## AVClimate Sensitivity during the Phanerozoic: Lessons for the Future\*

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#### **Abstract**

Understanding the link between the level of atmospheric  $CO_2$  and global surface temperature is profoundly important. Almost all estimates of climate sensitivity (typically defined as the warming caused by a  $CO_2$  doubling) have come from studying records spanning the last ~20 kyr Most of these studies find a modal climate sensitivity of ~3 °C, with a long probability tail at the high end. These studies have been vital for informing climate change issues, but a limitation is that they are calibrated to a present-day or cooler-than-present-day Earth; investigation of deep-time records (older than 2 Myr ago) are required to understand the dynamics of a globally warm Earth.

Ancient levels of CO<sub>2</sub> can be estimated from long-term carbon cycle models, which quantitatively track the major sources and sinks of atmospheric CO<sub>2</sub> over these timescales, or by proxy indicators. A synthesis of CO<sub>2</sub> estimates from carbon cycle models and from proxies shows a strong, first-order fit between CO<sub>2</sub> and geologic indicators of temperature: continental ice sheets are common when CO<sub>2</sub> drops below 500 ppm and absent when CO<sub>2</sub> exceeds 1000 ppm.

One parameter in most long-term carbon cycle models is climate sensitivity because the weathering of Ca and Mg-rich silicate rocks, which serves as a long-term sink for  $CO_2$ , is sensitive to temperature. Thus, it is possible to estimate long-term climate sensitivity by adjusting the climate sensitivity parameter in the long-term carbon cycle models until the model estimates of  $CO_2$  best match the independent proxy estimates of  $CO_2$ . Over the past 420 Myr, the modal climate sensitivity is ~3 °C; a sensitivity of < 1.5 °C is highly improbable, while sensitivities of 6+ °C cannot be excluded. Deep-time climate sensitivity thus matches the present-day, despite the two approaches capturing fundamentally different carbon cycle processes; a climate sensitivity of around 3 °C appears to be a robust feature of the Earth system, independent of temporal scaling. The geologic record generally supports a positive link between  $CO_2$  and temperature, and we should expect a climate sensitivity of 3 °C or more in the near future.

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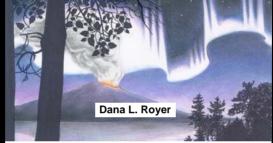
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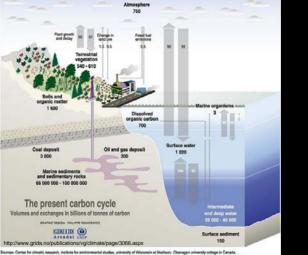
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# Climate sensitivity during the Phanerozoic: lessons for the future





#### Short-term carbon cycle (<10,000 yr)

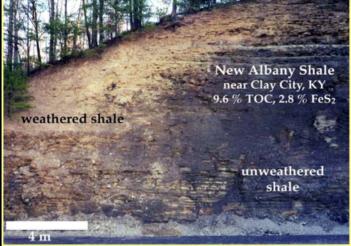


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#### Long-term carbon cycle (>100,000 yr) Volcanic Conversion of CO2 to dissolved HCO3- by CO2 Ca-Mg silicate weathering CO2 from sedimentary organic C weathering CaCO<sub>3</sub> CO2 and CH4 from Burial Metamorphism & Org C Deep Diagenesis Burial

Subduction of CaCO<sub>3</sub> & Organic C

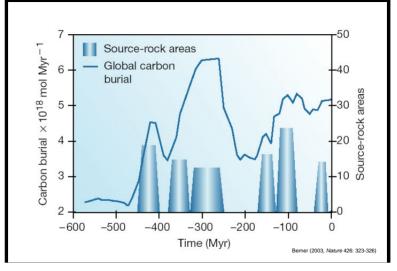
#### Weathering of organic carbon



#### Long-term carbon cycle (>100,000 yr) Volcanic Conversion of CO2 to dissolved HCO3- by CO2 Ca-Mg silicate weathering CO2 from sedimentary organic C weathering CaCO<sub>3</sub> CO2 and CH4 from Burial Metamorphism & Org C Deep Diagenesis Burial

Subduction of CaCO<sub>3</sub> & Organic C

## **Burial of organic carbon**



#### Phanerozoic record of CO<sub>2</sub> Time (Ma) 500 400 300 200 100 OSD Carb P K Tr Pg g 4000 Paleosols (n=206) • Stomata (n=171) △ Phytoplankton (n=184) △ Boron (n=35) 3000 O Liverworts (n=13) CO<sub>2</sub> (ppm) ▲ Sodium carbonates (n=2) 2000 1000 0 Sampling frequency n = 611 80 40 updated from Royer (2006, Geochim. Cosmochim. Acta 0

200

100

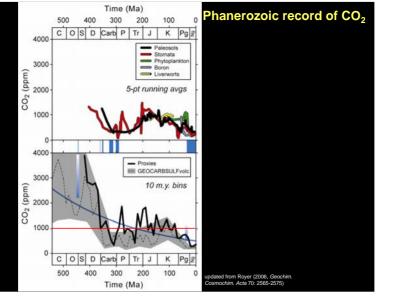
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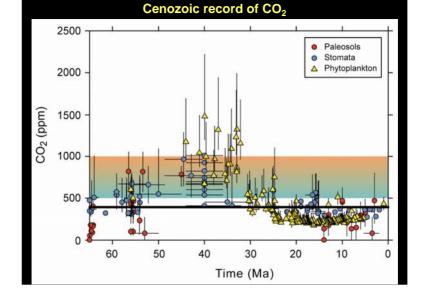
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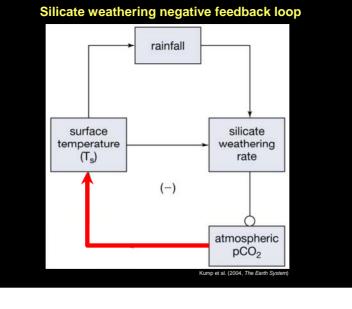
400

300 Time (Ma)

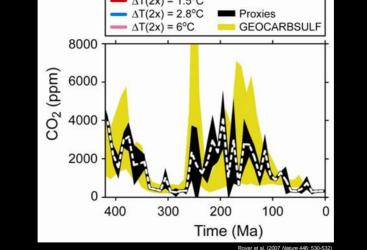
500



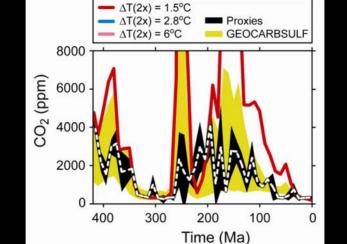




### Geologic constraints on climate sensitivity $\Delta T(2x) = 1.5^{\circ}C$

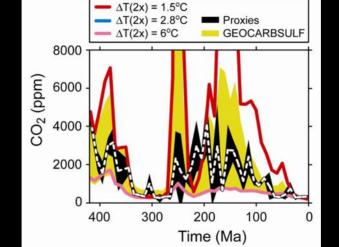


## Geologic constraints on climate sensitivity



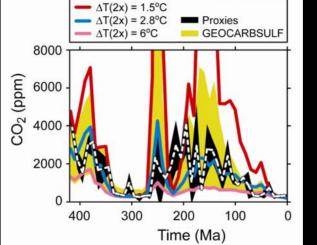
Royer et al. (2007 Nature 446: 530-532)

## Geologic constraints on climate sensitivity

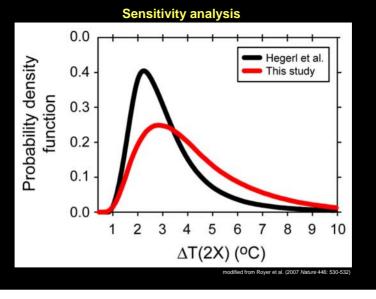


Royer et al. (2007 Nature 446: 530-532)

## Geologic constraints on climate sensitivity



Royer et al. (2007 Nature 446: 530-532)



**Presenter's Notes:** Residual variance: thick blue line = standard run; thin black line = varying the four factors simultaneously Where to next—time slices through time (do icy times have a higher DT(2x)?)

#### **Summary**

- There is a strong CO<sub>2</sub>-temperature coupling for much of the Phanerozoic. There is a threshold for nucleating ice sheets at the equivalent radiative forcing of ~500 ppmv (adjusted for changing luminosity through time). Above this CO<sub>2</sub> threshold, large continental ice sheets are typically absent.
- A model-data comparison for paleo-CO<sub>2</sub> indicates an average climate sensitivity of ~3 °C throughout the Phanerozoic; this calculated sensitivity is similar to calculations for the present-day, implying that an ~3 °C sensitivity is a robust feature of the Earth system.

#### Phanerozoic $\delta^{18}$ O shallow marine carbonate record 0 Marine carbonate 818O (% PDB) -8 10 m.y. bins ■:1 std. dev. linear fit: $r^2 = 0.78$ -12 600 Sampling frequency n = 4824low latitude sites 400 200 600 500 400 300 200 100

Time (Ma)

