

Syn depositional Tectonism and its Effects on Mississippian (Kinderhookian to Osagean) Lithostratigraphic Architecture: Part 1 – Based on Exposures in the Midcontinent USA*

Brian W. Wilhite¹, S.J. Mazzullo², Beau T. Morris², and Darwin Boardman II³

Search and Discovery Article #30207 (2011)

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*Adapted from oral presentation at AAPG Mid-Continent Section meeting, Oklahoma City, Oklahoma, October 1-4, 2011. Please see closely related article, [“Syn depositional Tectonism and its Effects on Mississippian \(Kinderhookian to Osagean\) Lithostratigraphic Architecture: Part 2 – Subsurface Occurrences in the Midcontinent USA”](#), Search and Discovery article #30208.

¹Woolsey Operating Co., Wichita, KS 67202

²Department of Geology, Wichita State University, Wichita, KS 67260 (sjmazzullo@gmail.com)

³Oklahoma State University, Stillwater, OK 74048

Abstract

Exposures of Kinderhookian to Osagean carbonates in Missouri and adjoining Arkansas and Oklahoma, deposited on the flank of the Ozark Uplift inboard of the Ouachita thrust-belt to the south, record aggradational followed by progradational sedimentation. Deposition responded to long-term tectono-eustasy. Early pulses of Ouachita collision resulted in areal shifts of fore-bulge arches and basins, and movement along associated faults during deposition. This syn depositional tectonism overprinted the fundamental lithostratigraphic and apparent sequence-stratigraphic architecture of this section, the manifestations of which include: (1) periodic uplift and unconformity formation within marine and subaerial meteoric environments; (2) northward progradation of relatively thick sections of strata that downlap regional unconformities and not maximum flooding surfaces; (3) the presence of allochthonous reefs that detached and slid downslope to the north within downlapping strata; (4) the formation of submarine erosionally-truncated folds at several horizons in the section, the axes of which uniformly trend E-W; and (5) regional down-dip (to the south) thinning that resulted from erosional truncation on uplifted, shallow-water arches and not from depositional thinning into deep water. The recognition of such processes and products has important bearing on the nature, types, and locations of petroleum reservoirs in the subsurface, and accordingly, on our exploration models for them.

Reference

Noble, P.J., 1995, Regional sedimentation patterns associated with the passive-to active-margin transition, Ouachita orogeny, southern midcontinent, USA, *in* K.S. Johnson, (ed.) Structural Styles in the Southern Midcontinent, 1992 Symposium: Oklahoma Geological Survey, Norman, Oklahoma, Circular No. 97, p. 99-112.

Website

Blakey, R.C., Paleogeography: Web accessed 8 November 2011, <http://www2.nau.edu/rcb7/>

**SYNDEPOSITIONAL TECTONISM AND ITS EFFECTS ON
MISSISSIPPIAN (KINDERHOOKIAN TO OSAGEAN)
LITHOSTRATIGRAPHIC ARCHITECTURE**

PART 1

***BASED ON EXPOSURES IN THE
MID-CONTINENT USA***

**B. W. Wilhite, S. J. Mazzullo, Beau T. Morris,
and Darwin R. Boardman**

Acknowledgements:

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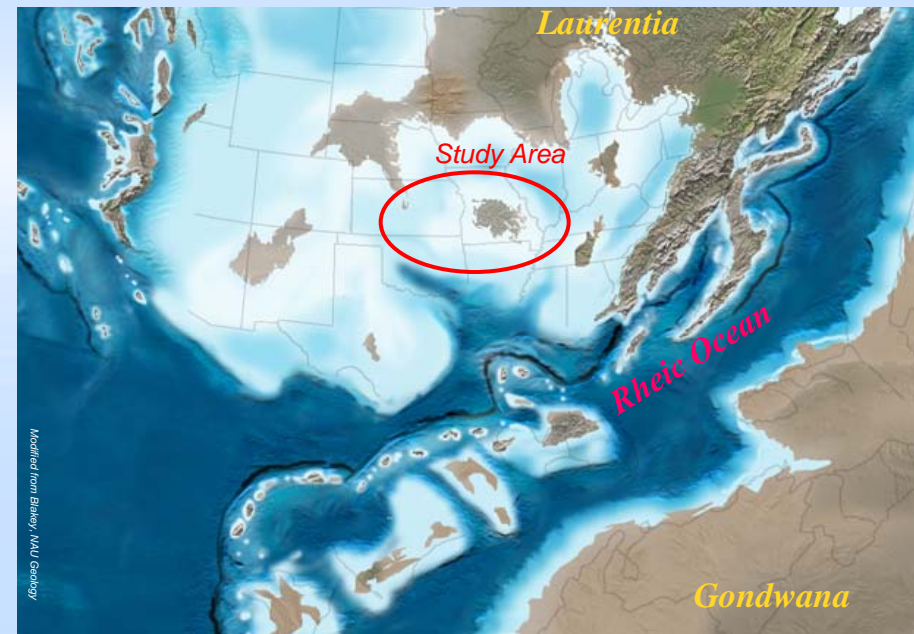
Current Belief...

The Rheic Ocean diachronously 'zippered' shut, from NE to SW, during plate convergence of Laurentia and Gondawana.

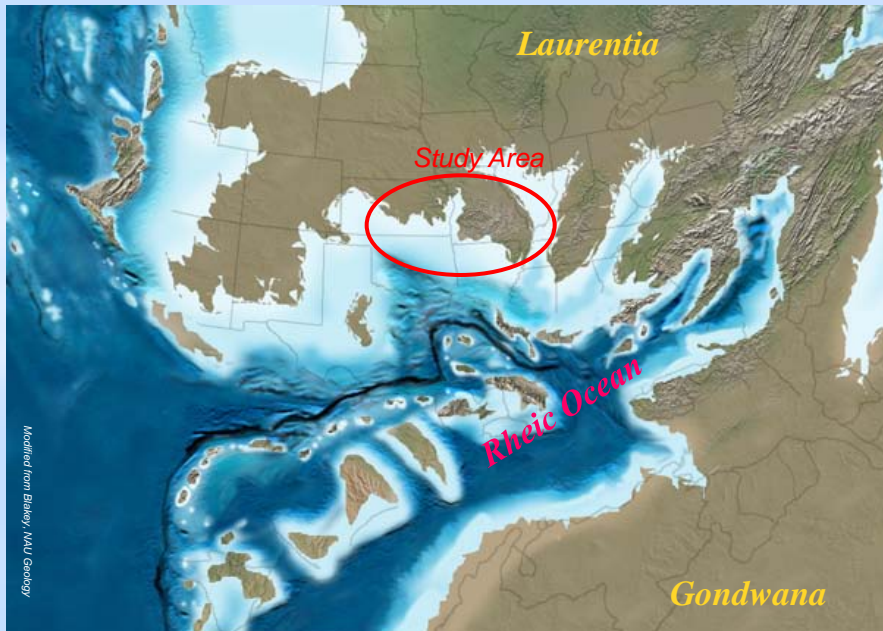
Timing of that crustal compression & onset of associated

**Mid-Continent tectonics is widely believed to be -
late-Mississippian (Mrmc-Chster) to
middle Pennsylvanian time.....**

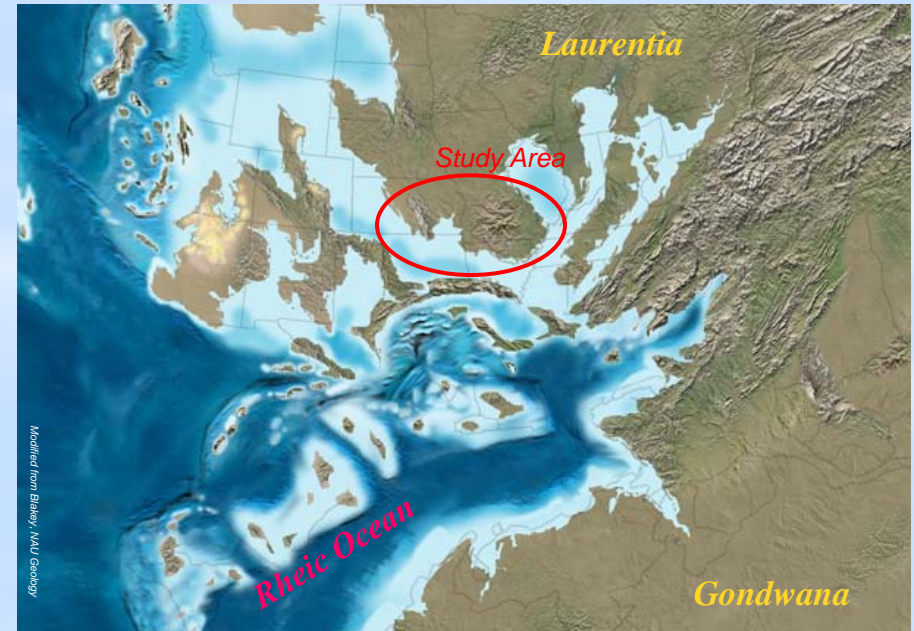
KINDERHOOK & OSAGE TIME: TECTONIC QUIESCENCE



LATE MERAMEC-CHESTER TIME: ONSET OF MID-CON TECTONICS & OUACHITA OROGENY



EARLY PENNSYLVANIAN TIME: CONTINUATION OF OUACHITA OROGENY



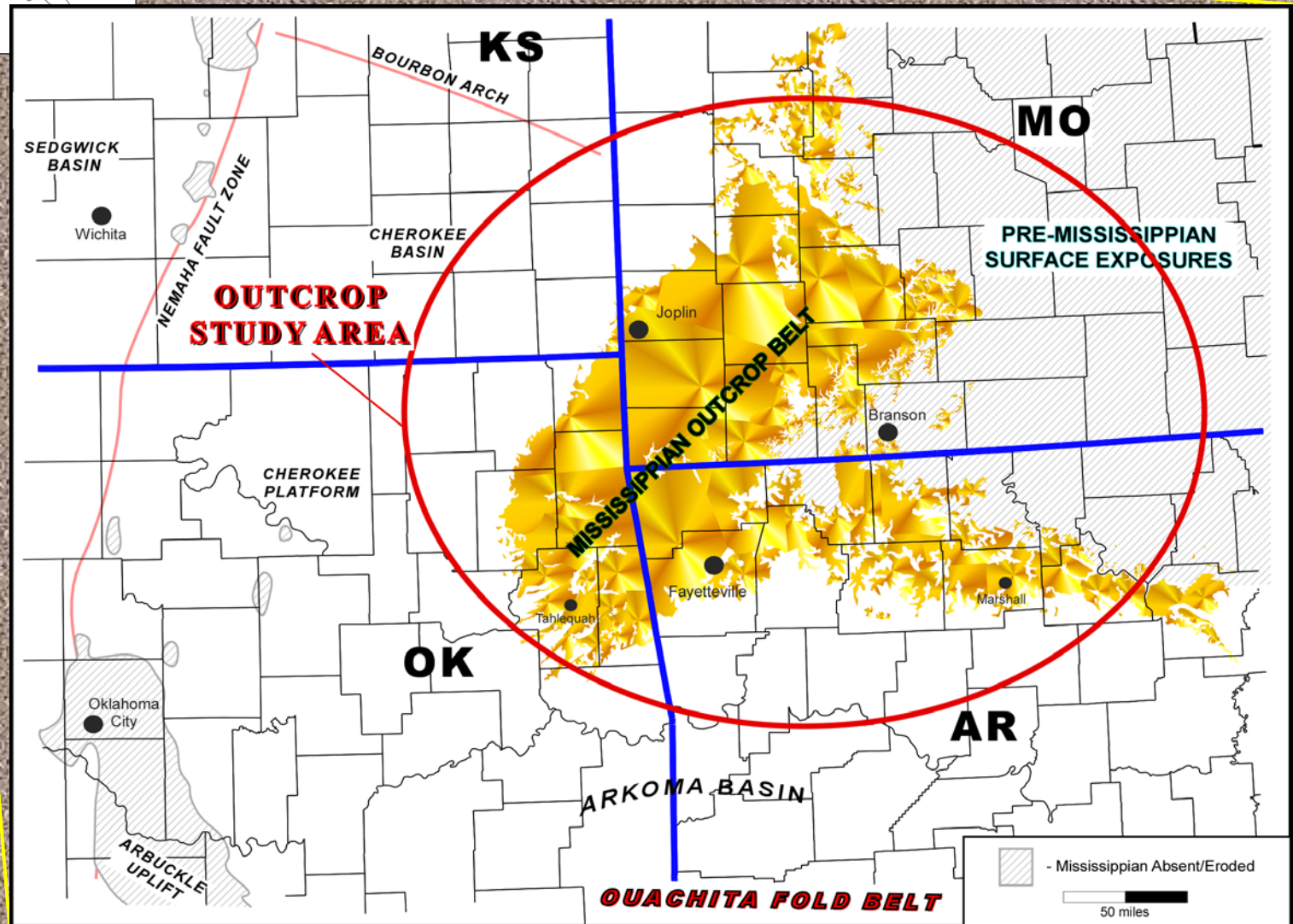
...a contention with which we do not agree.

Herein we'll illustrate
stratigraphic,
biostratigraphic,
facies,
& structural
evidence, from the outcrop, to suggest otherwise.

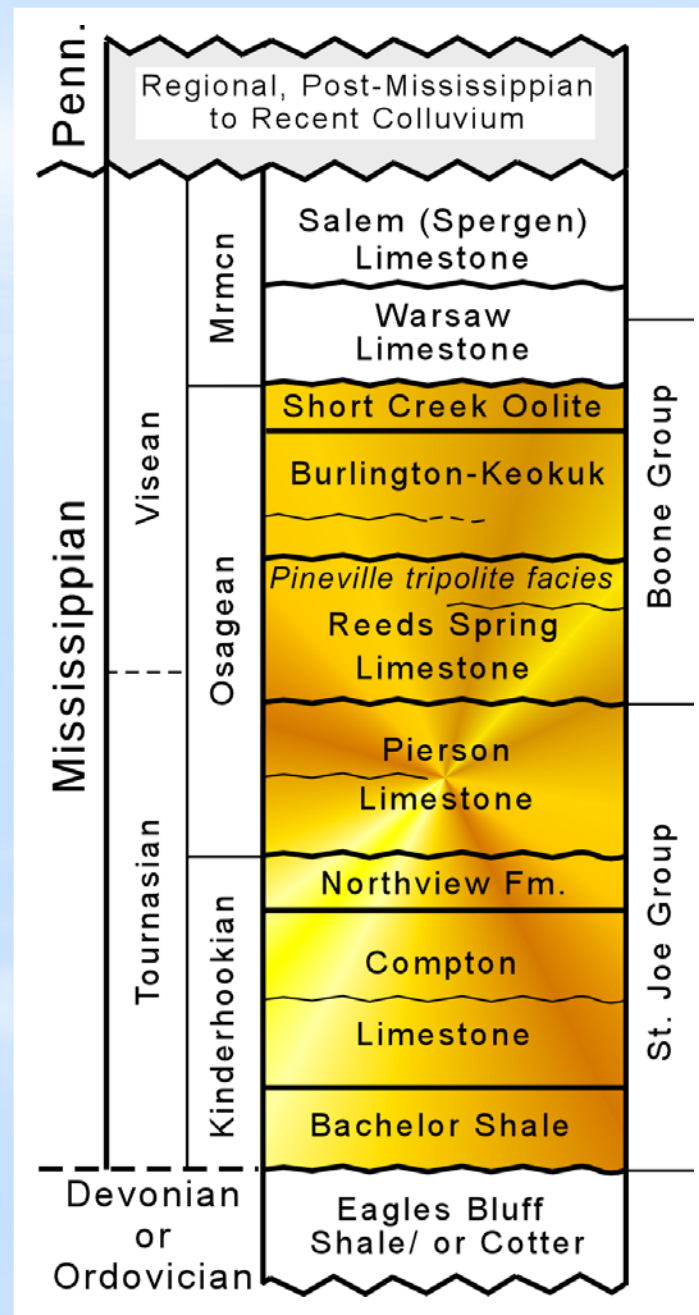
Rather, we'll show that *Ouachita tectonism
effecting the Mid-Continent was initiated in
Kinderhookian, and continued through Osage
time.*



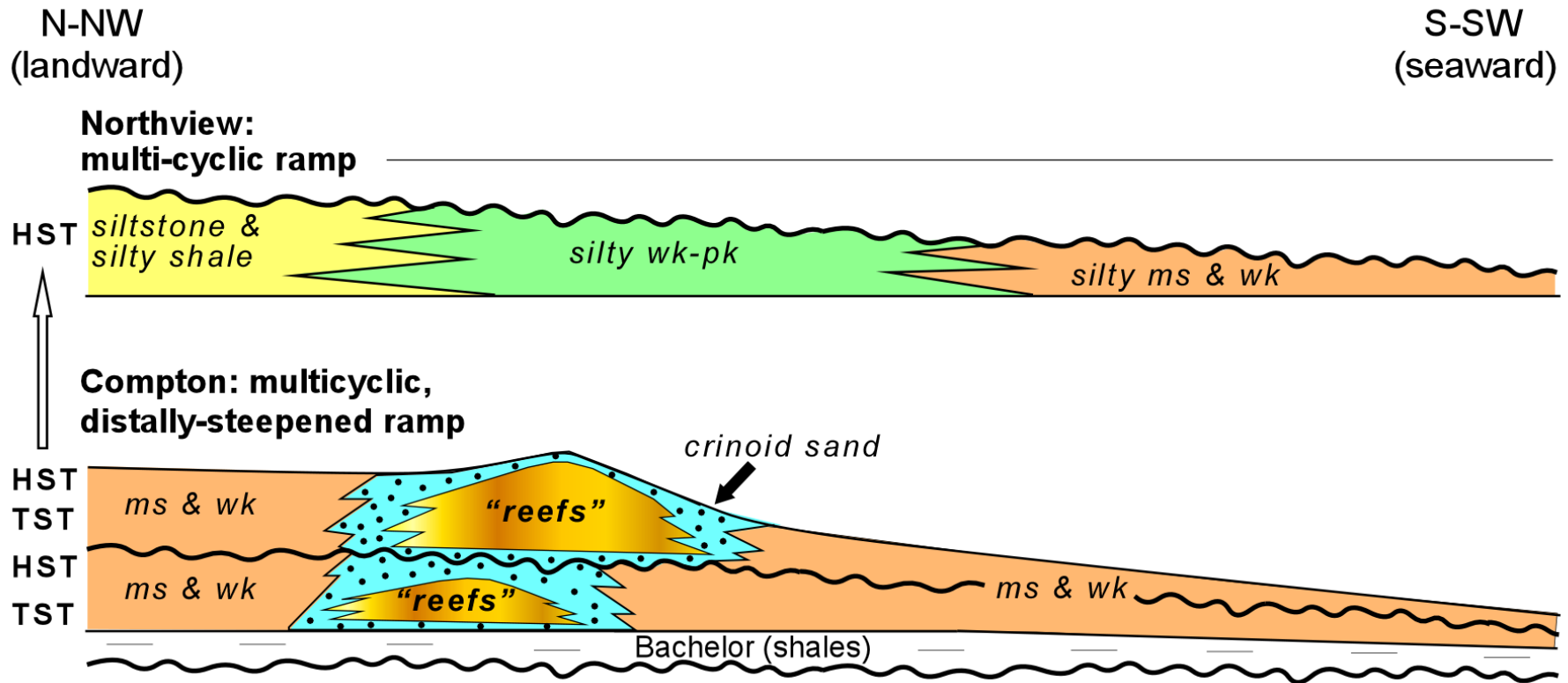
Area of Study



Stratigraphy of the Outcrop Area



Depositional systems of Kinderhookian COMPTON & NORTHVIEW



Multicyclic, shallow water ramps where platform architecture is aggradational.

Depositional systems of lower Osagean PIERSON LIMESTONE

N-NW
(landward)

S-SW
(seaward)

Pierson: heavily eroded, multi-cyclic, distally-steepened ramp

transitional Burlington-Keokuk type limestone

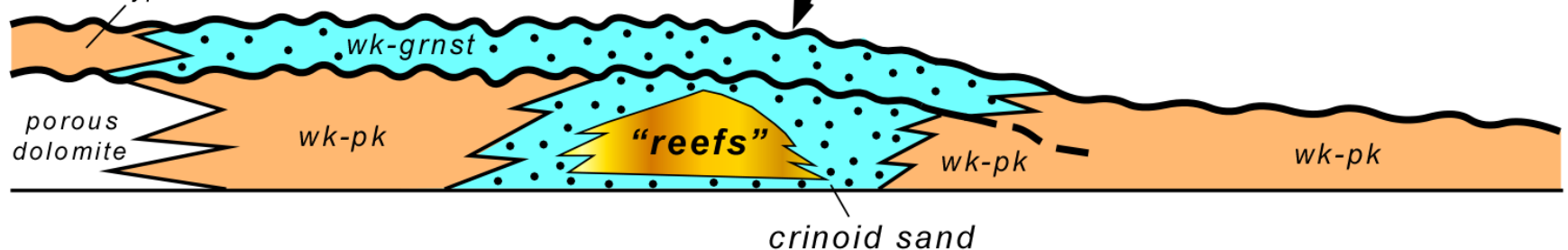
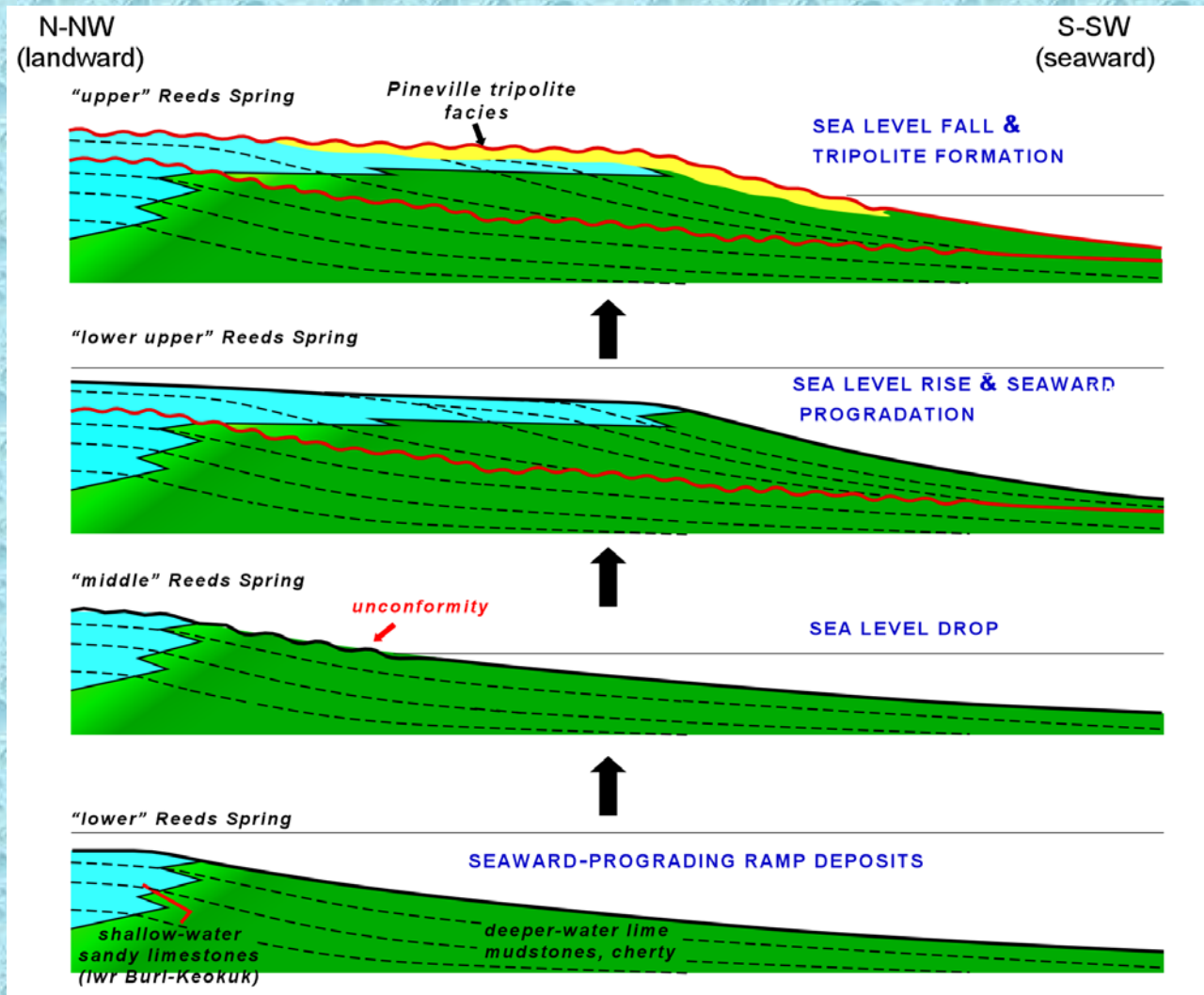


Figure 43. Inferred depositional model and platform architecture of the Pierson Formation in the field area.

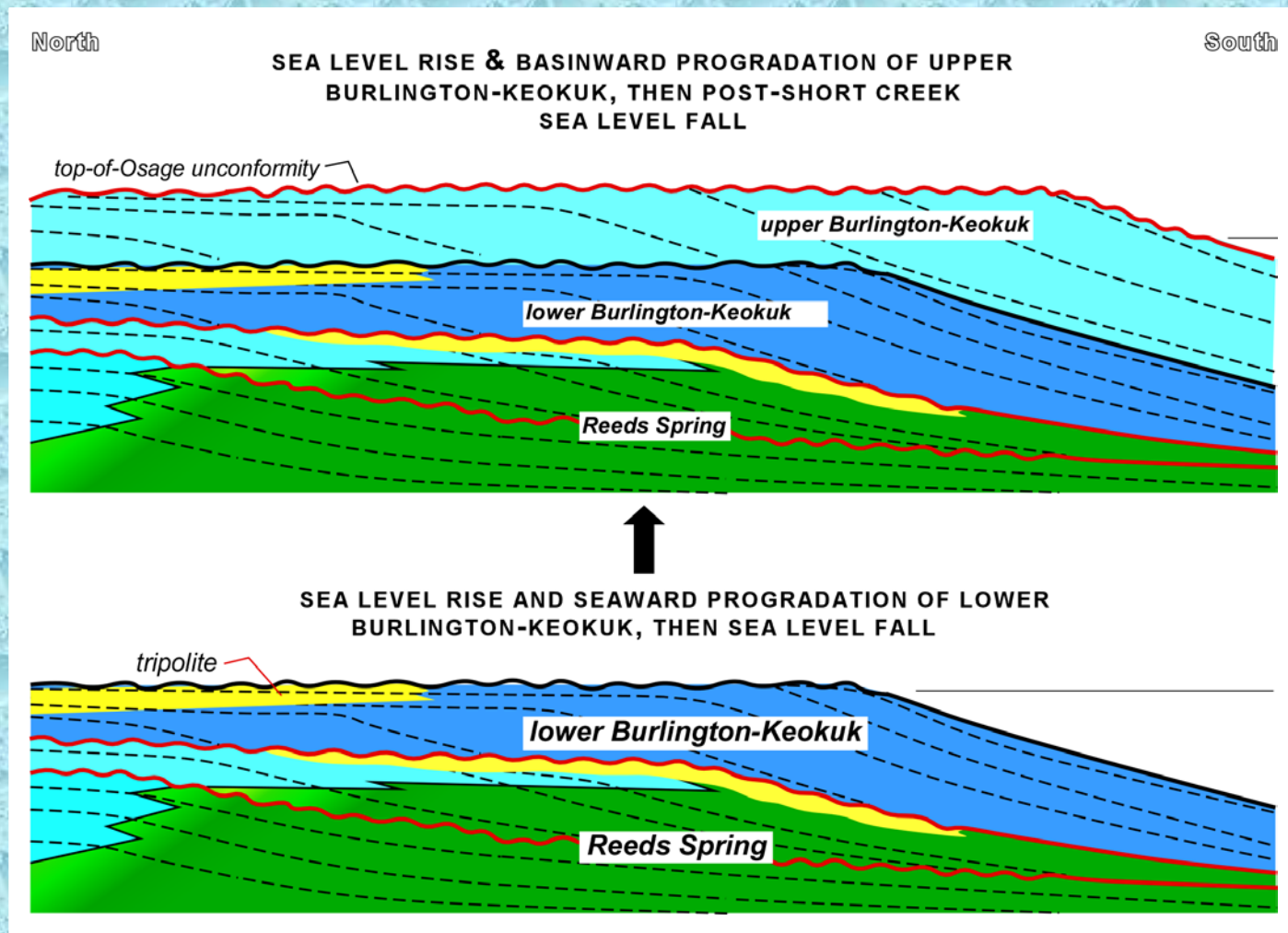
Multicyclic, *distally-steepened* shallow water ramps where platform architecture evolves to a progradational nature.

The Bachelor, Compton, Northview and Pierson are genetically related and make up the St. Joe Group

Depositional systems of Osagean REEDS SPRING & BURLINGTON-KEOKUK LIMESTONE



Deeper-water muds that grade, updip, into shallow water grainstones of the lower Burlington-Keokuk. Pronounced deepening of the ramp where platform architecture became strongly progradational.

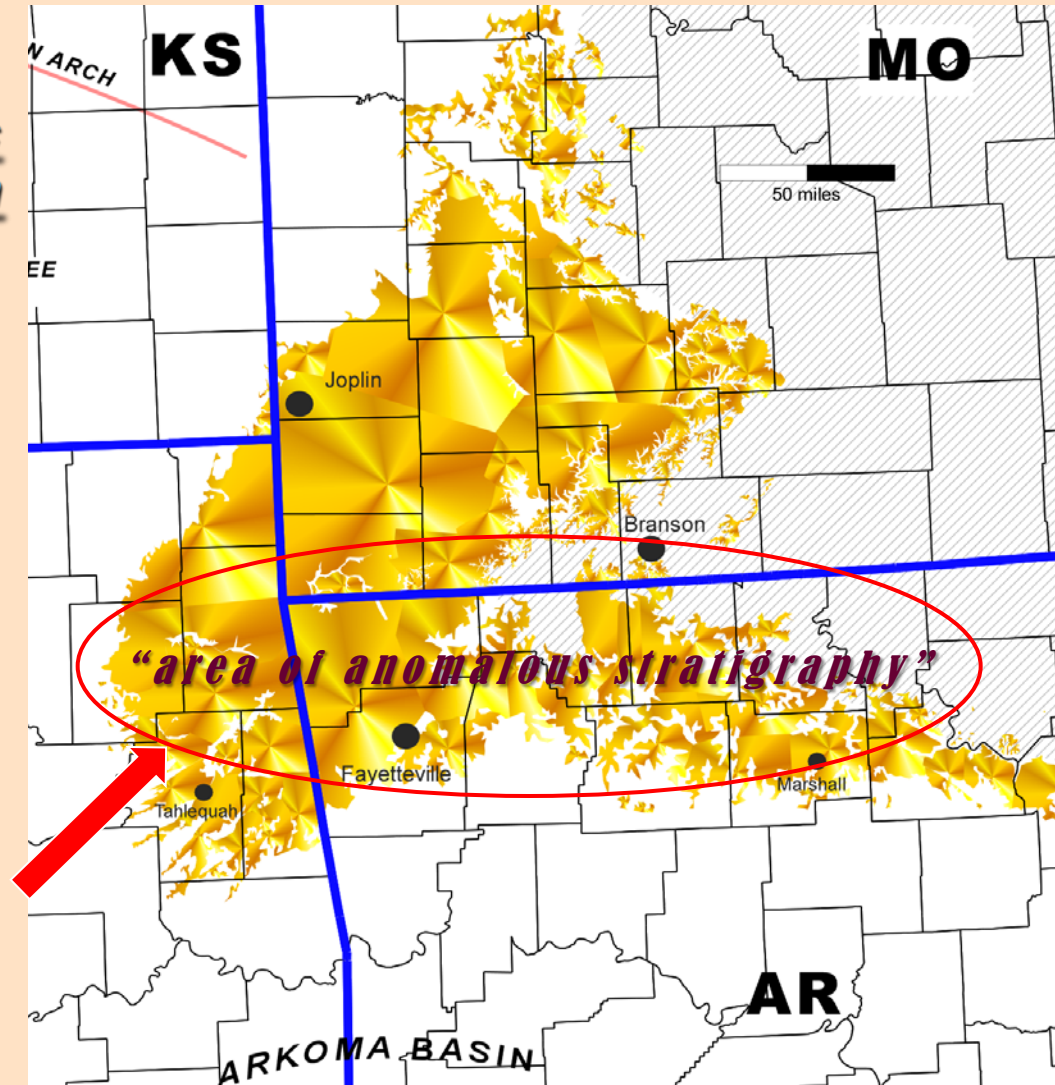


Final progradation of shallow water Burlington-Keokuk Limestone over deeper-water Reeds Spring Limestone.

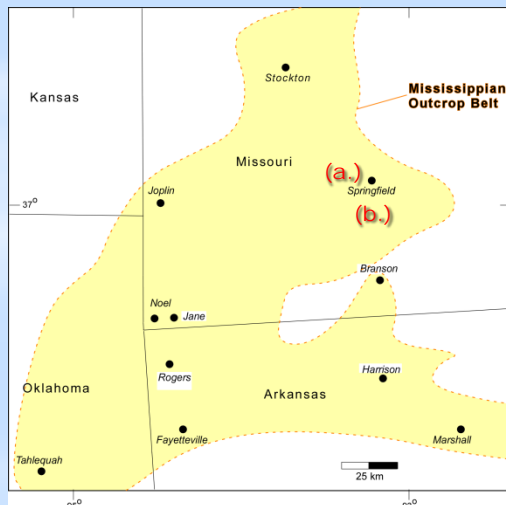
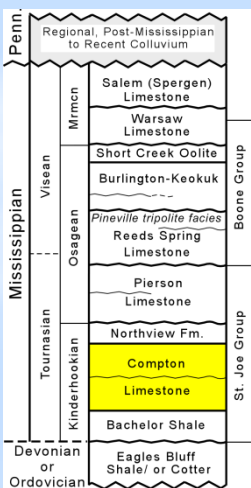
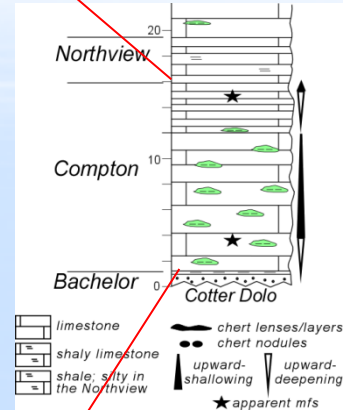
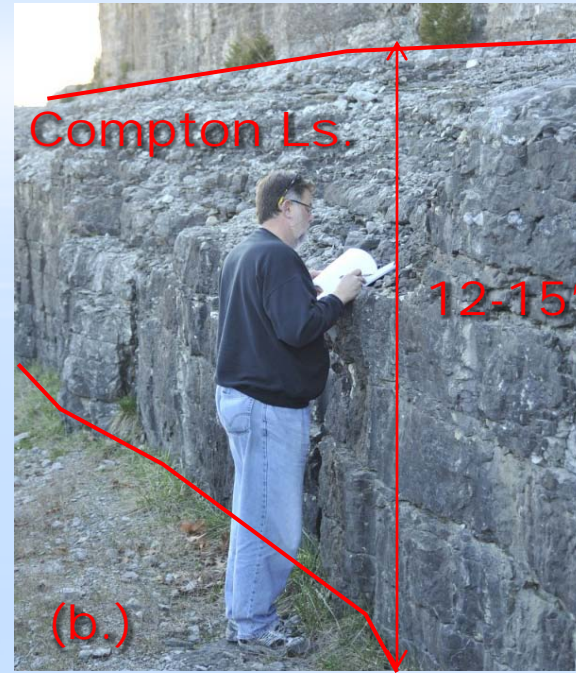
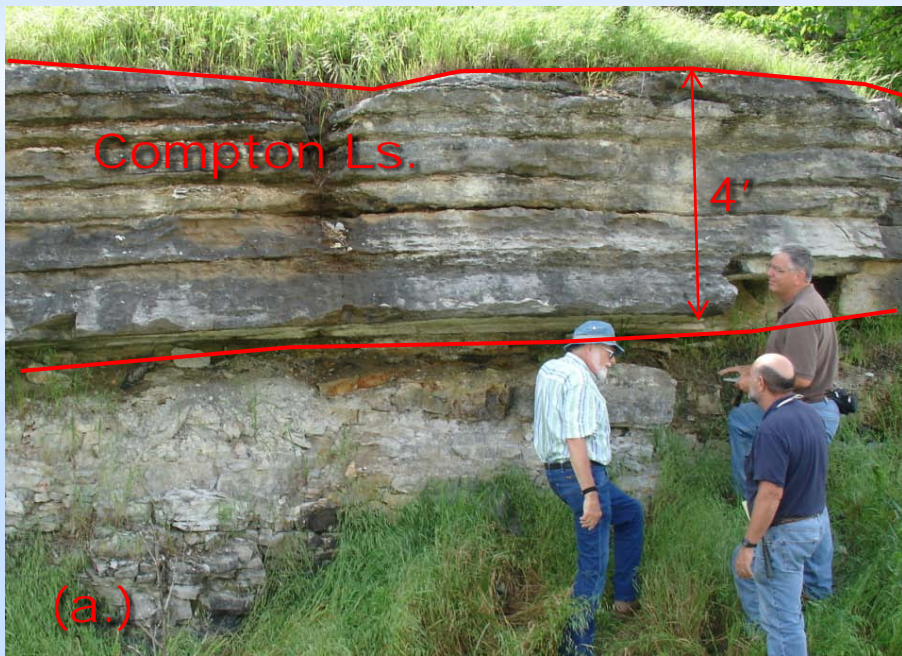
The aforementioned described depositional systems comprise the fundamental sequence-stratigraphic attributes of the lower Mississippian (KH-OS) and are recognized westward into the subsurface.

HOWEVER...

...the Compton through the Burlington-Keokuk show evidence of syndepositional tectonism in the southern part of the study area that overprinted the sequence-stratigraphic architecture of the rocks.



"NORMAL" COMPTON LIMESTONE

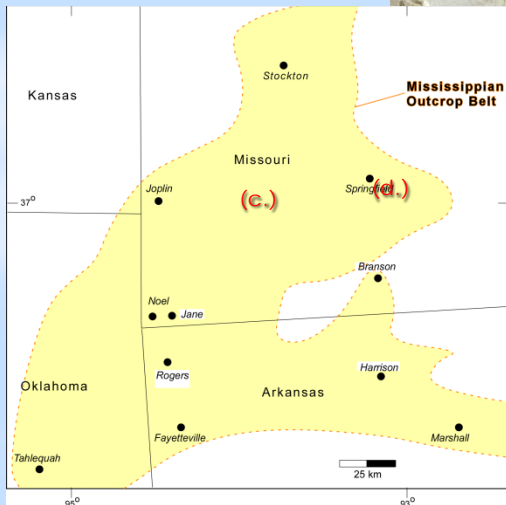
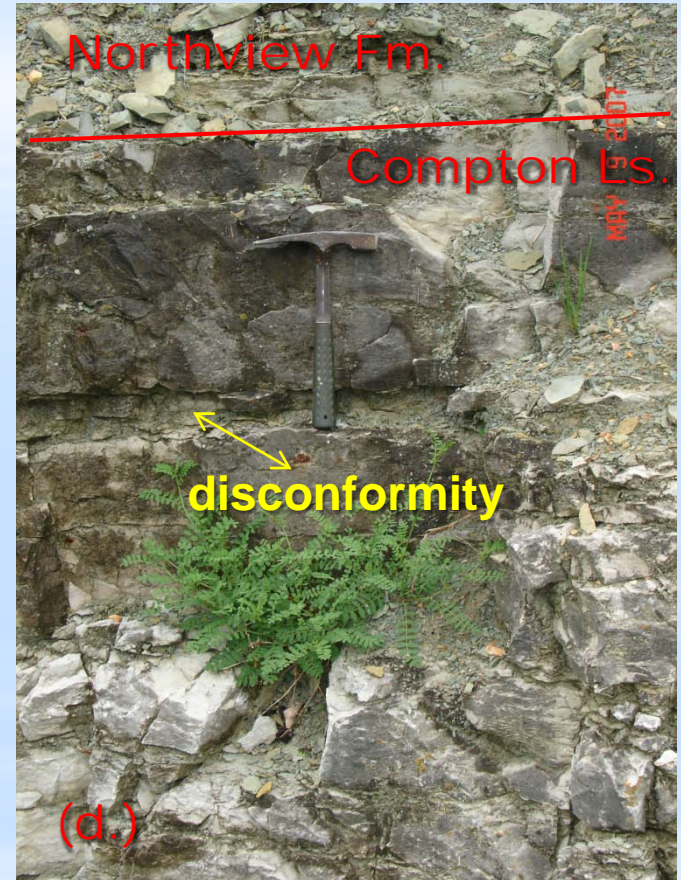


Lithology & Average Thickness

*Shallow water lime MS - PS
w/green shale wisps.
Average thickness 4 - 15 feet.*

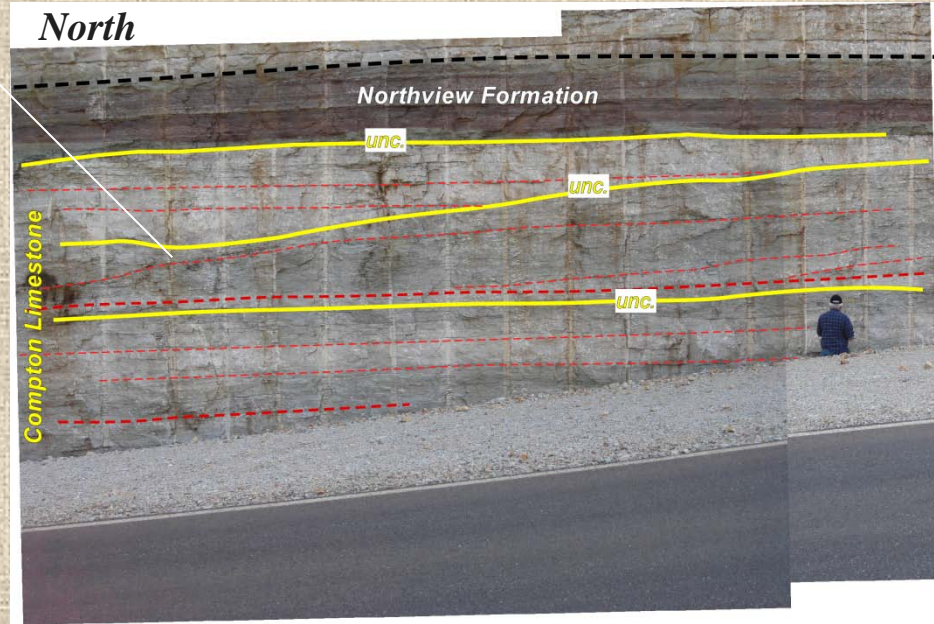
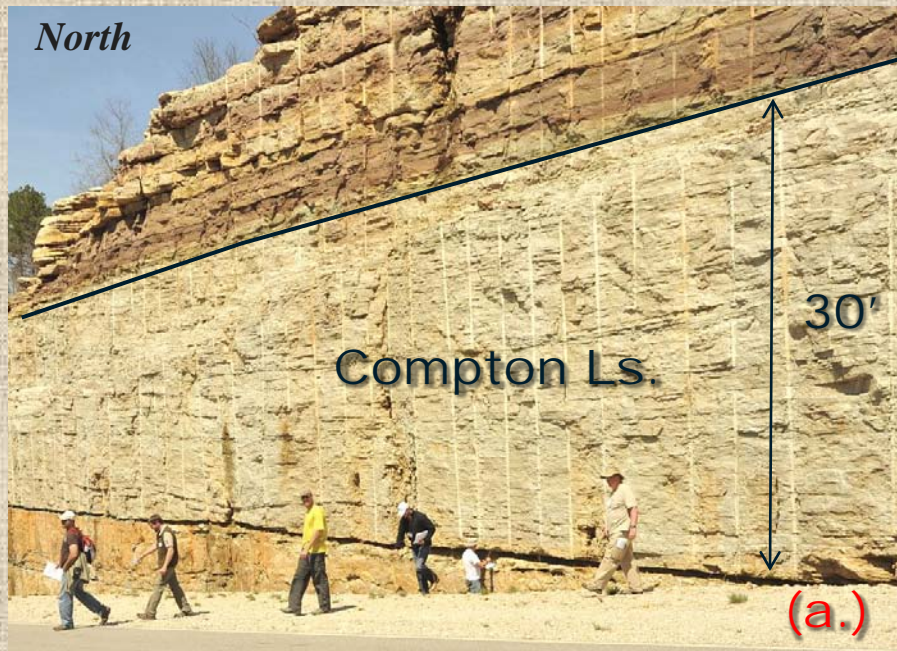
"NORMAL" COMPTON LIMESTONE

**Limestone-clast
conglomerate at
disconformity**

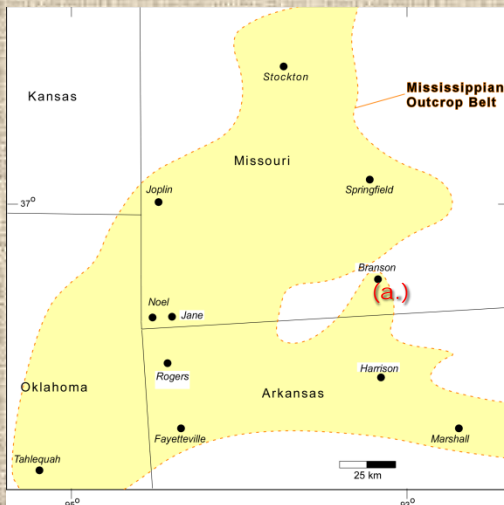


*Regional disconformity variously
indicated by limestone-clast
conglomerate, red and green shale, or
quartz sand*

"ANOMALOUS" COMPTON LIMESTONE



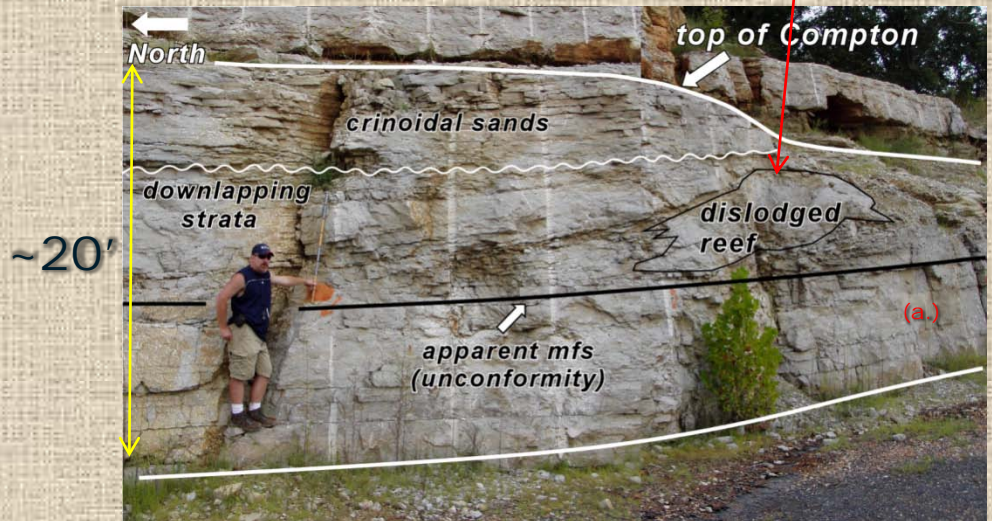
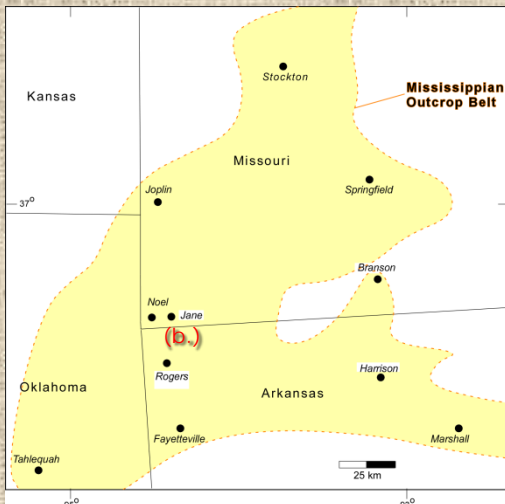
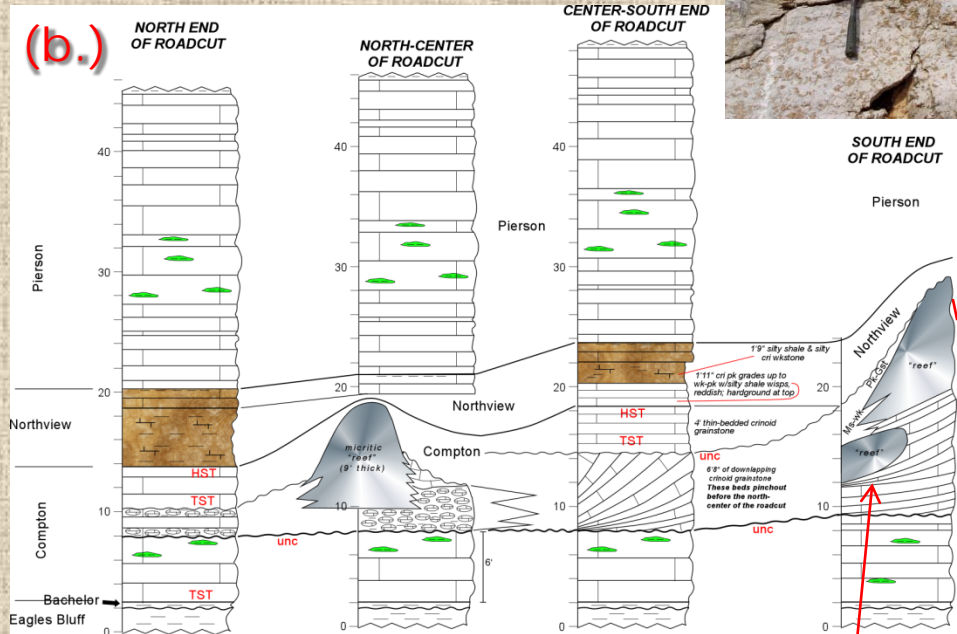
Here, the Compton is separated from the Northview by a low-angle, angular unconformity. Elsewhere, the upper Compton is part of an HST that includes the Northview.



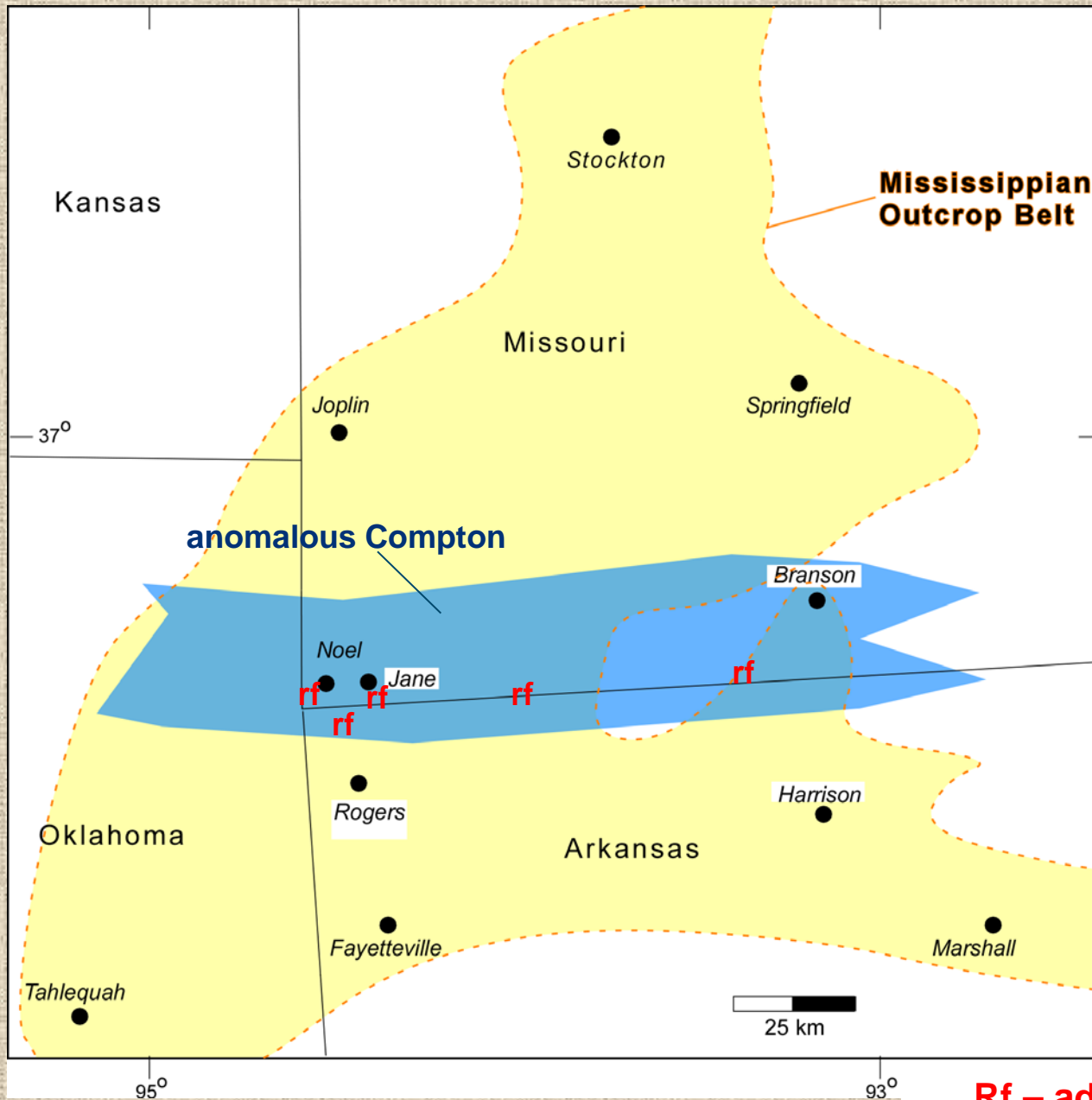
18 – 30+ feet of very shallow-water, coarse, crinoid grainstones with numerous unconformities and downlapping units that prograde northward.....counter to, presumed, regional paleodip!

“ANOMALOUS” COMPTON LIMESTONE

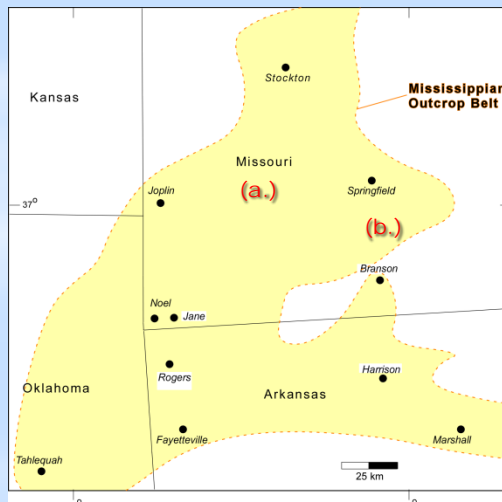
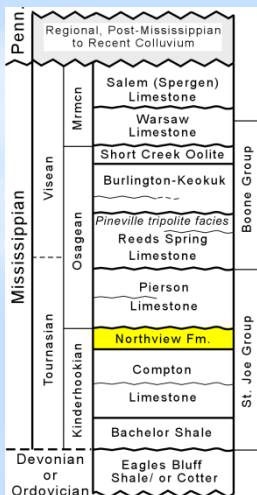
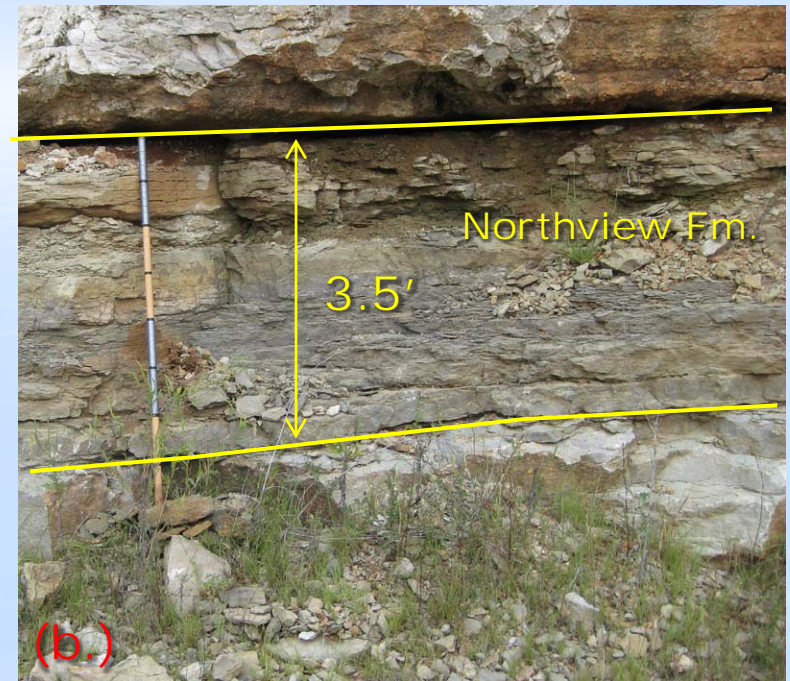
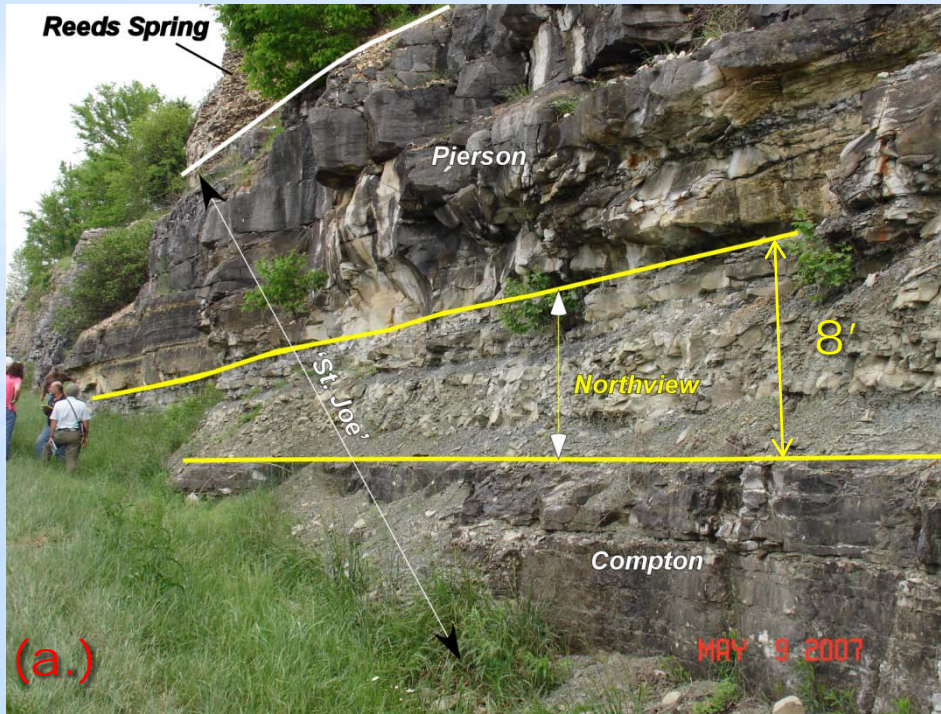
Micritic buildups occur within the crinoid sands and are common in this belt; they too downlap to the north (!) with multiple unconformities within the Compton – here, with evidence of vadose exposure (brown silt-filled vugs).



"ANOMALOUS" COMPTON BELT



"NORMAL" NORTHVIEW FM.



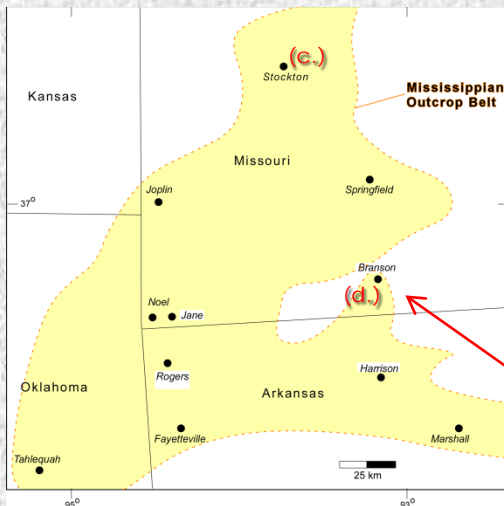
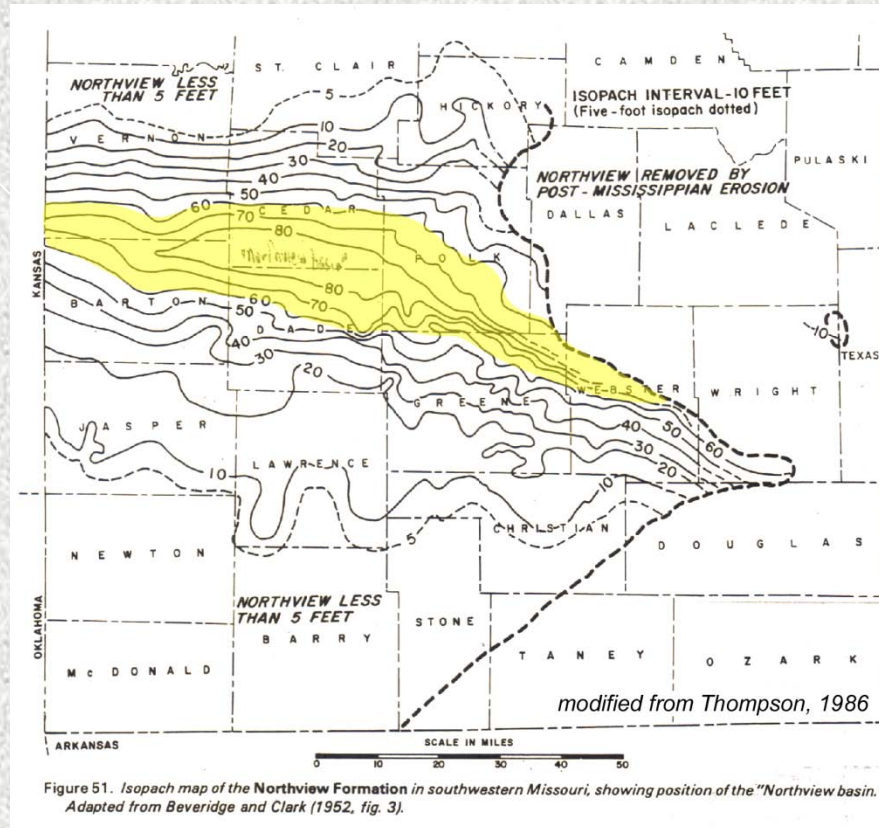
Lithology & Average Thickness

Mixed lithologies of shale, silty shales, siltstones, argillaceous limestone and limestone.

These deposits represent shallow (shallower than Cptn) water deposition.

Average thickness 2 - 8' feet.

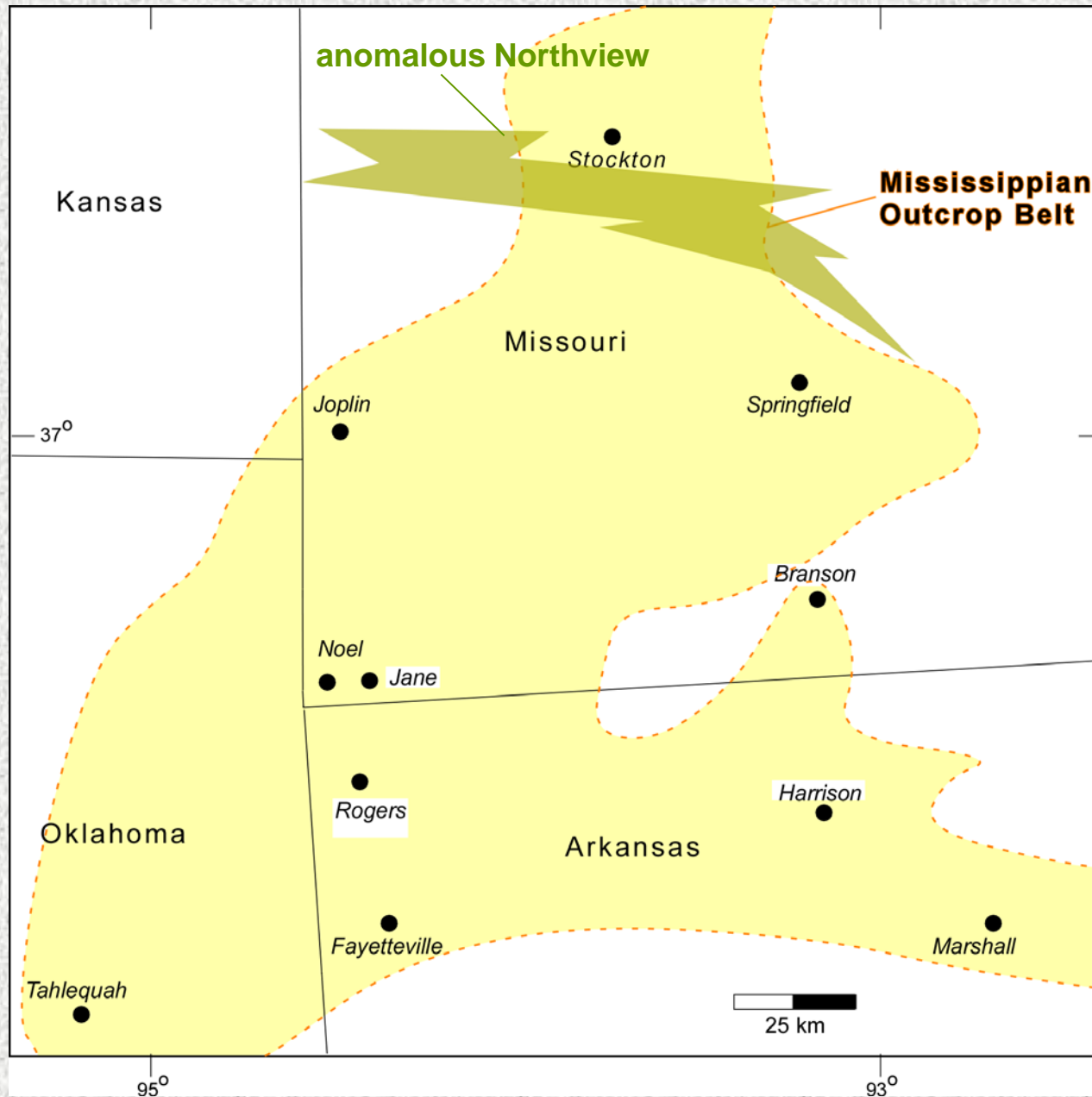
"ANOMALOUS" NORTHVIEW FORMATION



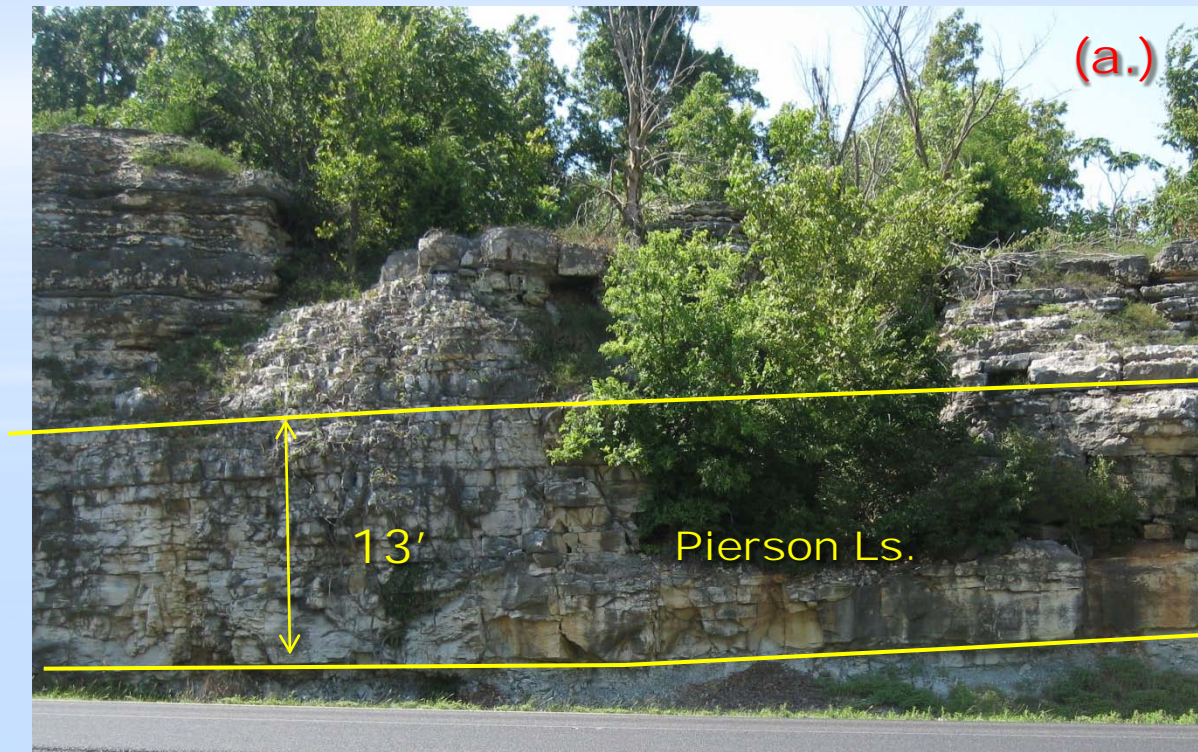
50 – 80+’ of siltstones and shaly siltstones of shallow-marine nearshore origin deposited in a, more-or-less, east-west trending trough.

Note, locality “d” in the adjoining map shows the area at which paleosols were recorded in the Northview Formation.

"ANOMALOUS" NORTHVIEW BELT

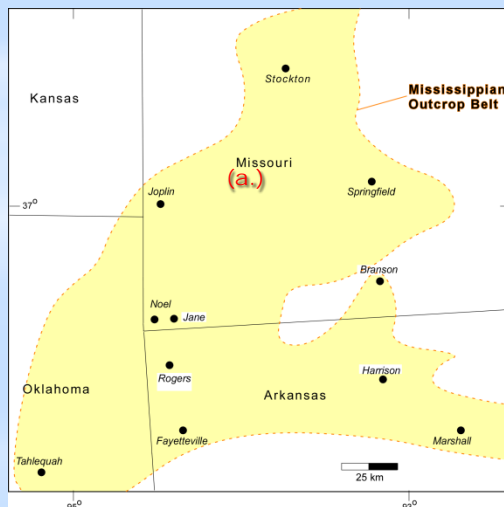
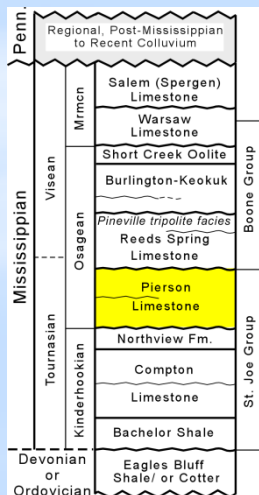


"NORMAL" PIERSON LIMESTONE

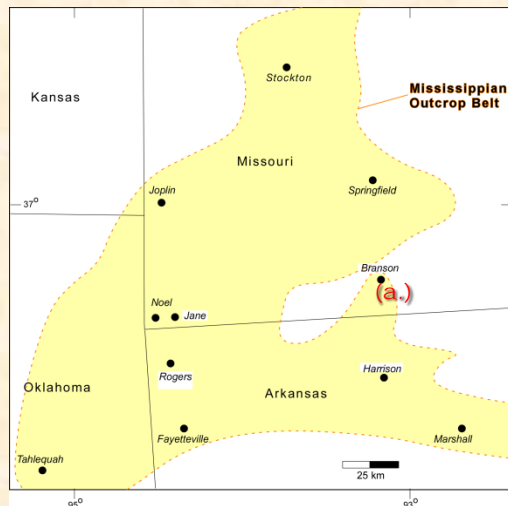
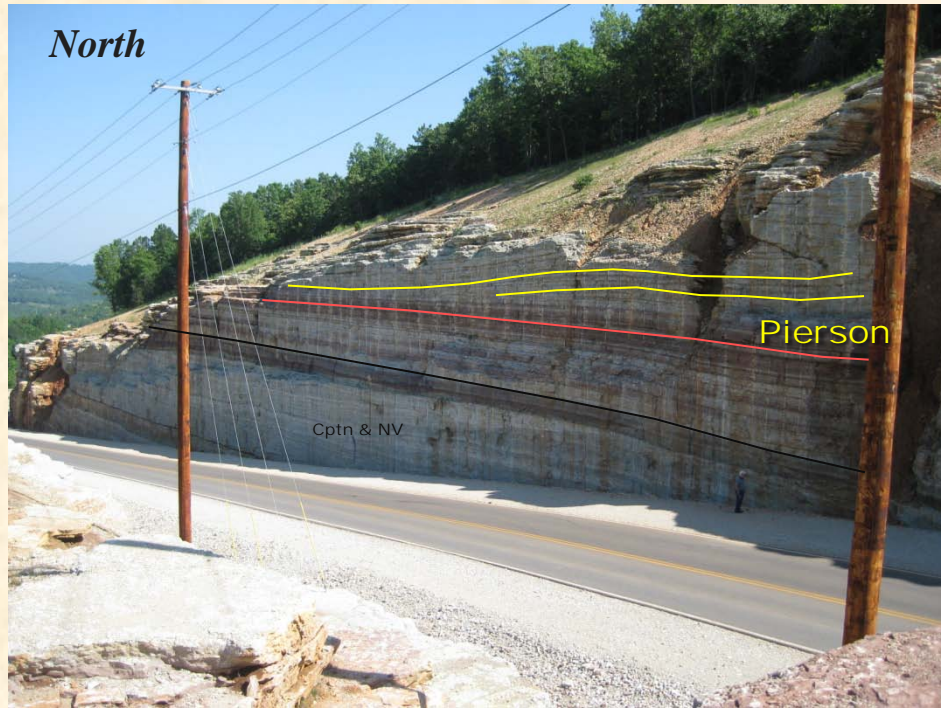
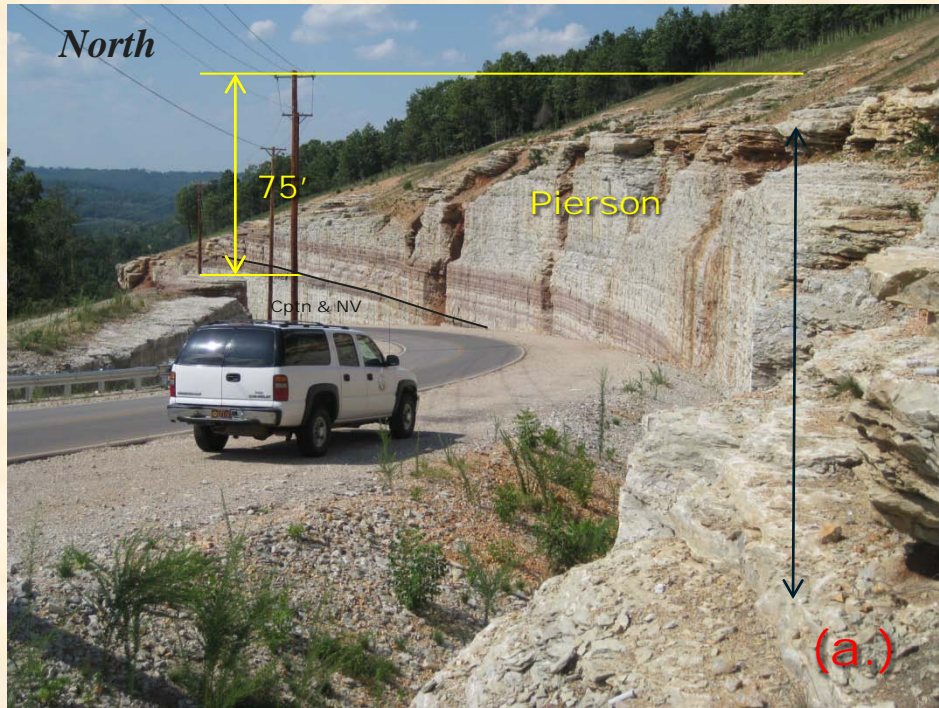


Lithology & Average Thickness

The Pierson is composed of relatively shallow-water lime MS-PS, and in the northern part of the study area, dolomitic. Average thickness 4 -18 feet.

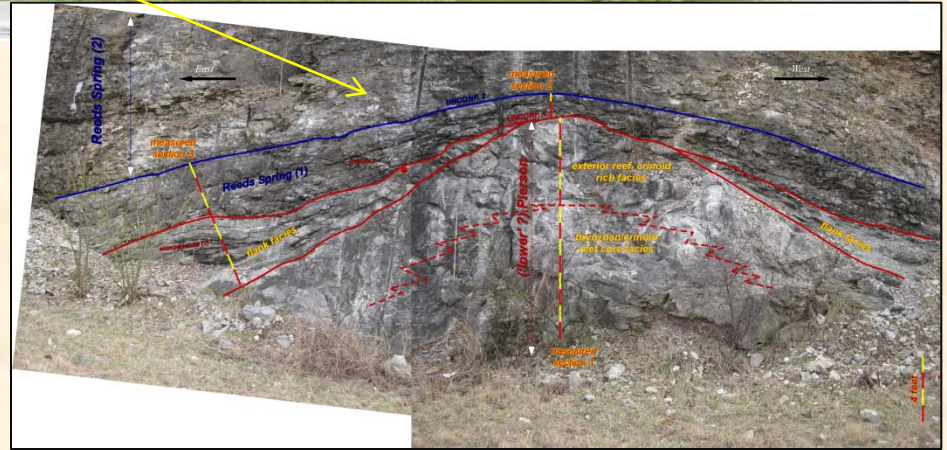
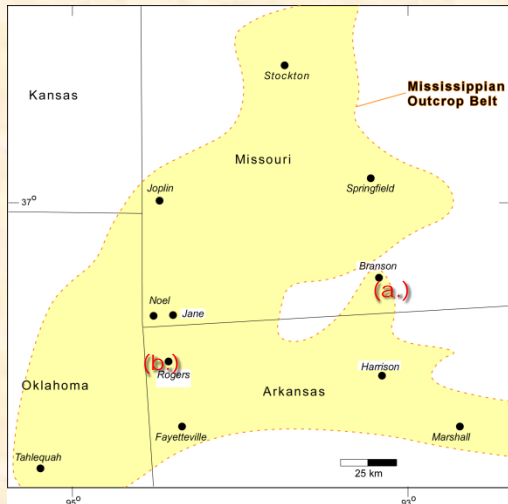
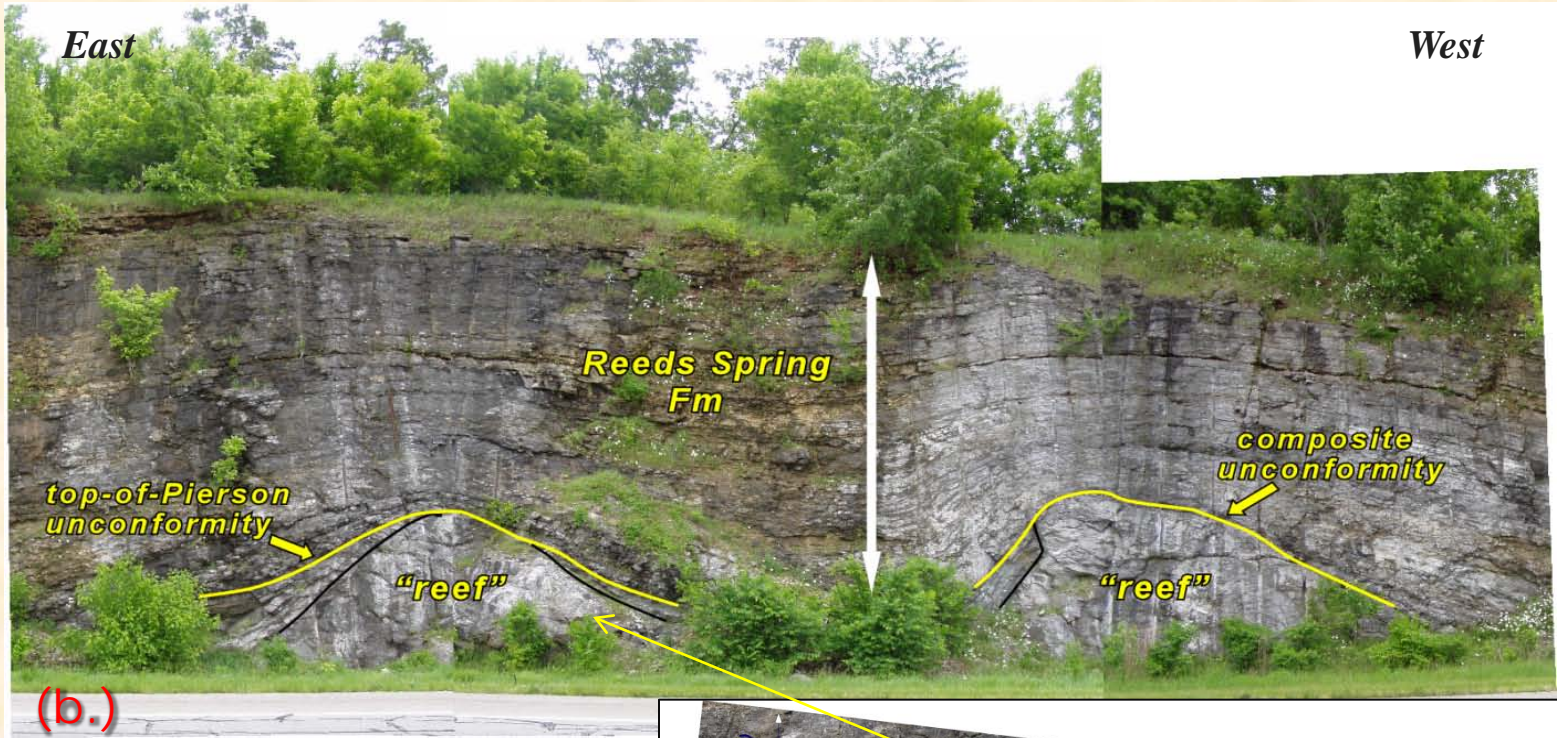


"ANOMALOUS" PIERSON LIMESTONE



*45 - 75+ feet of very shallow-water,
coarse-very coarse, crinoid
grainstones that downlap and
prograde northward.*

"ANOMALOUS" PIERSON LIMESTONE

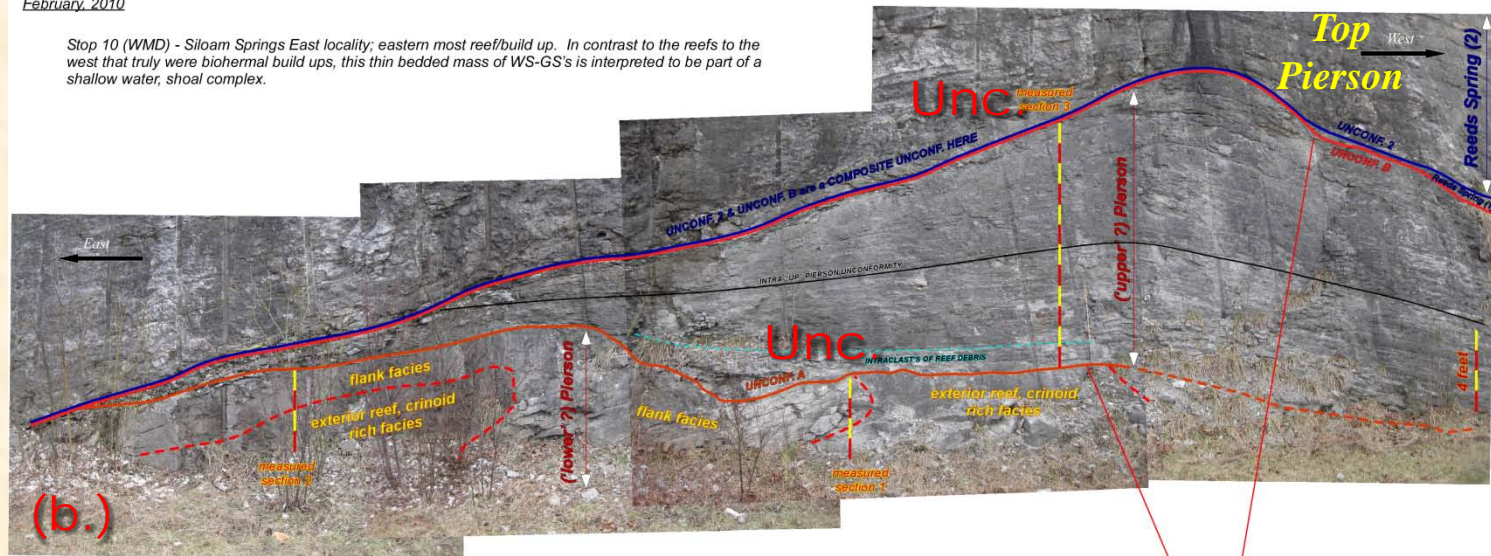


Bryozoan cored reefs with intra-formational unconformities.

"ANOMALOUS" PIERSON LIMESTONE

February, 2010

Stop 10 (WMD) - Siloam Springs East locality; eastern most reef/build up. In contrast to the reefs to the west that truly were biohermal build ups, this thin bedded mass of WS-GS's is interpreted to be part of a shallow water, shoal complex.



15 - 30' Reeds Spring "2": (thickness from photos) as described previously with the western & middle reef, with a sliver of Reeds Spring "1" on the western edge.

2' 6" - 9' 6" (from measured sections 3) thin to medium bedded, medium to dark gray alternating lithologies of wackestone and packstone with various sizes, small to large, crinoids and oxidized marcsite. Thin lenses and wisps of green to greenish gray shale are present however, less abundant than below. In contrast to the previously described unit, the chert here is different and more plentiful. The chert occurs as blebs and discontinuous bands (2-4" thick) of dark gray/bluish gray to black mottled chert with a relative abundance of crinoid debris. According to Beau, after scaling and walking out this section, the overall limestone lithology shallows upward in the section as the units become more of a packstone to grainstone. This unit is truncated by unconformities B & 2.

2' 6" - 3' 6" (from measured sections 3) thin bedded (2-3") medium to dark gray wackestone with lenses of packstones composed of various size, small to large, crinoids. Thin lenses and wisps of green to greenish gray shale are present and occur between the thin bedded limestones. Chert is sparsely disseminated throughout the section as thin 1-2" discontinuous bands of very dark gray chert with light gray interiors a very few crinoids. It was noted that the chert bands were truncated on the east side and hence, the intra-upper Pierson unconformity; submarine erosion?

1' 6" (from measured sections 3) thin to medium bedded (2-5") of medium gray crinoid wackestone to packstone with thin green to green-gray fossiliferous shale wisps, some shales are more continuous and up to 1/2 - 3/4" thick. This section grade upward into--

0' 10" - 1' 8" (from measured sections 3) thin bedded (2-3") of light to medium gray crinoid packstone to grainstone, medium grained (with some larger fragments) and oxidized marcsite. The thin limestone units are separated by 1/4" to 2" thick greenish-gray fossiliferous shale beds with crinoids; the basal shale on the unconformity is upwards of 3" thick. Throughout this interval there are quarter to softball sized clasts of reef crest and/or flank bed facies composed primarily of crinoid P.G. Some of the clasts are dark gray muddy wackestone; source? possibly the outer 2-3' rim of wackestone record-ward into--

ns 1&2) the flank beds between the measure sections consist of medium bedded (2-6") light s, all separated by thin greenish gray shale and shale wisps. Those overlying the reef in gray wackestone with various sized (some very large) crinoid grains, and some thin discontinuous what we feel is an oblique view at possibly 2 different reefs it was tough to outline accu-facies too are truncated by Unconformity A.

massive, non-bedded light gray crinoid packstone to grainstones, med to coarse grained, and tem reefs it is inferred that these facies are truncated (by Unconformity A) exterior or reef l above.

REEDS SPRING

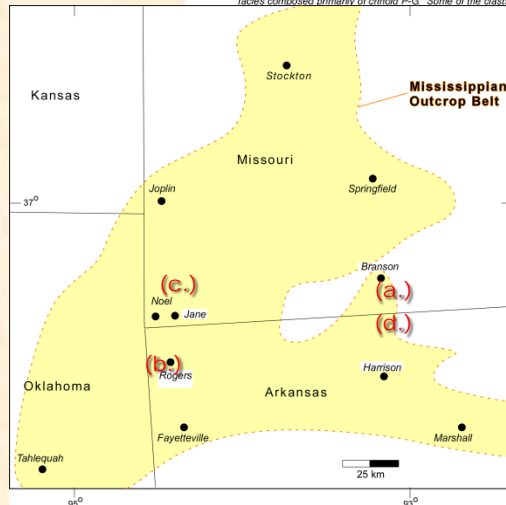
PIERSON

'upper' Pierson

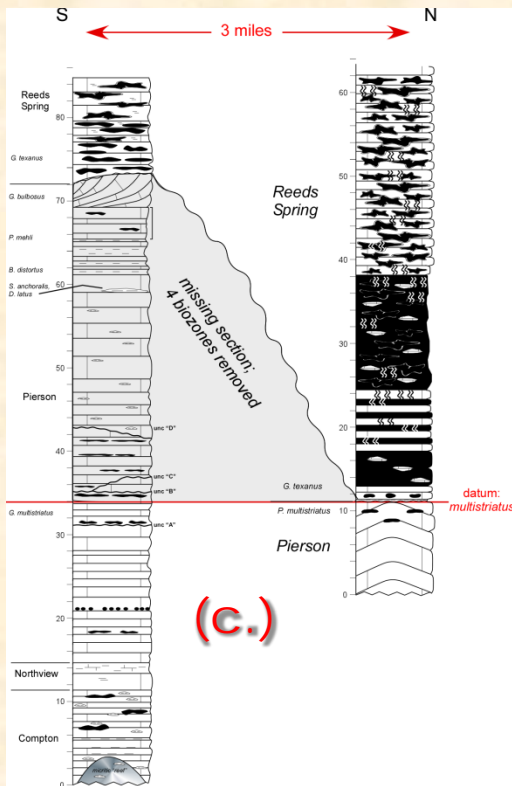
'lower' Pierson

Based on correlation and recognition of unit specific lithologies, it is our opinion and interpretation that Unconformities B and A, shown here, become one composite unconformity behind/beneath the talus slope just to the west of this photo and correlates to Unconformity 1 shown in the previous reef descriptions. Basis for this is, the lack of 'upper' Pierson lithologies* are seemingly missing to the west.

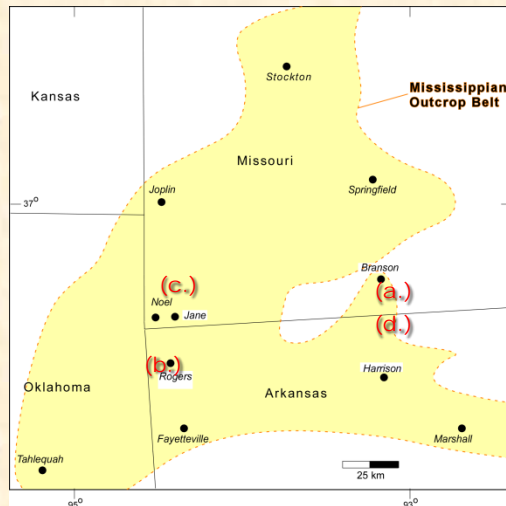
*Note: with the possible exception in the vicinity of the small reef between this locality and the western reefs, where crin PS-GS with black chert was found overlying reef flank facies and was capped by Reeds Spring.



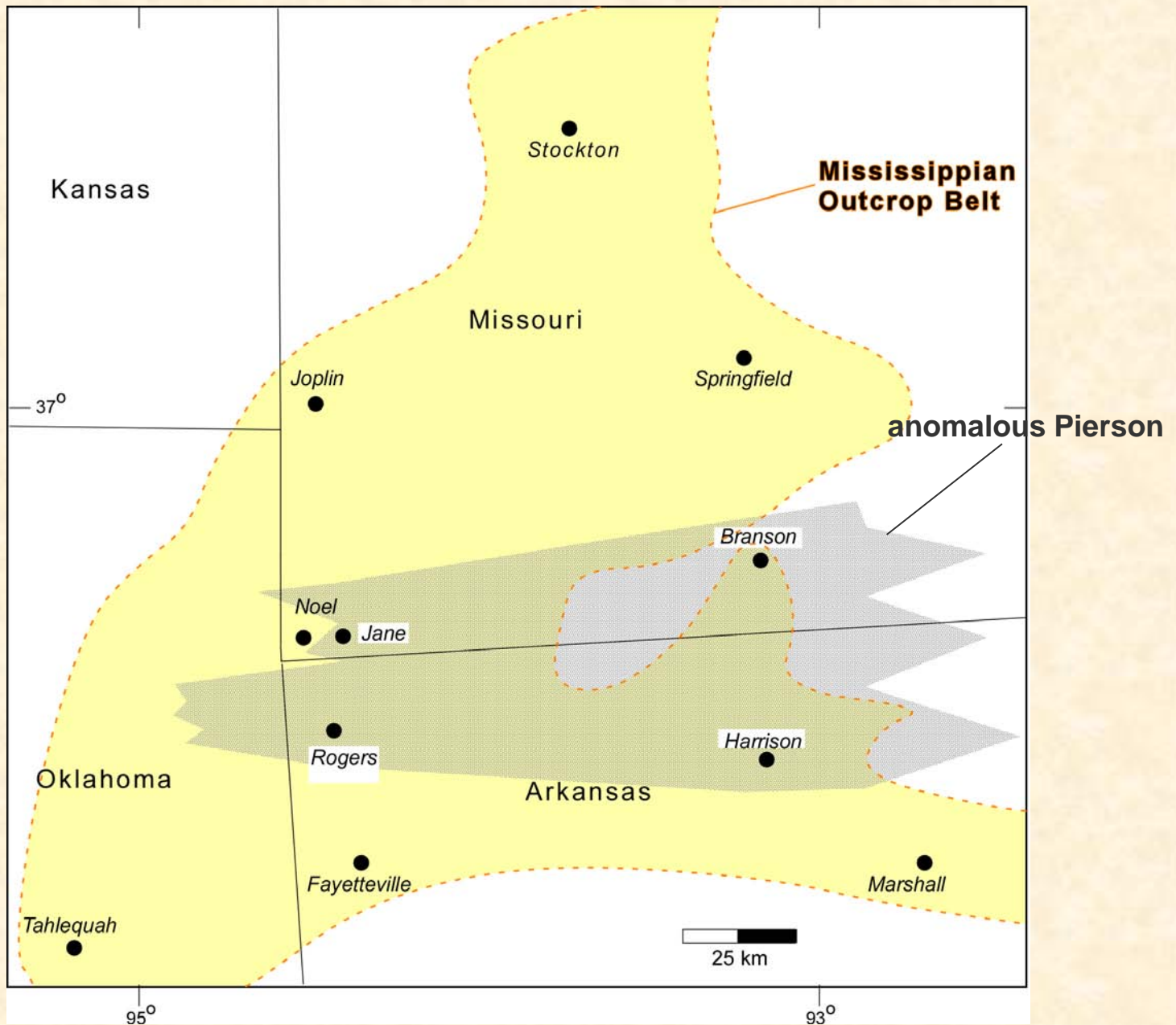
"ANOMALOUS" PIERSON LIMESTONE



Angular unconformities at the top of the Pierson; breached East-West trending anticlines!

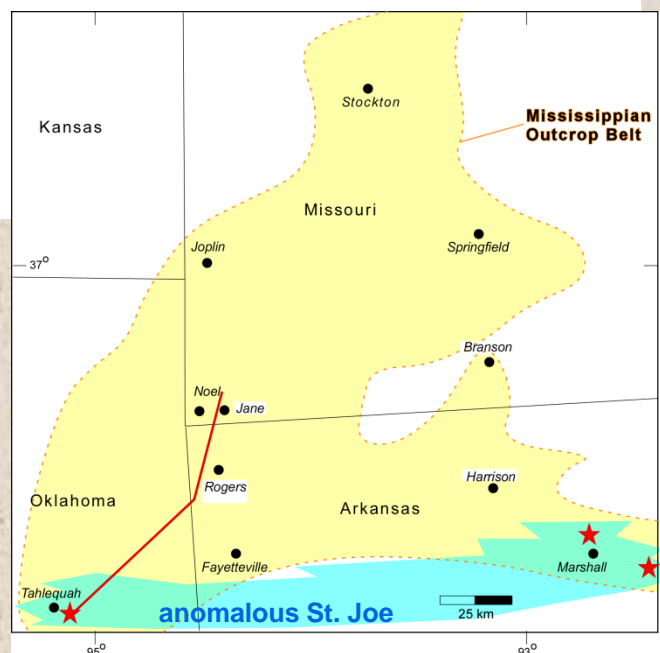
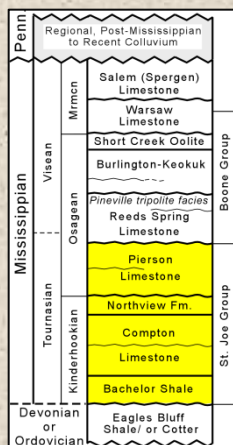
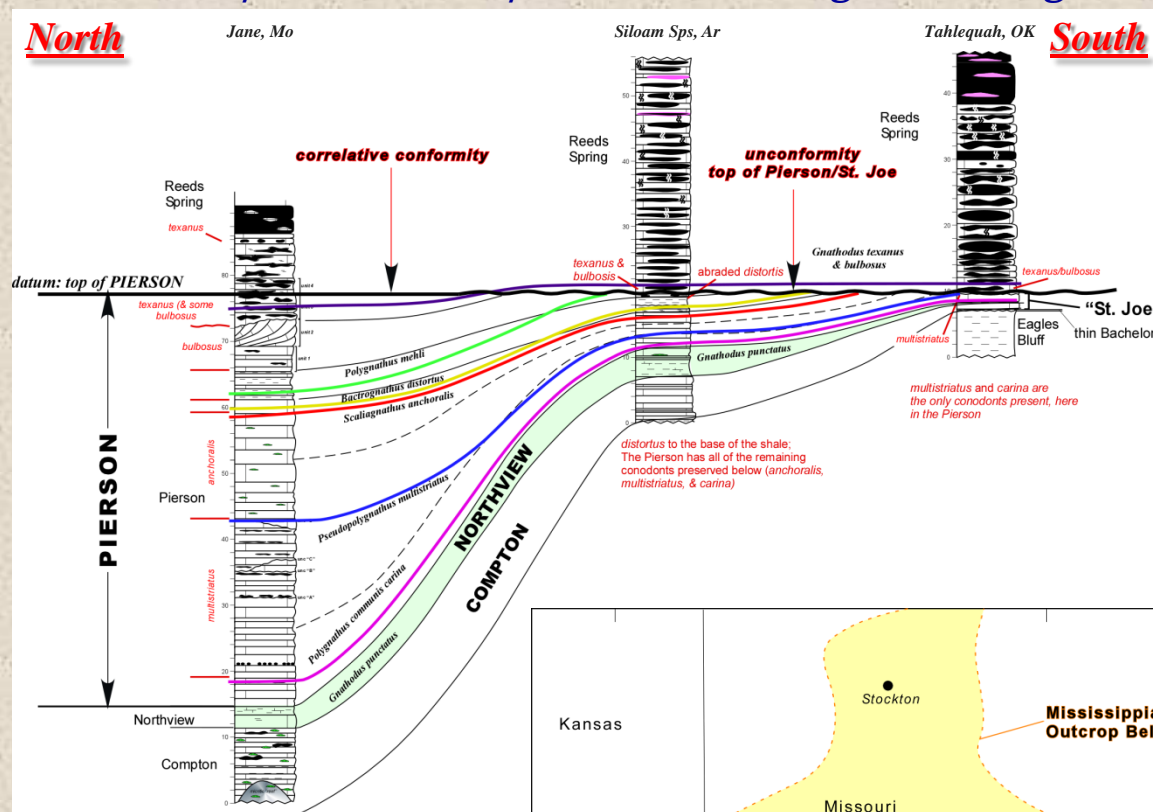


“ANOMALOUS” PIERSON BELT



"ANOMALOUS" ST. JOE (Bachelor, Compton, Northview, Pierson)

These Kinderhookian and basal Osagean units represent an overall lower order sequence composed of 3, regional, higher frequency T-R cycles.



Not only do the units of the St. Joe erosional thin to the south, they also depositional thin (avg. 3.5', total) and maintain their shallow-water characteristics!!

These units are NOT condensed. In addition, soil pisolites are typically seen at the top of the St. Joe, beneath the next sequence of Reeds Spring/Burlington-Keokuk.

....this greatly contrasts the existing models....

OUR AREA OF STUDY

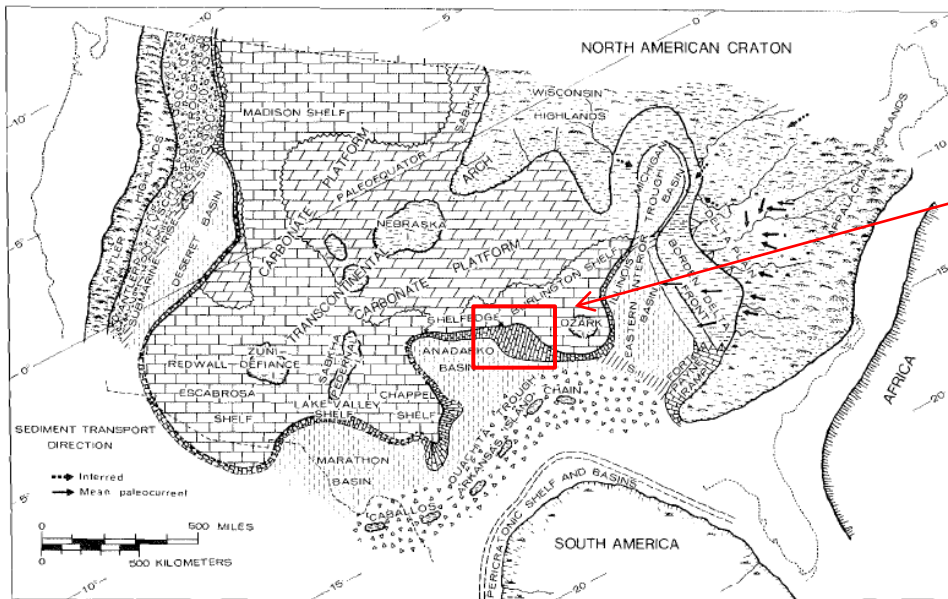
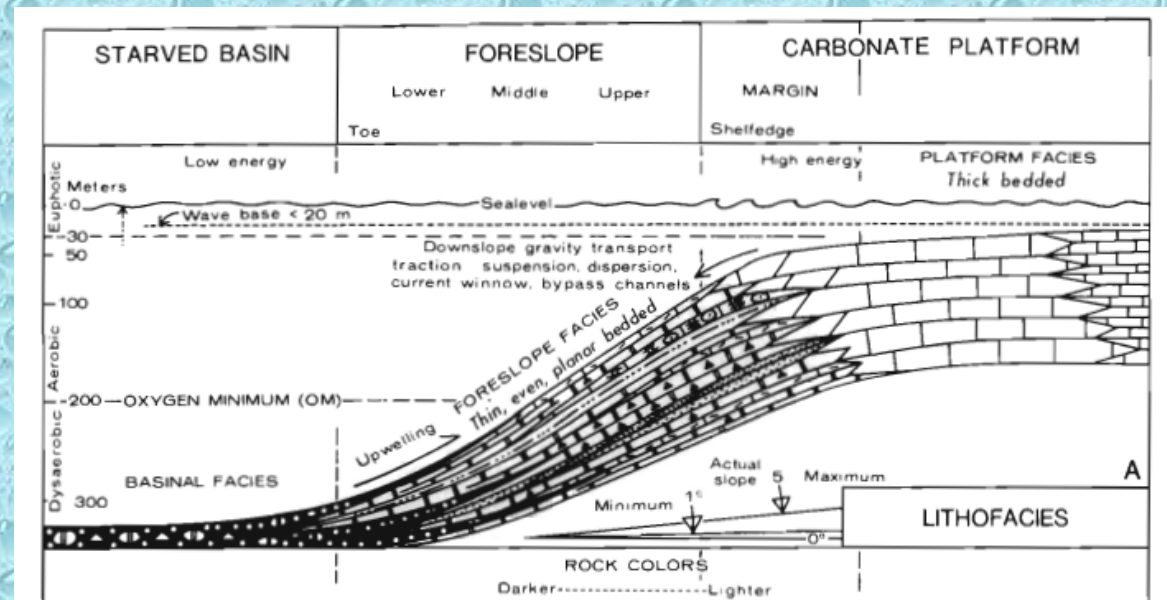
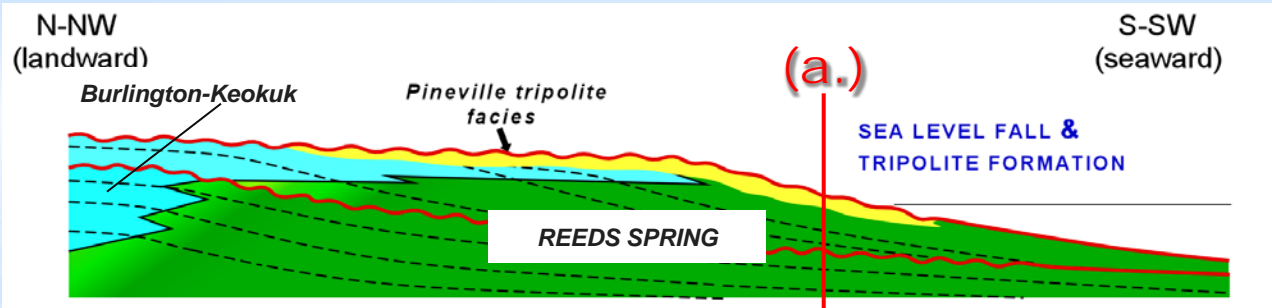


FIG. 5.—Paleogeographic and lithofacies map of conterminous United States at time of *anchoralis-latus* conodont Zone. Data unavailable in blank areas.

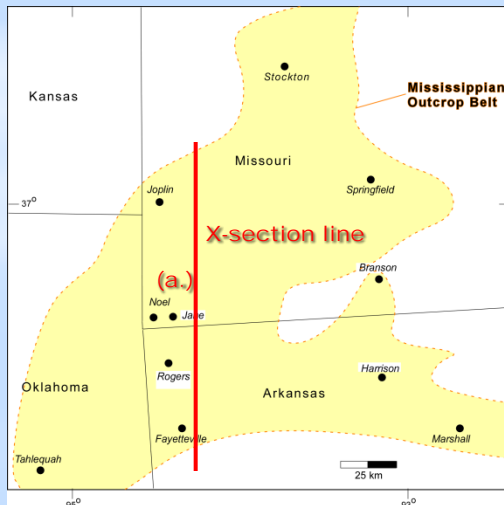
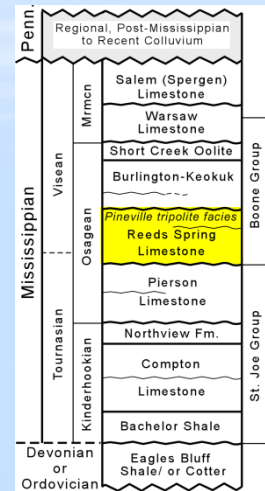
...the basin wasn't
to the south,
it was to the north
during KH & early
OS!



"NORMAL" REEDS SPRING LIMESTONE



Deep(er) water, cherty lime MS. Downdip equivalent to the shallow-water, Burlington-Keokuk. When exposed during Mississippian time substantial tripolite is developed with attendant meteoric exposure. Average thickness 75 -120 feet.

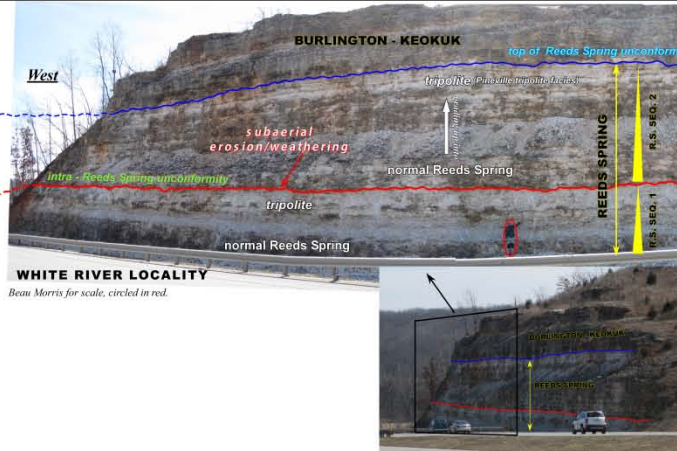


"ANOMALOUS" REEDS SPRING LIMESTONE

North (A)

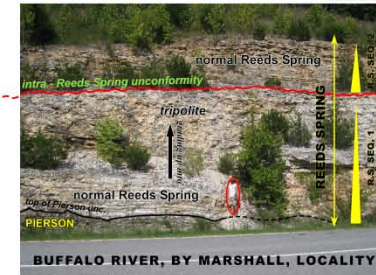


Robert Turner for scale, circled in red.



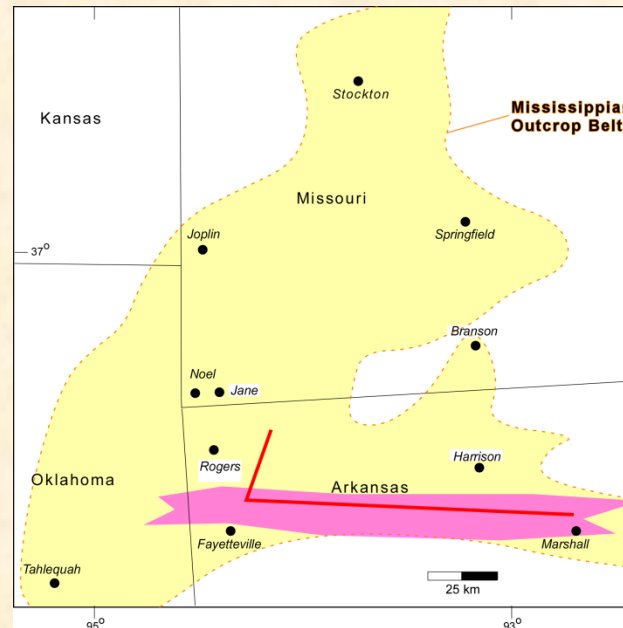
Beau Morris for scale, circled in red.

East (A')



Beau Morris for scale, circled in red.

Angular
unconformity
within the deep-
water Reeds
Spring; a
breached East-
West trending
anticline.



Evidence of
subaerial
exposure in deep-
water deposits in
our southernmost
Reeds Spring;
most basinward
exposures.

WHAT IS THE MECHANISM THAN CAN ACCOUNT FOR THE ANOMALOUS BELTS OF STRATIGRAPHIC CHANGE?

I.E.

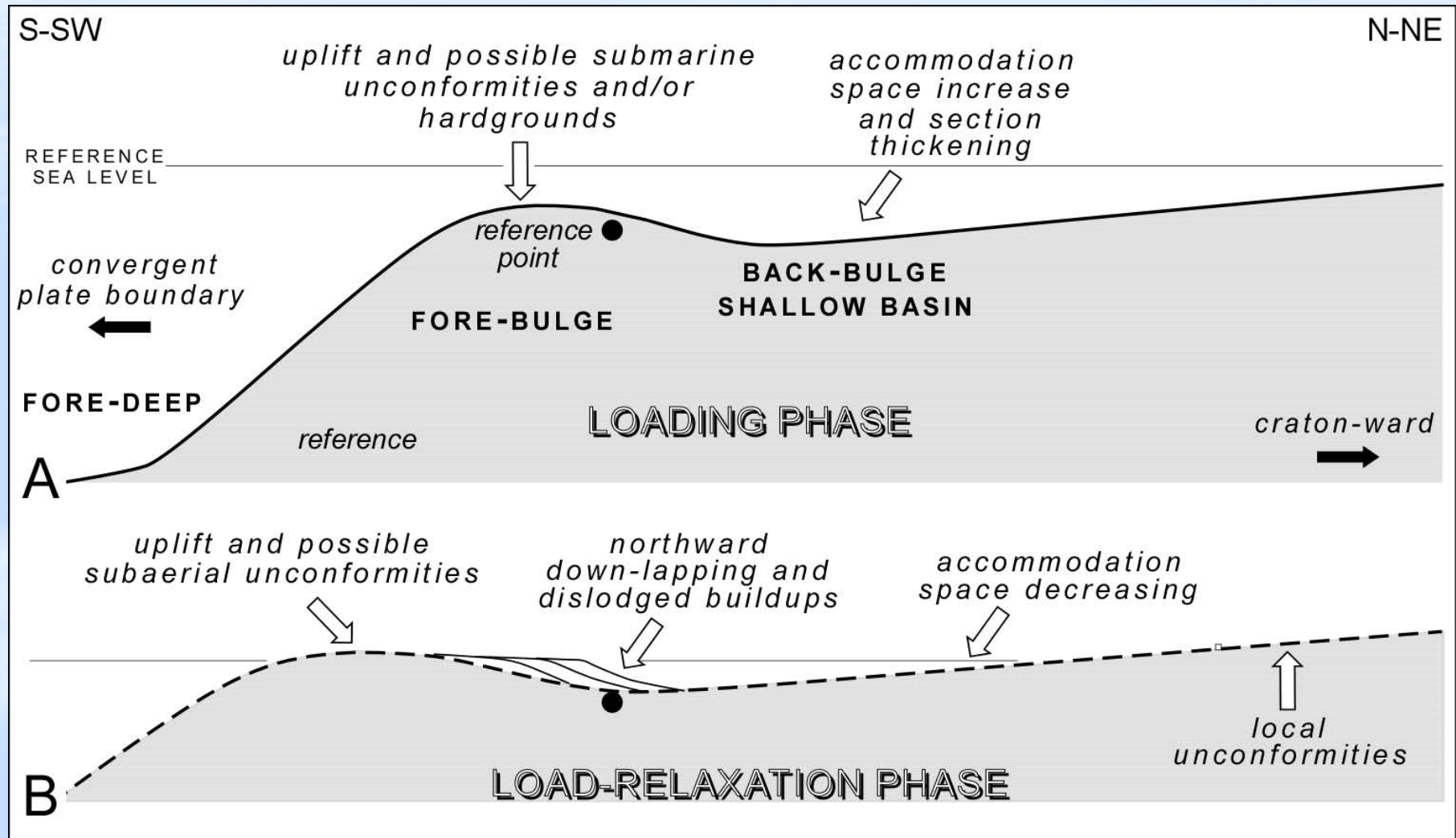
Relative Facies Changes in contrast to, and reversal of regional paleodip?

Anomalous Thickening & Northward Progradation (& allochthonous buildups)?

Multiple Unconformities, Syndepositional Folding & Resultant Non-Systematic Lithostratigraphic Architecture?

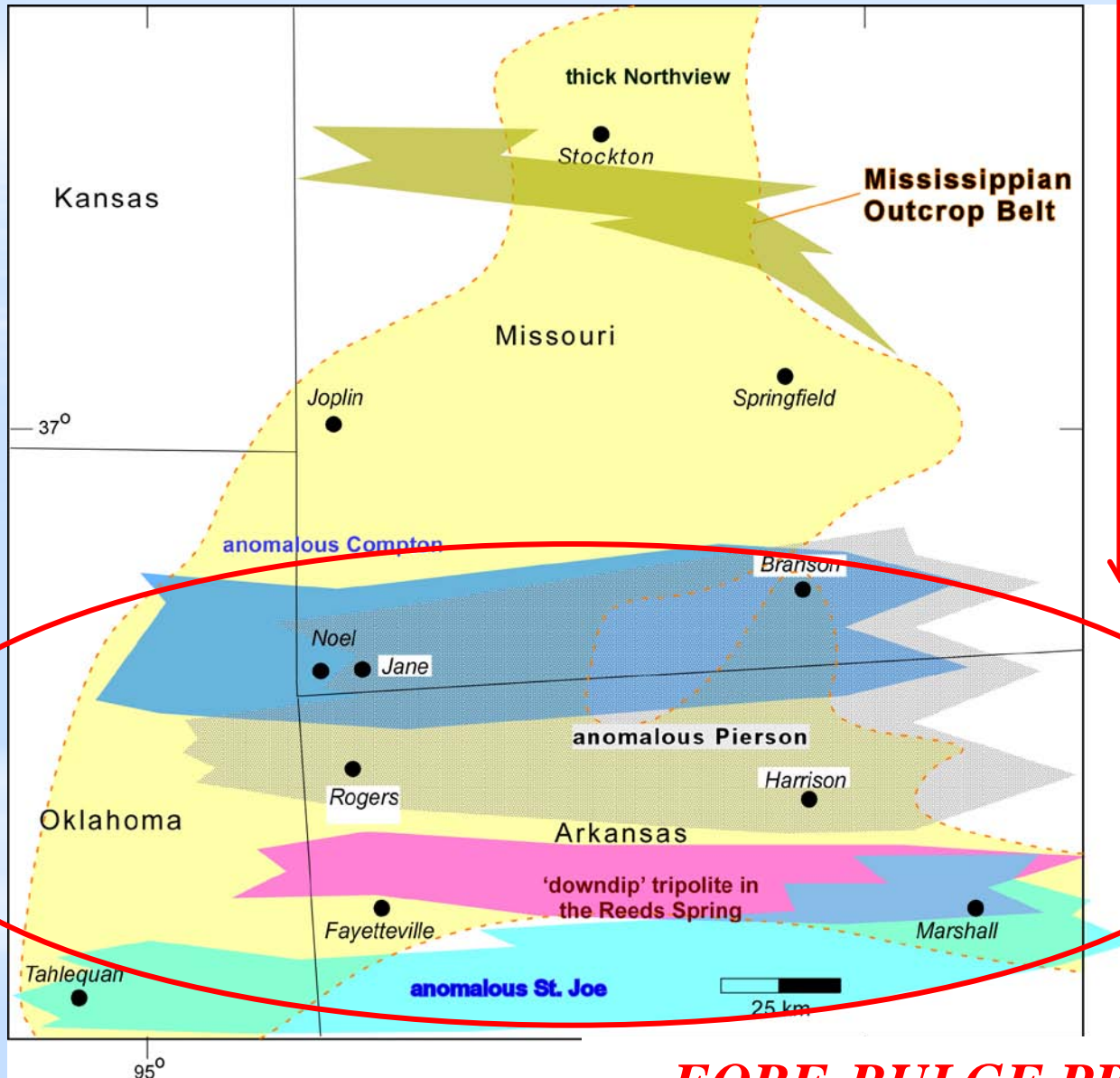


SYNDEPOSITIONAL TECTONISM RELATED TO FORE-BULGE VACILLATIONS IN RESPONSE TO OUACHITA OROGENY

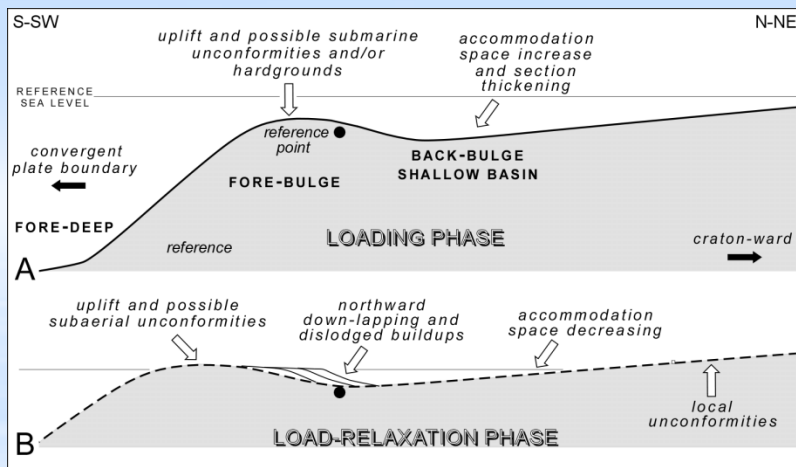


***Fore-bulge migration and back-stepping in response to
N-S compression associated with Ouachita Orogeny***

***BACK-BULGE
BASIN & SHELF***



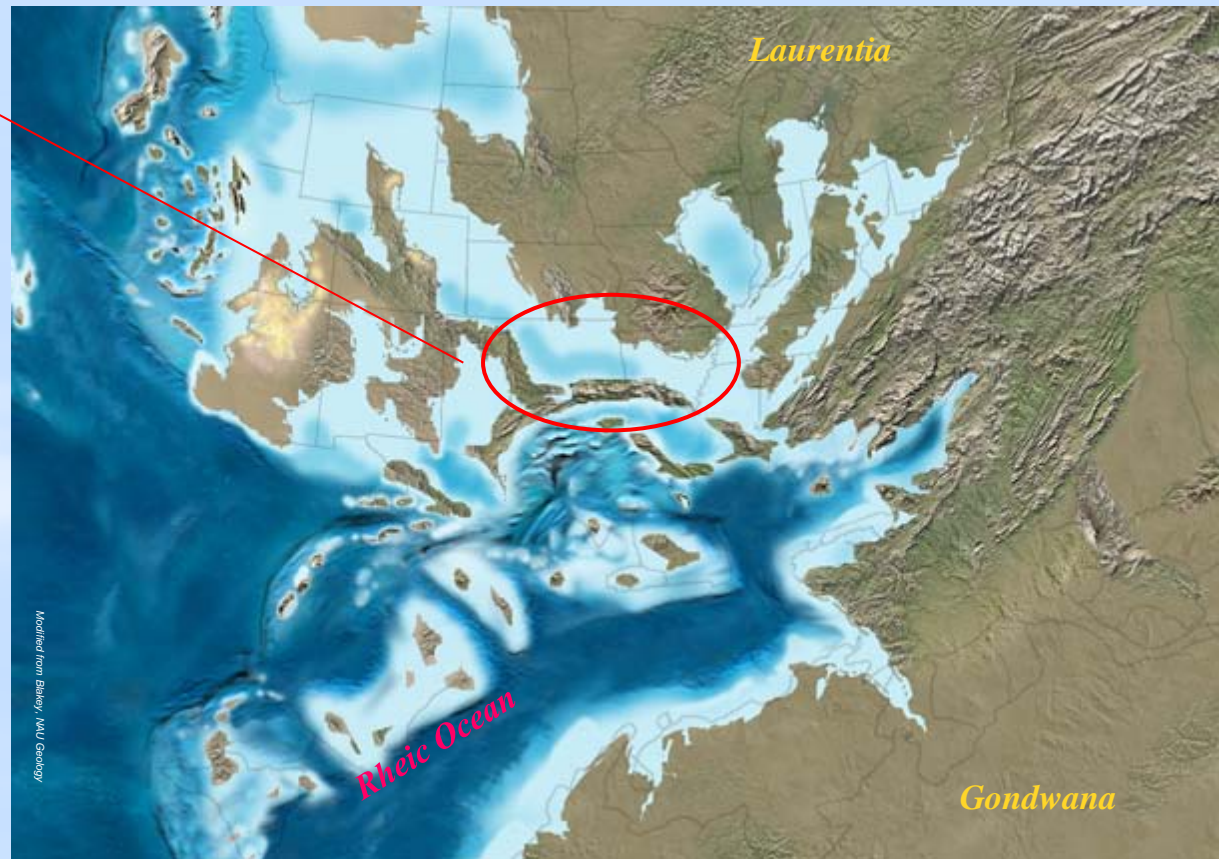
FORE-BULGE PROVINCE



The anomalous lithostratigraphy and syndepositional folding must incorporate the impending Ouachita island arc-continent collision as seen below.

More plausible of a scenario during KH – OS time

The fore-bulge belt in the outcrop area was ~100 km inboard of the Ouachita belt however, albeit mild, the tectonics were significant on deposition.



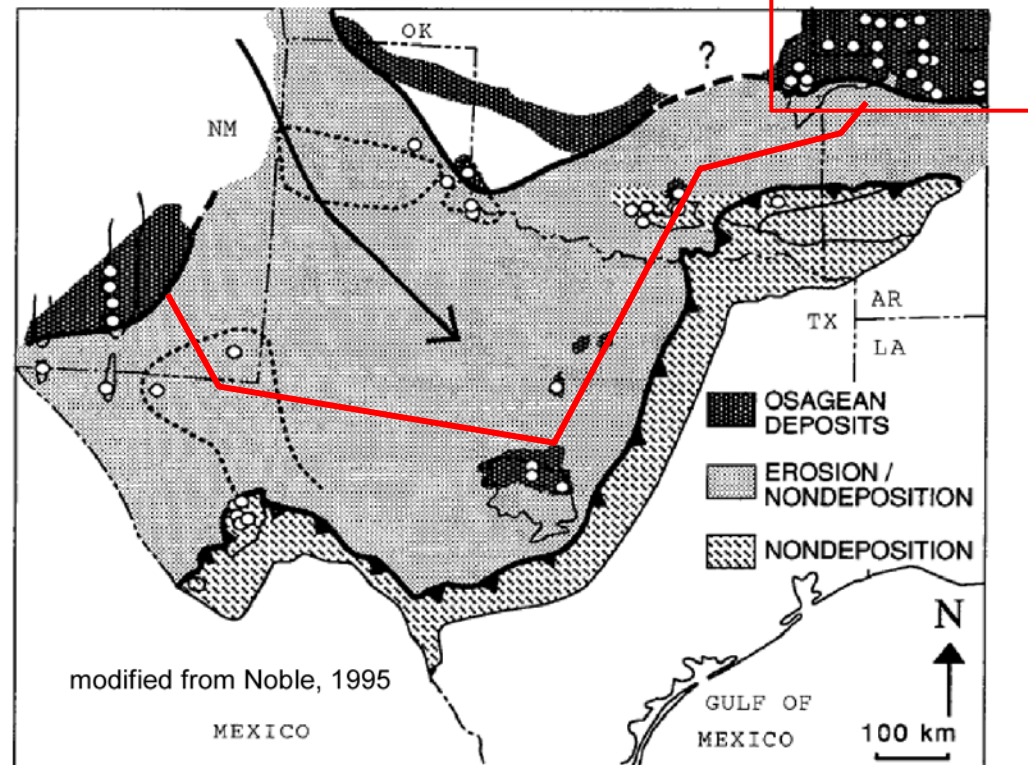
REWRITING BASIN HISTORY

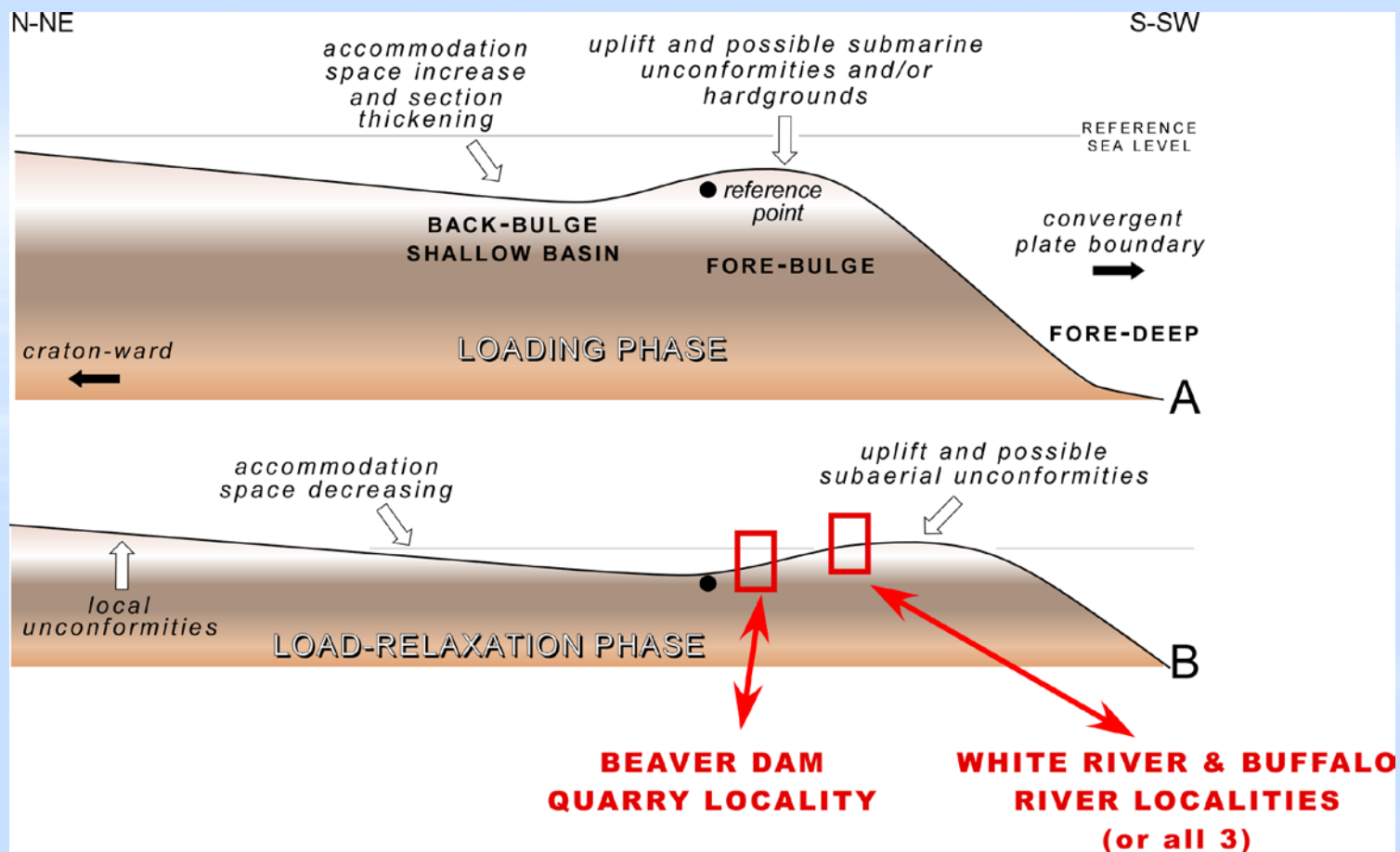
		OUACHITA FORELAND						OUACHITA BELT		
data type		Hueco & Franklin mtns.	Permian Basin	Palo Duro & Hardeman basins	Llano uplift	Lawrence uplift	Arbuckle Anticline	Ouachita frontal zone	Ouachita central zone	Marathon uplift
Period/Epoch		conodonts ¹	brachiopods ² conodonts ³	conodonts ⁴	conodonts ⁵	conodonts ⁶	Radiolaria ⁷ conodonts	Radiolaria ⁸ conodonts ⁹	conodonts ¹⁰	Radiolaria ¹¹ conodonts ¹²
Mississippian	Chesterian	Rancheria-Las Cruces	"Mississippian limestone"	"Chester"	Barnett Sh.		Caney Shale		Stanley Shale	Tesnus
	Meramecian		?	"Meramec" *	*		Sycamore *	*	*	*
	Osagean		?	"Osage" *			Woodford *	Woodford *	?	
	Kinderhookian		?		Chappel Ls.	Welden Ls.	Woodford *	Woodford *	Arkansas m	
Devonian		Percha *	Woodford *		Houy	Woodford *	Woodford *	Woodford *	Novaculite I	Caballos Novaculite
Silurian										
Ordovician				Ellenburger	Ellenburger					

modified from Noble, 1995

OUR AREA OF STUDY

The relationships seen at the outcrop, supported by conodont biostratigraphy, suggest that fore-bulge tectonics may be responsible for the absence of the KH-OS across the mid-continent, particularly adjacent to the Ouachita & Marathon thrust belt.





*-Based on the data and our interpretations, the Compton, Northview and Pierson were primarily back-bulge basin deposits where depositional architecture was aggradational.

*-Due to the strongly progradational nature of the (Burl-Keok) Reeds Spring, this formation extended beyond the fore-bulge area however, was effected by thrust loading and fore-bulge vacillations as is evident by the unconformities within the section.

SIGNIFICANCE IN EXPLORATION

Is there expression of this early, syndepositional, Mississippian fore-bulge tectonics in the subsurface? YES!

Both directly and indirectly.

- Understanding and recognition of paleo-geography is paramount in recognizing structurally positive areas and basins.
- Knowledge of syndepositional tectonics will aid in the recognition of correct correlations; to make that leap of faith.
- Understanding of the tectonics should aid in recognizing unconformities and associated porosity zones and traps.

These relationships are recognized in both subsurface and on seismic.



Conclusions:

- Ouachita orogeny began in early Mississippian, Kinderhookian time – much earlier than previously postulated.
- Regional tectonics overprinted the “normal” stratigraphy of the KH – OS units and negates a sequence-stratigraphic interpretation. Instead, the lithostratigraphic architecture of the rocks is solely tectonic origin as opposed to eustatic.
- Facies belts and lithostratigraphic changes coincide with syndepositional tectonics.
- Such changes can and should be recognized in the subsurface based on the outcrop model.

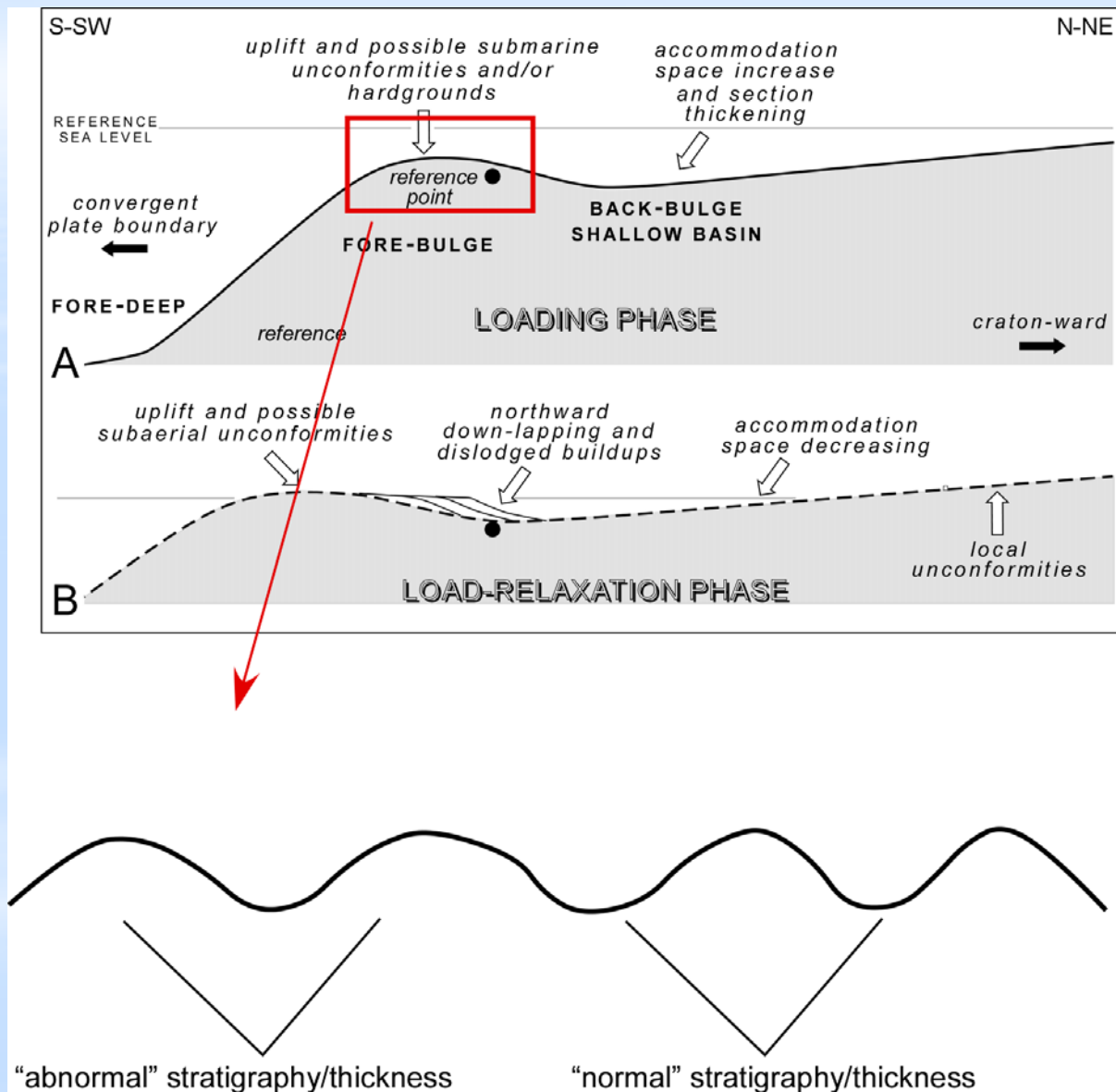


Thank You



WOOLSEY OPERATING COMPANY, LLC





WMO
STOP 44
STOCKFORD LINE
+ 0

<174.45FT>

WMO
STOP 45
TURNER SECTION
+ 0

<183.66FT>

WMO
STOP 46
CHESAPEAKE SECTION
+ 0

<180.214FT>

WMO
STOP 47
CHESAPEAKE MOORE
+ 0

<205.034FT>

WMO
STOP 48
BOKARIN FINDER
+ 0

<154.855FT>

WMO
STOP 49
GREATER 10
+ 0

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WMO
STOP 49
LOWER SECTION
+ 0

<142.831FT>

WMO
STOP 15A
SALOME SOUTH
+ 0

<120.002FT>

WMO
STOP 15
NO HEADFOLLOW DET.
+ 0

PIERSON LIMESTONE COMPLEXITY: A BACK-BULGE BASIN DEPOSIT??

***THE PIERSON THEN EVOLVES INTO A DISTALLY-STEEPENED RAMP WHERE SHALLOW-WATER SANDS PROGRADE INTO DEEPER WATER MUDSTONES, IN BOTH A SOUTHWARD & NORTHWARD(!) DIRECTION DURING upper multistriatus time to bulbosus time.**

***UPPER OSAGE (bulbosus & up) TIME THE BURLINGTON-KEOKUK & REEDS SPRING PROPER, ARE STRONGLY PROGRADATIONAL IN A SOUTHWARD DIRECTION.**

***LOW RELIEF, SHALLOW WATER RAMP WHERE PLATFORM ARCHITECTURE WAS AGGRADATIONAL DURING carina & lower multistriatus time.**



- Subsurface lithostratigraphy mimics and **DIRECTLY CORRELATES** to exposures in the Mississippian outcrop belt of NE Oklahoma, SW Missouri and NE Arkansas
- Outcrop-based models aid and are integral in interpreting depositional systems in the subsurface
- Depositional motifs and platform architecture are aggradational during Kinderhookian time and evolve to progradational during Osagean across the midcontinent, from the outcrop belt to the subsurface of Kansas and Oklahoma

