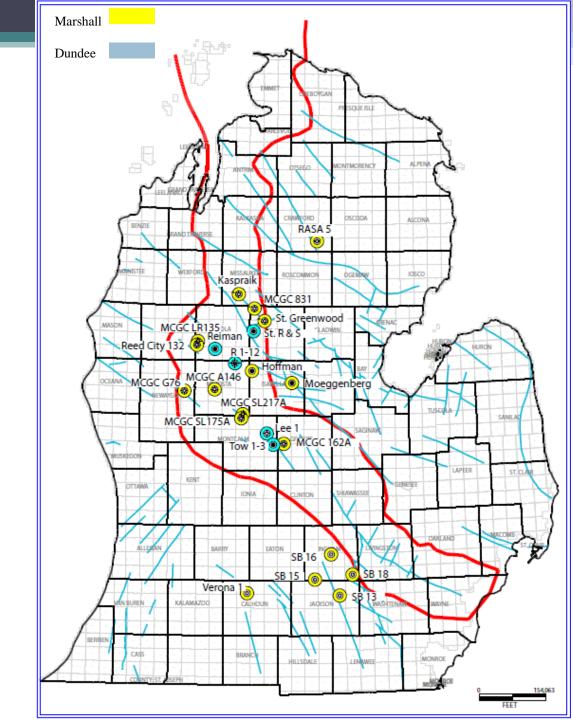
Results: K-Ar Age Dating

- Kaspraik275–268 Ma
- SB 15 313 294 Ma
- MGSC 831 –
 No Results
- Authigenic
 - Integrated?



Results: Dundee

- 5 Dundee Wells
 - 973-1203m
 - 11 thick sections
 - 203 fluid inclusion measurements
 - 37 stable isotope samples
 - Data for 33 stable isotope samples provided by Dr. David Barnes

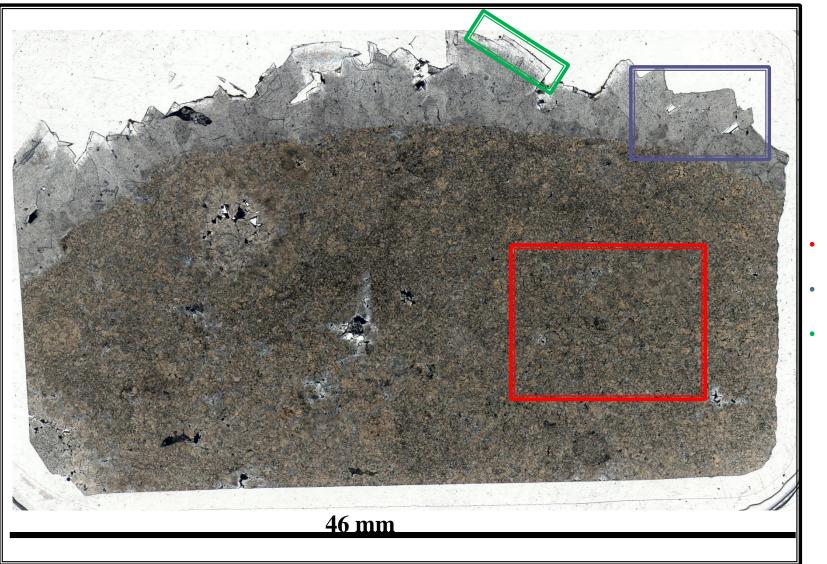


Results: Dundee Petrography

- 11 thick sections
- Secondary components:
 - Dolomite
 - Mosaic
 - Replacive
 - Frac/vug
 - Neoformed, saddle
 - Overgrowth Dolomite
 - Neoformed, on other phases
 - Calcite
 - Fracture filling
 - Anhydrite
 - Fracture filling
 - Organic Material
 - Coats dolomite

Results: Dundee Petrography

Dolomite phases Lee 3466'

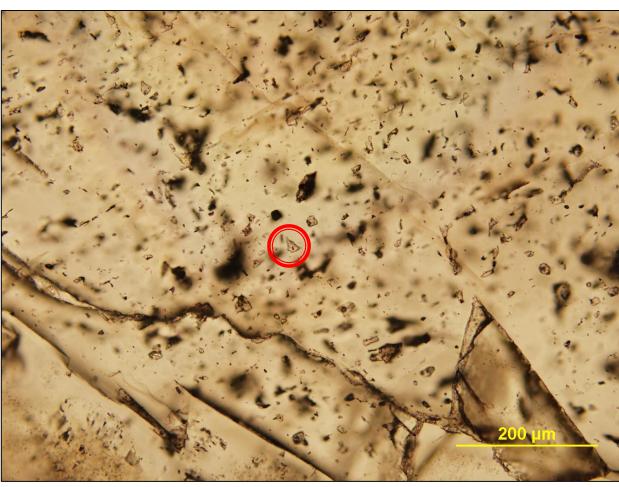


- Mosaic
- Frac/Vug
- Overgrowth on saddle frac/vug

Methods: Fluid Inclusion Analysis

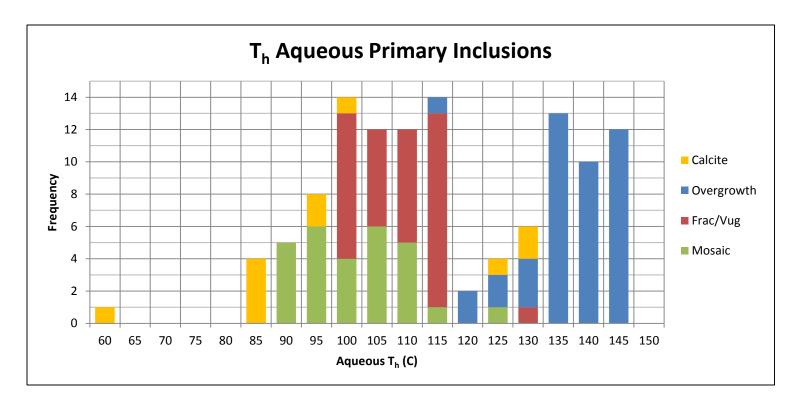
Fluid inclusions

- Small amount of fluid trapped in a crystal
 - Primary
 - Secondary
- Record information about conditions during entrapment
 - T_I
 - Homogenization temp. of 2 phase inclusions
 - Minimum temp. or entrapment
 - T_m
 - · Final melting temp.
 - · Salinity of fluid
 - T_i
 - · Eutectic temp.
 - · Salt system of fluid



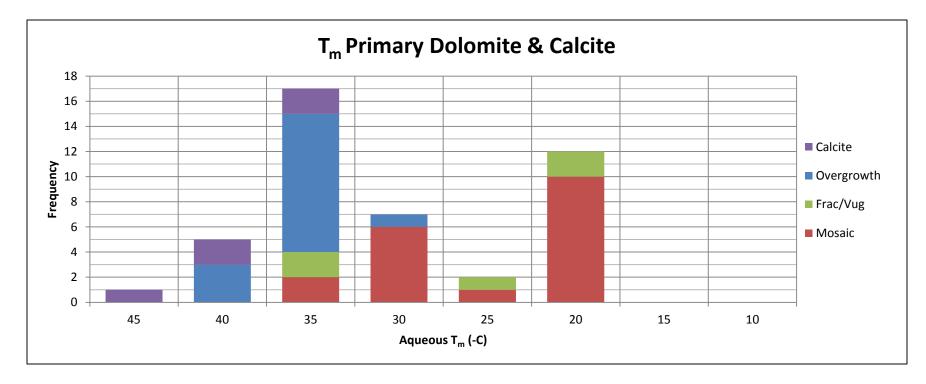
Results: Fluid Inclusion Analysis

- Average T_h for Primary FI's varied by phase
 - Mosaic-95° C
 - Frac/Vug-106.5° C
 - Overgrowth–134° C
 - Calcite-96° C



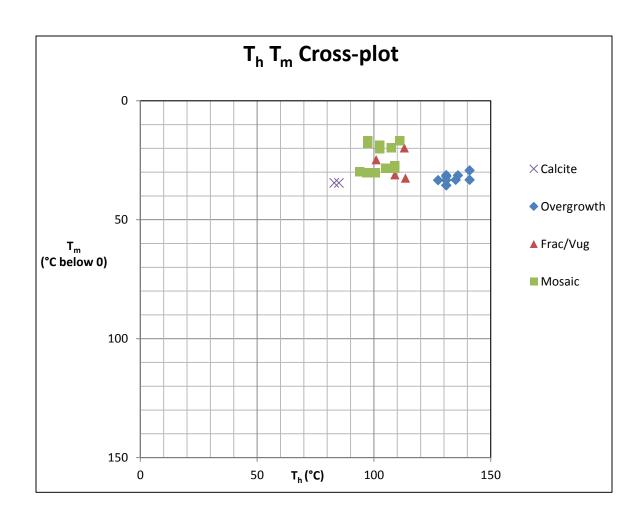
Results: Fluid Inclusion Analysis

- Average T_m for primary FI's varied by phase
 - Mosaic- -22.5° C, ~24% Na Cl equiv.
 - Frac/Vug- -25.6° C, ~26% Na Cl equiv.
 - Overgrowth -33.3° C ~31% Na Cl equiv.
 - Calcite -36.1° C, ~33% Na Cl equiv.
- Na Cl equivalent salinities are ~7-10 x modern sea water



Results: Fluid Inclusion Analysis

- Quality control
 - Phases cluster in cross-plot



Results: Fluid Inclusion Analysis

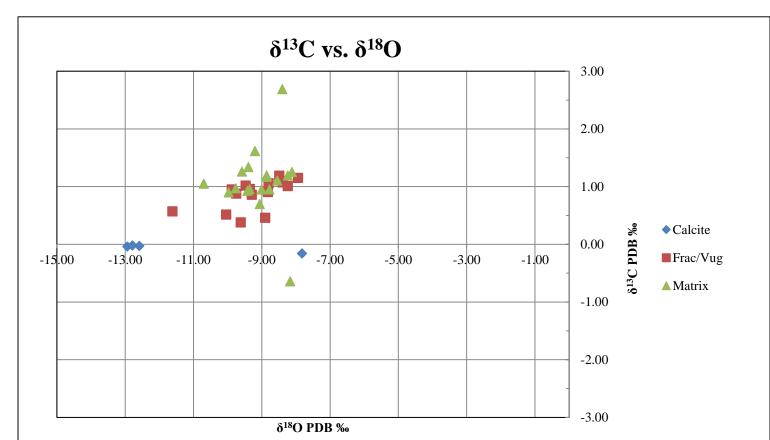
- 9 T_e
 measurements
 recorded as
 ranges
 - □ -45 to -65° C
 - NaCl-CaC1₂-MgCl₂-H₂ salt system
 - Luczaj et al., 2006

Well	Depth	FIA	Phase	T _e (-°C)
Lee	3466'	1	Overgrowth	65-55
Lee	3466'	1	Overgrowth	65-55
Lee	3466'	1	Overgrowth	65-55
Lee	3466'	1	Overgrowth	65-55
R1-12	3948'	1	Mosaic	65-50
R1-12	3948'	1	Mosaic	65-50
Tow	3191'	1	Mosaic	55-45
Tow	3191'	2	Frac/Vug	60-50
Tow	3191'	2	Frac/Vug	60-50

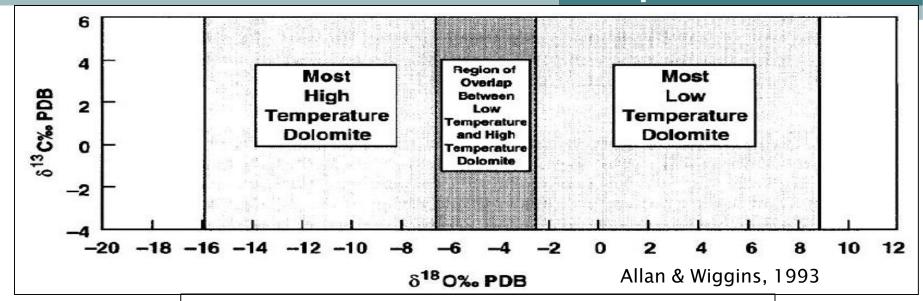
Results: Stable Isotopes

- ¹⁸O and ¹³C measurements for 37 Dundee samples
 - 18 mosaic, 15 frac/vug, and 4 calcite

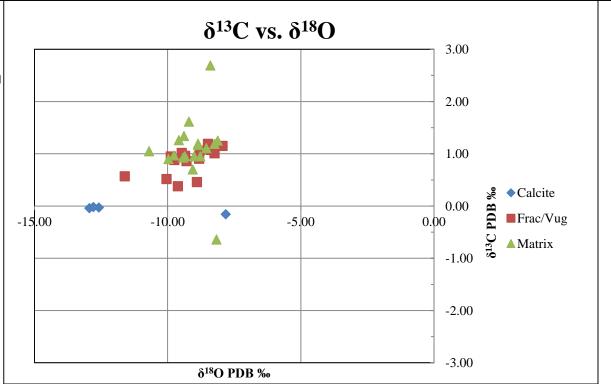
- Mosaic Avg. (PDB)
 - $-9.08 \delta^{18}O$
 - $+1.09 \delta^{13}C$
- Frac/vug Avg. (PDB)
 - $-9.24 \delta^{18}O$
 - $^{\text{\tiny 0}}$ +0.87 $\delta^{13}C$
- Calcite Avg. (PDB)
 - $-11.54 \delta^{18}O$
 - $^{\circ}$ -0.06 δ^{13} C



Results: Stable Isotopes



Results are indicative of high temperature inorganic dolomitization



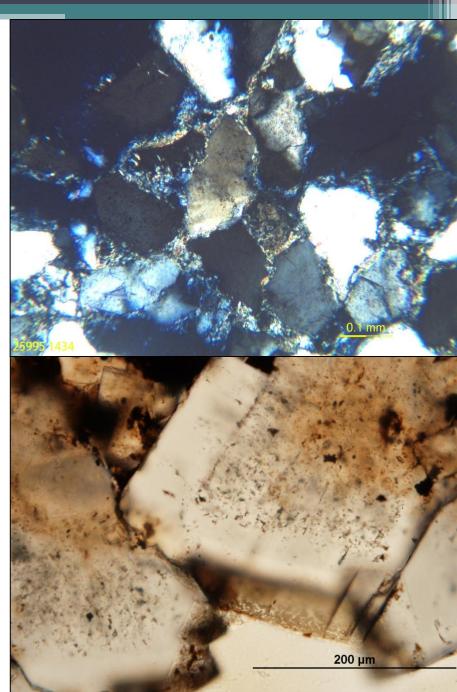
1. What authigenic minerals?

Marshall

- K-feldspar, quartz, ankerite, calcite, kaolinite, chlorite
 - Kaolinite replaced K-feldspar
- Illite
 - Illite replaced kaolinite

Dundee

- Dolomite
 - Mosaic dolomite
 - Frac/vug dolomite
 - Overgrowth dolomite
- Calcite
- Anhydrite



2. Formation Cond.

Marshall Sandstone

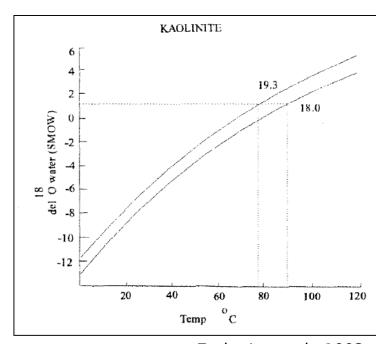
 Has experienced <u>temperatures > 80° C</u> during <u>migration of deep basinal fluids</u> through the formation

Dundee Limestone

 Dolomitization occurred at temperatures of <u>90 to</u> <u>145° C during the migration of deep basinal fluids</u> <u>through the formation</u>

2. Formation Cond.

- Marshall Sandstone
 - Temperatures >80° C
 - Burial (>2-4km) diagenetic assemblage
 - Zacharias et al., 1993
 - Chlorite: 55-85° C
 - Ankerite: 75–90° C
 - Kaolinite: 80-90° C
 - Quartz overgrowths
 - Uncommon below 80° C (Wilkinson & Haszeldine, 2002)



Zacharias et al., 1993

- Illite
 - Formation sandstone uncommon below 100° C (Wilkinson & Haszeldine, 2002)
 - Replaced kaolinite
 - Reaction related to temp. increase (Meunier & Velde, 2004)

- Authigenic illite formation in <u>sandstone</u>
 - Requires high flow rates of K+ rich fluid (Wilkinson & Haszeldine 2002)
 - External K⁺ source needed (Meunier & Velde, 2004)
 - Formation generally attributed to hydrothermal fluid flow related to tectonic events (Lanson et al., 1996)

Dundee Limestone

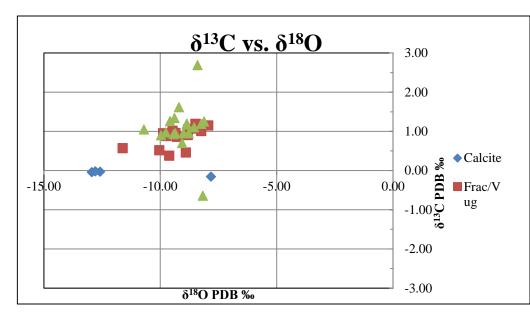
 Dolomitization occurred at temperatures of <u>90 to</u> <u>145° C</u> during the <u>migration of deep</u> <u>basinal fluids through</u> the formation

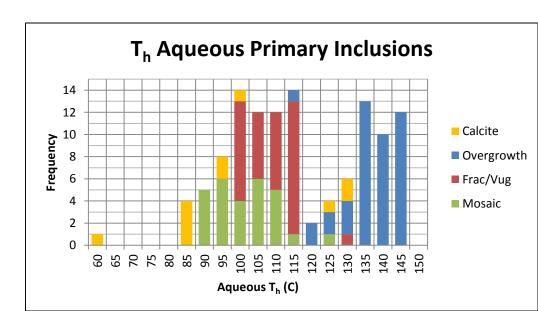


Mosaic: 90–110 ° C

• Frac/vug: 100–115° C

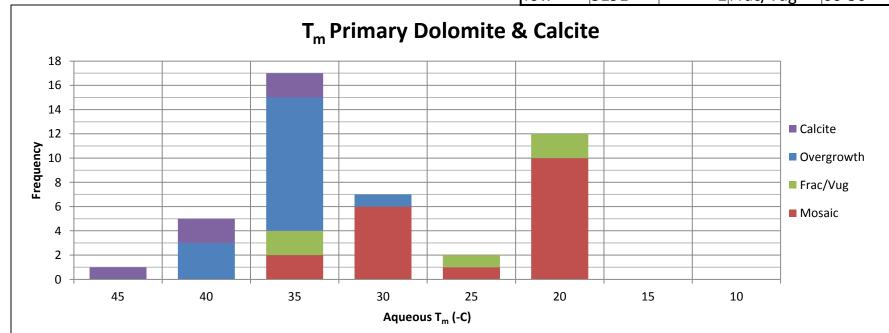
Overgrowth:130–145° C





- Dolomitizing fluid
 - □ ~7–10 x salinity modern sea water
 - NaCl-CaCl₂-MgCl₂-H₂ salt system

Well	Depth	FIA	Phase	T _e (-°C)
Lee	3466'	1	Overgrowth	65-55
Lee	3466'	1	Overgrowth	65-55
Lee	3466'	1	Overgrowth	65-55
Lee	3466'	1	Overgrowth	65-55
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R1-12	3948'	1	Mosaic	65-50
Tow	3191'	1	Mosaic	55-45
Tow	3191'	2	Frac/Vug	60-50
Tow	3191'		Frac/Vug	60-50



- Formation fluid isotopic composition
 - Dolomitizing fluid estimate
 - -0.5 to $+2.00 \delta^{18}$ O SMOW

Phase	Avg δ ¹⁸ O ‰ PDB	Avg δ ¹⁸ O ‰ SMOW	Formation Temp °C	Est. Water δ180 ‰ SMOW Low Temp	Est. Water High δ180 ‰ SMOW High Temp	
Calcite	Without 3195					
	-12.77	+19.03	80-100	+0.6	+1.8	
	With 3195					
	-11.54	+21.90	80-100	+2.3	+4.6	
Frac/Vug	-9.24	+21.39	100 to 115	+0.6	+2.3	
Mosaic	-9.08	+21.56	90 to 110	-0.50	+2.00	

- Marshall brine measurement
 - $+0.8 \delta^{18}O$ SMOW (Zacharias et al., 1993)

- Evidence for deep basinal brine migration
 - High salinity fluids
 - Marshall brine genetically related to Silurian brines (Ma et al., 2005)
 - Dundee fluids ~7-10 more saline than seawater
 - Fluids in both formations have interacted with basement
 - Marshall brine He (Ma et al., 2005)
 - Dundee brines mantle-like signature (Ma et al., 2009)
 - Illitization and dolomitization require chemical material from an outside source
 - External K+ source needed for illitization (Meunier & Velde, 2004)
 - Outside source of Mg needed for dolomitization (Allan & Wiggins, 1993)

3. When did they form?

- Marshall (dep. age ~345 Ma)
 - □ ~280 Ma
 - Single period of illitization in the Permian?
 - Multiple illitization events?
- Dundee (dep. age ~390 Ma)
 - Cogenetic with Marshall?
 - Multiple episodes?

4. Comparison with Burial

	Mar	shall	Dundee		
	Shallowest	Deepest	Shallowest	Deepest	
Depth (m)	14m	477m	973m	1203m	
Current Temp (C)	15°	25°	35°	40°	
0.5km Add. Burial Temp (C)	37°	47°	57°	62°	
1km Add. Burial Temp (C)	47°	57°	68°	73°	
2km Add.Burial Temp (C)	69°	79°	91°	95°	
Illite/Dolomite Formation Temp.(C)	>80 °		130-145°		

- Features can not be explained by burial
 - Temperature
 - Require >2km burial
 - Chemical mass
 - Require external source
 - Deep basin brine signature

Geothermal gradient of 22° C Current assumes 15° C surface temp Add. Burial assumes 25° C surface temp

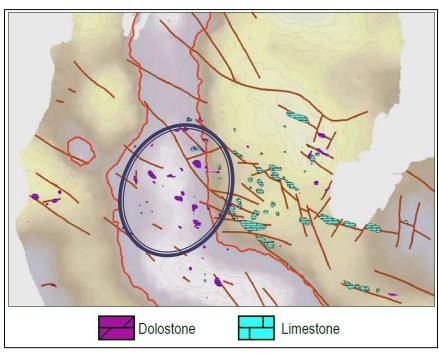
5. Distribution

- Are anomalous paleo-temperature indicators spatially related to the rift and known structures?
 - Marshall
 - Distribution of illite in the Marshall coincides w/ Dundee dolomitization and known structure

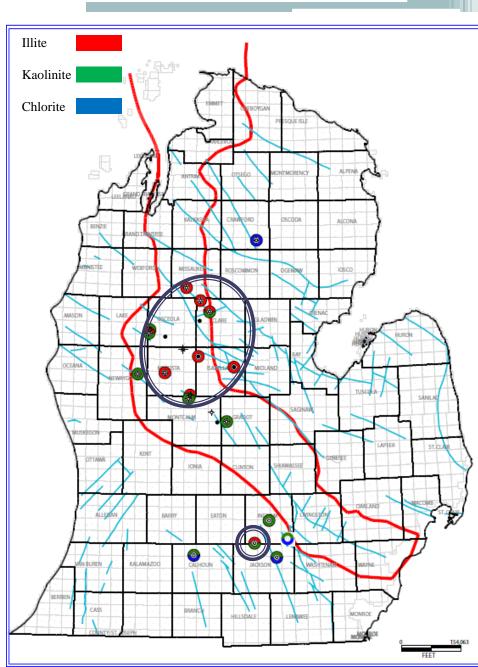
Dundee

 Dolomitzation common within the MMR gravity anomaly and near structure

5. Distribution



Barnes et al., 2008

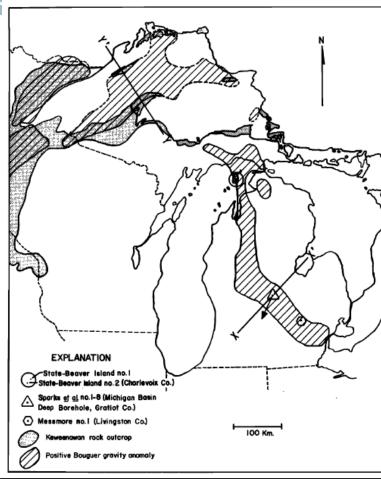


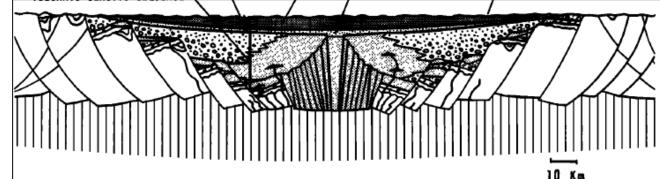
Mechanism

- Hydrothermal Fluid Flow & Alteration
 - Burial alone would not provide the heat energy and chemical mass required
 - Need vehicle to be transported to the formation
 - Structural control is noted
 - Hydrothermal fluids would preferentially travel along faults, fractures, and along anticlines
 - MMR
 - Distribution of anomalous features (this study, Girard & Barnes 1995)
 - Shallow formations w/ deep basinal brines
 - Significant vertical flow component in the basin (Ma et al., 2005)
- Hydrothermal fluid flow results from tectonic events?

Mechanism

- Appalachian Tectonic Events
 - Have affected the subsidence history of the basin (Howell & Van der Pluijm, 1999)
 - Hydrothermal illitization in the Appalachian basin and Mississippi Valley strata (Elliot & Aronson, 1987; Hay et al., 1988; Clauer et al., 2013)
 - Reactivation of MMR related basement rooted faults
 - Girard & Barnes, 1995; Castro et al., 2009; Ma et al., 2009
 - Mantle-like signature in basin brines (Castro et al., 2009)
 - Illite ages comparable
 - Marshall ~280 Ma
 - Alleghenian (320–260 Ma)
 - Ordovician St. Peter ~370–320 Ma (Girard & Barnes, 1995)
 - Integrated Acadian (420–360 Ma) & Alleghenian (320–260 Ma)





Conclusions

- Evidence for anomalous paleo-temperatures is present in the Marshall Sandstone and Dundee Limestone
- This study and others indicate that hydrothermal fluid flow is the best explanation
 - Deep basinal fluids travelling along faults and fracture zones
- Distribution of anomalous features and basin brine chemistry suggest a connection with the MMR
 - Hydrothermal events maybe related to reactivation of MMR faults
- Illite K-Ar ages are comparable to Late Paleozoic Appalachian orogenies

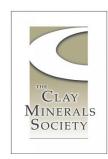
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Thank you

Questions?

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