

AAPG HEDBERG RESEARCH CONFERENCE
“NATURAL GAS GEOCHEMISTRY: RECENT DEVELOPMENTS, APPLICATIONS, AND
TECHNOLOGIES”
MAY 9-12, 2011 — BEIJING, CHINA

Geochemical Characteristics of the Abiogenic Alkane Gases in the Songliao Basin, China

Jinxing Dai^{a,*}, Yunyan Ni^a, Jian Li^b, Xia Luo^a, Guoyi Hu^a, Xiaoqi Wu^a, Shipeng Huang^a,
Fengrong Liao^a

^aResearch Institute of Petroleum Exploration and Development, PetroChina, Beijing, China

^bResearch Institute of Petroleum Exploration and Development-Langfang Branch, PetroChina, Langfang, Hebei,
China

Although it is found that in many areas there may be alkane gases with abiogenic origin recently, e.g., hydrothermal fluids from the midocean ridges, seepages from the ultramafic rocks, and fluids from the Precambrian shields, there are few reports about commercial abiogenic alkane gas reservoirs. Recently a series of abiogenic alkane gas fields (reservoirs) have been found in the Songliao Basin, Northeast China, which are located east to the paleo-Central Uplift in the eastern Songliao Basin, e.g., Xujiaweizi Fault Depression east to the Daqing Oil Field, Changling Fault Depression in the southern basin and Chaghanhua Fault Depression. There are five gas fields with gas reserves over $50 \times 10^8 \text{ m}^3$ including Xingcheng, Shengping, Changde, and Changling-Songnan gas fields. Among them, Xingcheng gas field has been the largest one. Reservoirs of these gas fields mainly consist of the volcanic rocks from the Lower Cretaceous Yingcheng Formation and their distribution is closely related to large faults.

Xingcheng gas field is located in the Xujiaweizi Fault Depression near the Xuzhong Fault (Figure 1). It is a volcanic gas reservoir, with gas-bearing area of 62.6 km^2 and recoverable gas reserves of $24.7 \times 10^9 \text{ m}^3$. Its reservoirs are dominated by tuffs and rhyolites from the first and fourth members of Yingcheng Formation, Lower Cretaceous. Alkane gases account for 94-98% (Table 1) and demonstrate a carbon isotopic reversal trend among C_1 - C_4 n-alkanes ($\delta^{13}\text{C}_1 > \delta^{13}\text{C}_2 > \delta^{13}\text{C}_3 > \delta^{13}\text{C}_4$).

Both Changde and Shengping gas fields are located in the Xujiaweizi Fault Depression, with proved geological gas reserves of $182 \times 10^8 \text{ m}^3$. In contrast to the Xingcheng, Shengping and Changling gas fields where reservoirs mainly consist of volcanic rocks from Yingcheng Formation, reservoirs of the Changde gas field are dominated by fine sandstones of the Lower Cretaceous Dengloulou Formation. Changde gas field is related to the Xuxi Fault (Figure 1), and its gases belong to dry gas, with CH_4 ranging from 92.06 to 95.11% and C_{2+} from 0.89 to 1.87%. In a similar fashion to the Xingcheng gas field, natural gases in the Changde gas field also show a carbon isotopic reversal trend among C_1 - C_4 n-alkanes ($\delta^{13}\text{C}_1 > \delta^{13}\text{C}_2 > \delta^{13}\text{C}_3 > \delta^{13}\text{C}_4$) (Table 1). Shengping gas field lies north to the Xingcheng gas field, and is also related to the Xuzhong Fault. Its reservoirs mainly consist of rhyolites from the Yingcheng Formation, and its gases belong to dry gas with CH_4 ranging from 92.0 to 94.6 % and C_{2+} from 1.4 to 1.9 %. Natural gases in the

Shengping gas field also show a carbon isotopic reversal trend among C₁-C₄ n-alkanes ($\delta^{13}C_1 > \delta^{13}C_2 > \delta^{13}C_3 > \delta^{13}C_4$) (Table 1).

Changling-Songnan gas field is located in the Chaghanhua Fault Depression in the southern Central Fault Depression, Songliao Basin. Proved geological gas reserves are more than $1000 \times 10^8 \text{ m}^3$, and CO₂ are expected to be $257 \times 10^8 \text{ m}^3$. Reservoirs mainly consist of tuffs and rhyolites of Yingcheng and Denglouku formations. Central Fault Depression is characterized by strong tectonic activities, multistage volcanic activities and widespread large faults. In contrast to the low content of CO₂ in the Xingcheng, Shengping and Changde gas field, CO₂ ranges from 10.16% to 98.70% in the Changling-Songnan gas field and shows relatively heavy $\delta^{13}C$ reflecting the abiogenic mantle origin. Similarly, a carbon isotopic reversal trend among C₁-C₄ n-alkanes ($\delta^{13}C_1 > \delta^{13}C_2 > \delta^{13}C_3 > \delta^{13}C_4$) has also been found in this field (Table 1).

Hence, the above four abiogenic alkane gas fields are all characterized by a carbon isotopic reversal trend among C₁-C₄ n-alkanes ($\delta^{13}C_1 > \delta^{13}C_2 > \delta^{13}C_3 > \delta^{13}C_4$) and high R/Ra values. As demonstrated by previous studies, abiogenic gases from the fluid inclusions associated with the Khibiny massif on the Kola peninsula, Russia and Lost City hydrothermal field at 30°N on the Mid-Atlantic Ridge all show a significant depletion in ¹³C for C₂-C₄ with respect to C₁. The carbon isotopic reversal trend found in these four gas fields also reflects an abiogenic origin. In addition, the associated helium has R/Ra values varying from 0.6 to 3.9 (Table 1) implying the addition of considerable mantle origin ³He, which indicate that alkane gases in these gas fields are possibly related to abiogenic origin.

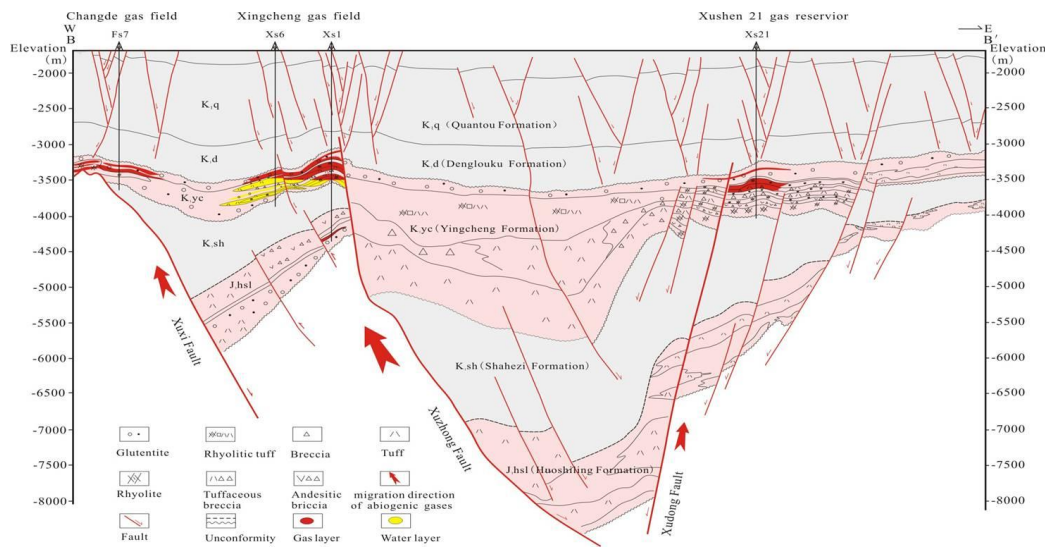


Figure 1 Structural section (B-B') of Xujiaweizi Fault Depression showing gas reservoirs.

Table 1 Main gas composition and isotopic ratios of natural gases from the Xingcheng, Shengping, Changde and Changling-Songnan gas fields

Field	Well	Strata	Depth (m)	Natural gas component (%)						$\delta^{13}\text{C}$ (‰), PDB					R/Ra
				CH ₄	C ₂ H ₆	C ₃ H ₈	C ₄ H ₁₀	CO ₂	N ₂	CH ₄	C ₂ H ₆	C ₃ H ₈	C ₄ H ₁₀	CO ₂	
Xingcheng	Xs1	K ₁ yc	3440-3750	92.62	2.62	0.78	0.26	2.25	1.43	-27.4	-32.3	-33.9	-34.7	-5.0	
	Xs1	J ₃ hsl	4446-4466	92.66	2.22	0.54	0.10	1.14	1.42	-29.7	-32.9	-34.3	-35.0	-5.9	1.1
	Xs1-1	K ₁ yc	3416-3424	93.94	2.13	0.40	0.08	1.17	1.86	-28.9	-32.6	-33.3	-34.1	-5.5	1.1
	Xs1-201	K ₁ yc	3328-3358	94.56	2.13	0.36	0.07	1.53	1.30	-28.6	-32.2	-34.0	-35.1	-7.6	1.2
	Xs1-4	K ₁ yc	3530-3540	94.46	2.05	0.48	0.23	1.38	1.38	-27.4	-31.8	-33.7	-34.4	-6.8	0.8
	Xs5	K ₁ yc	3611-3629	91.04	2.33	0.62	0.15	4.09	1.73	-28.6	-33.9	-34.4	-35.2	-5.1	0.9
	Xs6	K ₁ sh	3629-3637	95.77	2.39	0.49	0.11	0.28	0.86	-28.3	-33.2	-34.3	-34.6	-13.0	1.0
	Xs6-1	K ₁ yc	3613-3640	94.38	2.45	0.64	0.27	0.32	1.86	-26.9	-33.8	-34.2	-34.6	-8.2	1.2
	Xs6-102	K ₁ yc	3557-3580	94.46	2.23	0.22	0.11	0.20	0.71	-27.5	-29.3	-31.4	-31.4		
	Xs6-104	K ₁ yc	3505-3515	95.88	2.20	0.24	0.07	0.30	1.27	-27.9	-31.1	-32.8	-34.9	-15.9	1.3
	Xs6-2	K ₁ yc	3570-3759	95.41	2.25	0.47	0.10	0.21	1.53	-25.9	-32.4	-33.1	-33.7	-11.1	0.9
	Xs6-208	K ₁ yc	3542-3550	95.67	2.24	0.22	0.08	0.32	1.33	-28.3	-31.1	-33.5	-35.1	-14.8	1.2
	Xs603	K ₁ yc	3514-3521	95.48	2.17	0.29	0.08	0.45	1.47	-27.0	-30.4	-32.3	-34.3	-12.3	1.2
Changde	Fs1	K ₁ d		92.06	1.42	0.13	0.02	0	6.35	-18.9	-22.8	-25.3	-27.6	-18.9	
	Fs2	K ₁ d		93.87	0.74	0.11	0.04	0.0003	5.03	-17.4	-22.2	-30.5	-31.4	-16.5	0.6
	Fs5	K ₁ d		95.11	1.54	0.28	0.03	0.46	2.48	-27.1	-28.5	-30.8	-32.2	-16	1.5
Shengping	Shs2	K1d		94.60	1.60	0.30	0.10	0.20	3.30	-27.8	-29.1	-30.6	-30.8		
	Shs2-1	K1yc		92.70	1.50	0.20	0.30	2.60	2.90	-26.8	-29.1	-33.5	-36.5	-14.5	1.8
	Shs2-25	K1yc		92.70	1.40	0.30	0	2.60	2.90	-26.6	-28.8	-32.6	-35.7	-13.2	1.7
	Shsg2	K1yc		92.00	1.40	0.20	0	0.70	3.60	-27.2	-28.1	-32.7	-34.9	-14.8	1.8
Changling No.1	CS1	K1yc	3615	80.46	1.23	0.19	0	10.16	0	-24.2	-26.9	-27.2		-8.2	1.9
	CS1	K1yc	3594	71.40	1.79	0.11	0	22.56	4.14	-23	-26.3	-27.3		-6.8	2.1
	CS1-1	K1yc	3739	75.45	1.91	0.21	0	12.55	5.87	-22.2	-26.9	-27		-7.5	2.3
	CS1-2	K1yc	3697.0-3704	69.44	1.79	0.09	0	21.95	6.73	-24.1	-27.6	-27.2		-8.3	2.1
	CS1-2	K1yc		18.6	0.44			77.80	3.20	-29.9	-23.8	-23.6		-5.8	1.9
	CS2	K1d		4.20	0.40			94.00	1.40	-19.3	-24.6	-24.2		-6.7	
	CS2	K1yc	3791.6-3809	2.25	0.18	0.01	0	96.48	1.08	-17.5	-26.2	-26		-5.0	2.1
	CS6	K1yc		0.40				98.70	0.90	-25.1	-29.6	-30.9		-6.3	3.9
	YS1	K1yc	3544.4-3574	71.72	1.22	0.05	0.03	20.74	5.83	-23.6	-26.4	-26.4		-7.7	
Songnan	YS1	K1yc								-21.2	-26.5	-26.7		-7.9	
	YS101	K1yc	3824.0-3833	71.96	0.84	0	0	21.51	5.59						
	YS101	K1yc	3773.5-3792	69.02		0.05	0	24.75	5.86						

References:

- Abrajano, T.A., Sturchio, N.C., Bohlke, J.K., et al., Methane-hydrogen gas seeps, Zambales ophiolite, Philippines: deep or shallow origin? *Chem Geol*, 1988, 71: 211-222.
- Des Marais, D.J., Donchin, J.H., Nehring, N.L., et al., Molecular carbon isotope evidence for the origin of geothermal hydrocarbon. *Nature*, 1981, 292: 826-828.
- Galimov, E.M. and Petersil, I., On Isotopic Composition of Carbon in Hydrocarbonic Gases Contained in Alkaline Rocks of Khibiny Lovozero and Illimaussak Massifs. *Doklady Akademii Nauk Sssr*, 1967, 176(4): 914-917.
- Jenden, P.D., Hilton, D.R., Kaplan, I.R., et al., Abiogenic hydrocarbons and mantle helium in oil and gas fields in *The Future of Energy Gases*, (ed. D.G Howell), U.S. Geological Survey Professional Paper, 1993, 1570: 31-56.
- Sherwood Lollar, B., Westgate, T.D., Ward, J.A., et al., Abiogenic formation of alkanes in the Earth's crust as a minor source for global hydrocarbon reservoirs, *Nature*, 2002, 416: 522-524.
- Welhan, J.A., Origins of methane in hydrothermal systems. *Chem Geol*, 1988, 71: 183-198.
- Yuen, G., Blair, N., Des Marais, D.J., Chang, S., Carbon isotope composition of low molecular weight hydrocarbons and monocarboxylic acids from Murchison meteorite. *Nature*, 1984, 307: 252-254.