What Happened to My Marine Reservoir? Implications of Falling Stage and Lowstand Fluvial Sediment Storage during "Sequence-Boundary" Scour for Sand Starvation of Coastal Marine Reservoirs*

John M. Holbrook¹ and Janok Bhattacharya²

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Abstract

Recent flume and field studies show that sediment may be stored by fluvial aggradation and lateral migration during regression, resulting in sediment starvation in coeval marine environments. In extreme cases, this may result in storage of all sediment within the fluvial system and complete starvation, or "autodetachment," of the contemporary marine shoreline. Complete sediment starvation of the shore is still theoretical, and likely rare. Several recent field studies, coupled with new ideas regarding the scour processes of sequence boundaries, however, suggest that significant falling and lowstand fluvial sand storage commonly results in diminished to near total reduction of marine reservoir sand.

Newer views on scour of the "subaerial unconformity" sequence boundary show that it does not actually record a surface of exposure and near-complete bypass of sediment at lowstand as originally presumed, but rather records a composite surface formed by lateral migration and incision of rivers that 'carve-and-cover' the subaerial unconformity throughout regression. This carve-and-cover process means that fluvial sediment is deposited above this surface throughout the regressive phase. Because transport of sand lags transport of suspended load, regressive fluvial sediments disproportionally sequester the sandy fraction. Coastal Quaternary systems and the Cretaceous of the Western Interior provide several examples where coastal systems were deprived of sandy sediment to varying degrees during regression. Sand starvation ranges from minimal, resulting in prominent regressive coastal sand reservoirs, to near-complete, in which case lowstand terminal estuaries with negligible coastal sand deposition result. Partial to near-complete "sand autodetachment", in which there is sand starvation of marine reservoirs, appears more pronounced where regressive slopes are low, compared to river profiles and where base-level remains relatively stable during regression. Such minimally incised systems are common where stable base-level promotes lateral migration of channels during falling and lowstand stage, enhancing fluvial sand storage.

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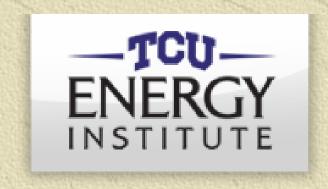
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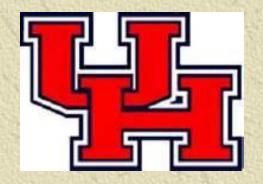
Implications of Falling Stage and Lowstand Fluvial Sediment Storage during "Sequence-Boundary" Scour for Sand Starvation of Coastal Marine reservoirs

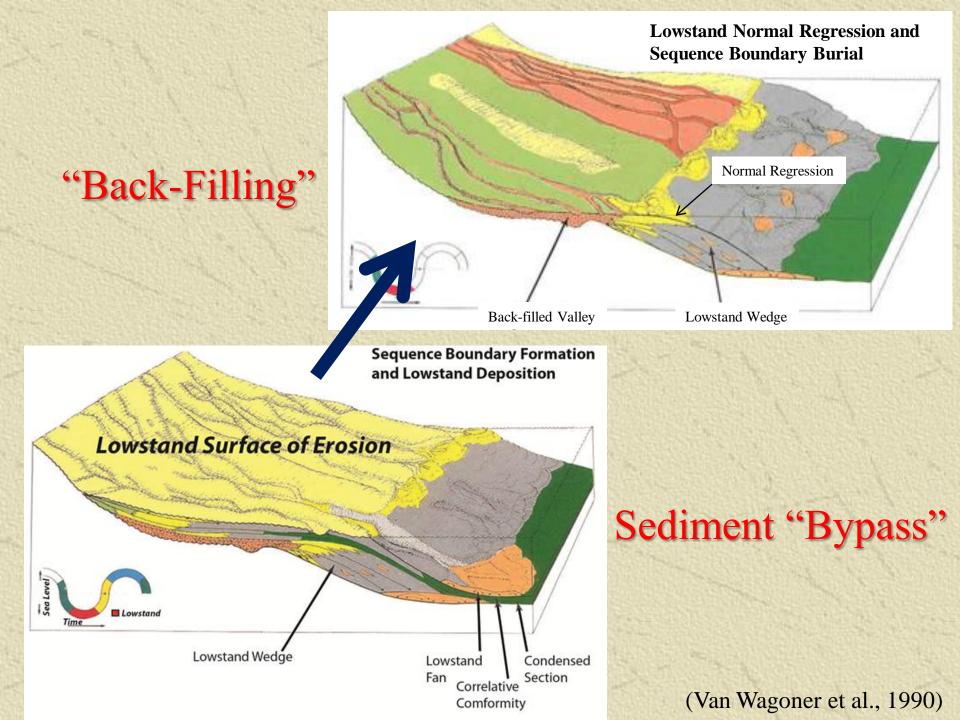
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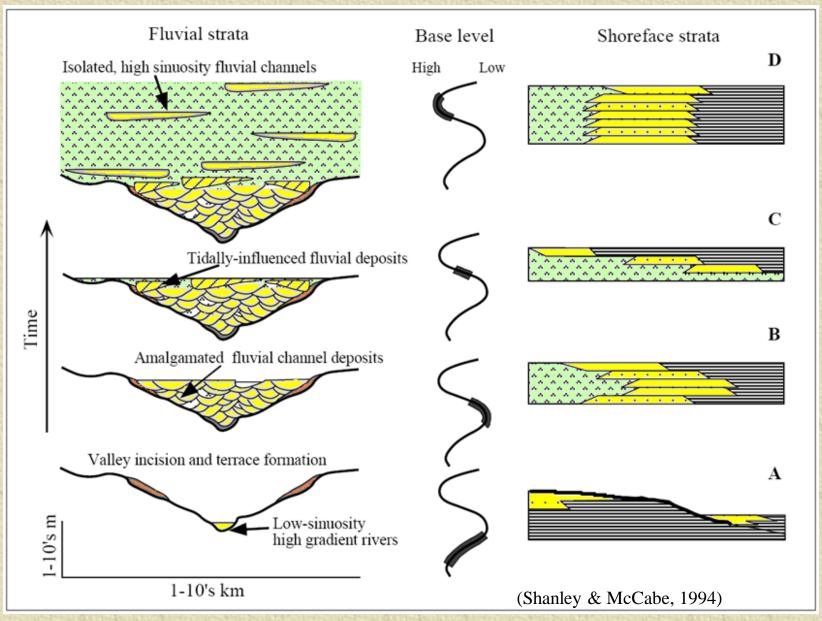


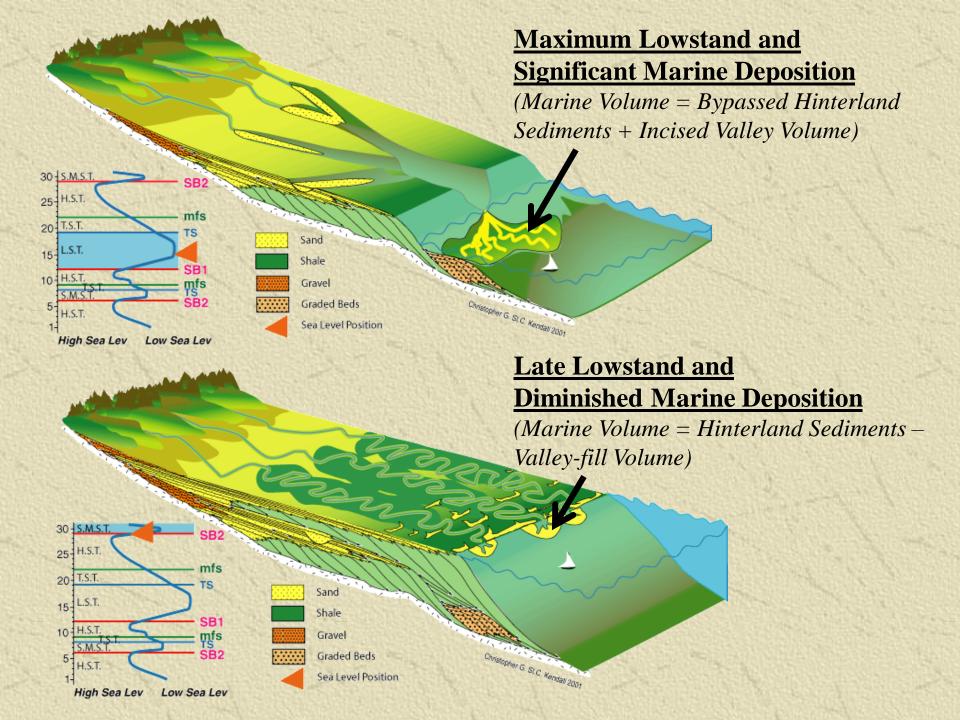
Janok Bhattacharya University of Houston



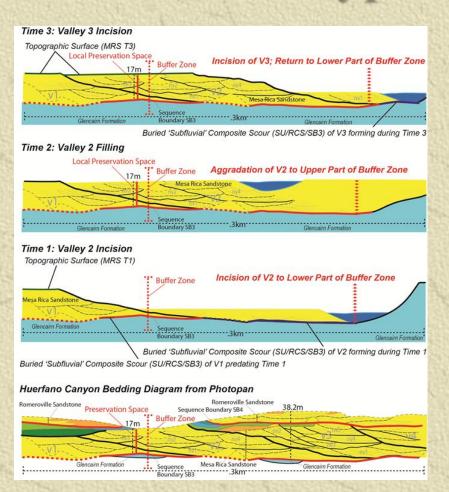


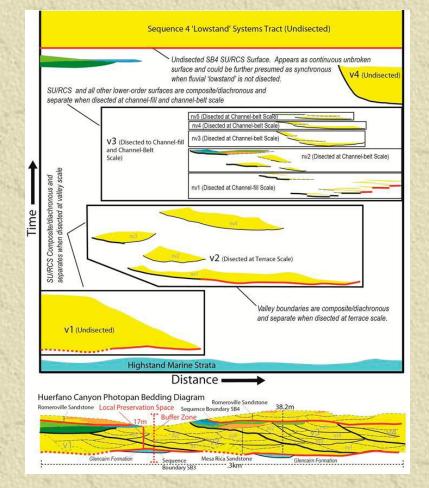
Fluvial Sequence Stratigraphy



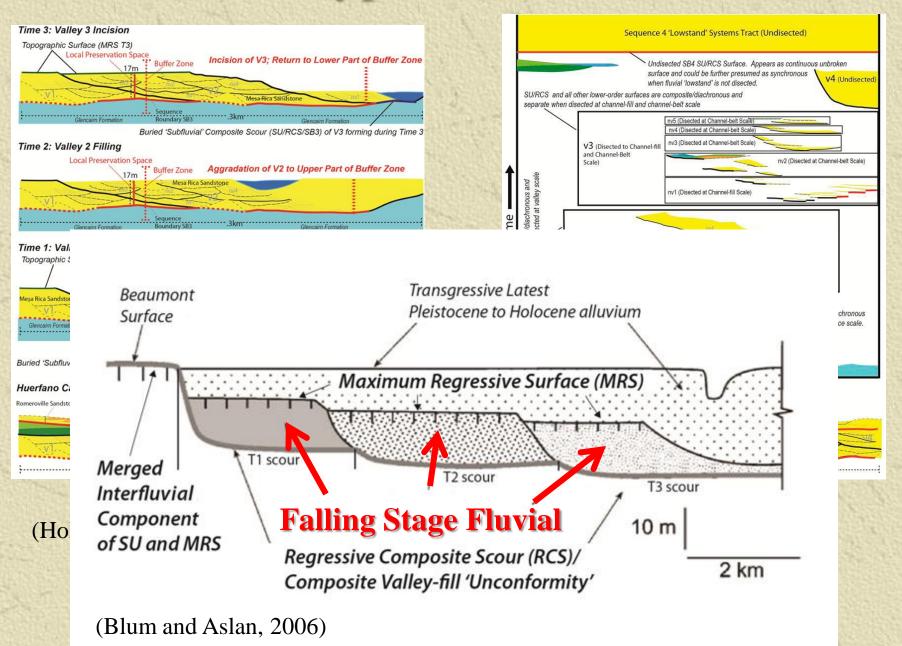


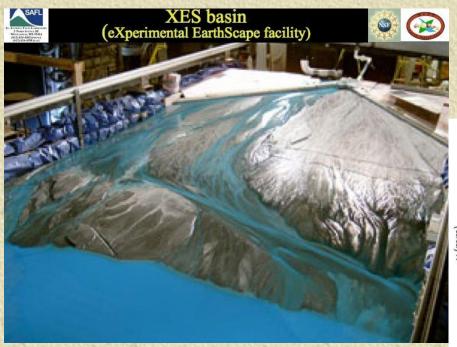
A Bypass Alternative



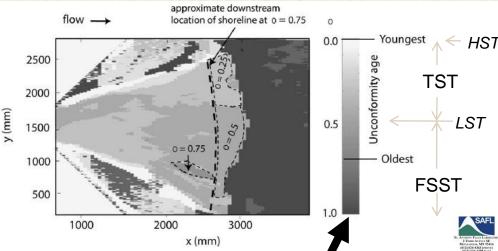


A Bypass Alternative





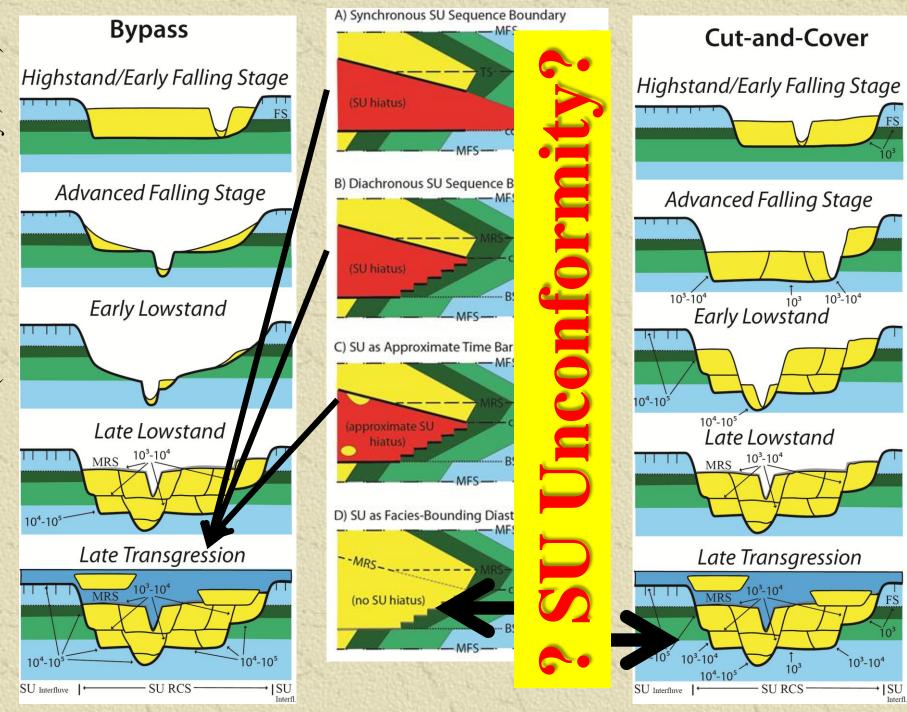
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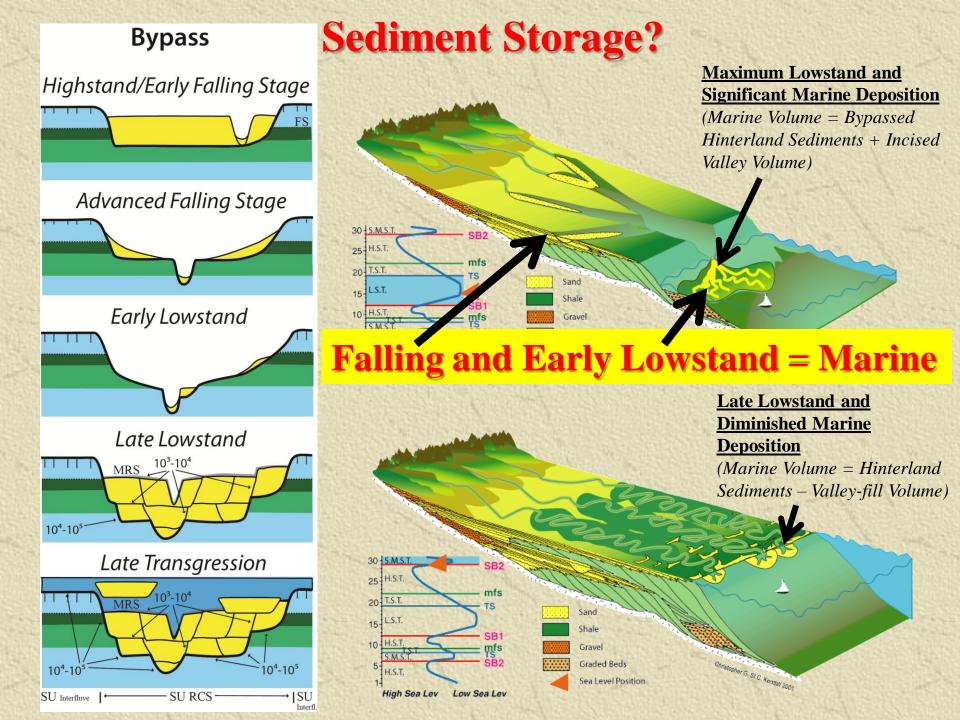


XES basin (eXperimental EarthScape facility) (Strong and Paola, 2008) **AUTOGENIC "NOISE"** "Noisy" sequence boundaries 1-2 channels thick

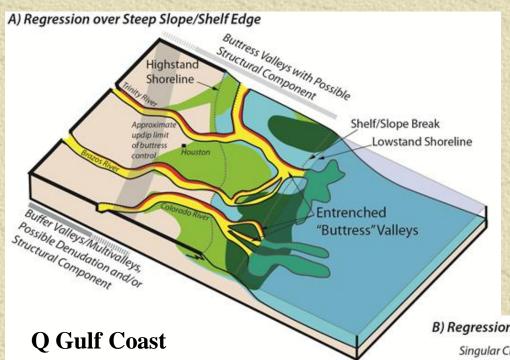
(Y,Z) From (0,0) To (1515,1400)

Sequence boundary formed over 75% of the entire sea level cycle!!!



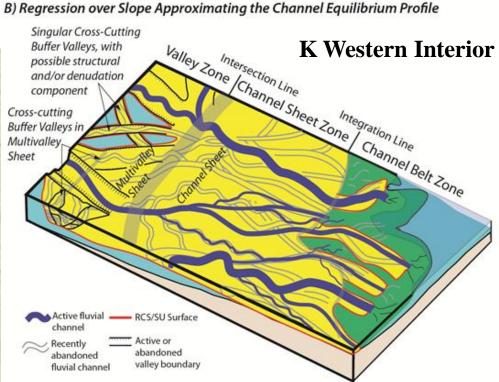


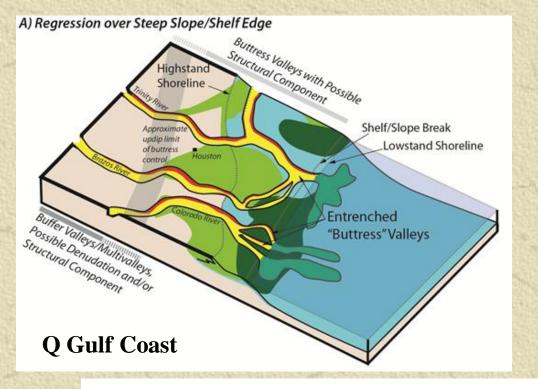
Sediment Storage? Cut-and-Cover Highstand/Early Falling Stage Maximum Regressive Surface (MRS) (Topographic surface) Falling and Early Lowstand = Fluvial and Marine Advanced Falling Stage T1 Sand Limit $10^{3}-10^{4}$ 10³ Early Lowstand Up-dip valleys and Older Fluvial Mesa Rica Multivalleys Sandstone above Younger Glencairn/Mesa Rica Deltaic strata across SU/RCS/SB3.1 104-105 T2 Sand Limit 10⁴-10⁵ Late Lowstand MRS Late Transgression Active fluvial channel 103-104 Recently Active or abandoned abandoned fluvial channel valley boundary 103-104 103-104 10^{3} SU RCS SU Interfluve · ISU



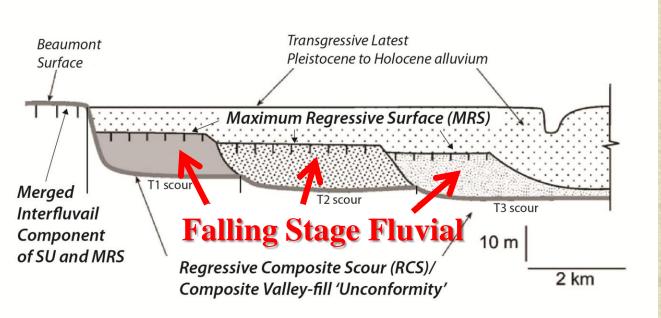
Narrow "Buttress" Valleys and Minimal Storage

"Buffer" Valleys with Lateral Planation and Significant Storage





Narrow "Buttress" Valleys and Minimal Storage

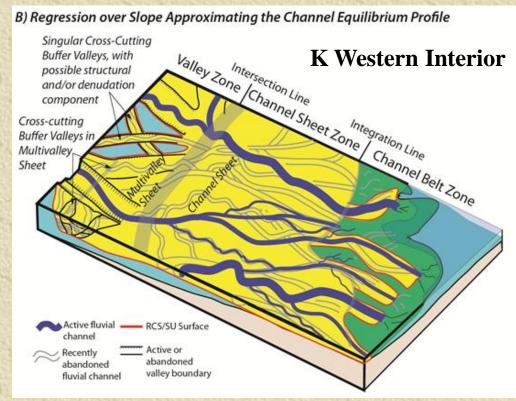


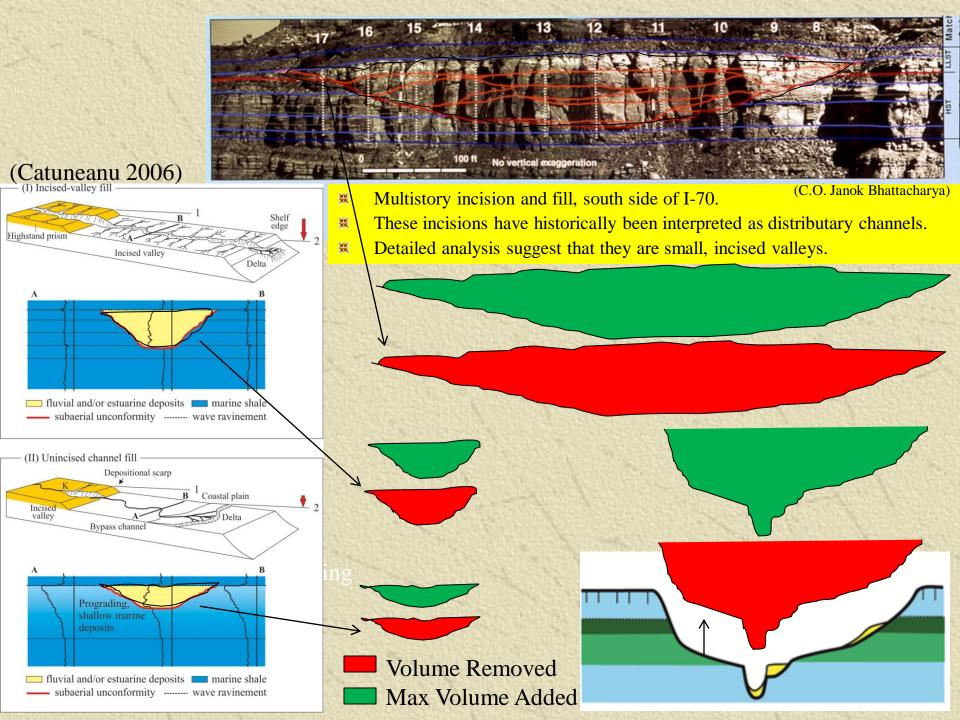
Maximum Regressive Surface (MRS) (Topographic surface) SU/RCS/SB3.1 (Buried subfluvial erosional surface) OK T1 Sand Limit Up-dip valleys and Older Fluvial Mesa Rica Multivalleys Sandstone above Younger Glencairn/Mesa Rica Deltaic strata across SU/RCS/SB3.1 T2 Sand Limit Active fluvial channel Active or Recently abandoned abandoned fluvial channel valley boundary

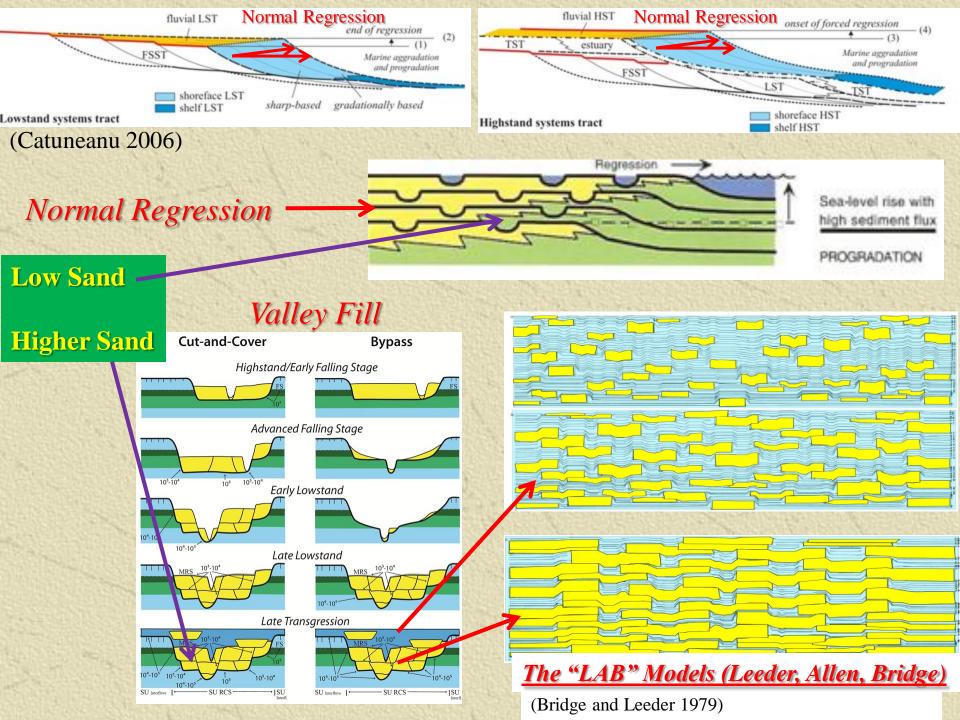
"Buffer" Valleys with Lateral Planation and Significant Storage

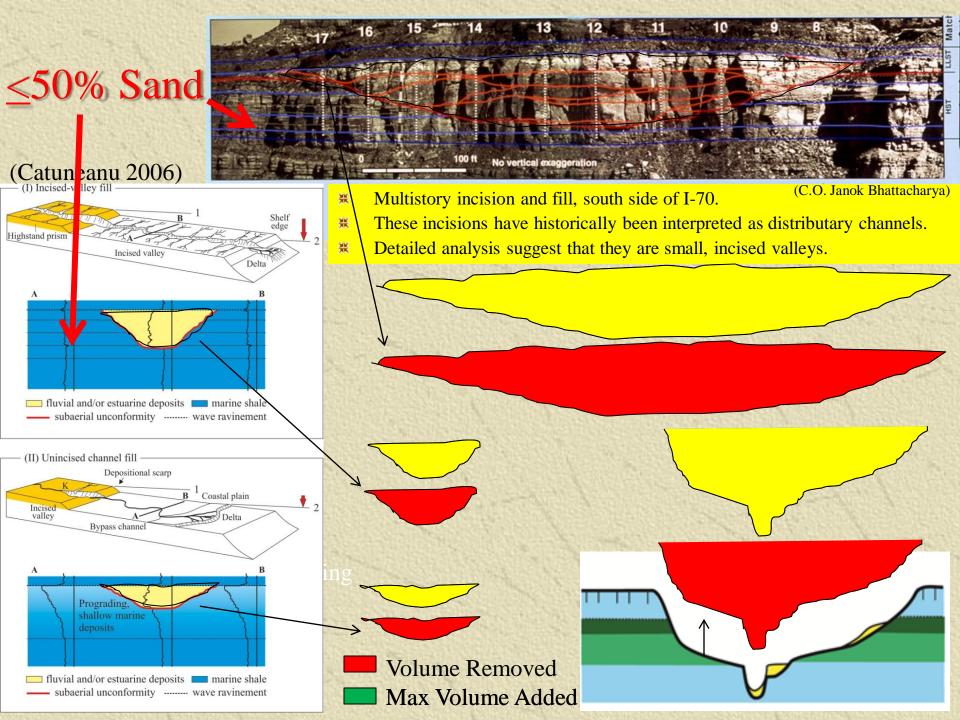
Falling Stage

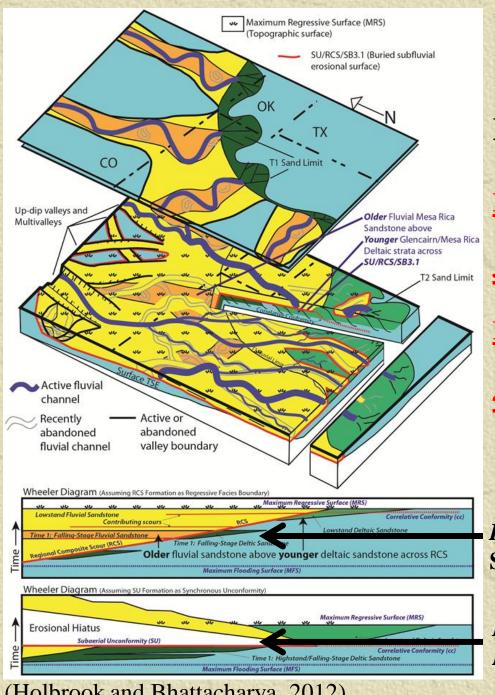
Maximum Lowstand











Mesa Rica and Glencairn Fms

Fluvial Volume

 $14m \times 37500 \text{km}^2 =$

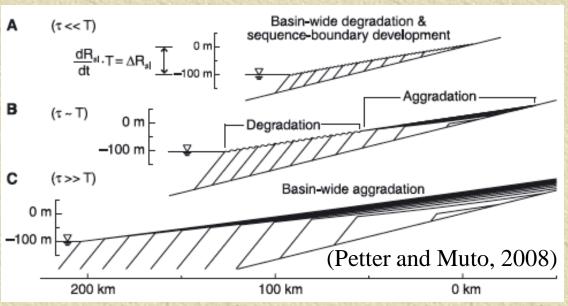
500km³

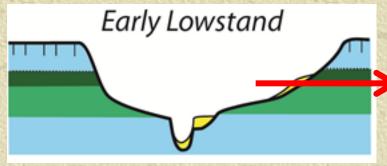
Marine Volume at (50% sand) $150 \text{km}^3 \text{ (Falling)} + 250 \text{km}^3 \text{ (Lowstand)} =$ $400km^3$

If SU is Cut-and-Cover Diastem Sand storage equal for Fluvial and Marine

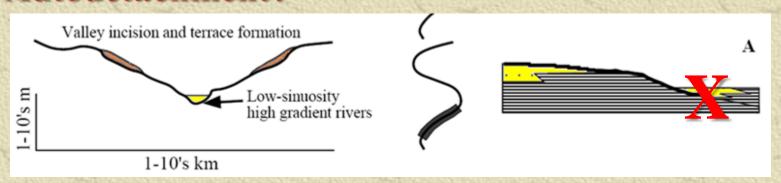
If SU is Bypass Unconformity Marine Sand, Minor Fluvial Sand Storage

Autodetachment?

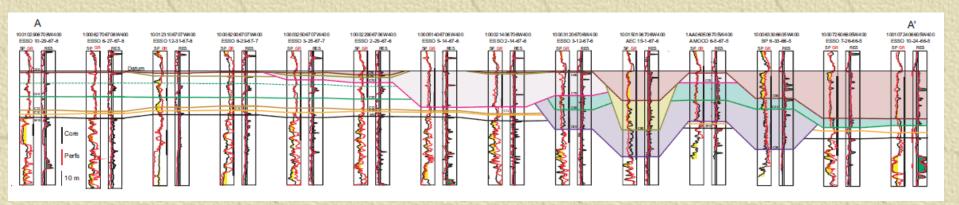


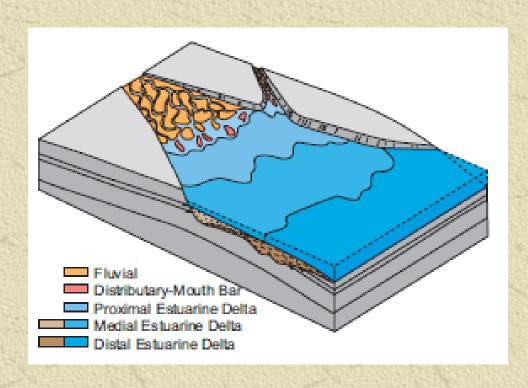


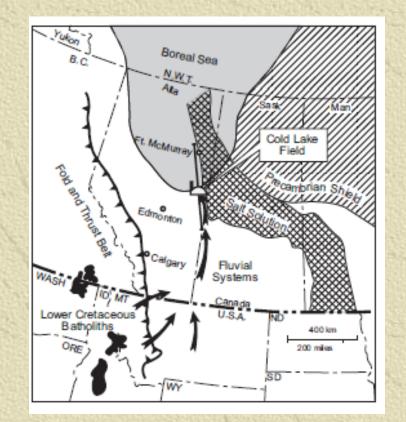
Sand Autodetachment?



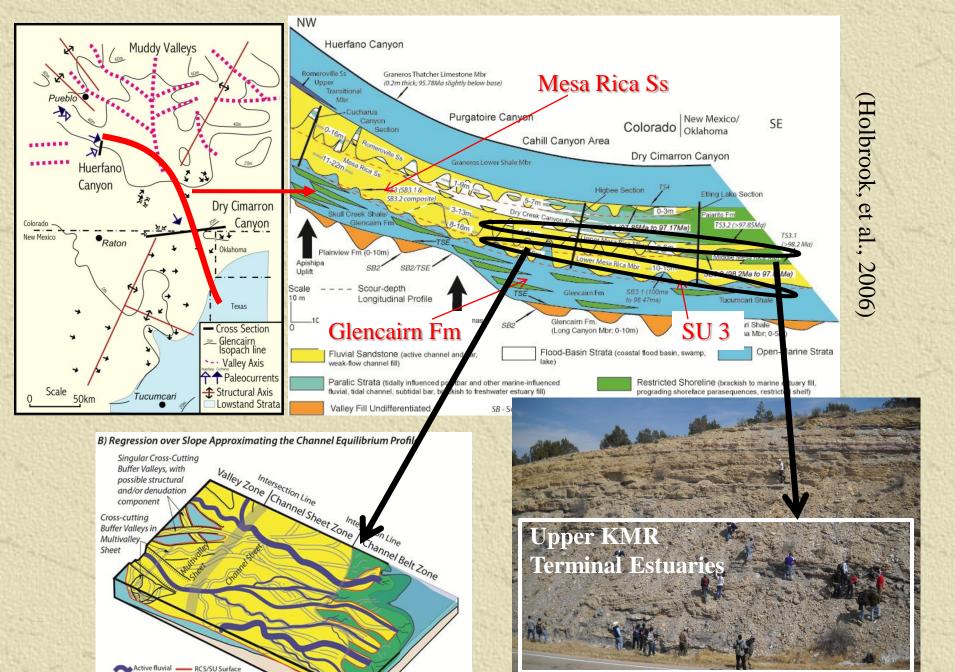
Terminal Estuaries in the K Clearwater Fm, Canada







(Feldman et al., 2008)



channel Recently

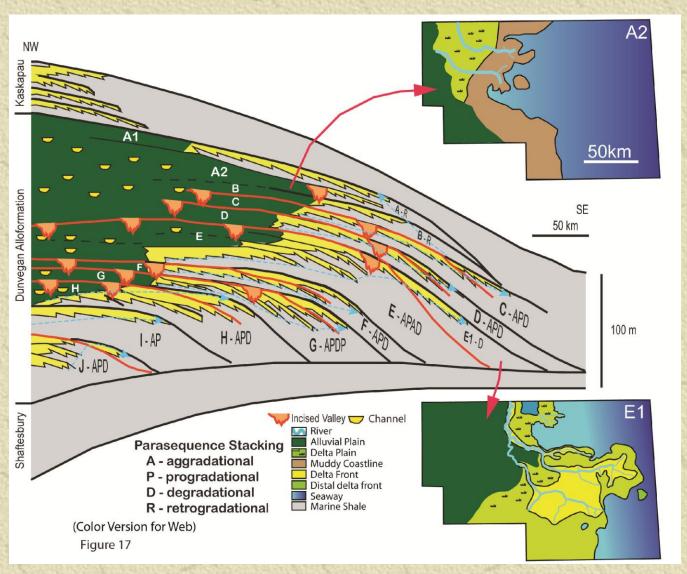
abandoned

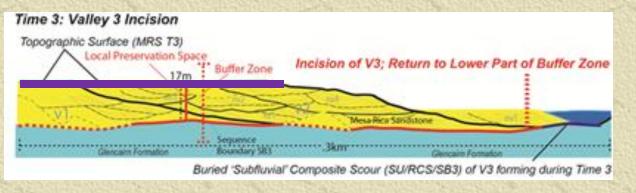
fluvial channel

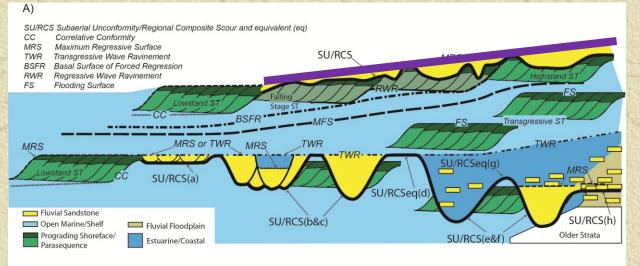
abandoned

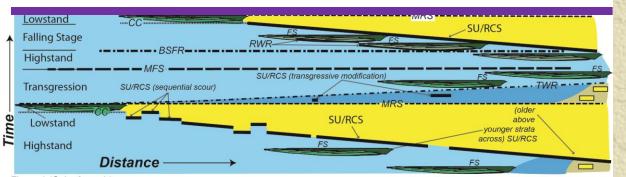
valley boundary

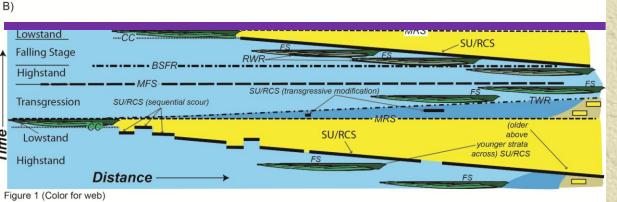
Variation in Shoreline Sand Preservation K Dunvegan Fm, Alberta

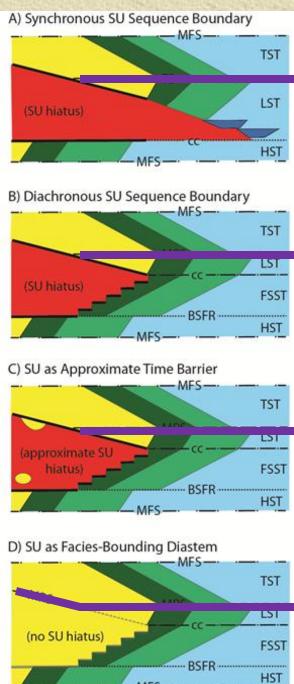


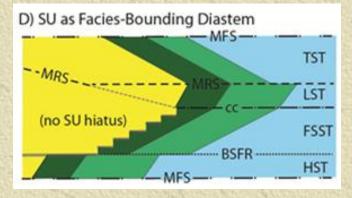






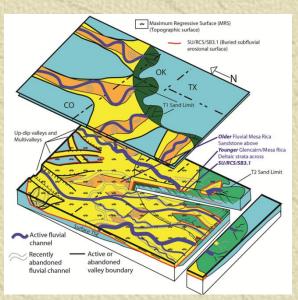






Conclusions

SU is probably not a bypass surface or unconformity



Falling and lowstand fluvial sand storage can be high to the expense, and possible autodetachment, of marine reservoirs

MRS is probably a better correlation surface than given credit

